

An adaptive vertical coordinate system for a coastal flows discontinuous Galerkin finite element model

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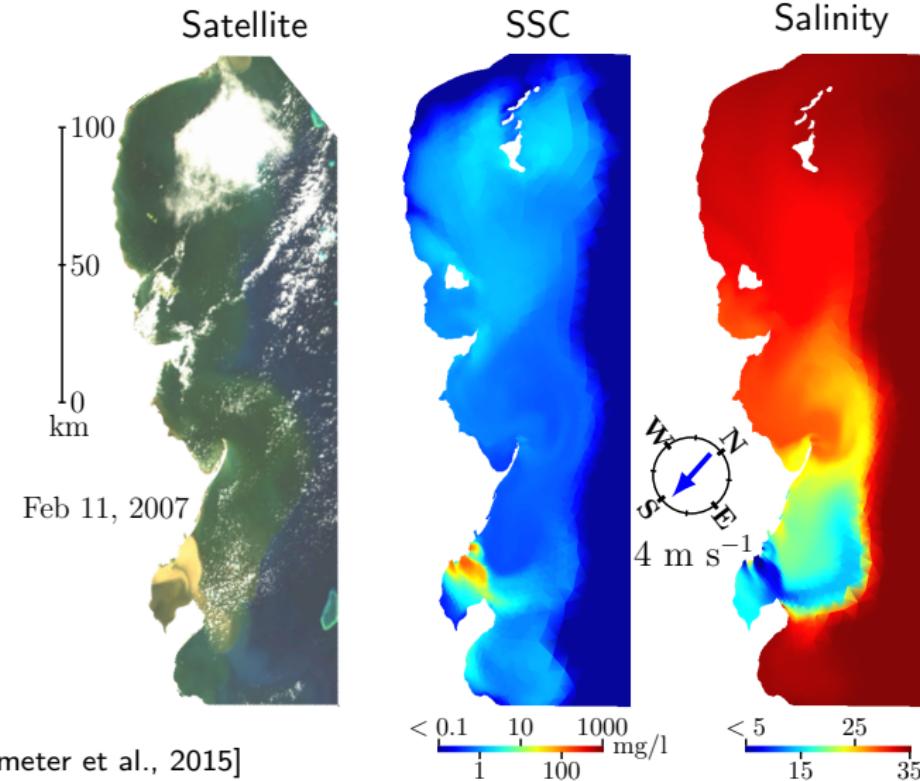
JONSMOD 2016

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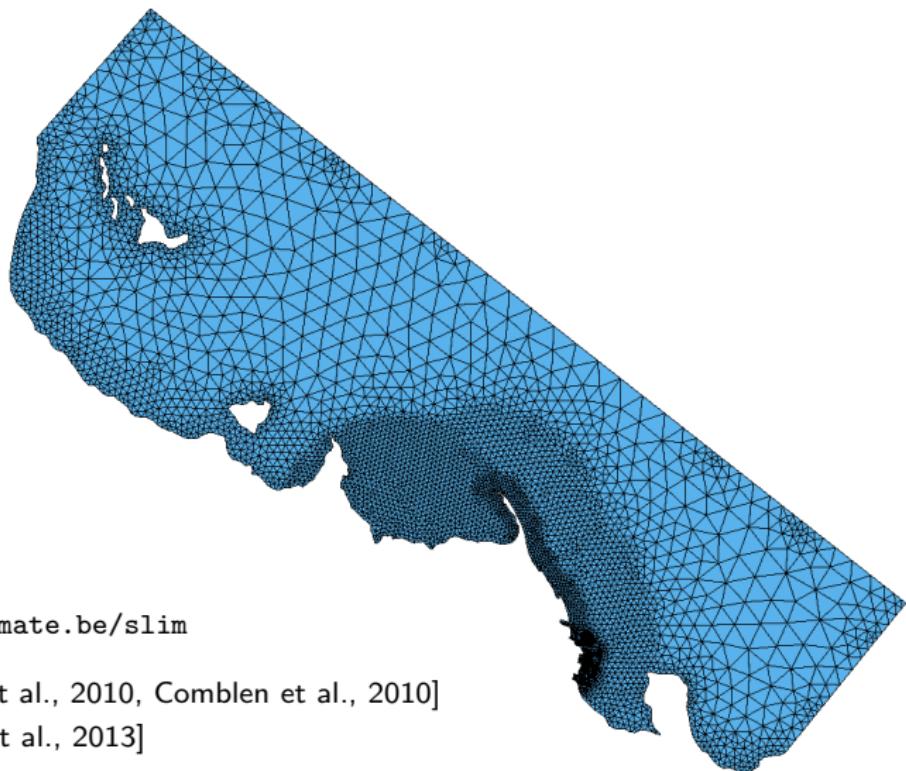
SLIM 3D

- Burdekin river flood (Australia)



