IHE 2018

Density dependent groundwater flow in the coastal zone

Gualbert Oude Essink, PhD

Lecture set-up:

- · PowerPoint sheets
- · Lecture Notes
- Practicals numerical modelling

http://freshsalt.deltares.nl





Deltares
Unit Subsurface and Groundwater Systems
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13-14-15 June 2018

Introduction

Curriculum Vitae

Delft University of Technology, Civil Engineering: till 1997

Ph.D.-thesis: Impact of sea level rise on groundwater flow regimes

- Utrecht University, Earth Sciences: till 2002
- Free University of Amsterdam, Earth Sciences: till 2004
- Deltares
- Utrecht University (Associate Professor): from 2014

Qualifications:

- · Groundwater resources management
- Density-dependent groundwater flow and coupled solute transport
- Salt water intrusion in coastal aquifers
- Assessment of climate change on groundwater resources
- Numerical Modeling
- · Teaching and training

http://freshsalt.deltares.nl
Deltares: gualbert.oudeessink@deltares.nl

Colleagues at Deltares Groundwater in the Coastal Zone

http://zoetzout.deltares.nl http://freshsalt.deltares.nl







Gualbert Oude Essink Joost Delsman

Pieter Pauw









Perry de Louw

Esther van Baaren

Jarno Verkaik

iviarta Fanec

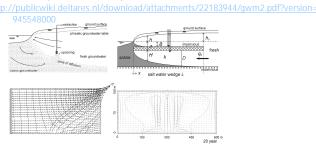


Research on groundwater in the coastal zone

- 18 years experience in modelling variable-density dependent groundwater flow and coupled solute transport in the coastal zone
- Incorporating monitoring campaigns results in numerical modeling tools
- Research on new fresh-saline phenomenae: salty seepage boils and shallow freshwater lenses in saline environments
- Knowledge on creating 3D initial chloride distribution, based on geostatistics and geophysical data (analyses, VES, borehole measures, AEM)
- Quantifying effects of climate change and sea level rise on fresh groundwater resources
- Developing adaptive and mitigative measures to stop salinization in the coastal groundwater system (e.g. ASR, MAR: fresh keeper, coastal collectors, freshwater storage underground)

Lecture notes, practicals and ppt on freshsalt.deltares.nl

1. Density dependent groundwater flow



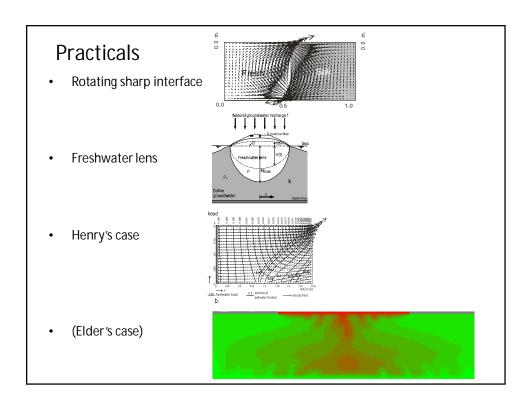
2. Groundwater modelling



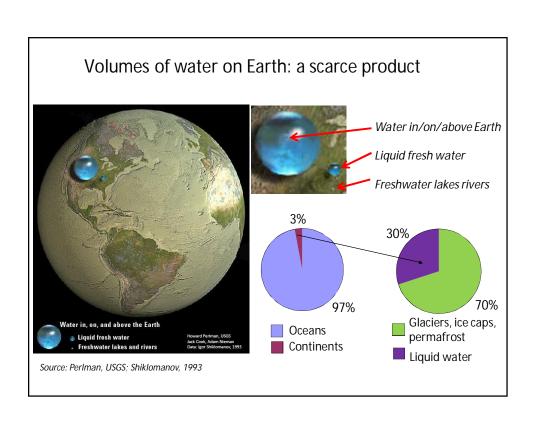
Practicals numeriacl modelling

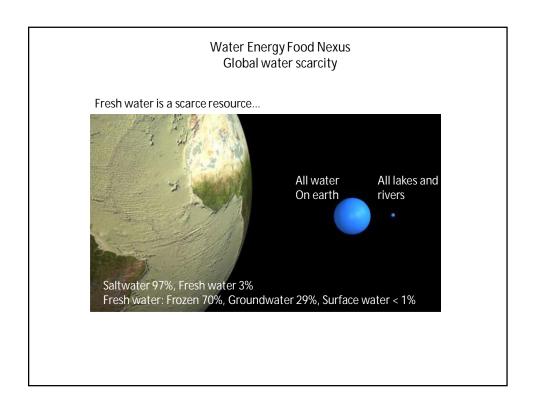
- PMWIN
- SEAWAT
- Cases:
 - Rotating sharp interface
 - Freshwater lens
 - Henry's case
 - (Elder's case)
- Setup practicals:
 - work in small groups of two persons
 - short report of findings (make screenshots)
 - deliver within one week after finish last lectures

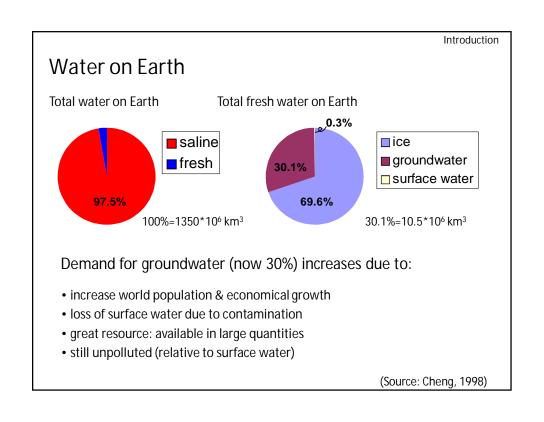
http://freshsalt.deltares.nl Deltares: gualbert.oudeessink@deltares.nl

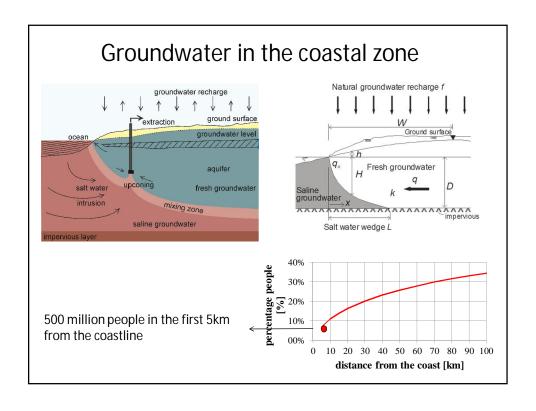






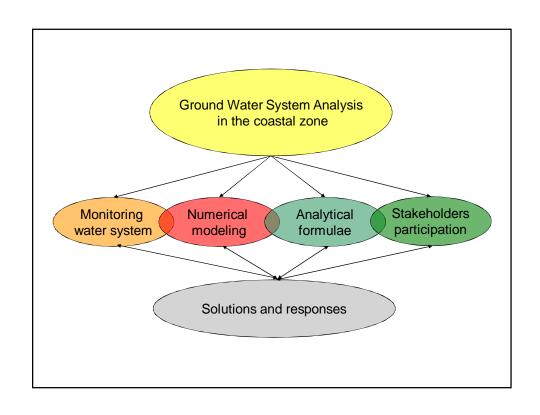


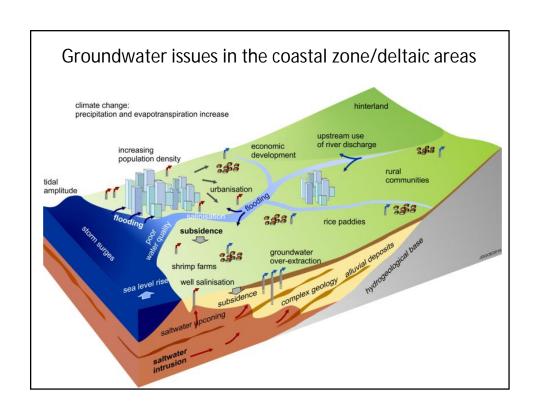


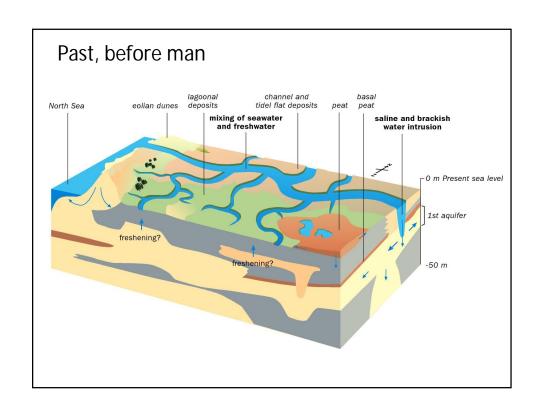


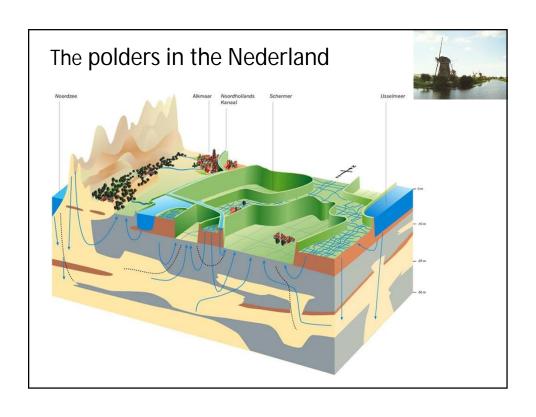
Topics of density driven groundwater flow

- 1. Introduction
 - water on earth
 - salt water intrusion
 - freshwater head
- 2. Interface between fresh and saline groundwater
 - analytical formulae (Badon Ghyben-Herzberg)
 - upconing example
- 3. Numerical modelling
 - mathematical background
 - Benchmark problems: Henry, Elder, Hydrocoin, etc.
- 4. Case-studies
 - hypothetical cases
 - 2D, 3D cases
 - real cases (Dutch coastal zone)









Groundwater in the future

We have to cope which...:

- We have to cope which...:
- · Groundwater extractions
- Development energy use/production (heat-cold)
- Climate change
- Land subsidence
- Development spatial land use
- Politics, Policy & Watermanagement

Direct anthopogenic influence on groundwater is more important than climate effect

Salt Water Intrusion Meeting, since 1968

Salt Water Intrusion Meeting, since 1968



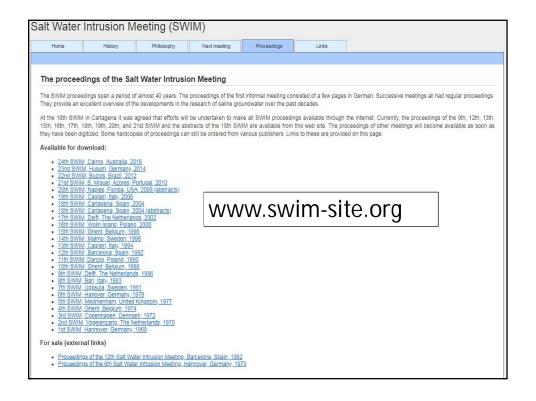
http://www.swim-site.org/

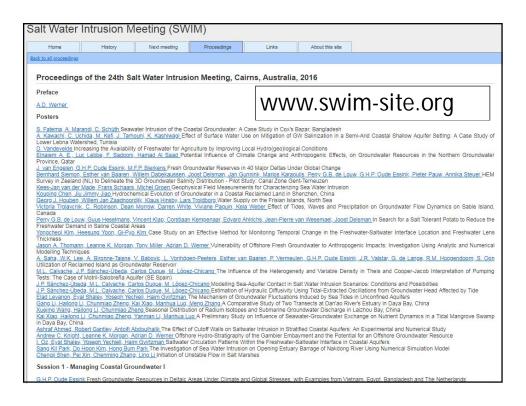
Themes

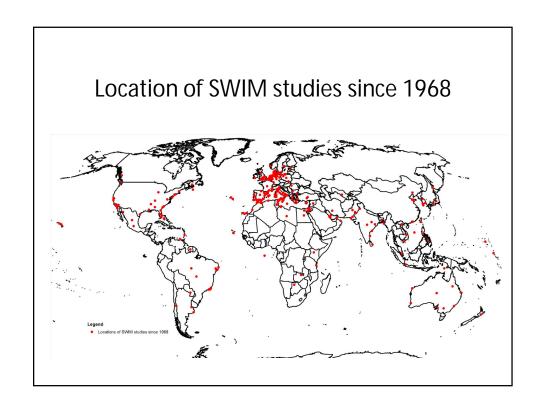
- · Water systeem analysis
- Monitoring
- Modelling
- Effects
- Solutions

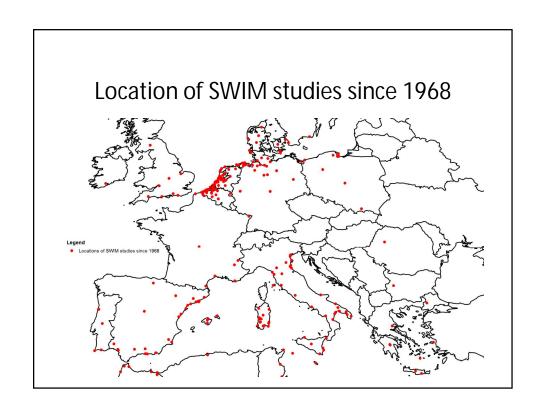




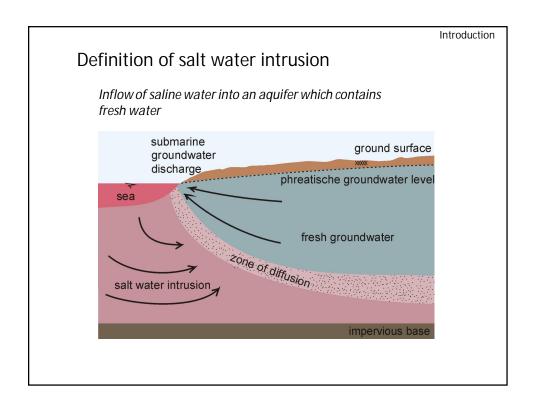








Introduction SWI



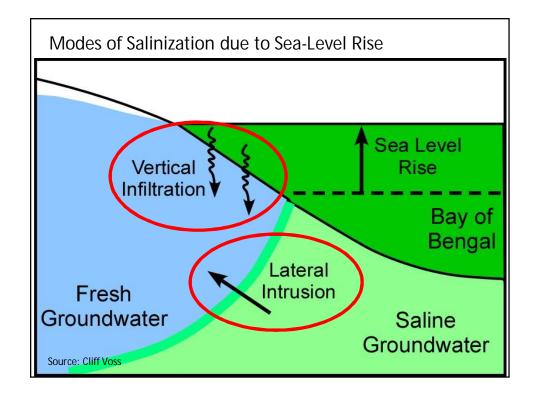
Origin of saline groundwater in the subsoil

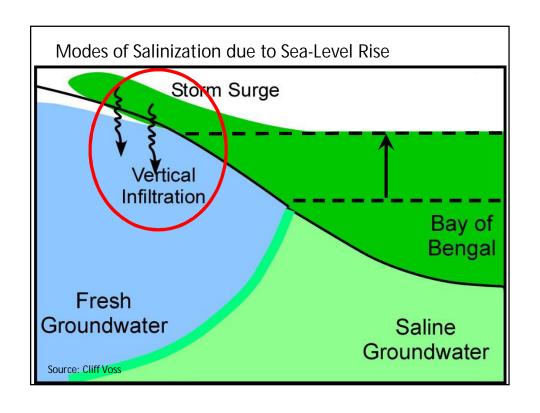
Geological causes:

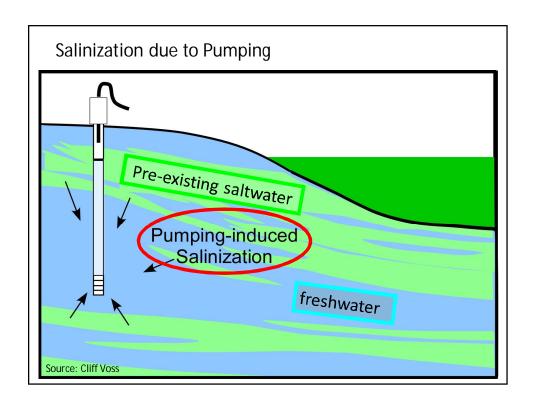
- -marine deposits during geological times
- -trans- and regressions in coastal areas (deltas)
- -salt/brine dome

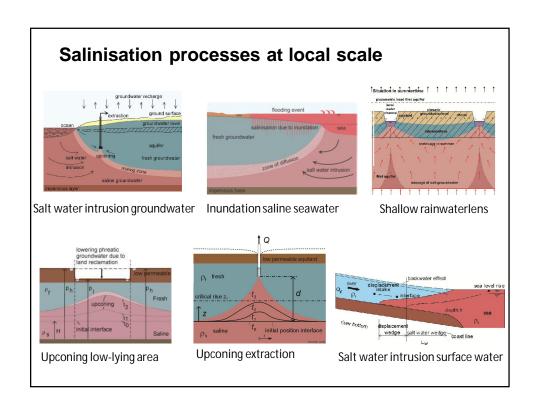
Anthropogenic causes:

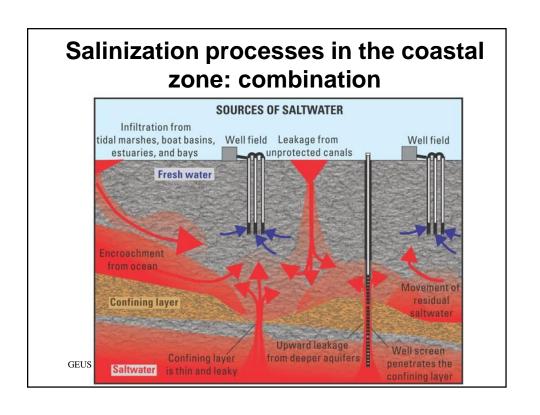
- -agriculture/irrigation (salt damage Middle East & Australia)
- -upconing under extraction wells throughout the world
- -upconing under low-lying areas (e.g. Dutch polders)

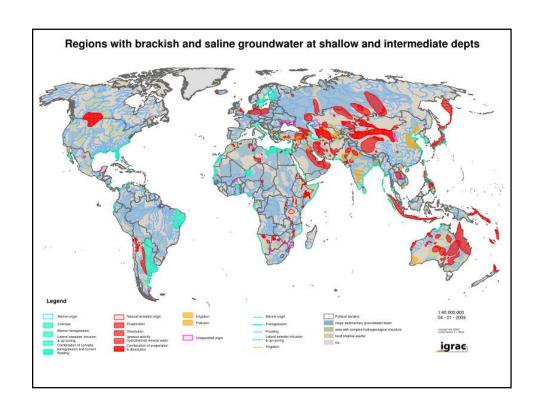


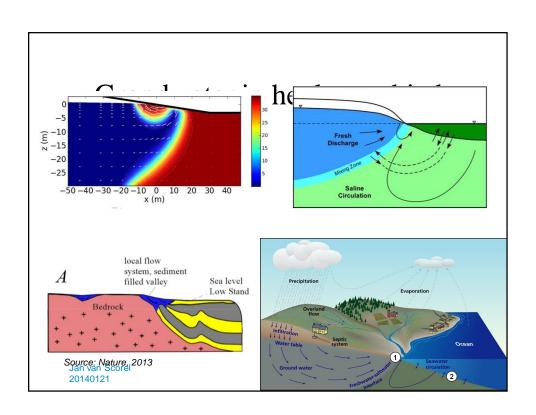


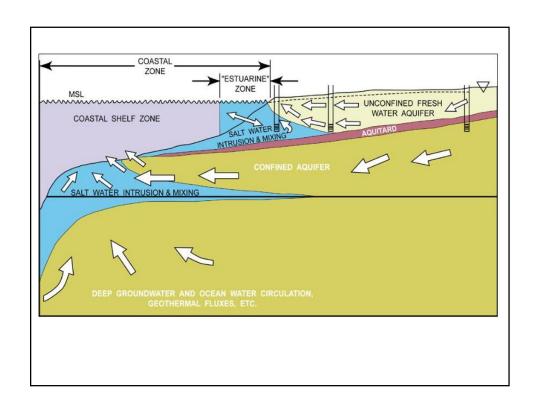


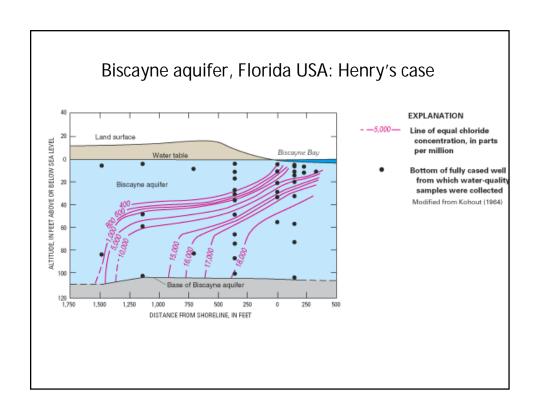


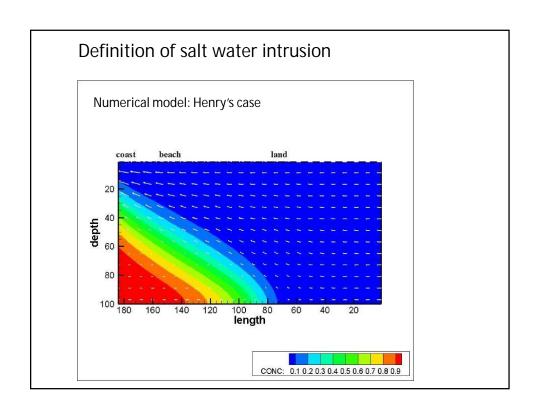


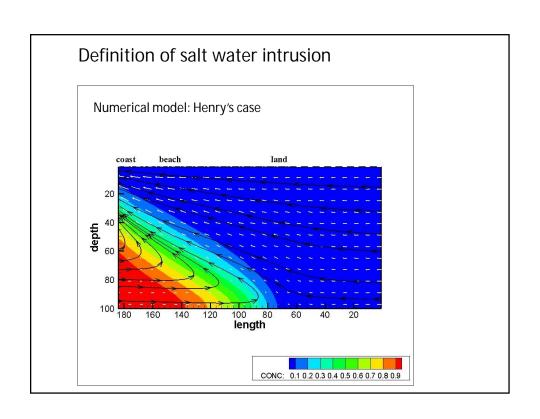


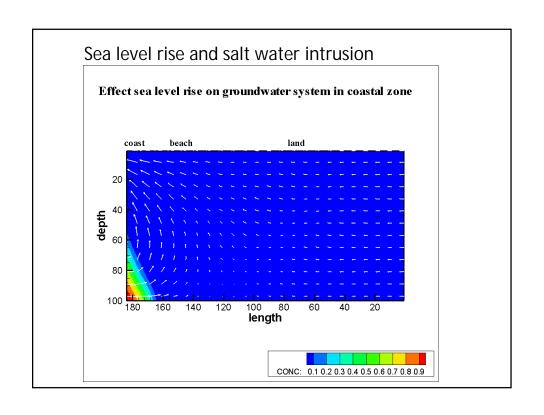


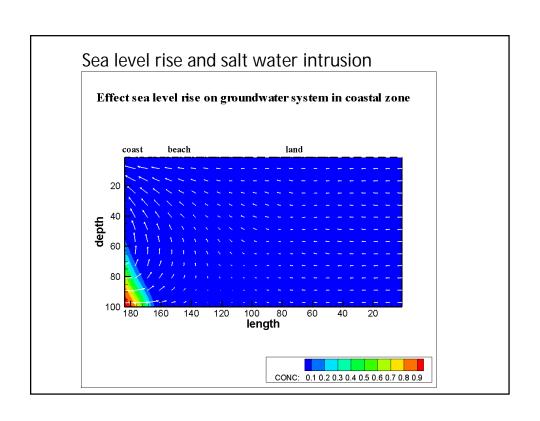


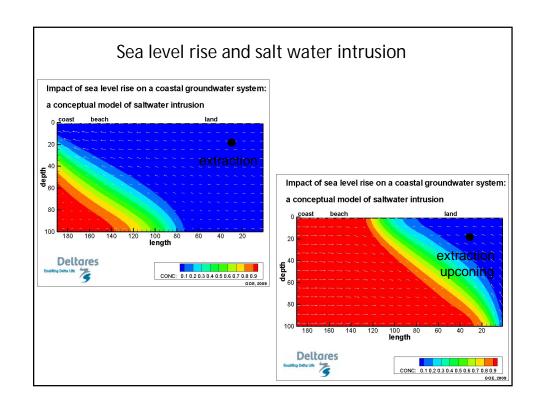


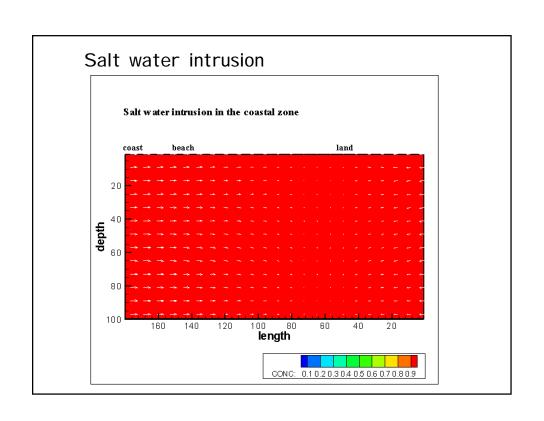


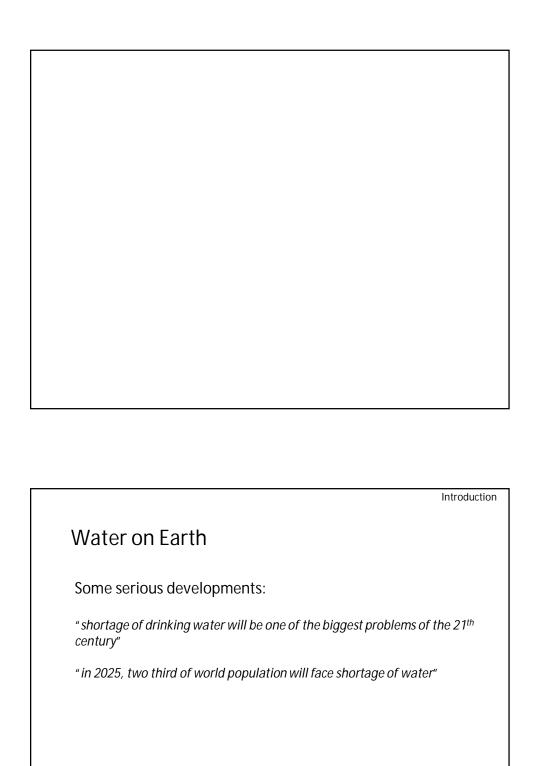


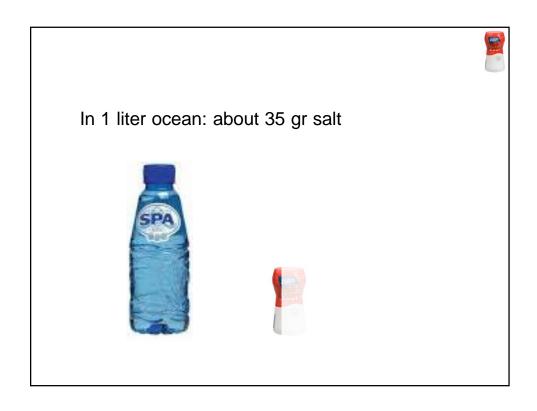




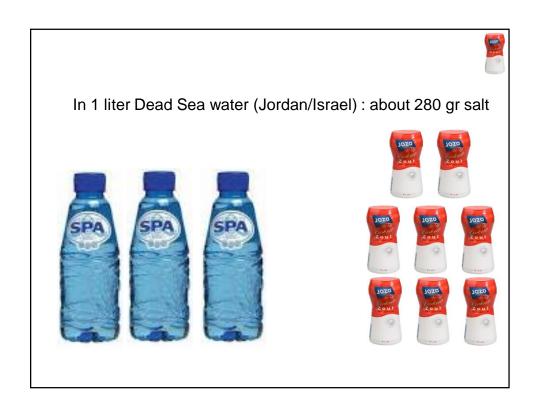


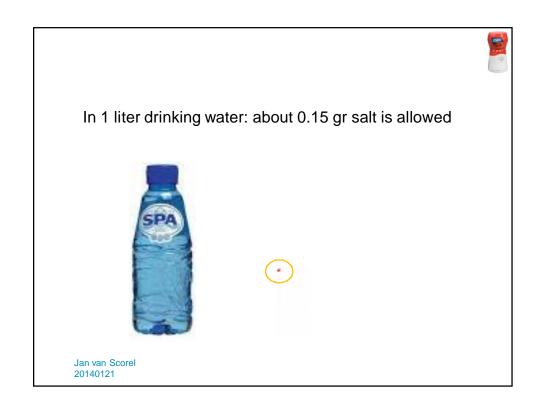


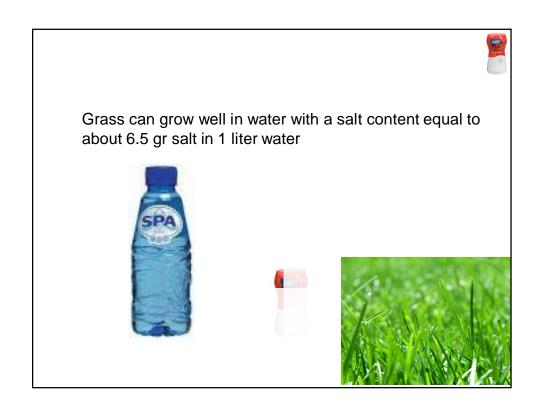












Fresh-brackish-saline groundwater [mg/L] Negative ions 19000 SO₄-2 2700 HCO 3 140 65 Total negative ions 21905 **Positive ions** 10600 Na⁺ Mg +2 1270 Ca +2 400 380 Total positive ions 12650 Total Disssolved Solids (TDS) 34555

Definition fresh-brackish-saline groundwater

Main type of groundwater	Chloride concentration [mg Cl ⁻ /L]	
oligohaline	0 - 5	
oligohaline-fresh	5 - 30	
fresh	30 - 150	
fresh-brackish	150 -300	
brackish	300 - 1000	
brackish-saline	1000 - 10.000	
saline	10.000 - 20.000	
hyperhaline or brine	≥20.000	

Туре	[mS/cm]	[mg TDS/L]	Drinking- or irrigation water	
Non-saline or fresh water	<0.7	<500	Drinking and irrigation water	
Slightly saline	0.7 - 2	500-1.500	Irrigation wate	
Moderately saline	2 - 10	1.500-7.000	Primary drainage water and groundwate	
Highly saline	10 - 25	7.000-15.000	Secondary drainage water and groundwater	
Very highly saline	25 - 45	15.000-35.000	Seawater is about 35000 TDS mg/L	
Brine	>45	>35.000	n.a.	

EOS

Examples of equations of state

$$\rho_{(S,T)} = 1000 + 0.8054S - 0.0065(T - 4 + 0.2214S)^{2}$$
 Linear (concentration) C.