SLIM 3D: domain discretization

- unstructured in the horizontal direction



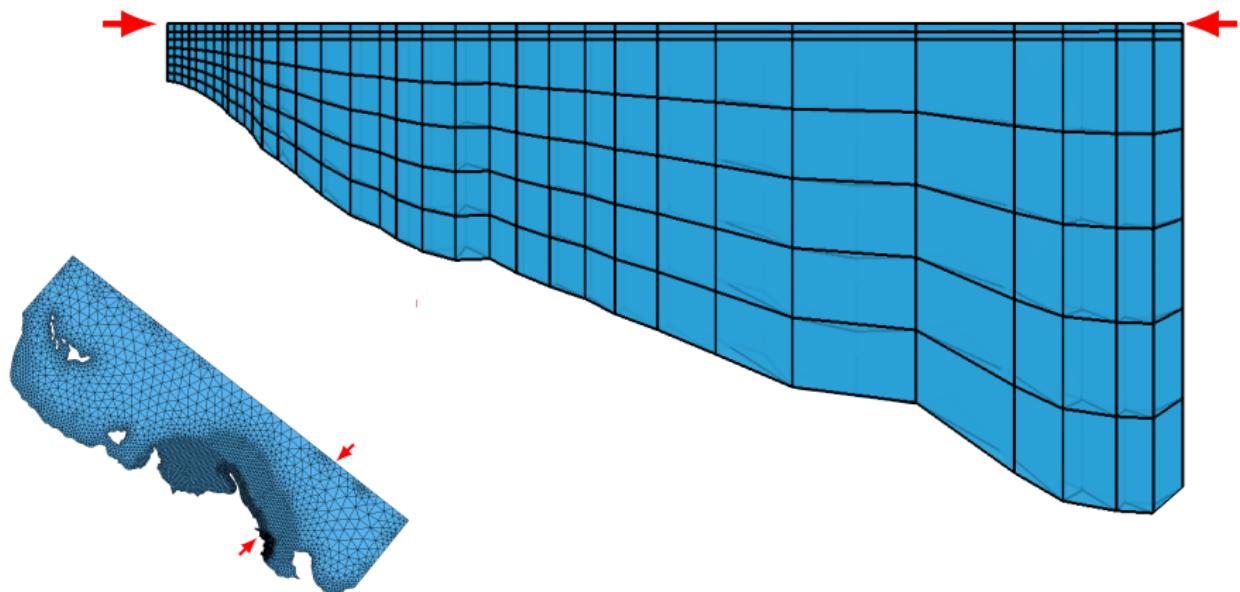
www.climate.be/slim

[Blaise et al., 2010, Comblen et al., 2010]

[Kärnä et al., 2013]

SLIM 3D: domain discretization

- unstructured in the horizontal direction
- structured in the vertical direction: prismatic elements



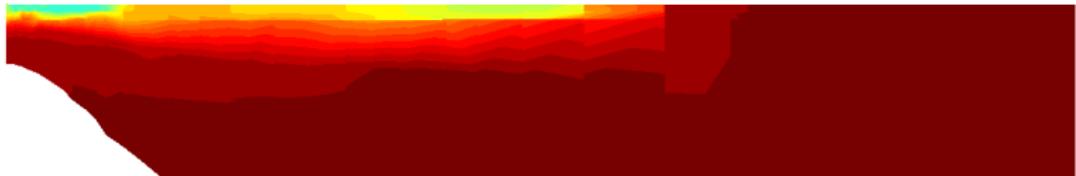
Layers distribution

Salinity Profile

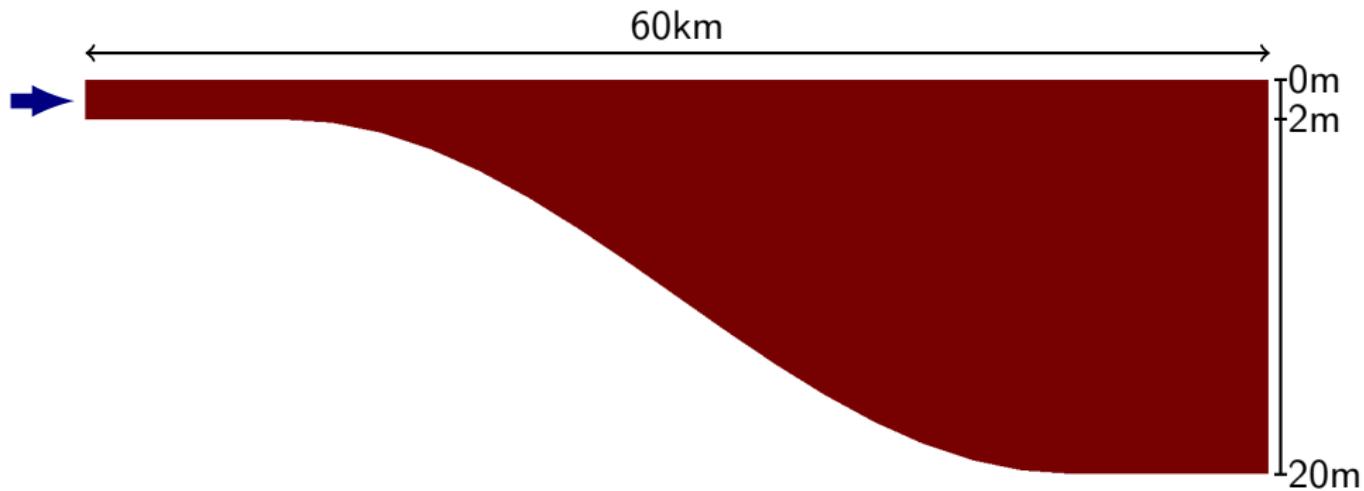
- 5 σ -layers



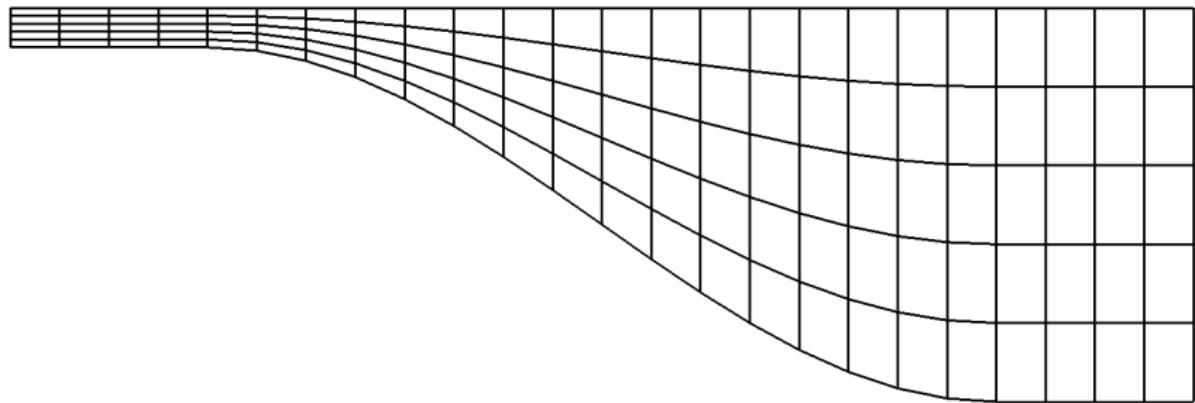
- 2 z-layers, followed by 5 σ -layers



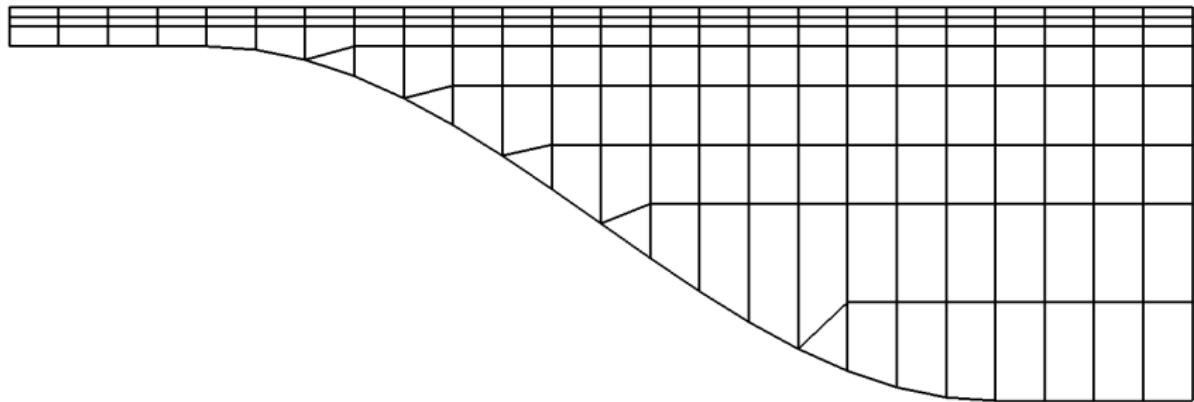
Test case: Idealized river plume



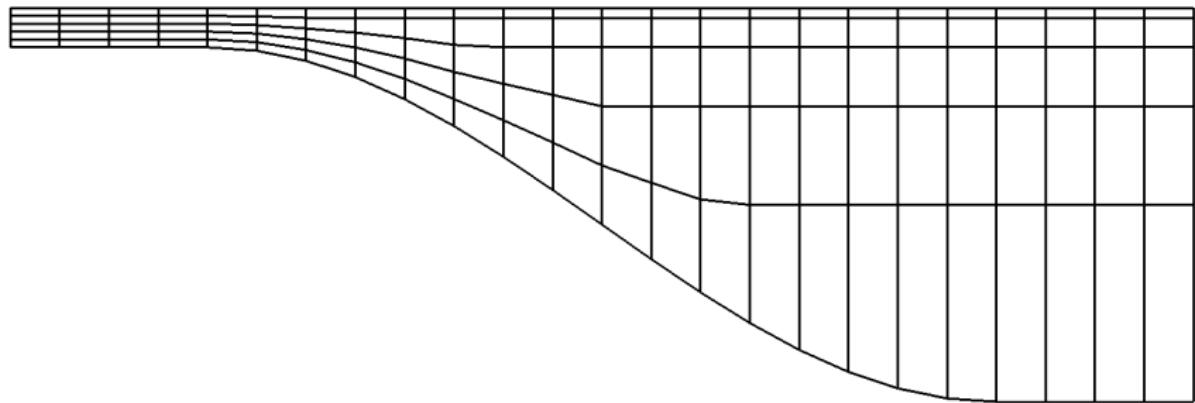