Linear (concentration)

Linear (concentration)
$$\rho_{(C)} = \rho_f [1 + \alpha \, \frac{C_i}{C_s}] \quad \text{where a=relative density difference}$$

Linear (temperature)

$$\rho_{\scriptscriptstyle (T)} = \rho_f [1 - \beta (T - T')]$$

Exponential (temperature, pressure, salt)

$$\rho_{(T,p,\omega)} = \rho_f e^{-\alpha(T-T_0)+\beta(p-p_0)+\gamma\omega}$$

Equation of state (SEAWAT)

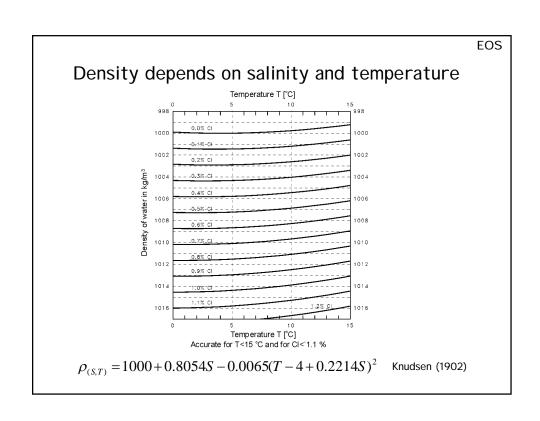
$$\rho_{i,j,k} = \rho_f + \frac{\partial \rho}{\partial C} C_{i,j,k}$$

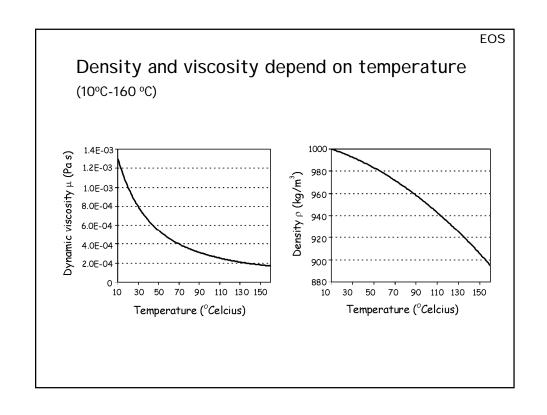
e.g.:

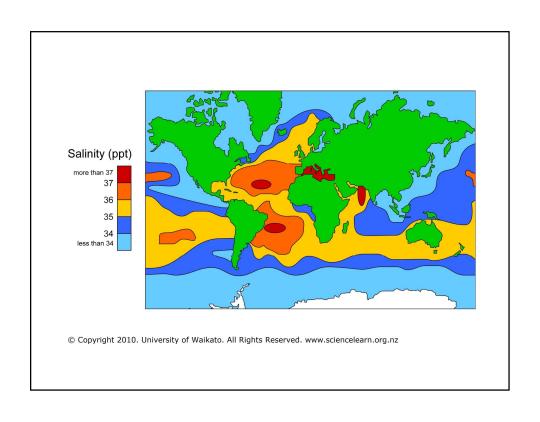
1. conc=35 TDS g/I: DRHODC=0.7143

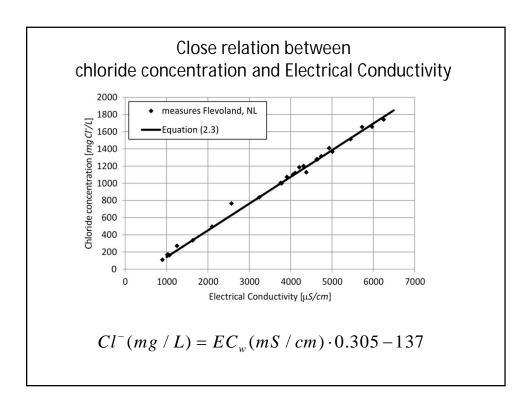
2. conc=19000 mg CI-/I: DRHODC=0.001316 (as 1025=1000+0.001316*19000)

3. conc=1: DRHODC=25 (example practicals)









Close relation between chloride concentration and Electrical Conductivity

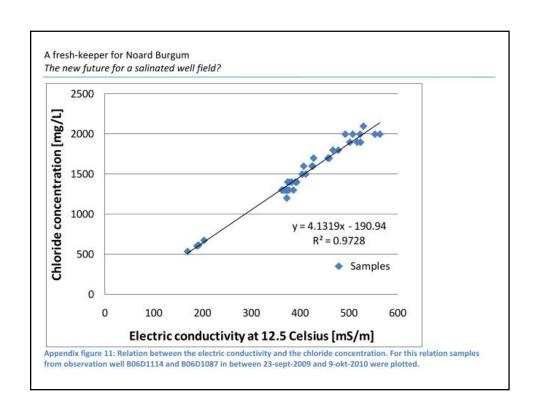
$$10^6 \,\mu\text{S/cm} = 10^3 \,m\text{S/cm} = 1 \,\text{S/cm}$$

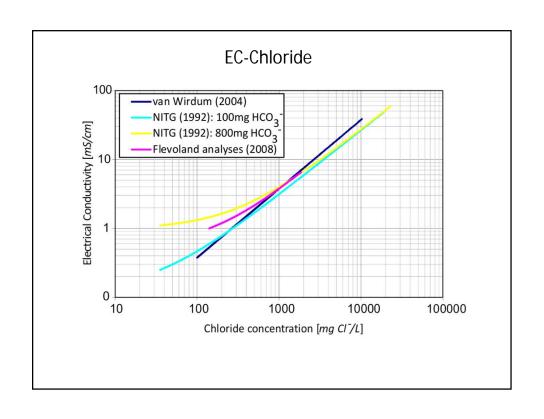
 $1 \,\mu\text{S/cm} = 100 \,\mu\text{S/m}$

ocean water:

- ~19000 mg CI-/L or ~34555 mg TDS/L
- ~5 S/m or ~48 mS/cm

the ratio Cl⁻ over TDS equal to ~0.554, under stable normal seawater environments



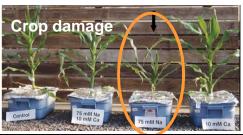


Salt in water is a problem







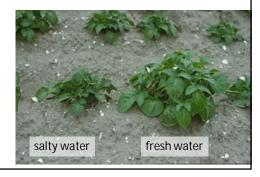




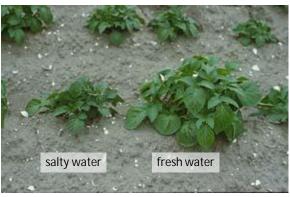
Introduction

Salt in water is a problem for different water management sectors:

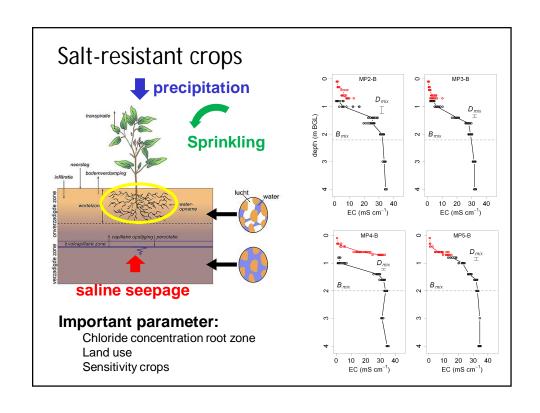
- -drinking water:
 - •taste (100-300 mg Cl-/l)
 - •long term health effect
 - •norm: EC& WHO=150 mg Cl⁻/I (live stock=1500 mg Cl⁻/I)
- -industry:
 - •corrosion pipes
 - preparation food
- -irrigation/agriculture:
 - •production crops
 - •salt damage



Effects salinisation: salt damage



Source: Proefstation voor de Akkerbouw en Groenteteelt, Lelystad



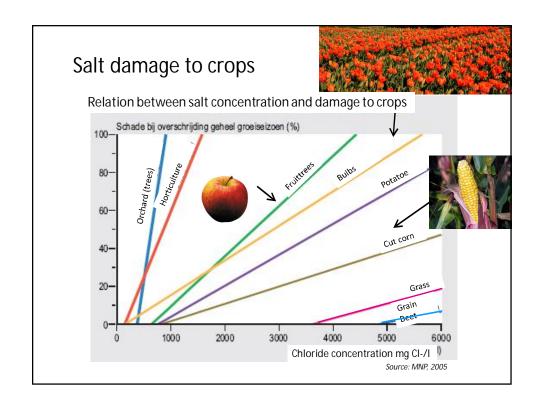
Salt damage to crops

Important parameters:

- Chloride concentration in the root zone
- Land use
- Sensitivity crops

Land use	Threshold root zone (mg CI-/I)	value	Gradient root zone (-)
Grass		3606	0.0078
Potatoes		756	0.0163
Beet		4831	0.0057
Grains		4831	0.0058
Horticulture		1337	0.0141
Orchard (trees)		642	0.0264
Bulb		153	0.0182

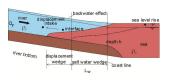
Source: Roest et al., 2003 en Haskoning

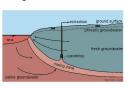


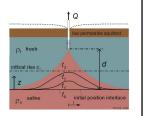
	Soil moisture		Irrigation water	
	Limi	Gradient	Limit	Gradient
Crop	mg/I CI	%/mg/I CI	mg/I CI	%/mg/I CI
Potatoe	756	0.0163	202	0.0610
Grass	3606	0.0078	962	0.0294
Sugar beat	4831	0.0057	1288	0.0212
Cut Corn	815	0.0091	217	0.0343
Grains	4831	0.0058	1288	0.0218
Fruit trees	642	0.0264	171	0.0991
Orchard (trees)	378	0.1890	101	0.7086
Vegetables	917	0.0158	245	0.0591
Horticulture	1337	0.0141	356	0.0527
Bulbs	153	0.0182	41	0.0683

Why is salinisation a pressing problem?

- 30% of world population lives <70 km from coastline
 economic and tourist activities increase
- enormous increase in extraction
- irreversible process
- increase saltwater intrusion problem world-wide:
 - upconing
 - salt water wedge
 - decrease outflow q₀
- climate change:
 - sea level rise
 - natural groundwater recharge







Processes that accelerate salt water intrusion:

- Sea level rise
- Land subsidence
- Human activities

Threats for:

- drinking water supply in dunes: upconing of saline groundwater decrease of fresh groundwater resources recharge areas reduction
- agriculture: salt damage to crops: salt load and seepage
- water management low-lying areas: flushing water channels
- ecology



The water footprint of products

global averages

1 kg wheat $1 \text{ m}^3 \text{ water}$

1 kg rice 3 m³ water

1 kg milk 1 m³ water

1 kg cheese 5 m³ water

1 kg pork 5 m³ water

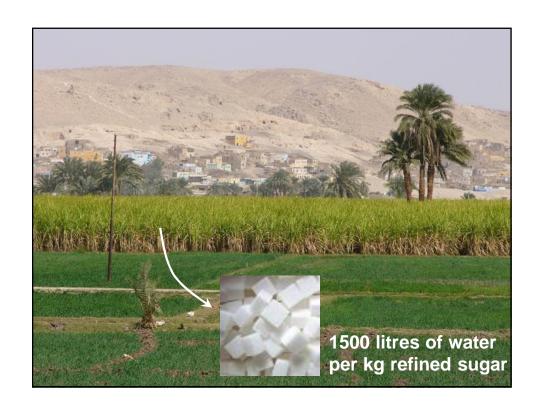
1 kg beef 15 m³ water

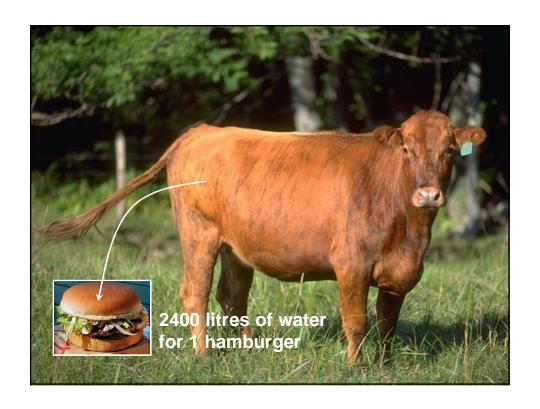


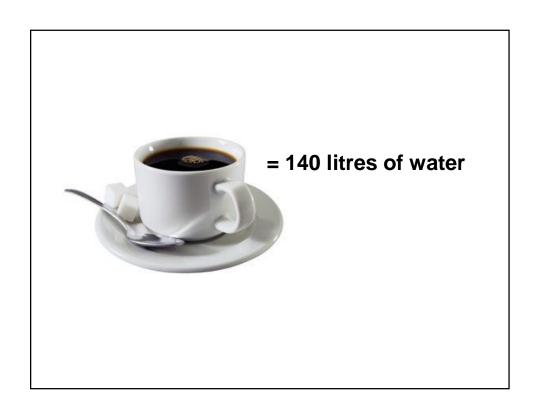


[Hoekstra & Chapagain, 2008]











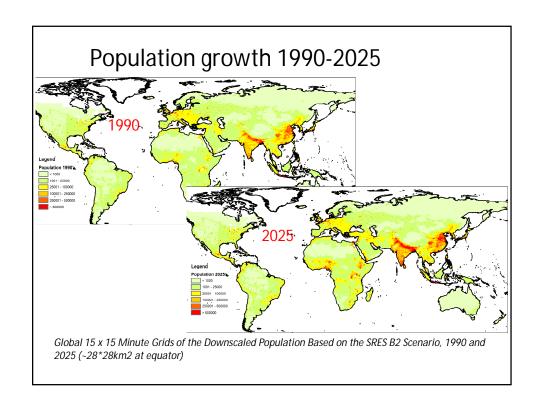
Introduction

Question:

Demand fresh water per capita per day in the Netherlands?:

- a. 10 litre/day
- b. 25 litre/day
- c. 100 litre/day d. 200 litre/day

Water withdrawal as % of total available water 1995 from 20 % to 10 %



Introduction

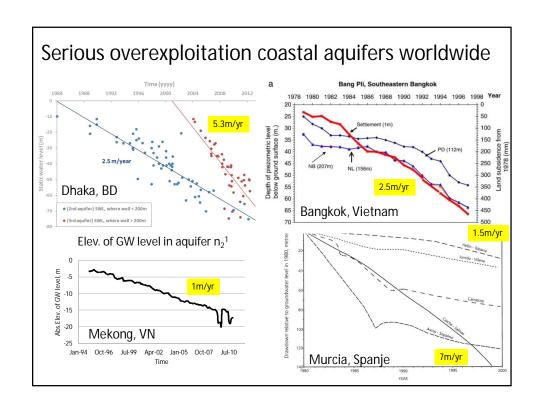
Reasons and drawbacks of using groundwater

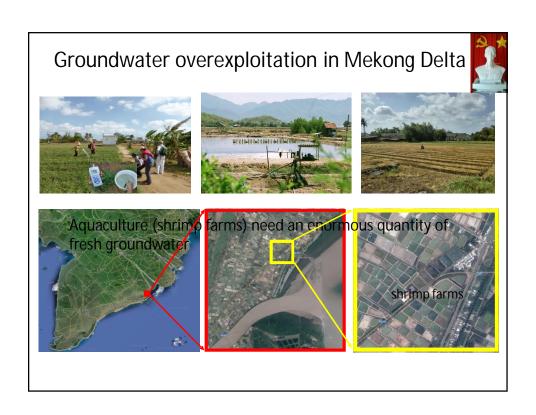
Advantage:

- -no seasonal effects
- -high quality
- -low storage costs
- -large quantities
- -no spatial limitations

Disadvantage:

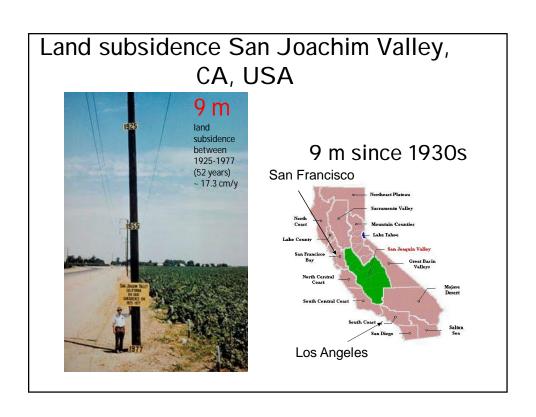
- -high extraction costs
- -local droughts
- -high mineral content
- -land subsidence....
- -salt water intrusion!





Land subsidence

Megacity	Maximum subsidence [m]	Date commenced
Shanghai	2.80	1921
Tokyo	5.00	1930's
Osaka	2.80	1935
Bangkok	1.60	1950's
Tianjin	2.60	1959
Jakarta	0.90	1978
Manila	0.40	1960
Los Angeles	9.00	1930's



What causes the land to subside?

Natural causes (geological processes):

- Loading of the earth's crust by ice sheets, sediment (delta's), the ocean/sea
- Compaction of older sediments after sedimentation

Anthropogenic causes (human-induced processes):

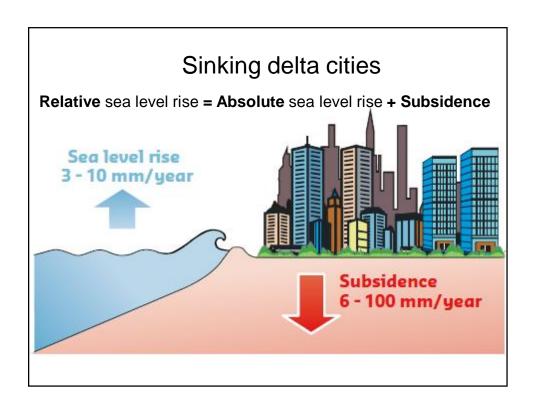
- Oil/gas extraction (usually relatively deep)
- Groundwater extraction (usually moderately deep)
- Drainage of soils ⇒ oxidation of peat, soil compaction

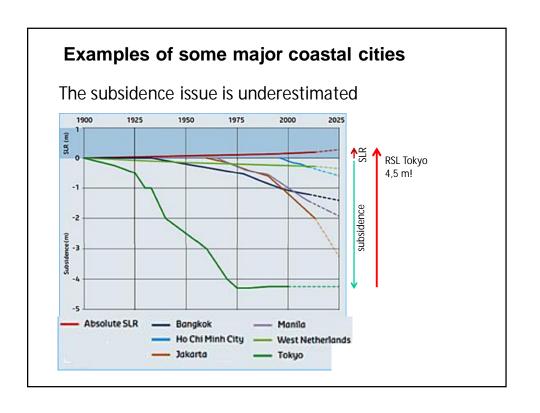
Why discriminating between human-induced and natural processes?

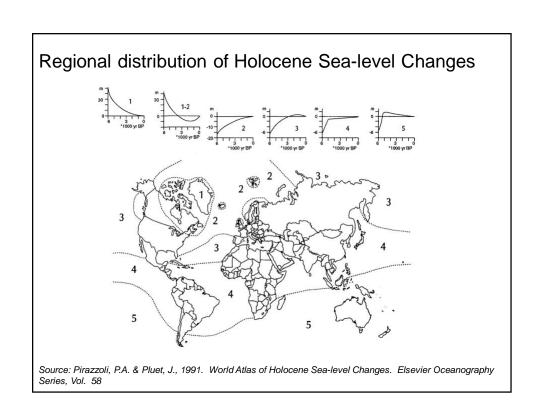
- Magnitude
- Cooping strategy (mitigation versus adaptation)

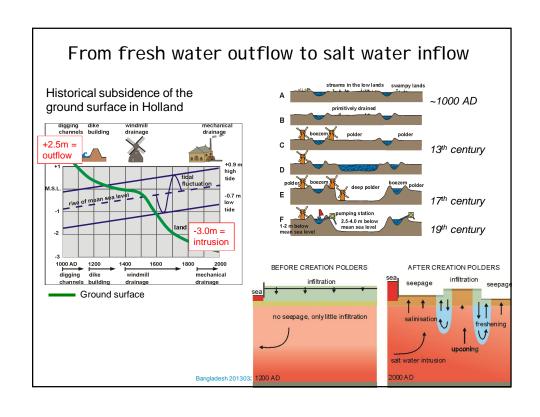


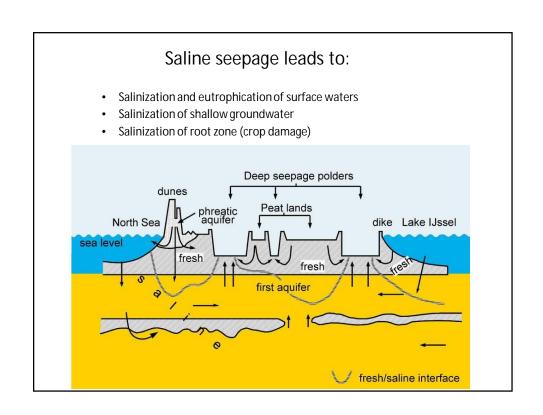


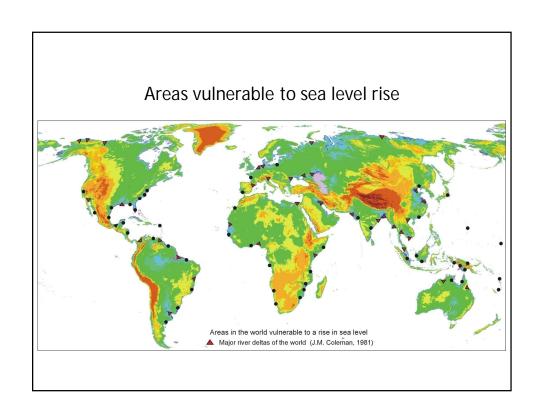




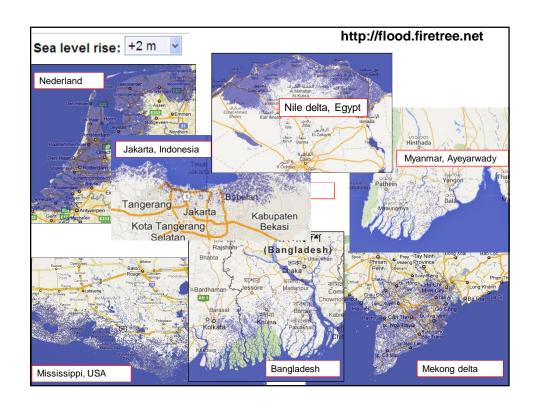


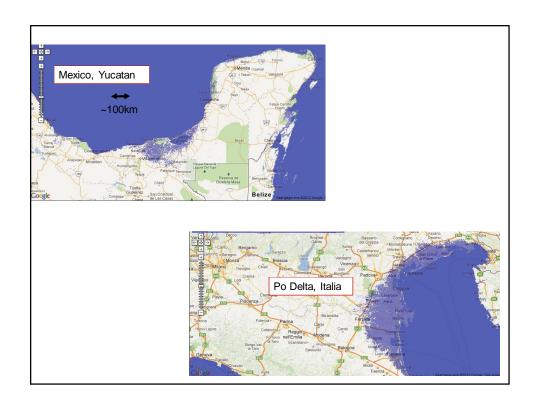


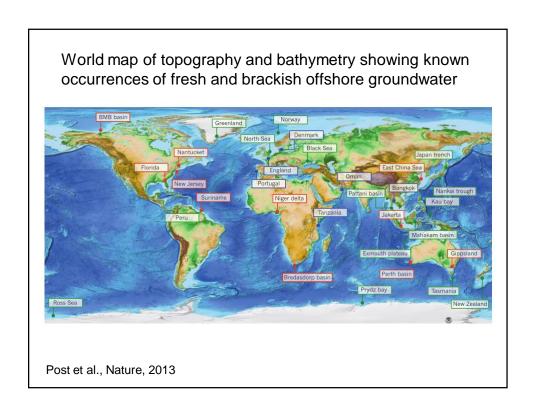


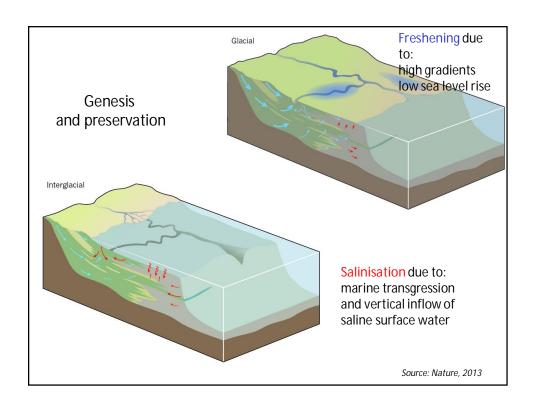


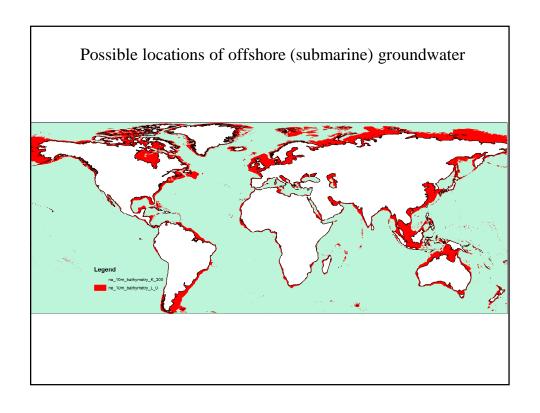


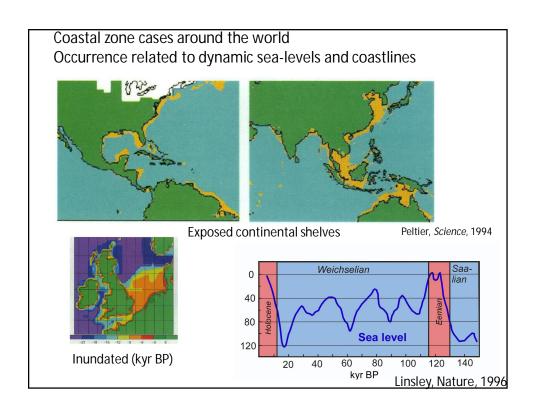


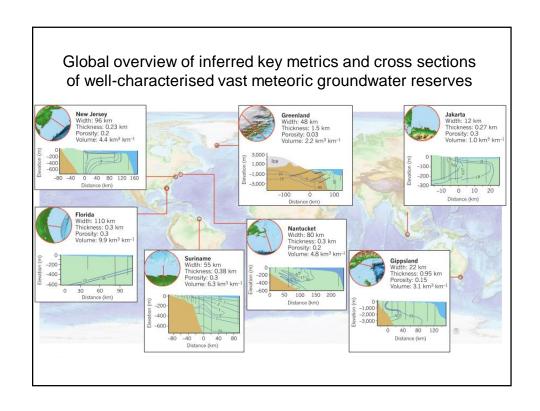




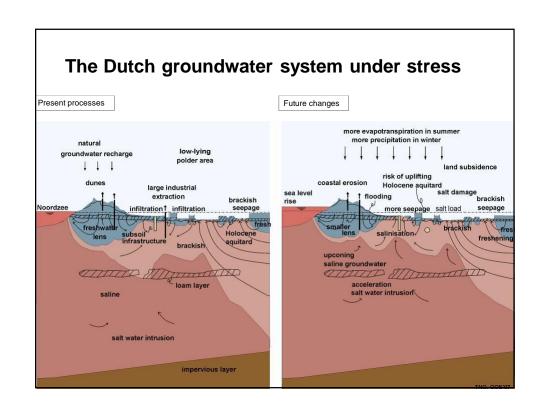


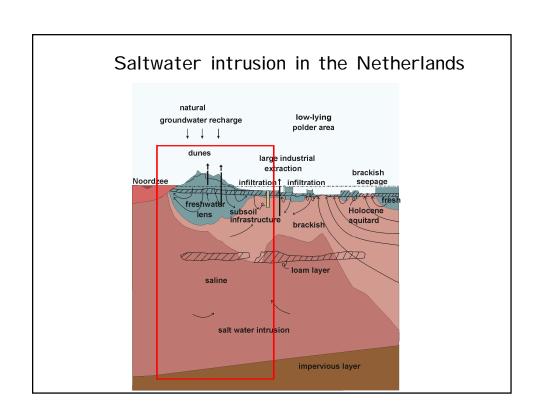




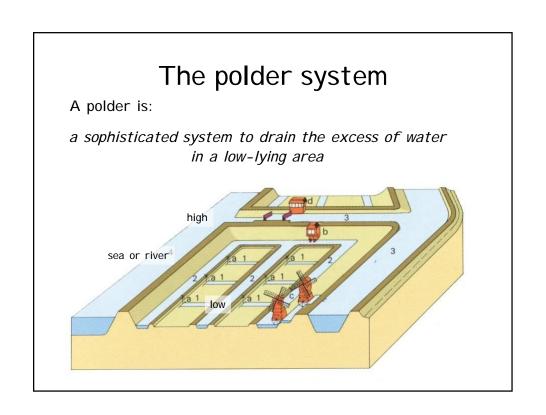


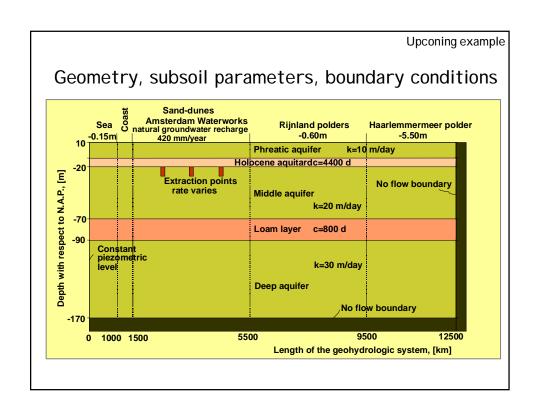
My first density dependent groundwater flow and solute transport model in 1990!