Idealized river plume: σ -layers



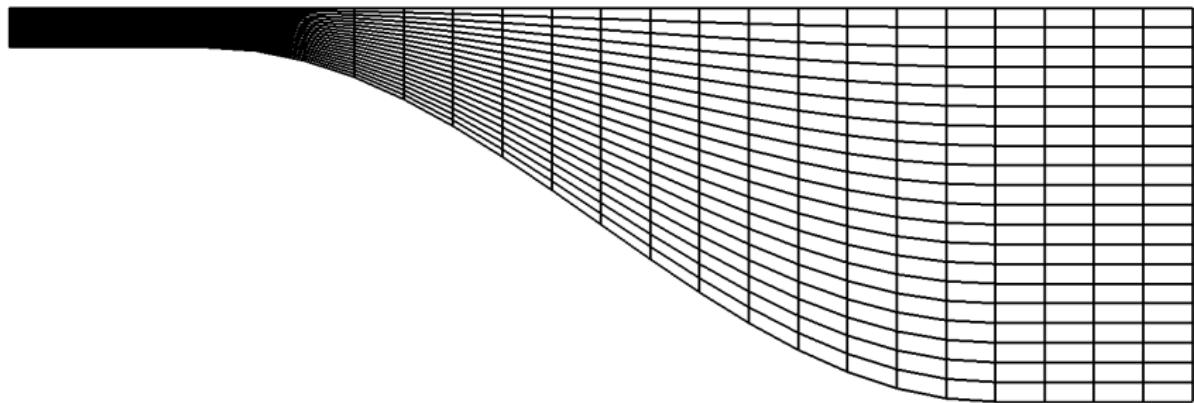
Idealized river plume: z-layers with shaved cells



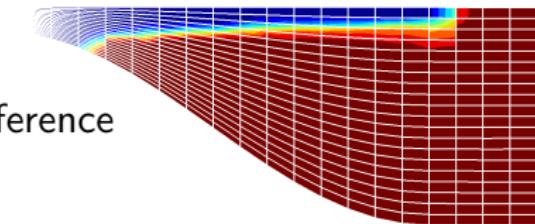
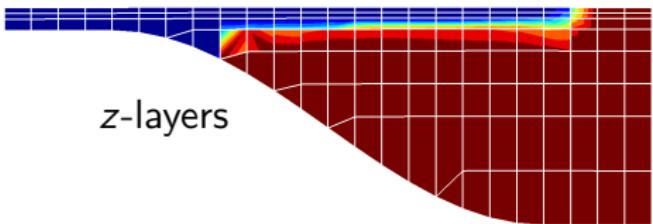
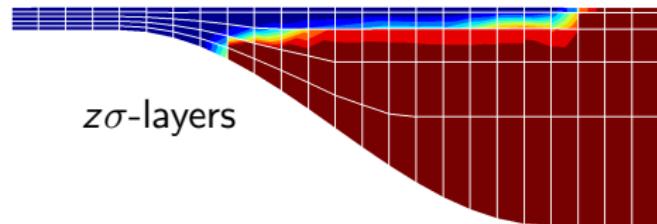
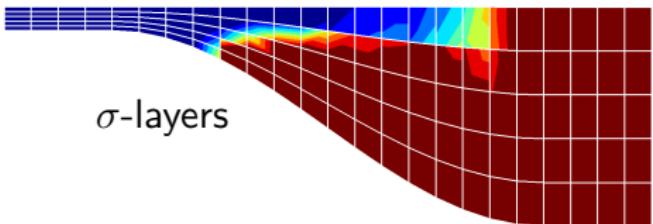
Idealized river plume: $z\sigma$ -layers