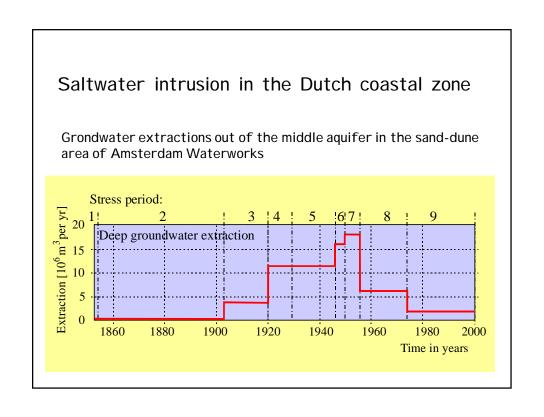


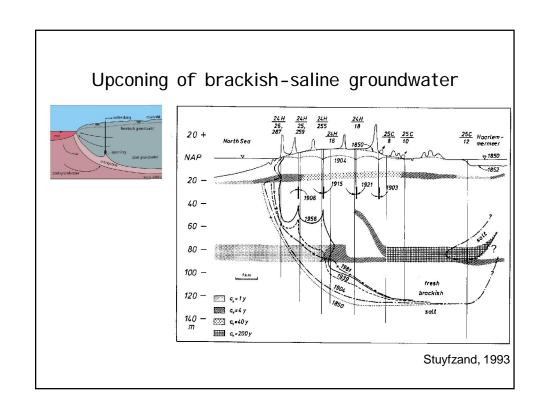


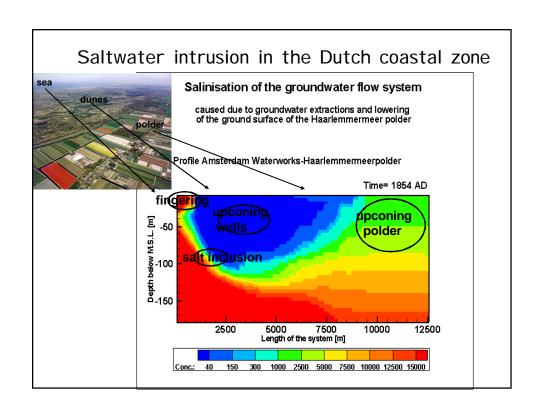
Saltwater intrusion in the Dutch coastal zone Position profile through Amsterdam Waterworks, Rijnland polders and Haarlemmermeer polder Haarlem Railway Ringvaart **North Sea** Specific profile Cruquius Amsterdam Waterworks M Hoofddorp Sand-dunes profile Haarlemmermeer city polder

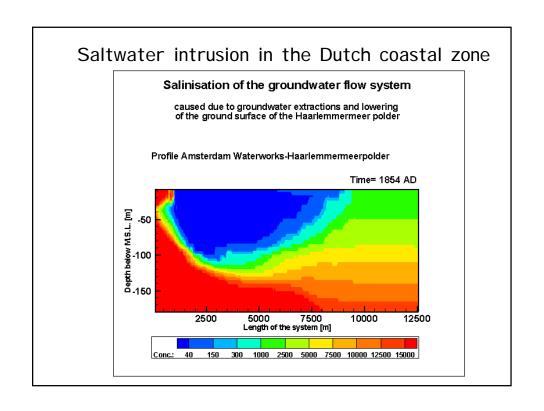








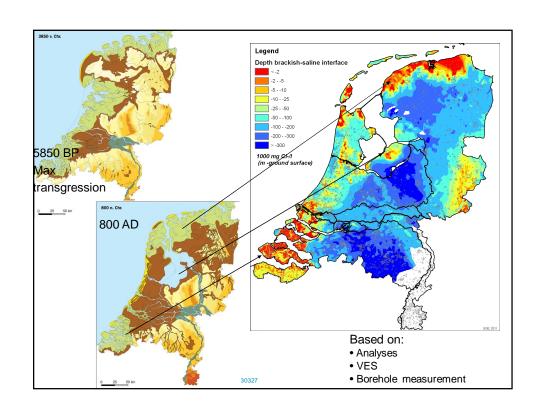


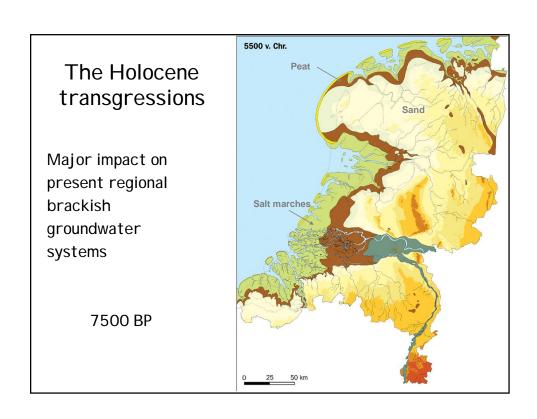


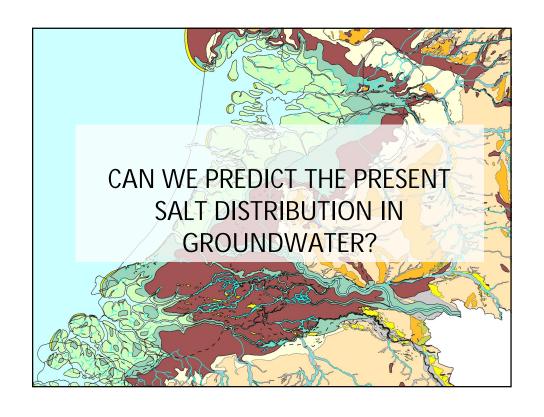
Palaeo hydrogeological modelling

Palaeo-modeling salt water intrusion during the Holocene: an application to the Netherlands

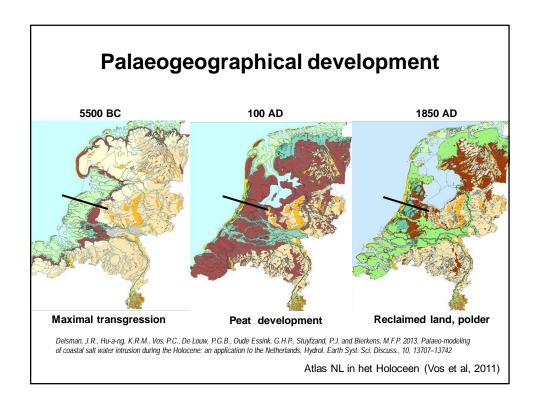
J.R. Delsman, K. Hu-a-ng, P.C. Vos, P.G.B. de Louw, G.H.P. Oude Essink and M.F.P. Bierkens









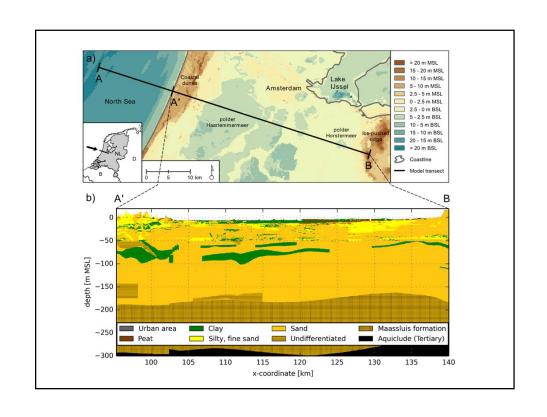


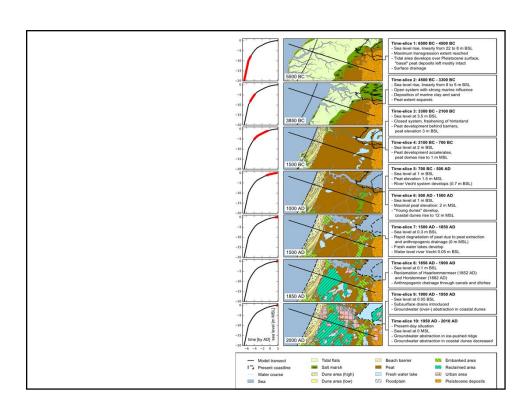
Occurrence of salt under the polder Haarlemmermeer

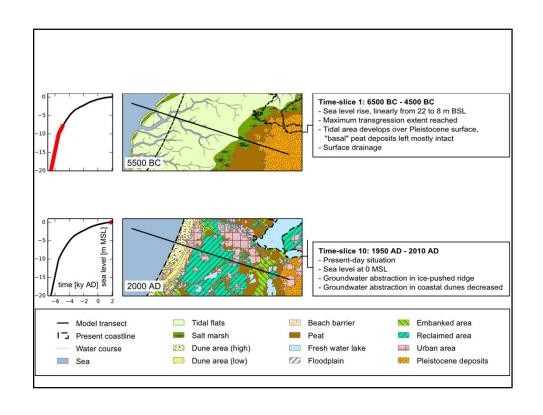
- Model profile Zandvoort Hoofddorp Hilversum
- Palaeogeographical development (Vos et al, 2011)
- 6500 BC 2010 AD
- · marine transgression
- · Peat development, peat degradation, drainage,

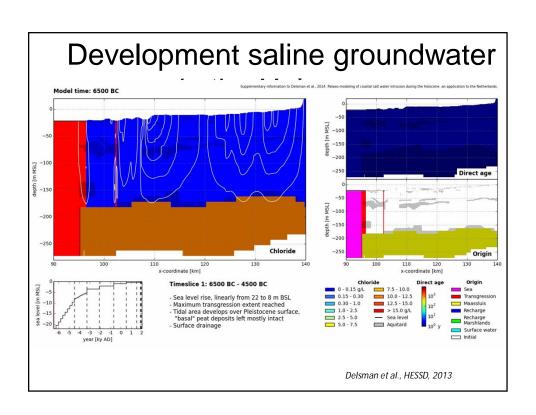


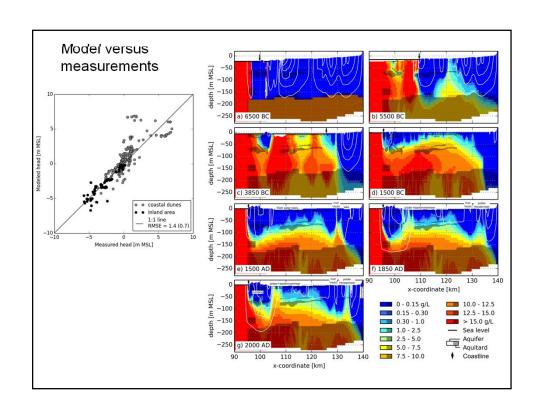
Delsman et al., HESS, 2013

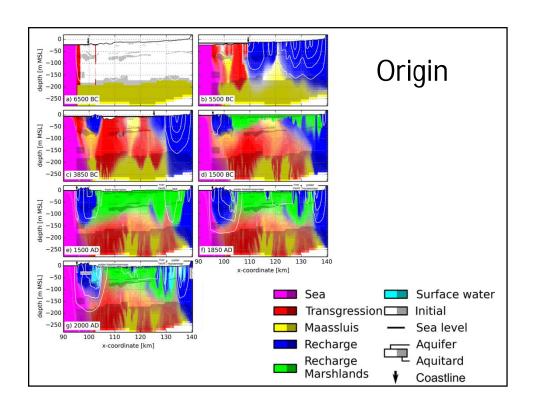


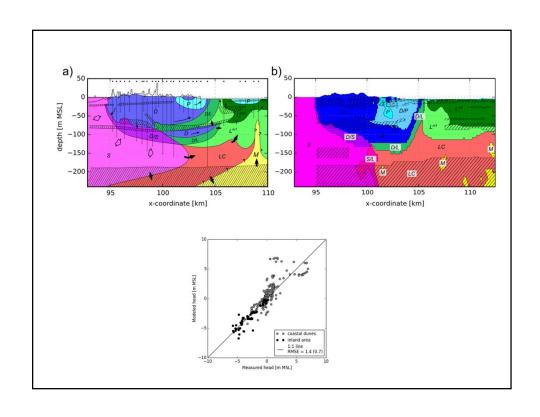


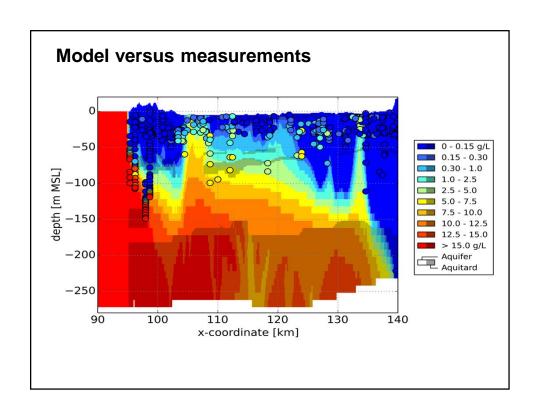




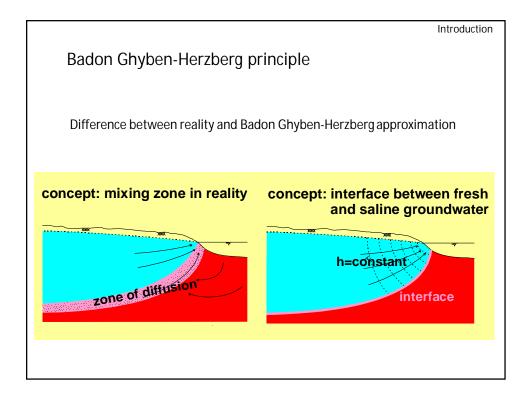


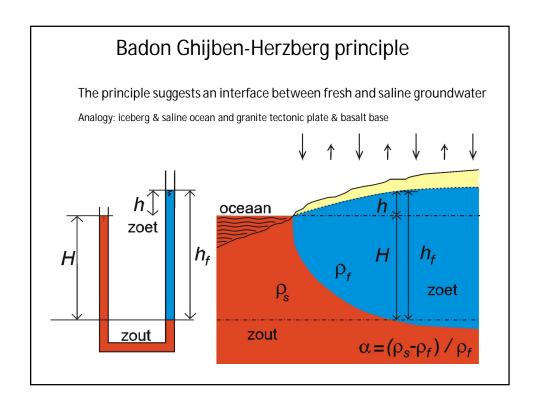


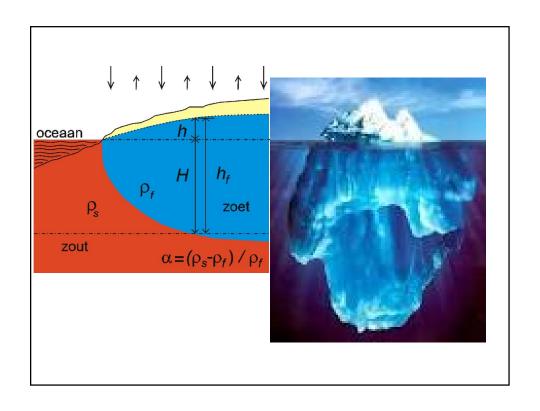




Sharp interface between fresh and saline groundwater





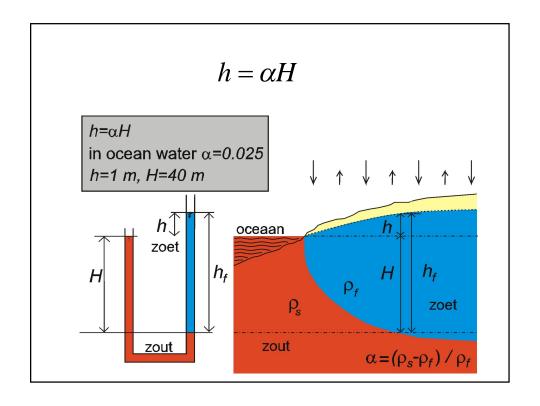


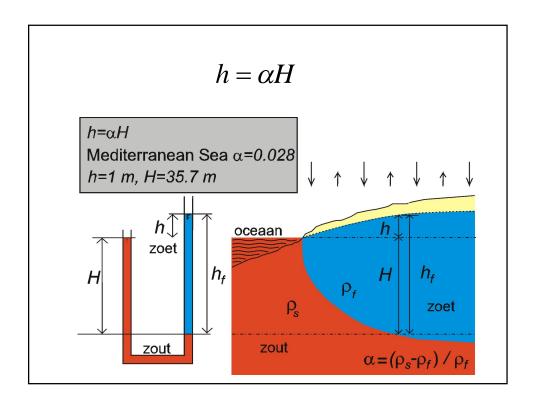
pressure saline groundwater=pressure fresh groundwater
$$\rho_s Hg = \rho_f (H+h)g$$

$$h = \frac{\rho_s - \rho_f}{\rho_f} H$$

$$h = \alpha H$$

$$h = \alpha H$$





Badon Ghyben-Herzberg principle

- gives analytical solutions (see later and lectures)
- educational
- interface is a simple approximation
- dispersion zone <10m
- relative simple geometries

Badon Ghyben-Herzberg principle

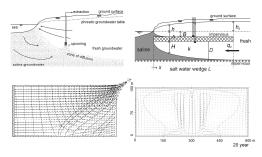
What is the case then $h\neq \alpha H$?

- 1. still dynamic situation
- 2. occurrence resistance layer
- 3. natural groundwater recharge not constant
- 4. relative density difference a is not ok
- 5. occurrence shallow bedrock
- 6. groundwater extractions

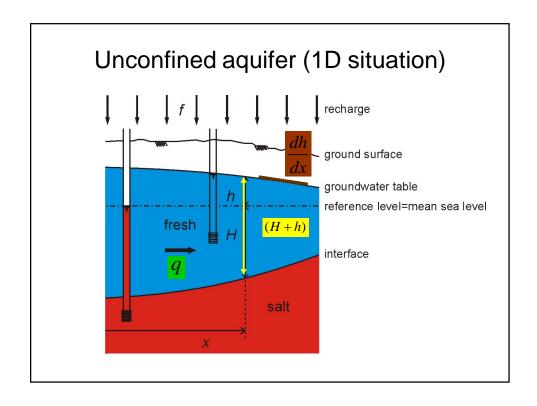
Analytical solutions

Analytical solutions

See lecture notes Density dependent groundwater flow (p. 29-48)



http://public.deltares.nl/display/FRESHSALT/Download



Unconfined aquifer (1D situation)

- (I) Darcy $q = -k(H+h)\frac{dh}{dx}$
- (II) Continuity dq = fdx
- (III) BGH $h = \alpha H$

Unconfined aquifer (1D situation)

$$dq = fdx$$
 integration $q = fx + C1$ gives

$$-k(H+h)\frac{dh}{dx} = fx + C1$$

$$h = \alpha H \rightarrow -k(H + \alpha H)\alpha \frac{dH}{dx} = fx + C1$$

$$HdH = -\frac{fx + C1}{k\alpha(1+\alpha)}dx$$

Unconfined aquifer (1D situation)

$$HdH = -\frac{fx + C1}{k\alpha(1+\alpha)}dx$$

integration gives

$$\frac{1}{2}H^{2} = \frac{-\frac{1}{2}fx^{2} - C1x + C2}{k\alpha(1+\alpha)}$$

$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}}$$

Unconfined aquifer (1D situation)

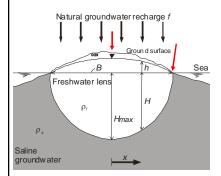
$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}}$$

$$h = \alpha H$$

$$q = fx + C1$$

Example 1: Elongated island

$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}} \qquad q = fx + C1$$

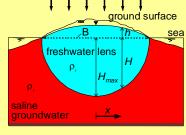


Boundary conditions

$$x = 0 : q = 0 \rightarrow C1 = 0$$

 $x = 0.5B : H = 0 \rightarrow C2 = fB^2 / 8$

Example of analytical solutions (I)



Depth of fresh-saline interface H

$$H = \sqrt{\frac{f(0.25B^2 - x^2)}{k\alpha(1+\alpha)}}$$

$$h = \alpha H$$

Maximal thickness lens

$$H_{\text{max}} = \frac{1}{2} B \sqrt{\frac{f}{k\alpha(1+\alpha)}}$$
 $V = \frac{1}{4} \pi (1+\alpha) H_{\text{max}} B n_e$

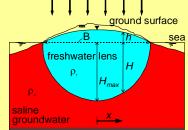
Volume lens

$$V = \frac{1}{4}\pi(1+\alpha)H_{\text{max}}Bn_e$$

Characteristic time $T = \frac{\text{volume of water in lens}}{\text{inflow of water}} = \frac{\pi n_e B}{8} \sqrt{\frac{(1+\alpha)}{kf\alpha}}$

Lecture notes p. 32

Example of analytical solutions (I)



Depth of fresh-saline interface H

B = 2000 m, f = 0.001 m/day

 $k = 10 \text{m/day}, \alpha = 0.025$

 $n_e = 0.35$

Maximal thickness lens

Volume lens (wrong in lectures notes)

$$H_{\text{max}} = 62.5 \text{m}, h_{\text{max}} = 1.56 \text{m}$$
 $V = 35203 \text{m}^3/\text{m}'$

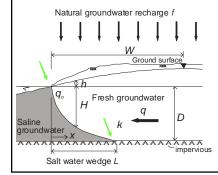
$$V = 35203 \text{m}^3/\text{m}^4$$

Characteristic time
$$T = \frac{35203}{2} days = 48.2 years$$

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Example 2: salt water wedge

$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}} \qquad q = fx + C1$$



Boundary conditions

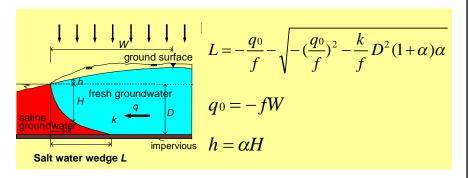
$$x = 0 : q = q_0 \rightarrow q_0 = -fW \rightarrow C1 = q_0$$

$$x = 0: H = 0 \rightarrow C2 = 0$$

Length of salt water wedge

$$x = L : H = D$$

Example of analytical solutions (II)



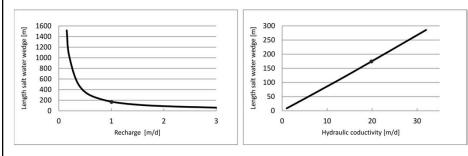
Example:

$$W = 3000 \text{m}, f = 0.001 \text{m/day}, \alpha = 0.020, k = 20 \text{m/day}, D = 50 \text{m}$$

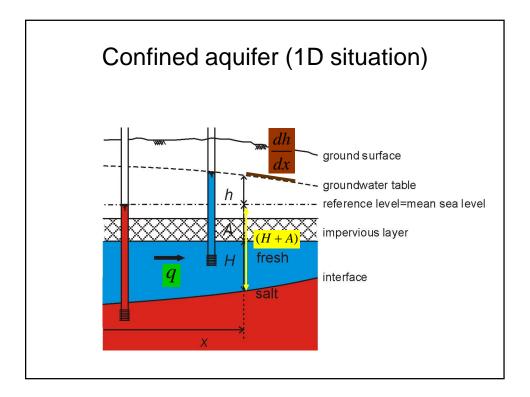
$$L = 175.1$$
m

Lecture notes p. 33

Length of the salt water wedge as a function of a. recharge and b. hydraulic conductivity



the dots resample with the example mentioned above



Confined aquifer (1D situation)

- (I) Darcy $q = -kH \frac{dh}{dx}$
- (II) Continuity $q = q_0$
- (III) BGH $h = \alpha(H + A)$

Confined aquifer (1D situation)

$$-kH\frac{dh}{dx} = q_0$$

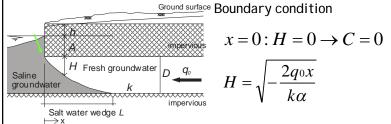
$$HdH = -\frac{q_0}{k\alpha}dx$$

integration gives
$$\frac{1}{2}H^2 = \frac{q_0x}{k\alpha} + C$$

$$H = \sqrt{-\frac{2q_0x}{k\alpha} + 2C}$$

Example 3: salt water wedge confined aquifer

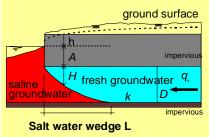
$$H = \sqrt{-\frac{2q_0x}{k\alpha} + 2C}$$



Length of salt water wedge x = L: H = D

$$L = -\frac{kD^2\alpha}{2q_0}$$

Example of analytical solutions (III)



Length of salt water wedge

$$H = \sqrt{-\frac{2q_0x}{k\alpha}}$$

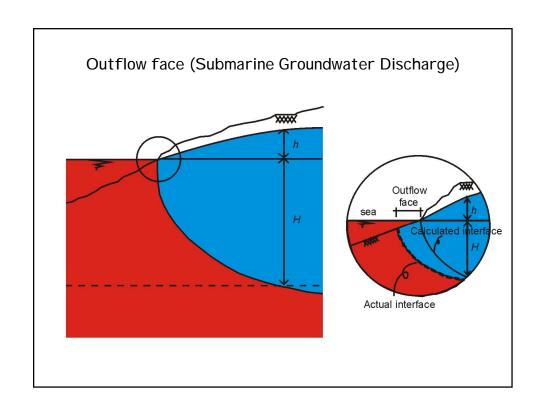
$$L = -\frac{kD^2\alpha}{2q_0}$$

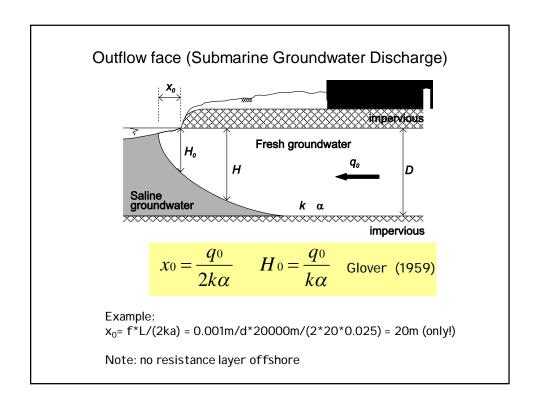
Example:

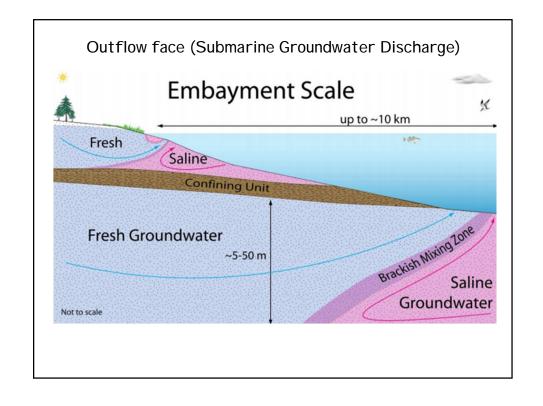
$$W = 2000 \text{m}, f = 0.001 \text{m/day}, \alpha = 0.025, k = 25 \text{m/day}, D = 40 \text{m}$$

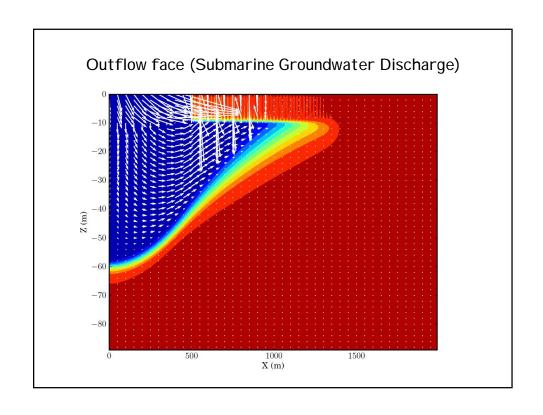
$$L = 250 \text{m}$$

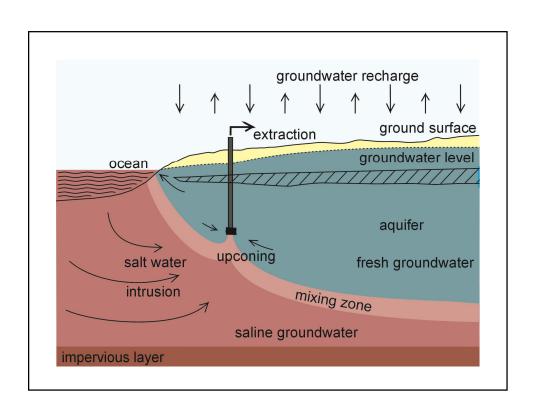
Lecture notes p. 35-36

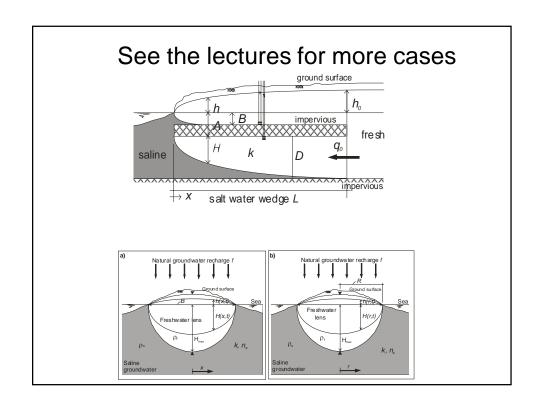




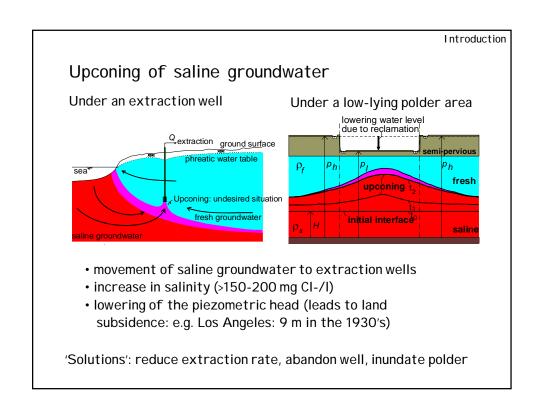


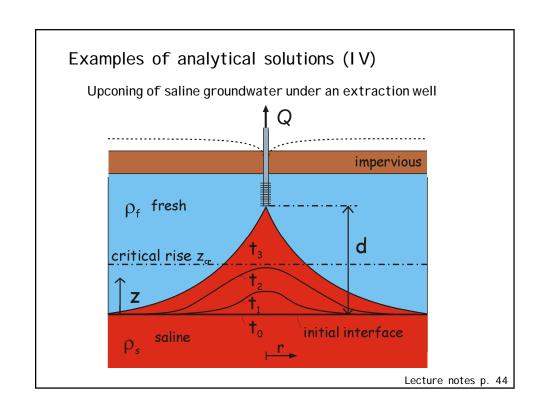


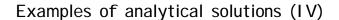




Upconing processes







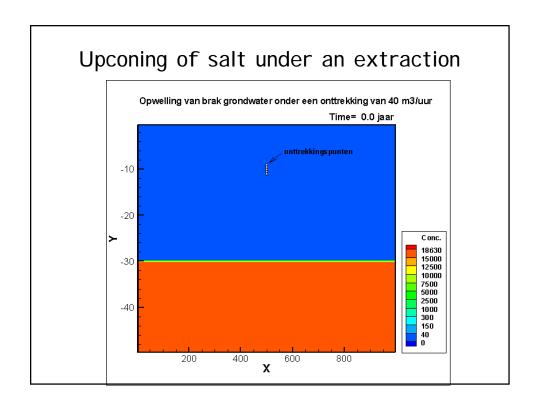


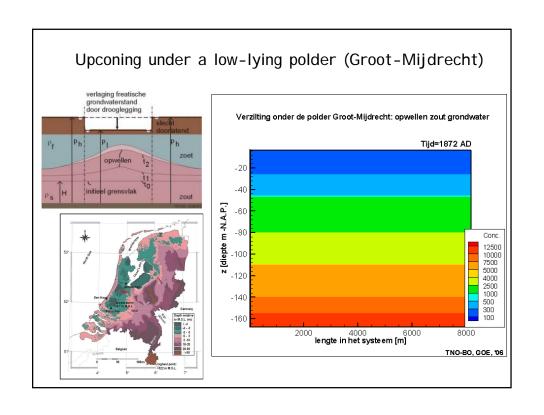
Upconing of saline groundwater under an extraction well

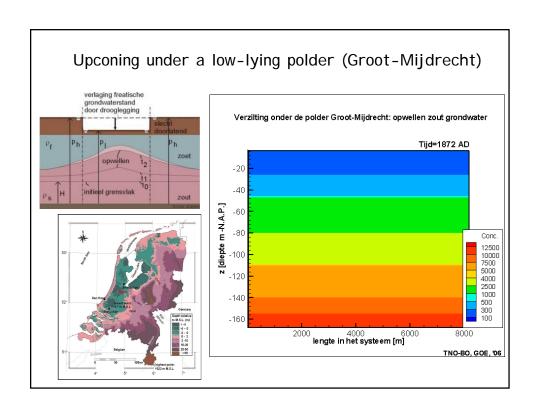
$$\begin{split} z(r,t) &= \frac{Q}{2\pi\alpha k_x d} \left[\frac{1}{(1+R'^2)^{1/2}} - \frac{1}{[(1+\gamma')^2 + R'^2]^{1/2}} \right] \\ R' &= \frac{r}{d} \frac{k_z}{k_x}^{1/2} \qquad \gamma' = \frac{\alpha k_z}{2n_e d} t \end{split}$$

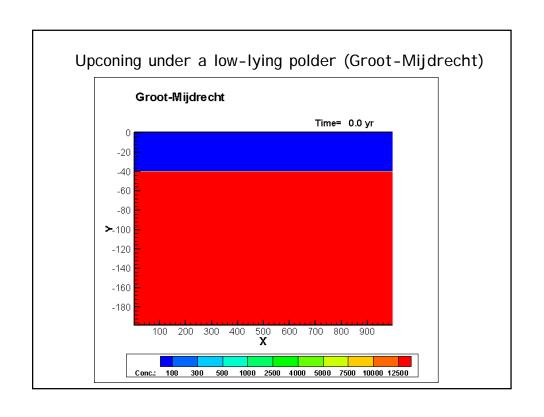
Dagan & Bear, 1968, J. Hydraul. Res 6, 1563-1573