Idealized river plume: reference σ -layers

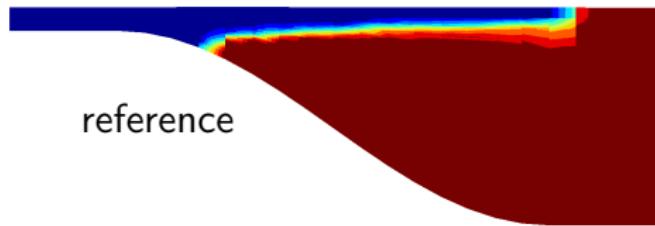
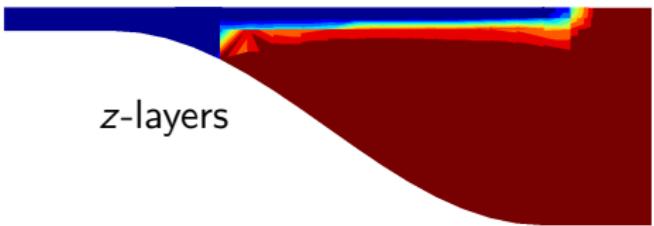
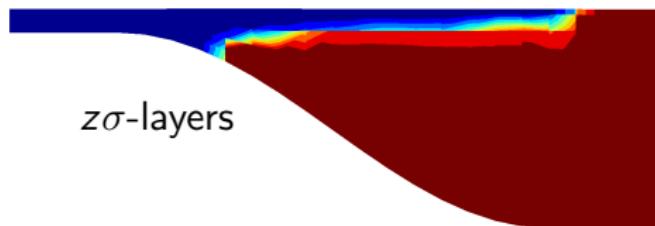
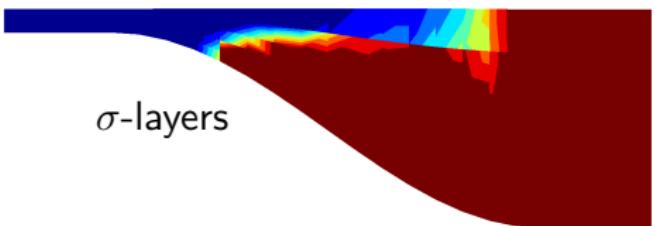


Idealized river plume



After 15 hours

Idealized river plume (mesh hidden)



After 15 hours

Non-uniform adaptive vertical grid

The grid should be adapted considering the:

- Stratification
- Shear
- Distance to the surface
- Distance to the bottom

Two options:

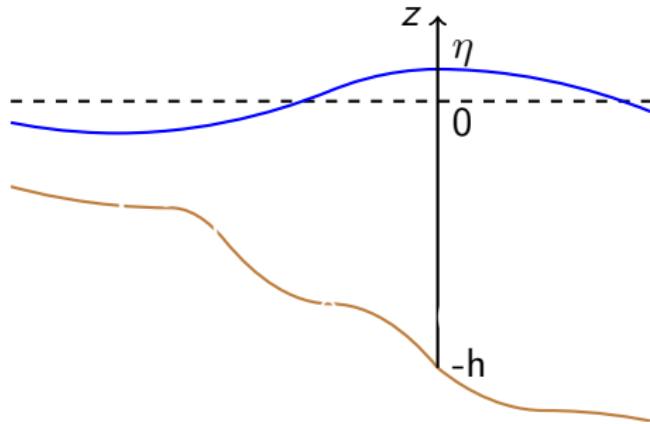
- Define the mesh resolution following an error function
- Diffuse the mesh following an error function

[Burchard and Beckers, 2004, Hofmeister et al., 2010]

Non-uniform adaptive vertical grid:

Moving mesh algorithm:

$$z_{\text{new}} = z + \frac{h+z}{h+\eta_{\text{old}}} (\eta - \eta_{\text{old}}) + dt v_z$$

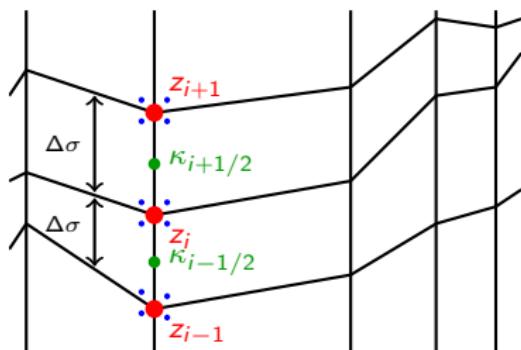


Non-uniform adaptive vertical grid:

Arbitrary moving mesh function:

$$v_z = \begin{cases} \frac{\partial}{\partial \sigma} \left(\kappa_z \frac{\partial z}{\partial \sigma} \right) \\ v_z(-h) = 0, \quad v_z(\eta_{\text{old}}) = 0 \end{cases}$$

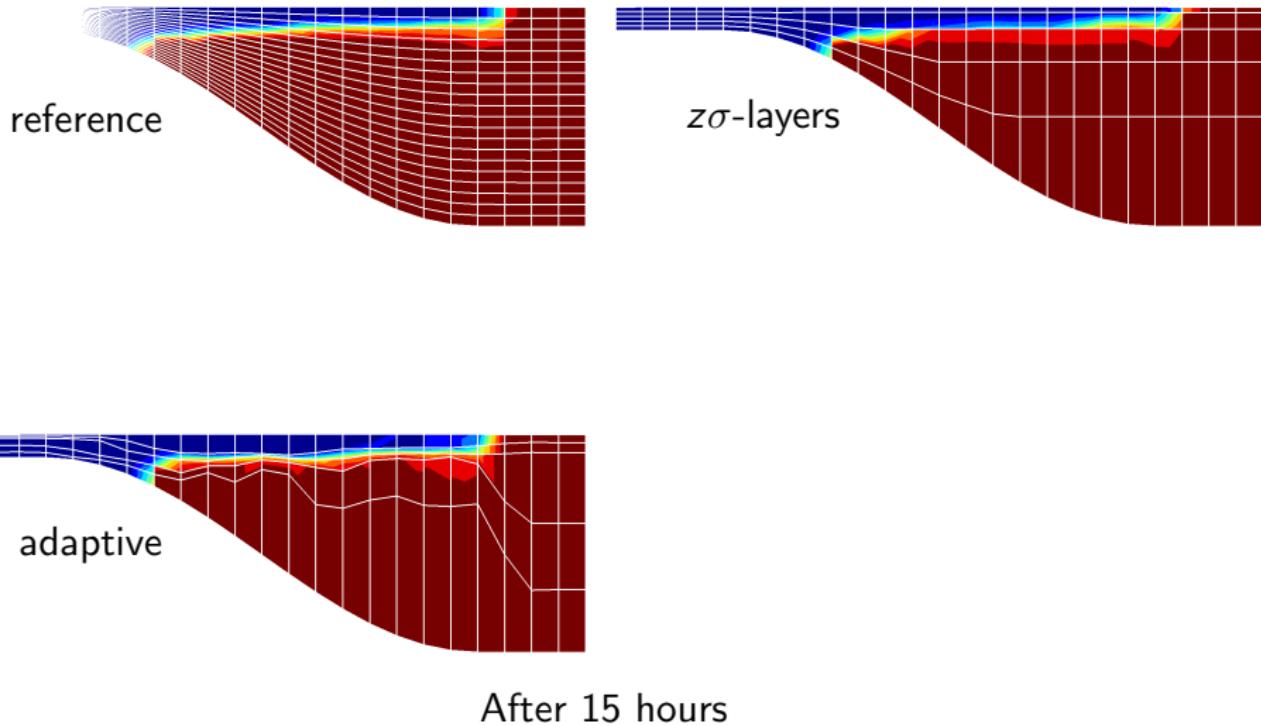
$$\kappa_z = \frac{1}{\tau} \left(c_{N^2} \kappa_{N^2} + c_{S^2} \kappa_{S^2} + c_d \kappa_d + c_b \kappa_b \right)$$



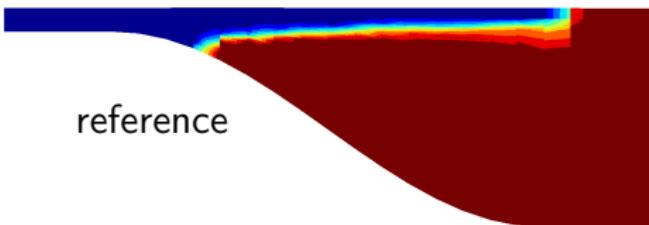
$$\kappa_{N^2} = \frac{[\rho]_v}{\rho_0}, \quad \kappa_{S^2} = \frac{[u]_v}{u_0},$$

$$\kappa_d = \frac{h}{d + d_0}, \quad \kappa_b = 1.$$

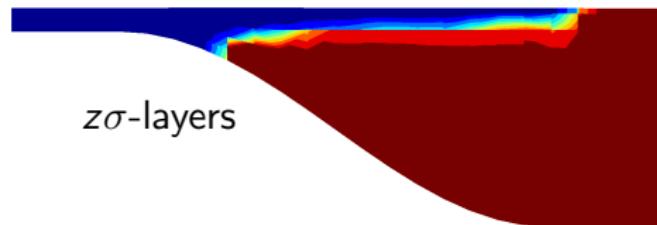
Idealized river plume



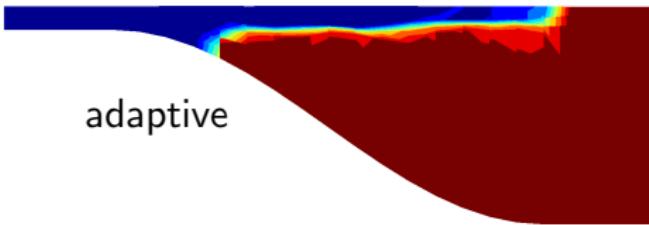
Idealized river plume (mesh hidden)