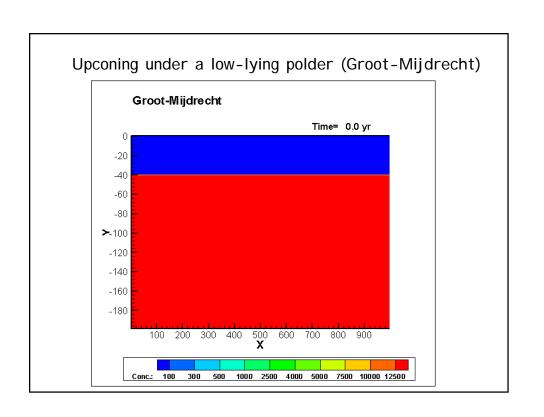
Lecture notes p. 44

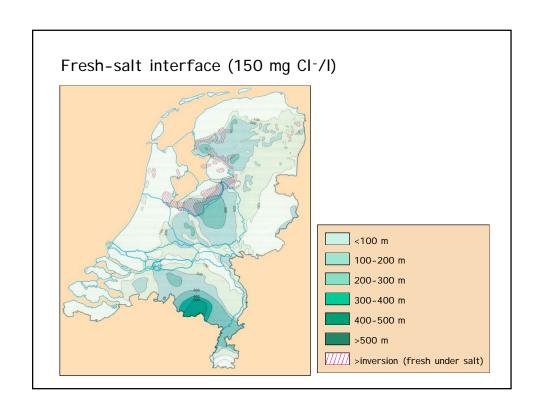


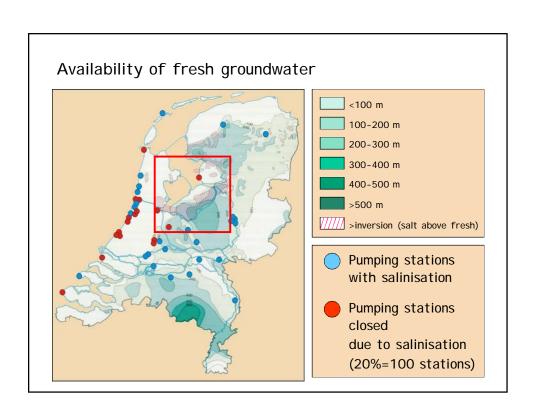


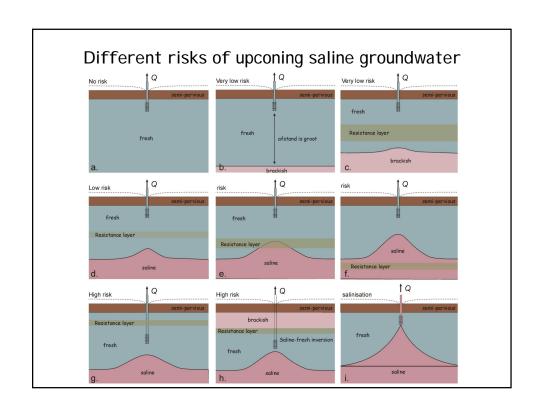


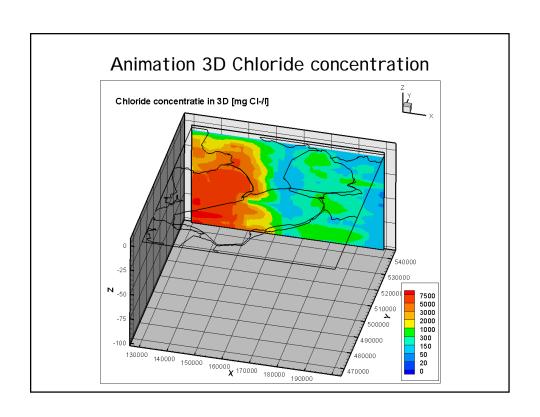


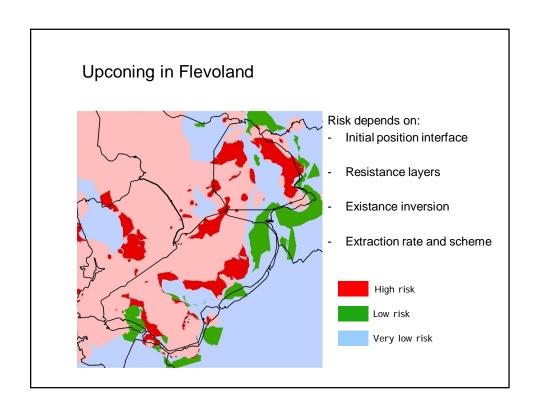












Compensating measures

Possible solutions to stop salt water intrusion:

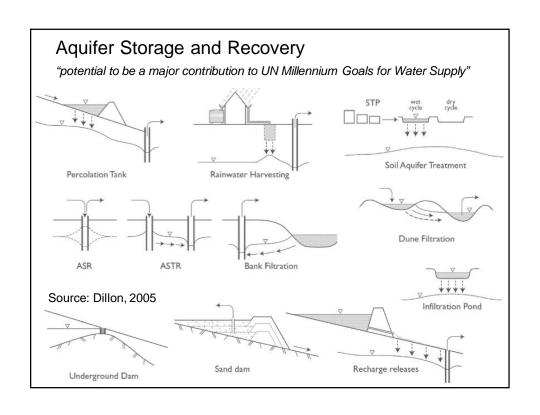
- Restriction of groundwater extractions through permits
- Co-operation between authorities and water users
- Desalinisation of saline water
- Technical countermeasures of salt water intrusion
 - six examples

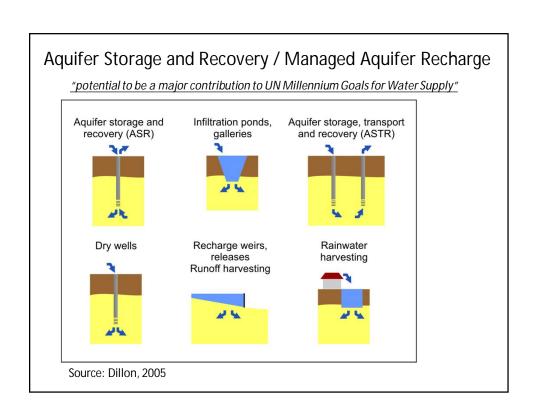
Tools to understand salt water intrusion:

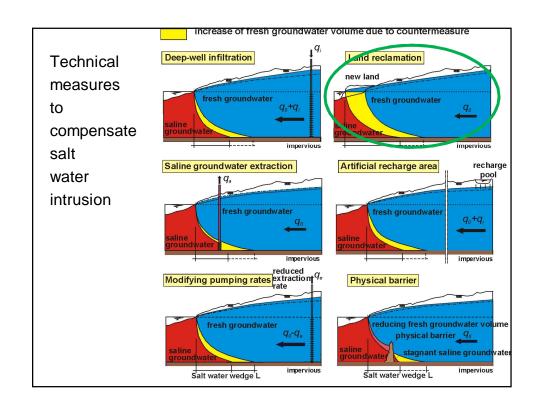
- Monitoring of salinities and piezometric levels
- Numerical modelling of salt water intrusion

Measures to compensate salt water intrusion

- 'The Fresh Holder'
- Extraction of saline/brackish groundwater
- Infiltration of fresh surface water
- Modifying pumping rates
- · Land reclamation in front of the coast
- Creating physical barriers (chrystallisation or biosealing)

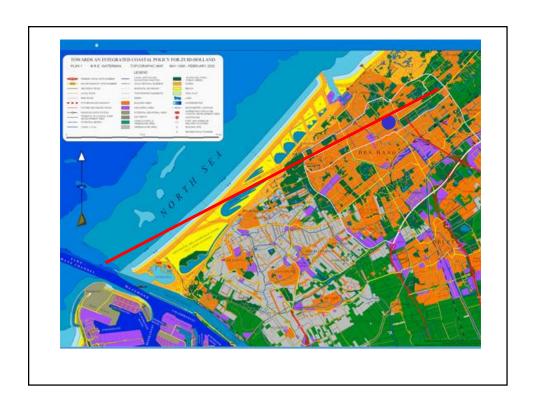




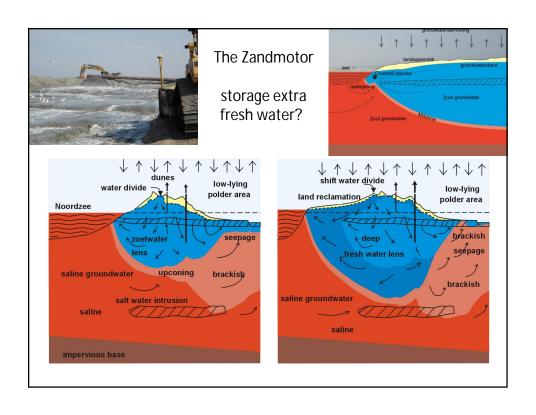


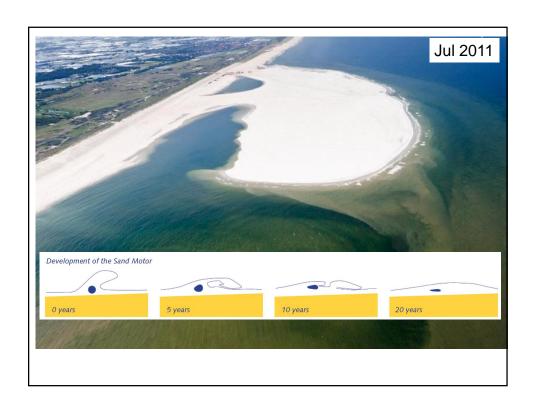


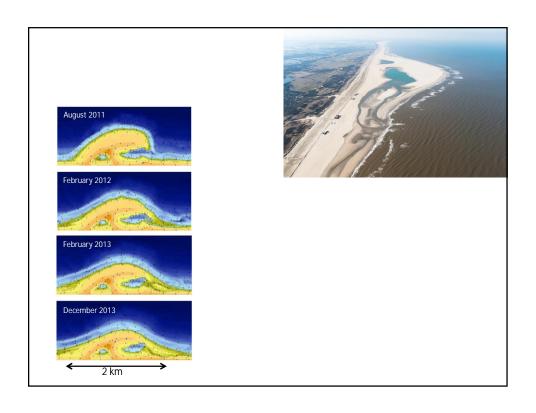


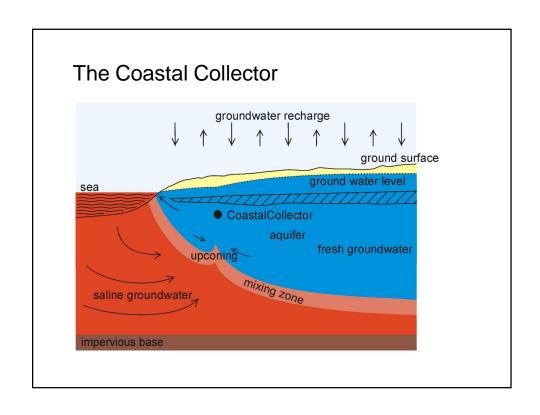


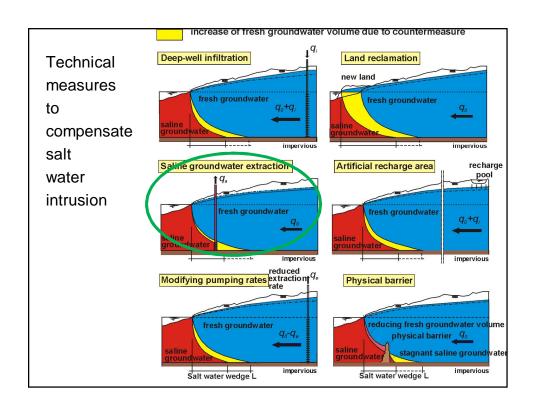


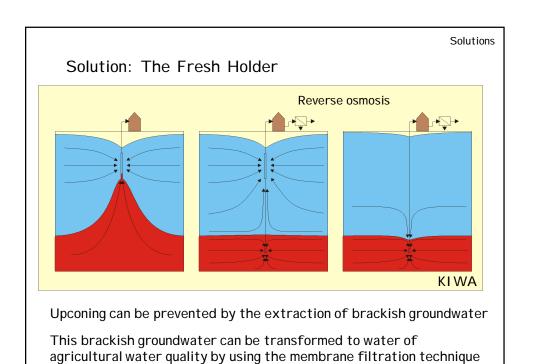


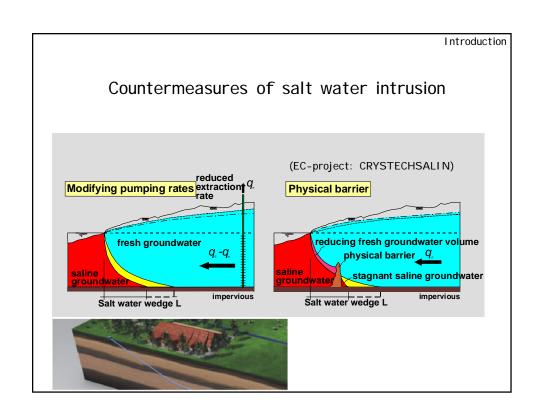


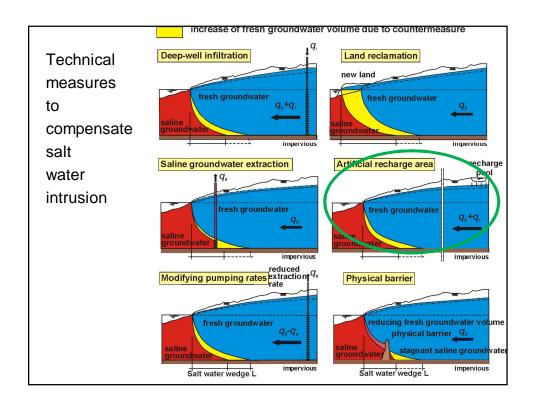






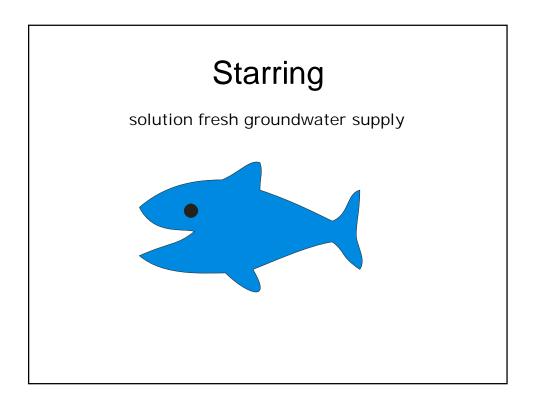






Base idea

Many local solutions for fresh groundwater supply can have regional impact



Starring

Local solution fresh groundwater supply



Starring

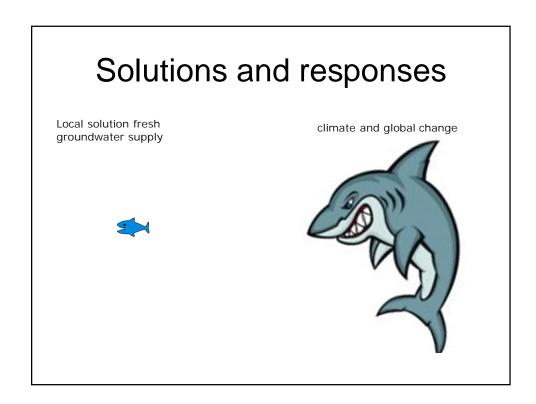
climate and global change



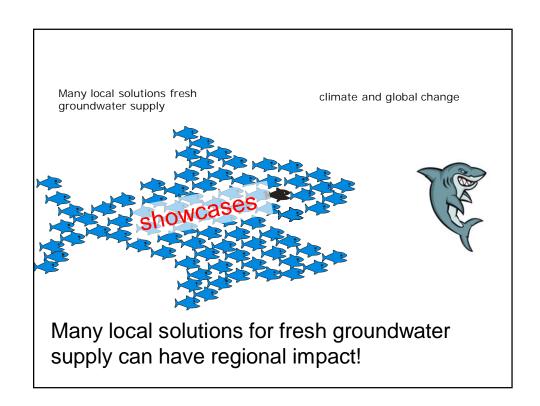
Starring

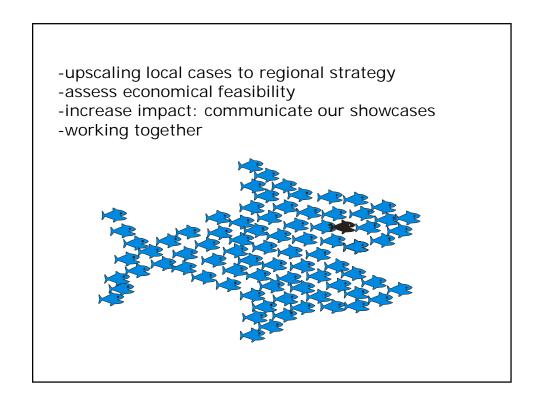
climate and global change

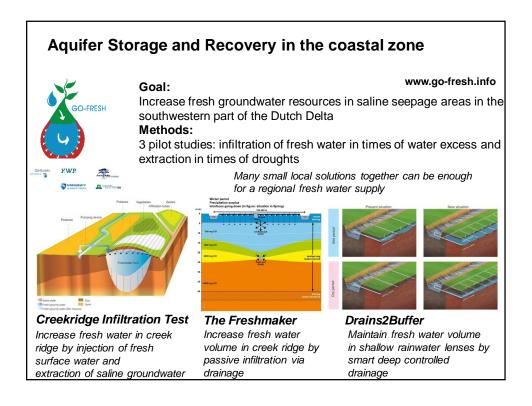




What should be the response?