reference



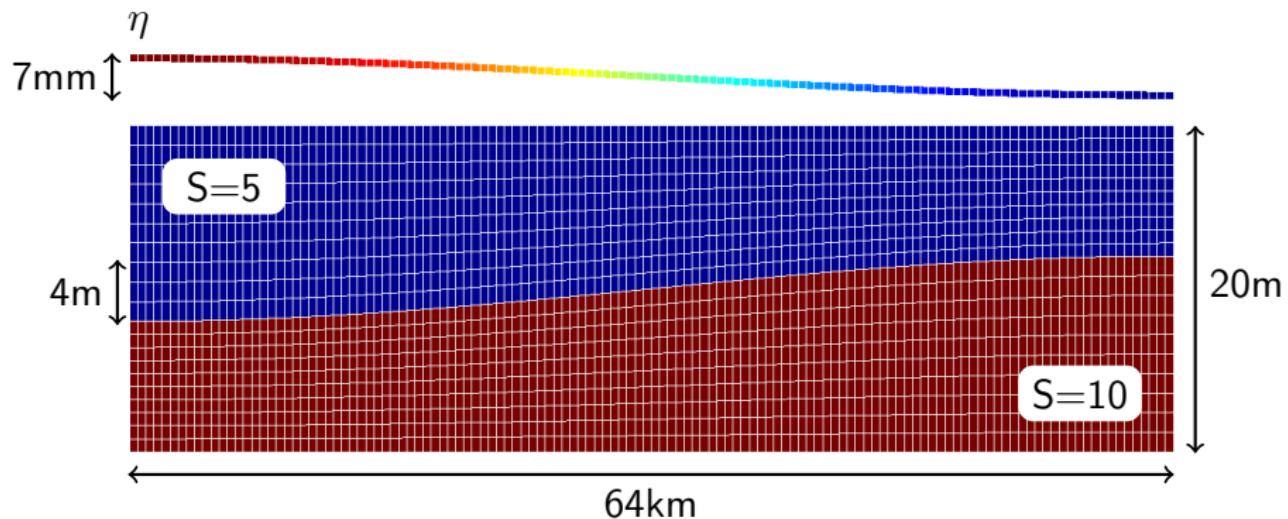
$z\sigma$ -layers



adaptive

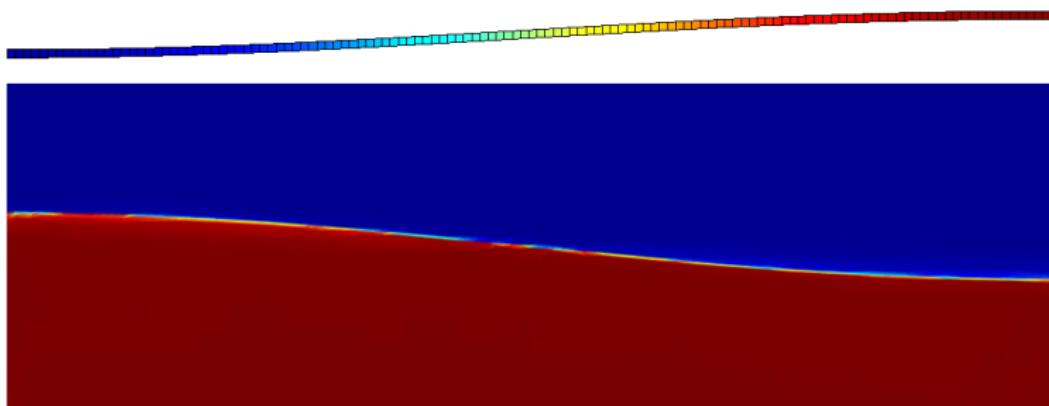
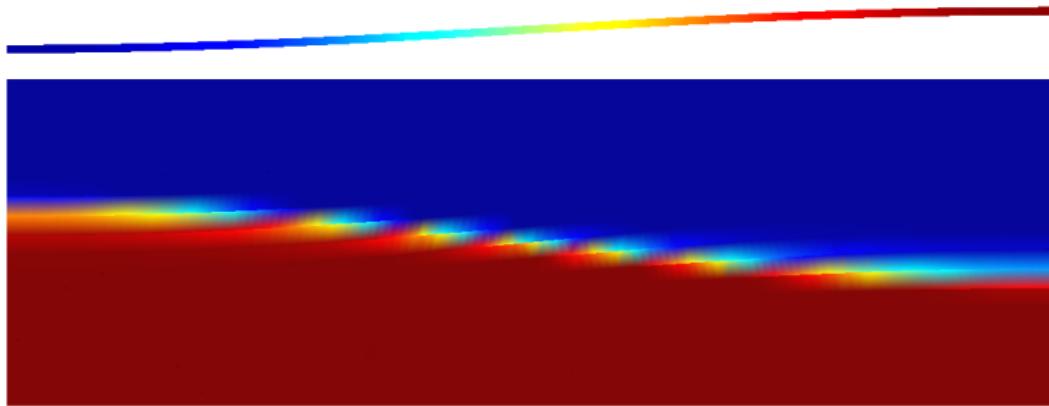
After 15 hours

Internal seiche



Oscillation period: $T = 3.4$ days

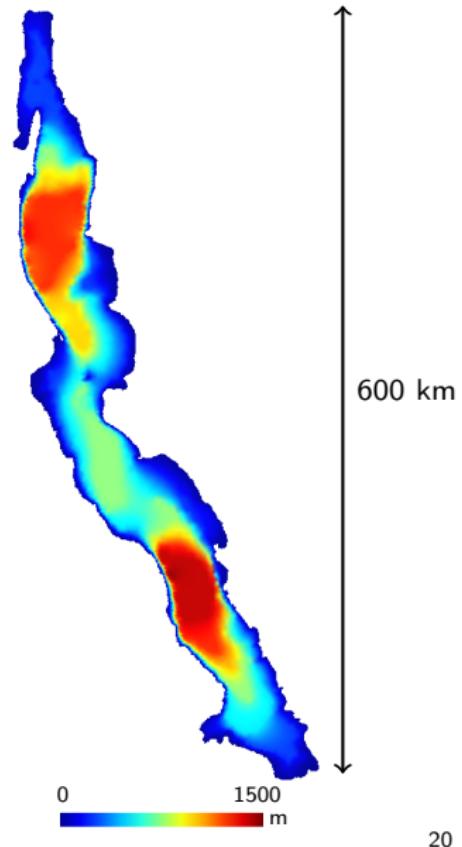
Internal seiche



After $T/2$

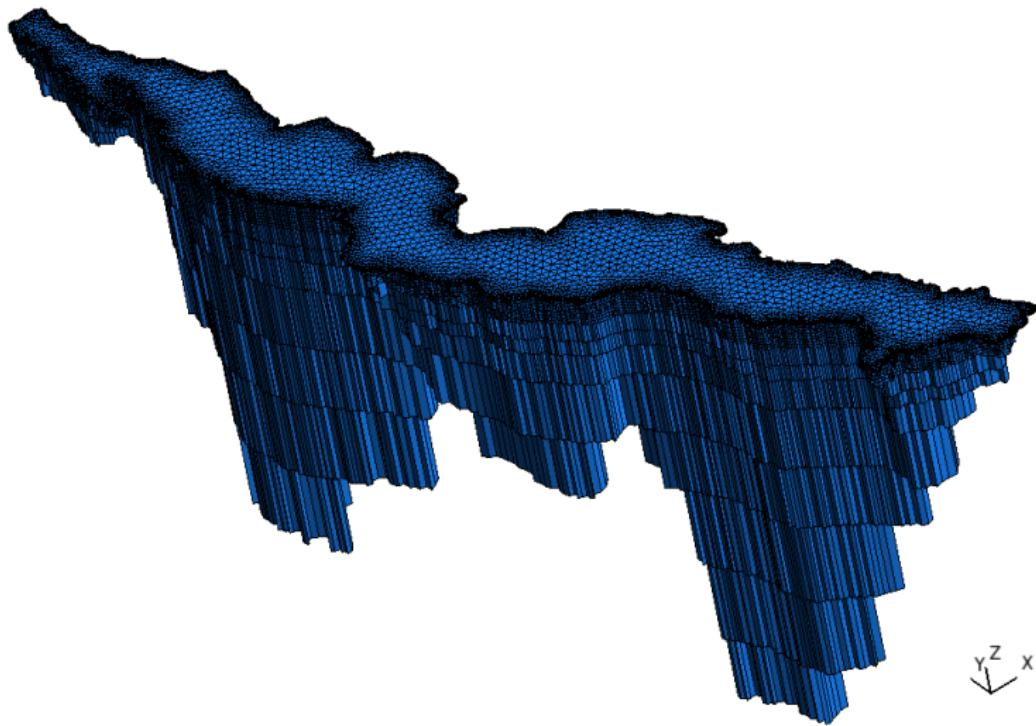
Lake Tanganyika

- Goal
 - Simulate the oscillations of the thermocline
 - Adaptive coordinates are necessary
- Difficulties
 - Deep lake
 - Strong bathymetry gradient



Lake Tanganyika

For this run, z-layers will be used.



Conclusions

- σ -layers
 - Plume dynamics depends on bathymetry
- $z\sigma$ -layers
 - + Simple and not noisy
 - + Good results on river plume test case (if enough z -layers)
 - Constant grid
 - Bad results on internal seiche test case
- Adaptive vertical layers
 - + Very good result for internal seiche
 - + Cheap implementation
 - Difficult to get best diffusivity parameters
 - Does not enhance the simulation for river plume
- Perspectives: test adaptive vertical coordinates on Lake Tanganyika!

Thank you for your attention !

References

- Sébastien Blaise, Richard Comblen, Vincent Legat, Jean-François Remacle, Eric Deleersnijder, and Jonathan Lambrechts. A discontinuous finite element baroclinic marine model on unstructured prismatic meshes. Part I: space discretization. *Ocean Dynamics*, 60(6):1371–1393, 2010.
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- Richard Hofmeister, Hans Burchard, and Jean-Marie Beckers. Non-uniform adaptive vertical grids for 3d numerical ocean models. *Ocean Modelling*, 33(1):70–86, 2010.
- Tuomas Kärnä, Vincent Legat, and Eric Deleersnijder. A baroclinic discontinuous Galerkin finite element model for coastal flows. *Ocean Modelling*, 61:1–20, 2013.