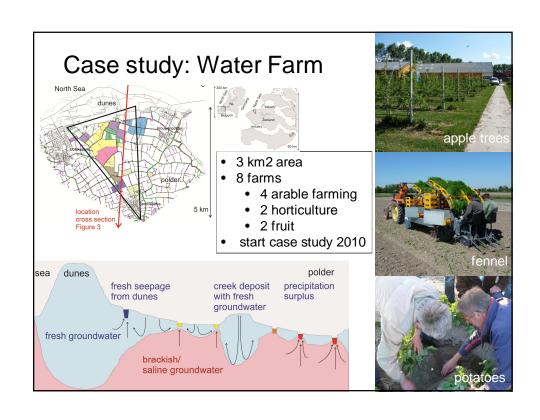


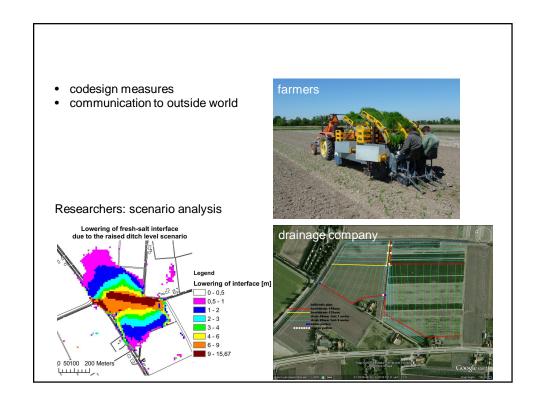
Problem statement

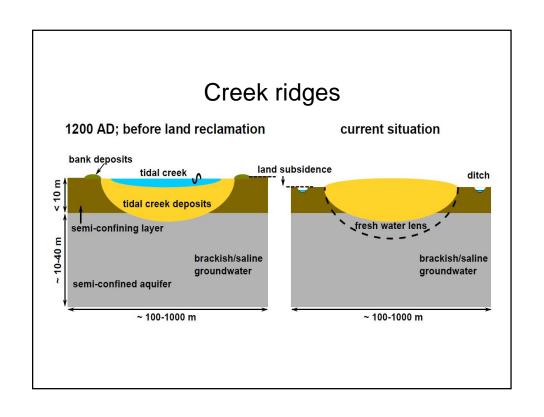
- Crop damage southwestern part of the Netherlands
- Fresh groundwater below creek ridges

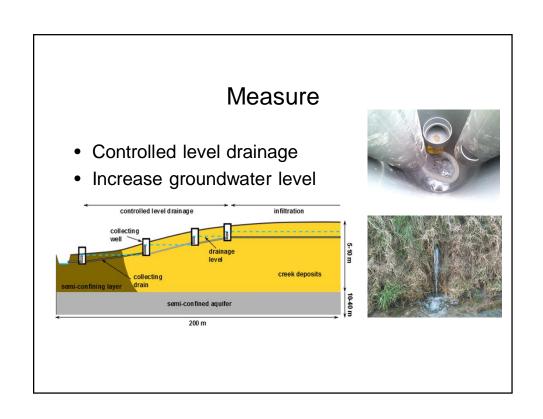


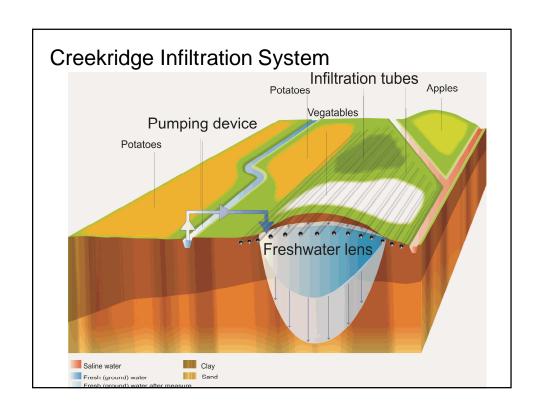


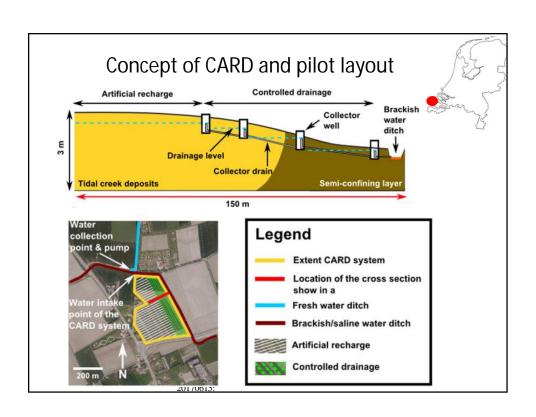




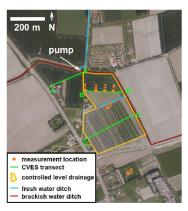








Installation of drainage and monitoring network









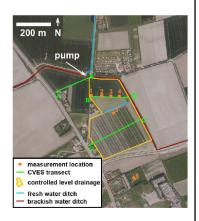


various types of field measurements

Different types of field measurements applied

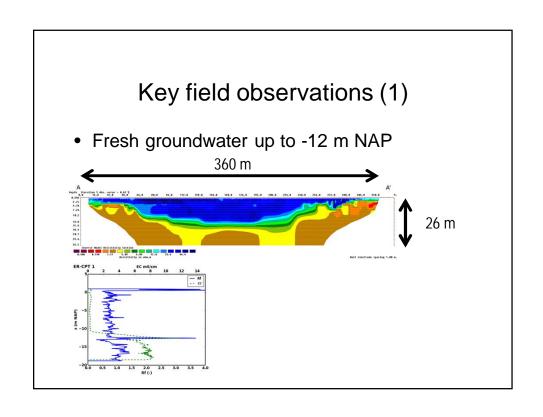
 $\begin{tabular}{lll} Measurement type & Purpose \\ \hline Pressure transducers^a & Groundwater levels \\ Sampling using & EC_{w20} \\ piezometer nest \\ SLIMFLEX^b & EC_{bulk} \\ CPT^c & Lithology and EC_{bulk} \\ CVES^d & EC_{bulk} \\ SMD^e & EC_{bulk} \\ \hline \end{tabular}$

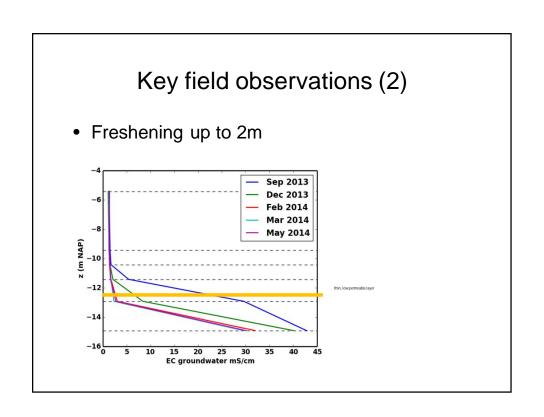
- a. Schlumberger, The Netherlands (type 'Diver')
- b. Deltares, The Netherlands
- · c. Fugro, The Netherlands
- · d. ABEM, Sweden
- e. Imageau, France

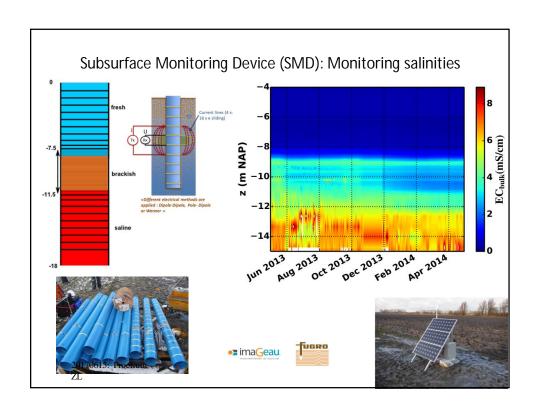


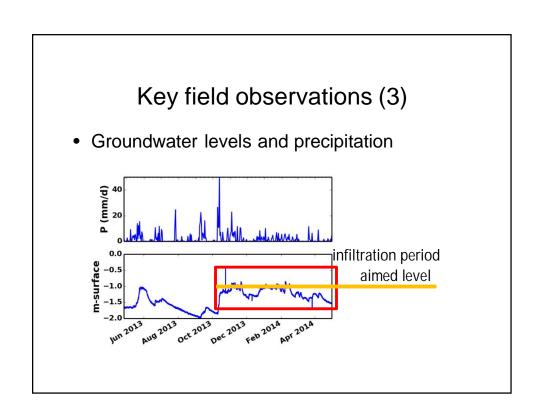


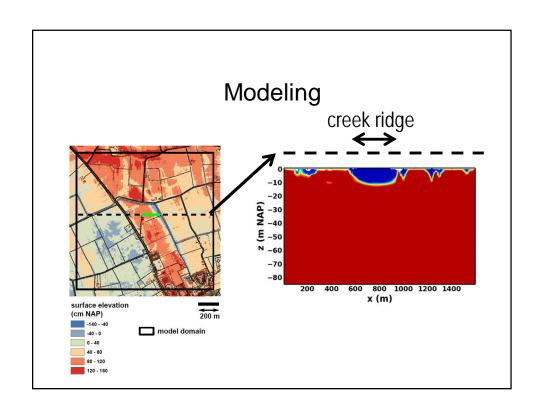


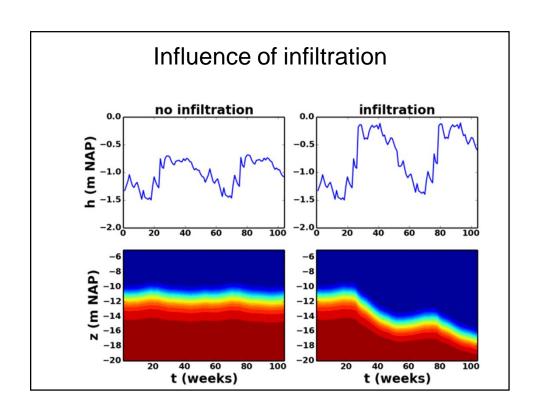








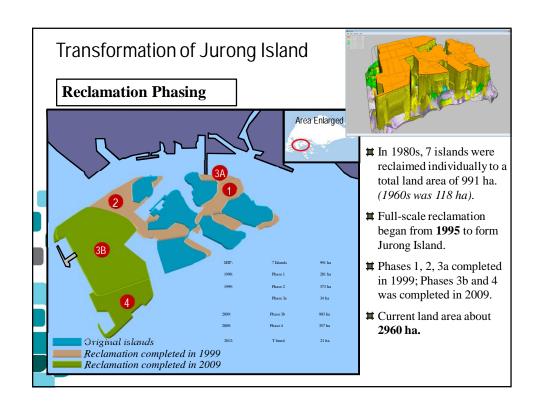


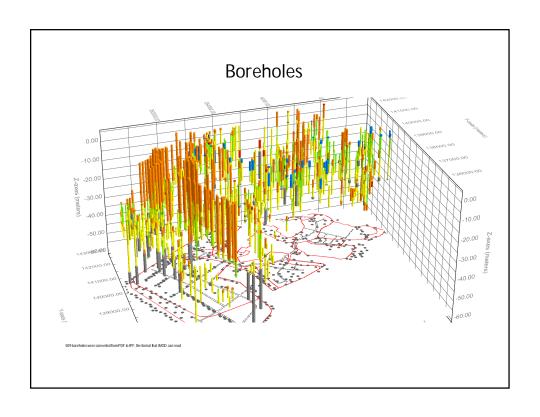


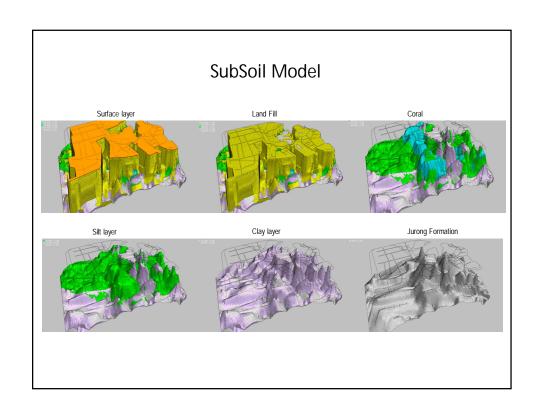
Singapore Jurong Island

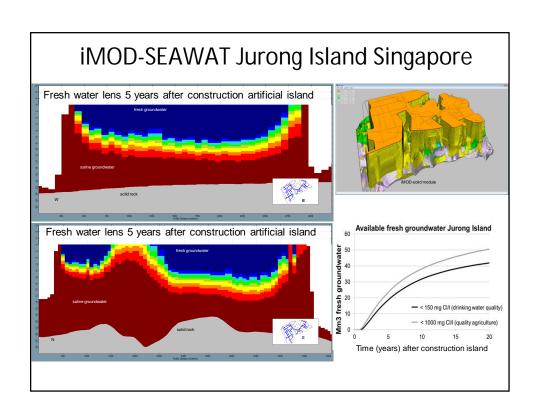
Aquifer Storage and Recovery

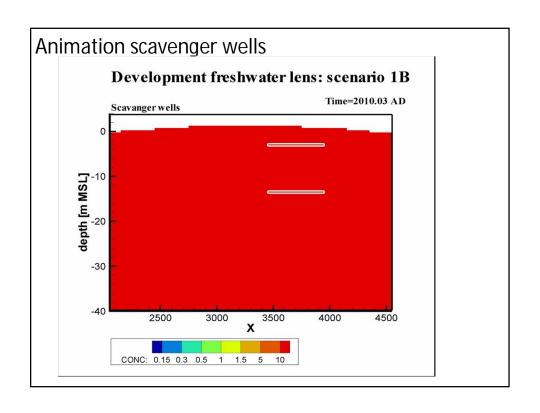


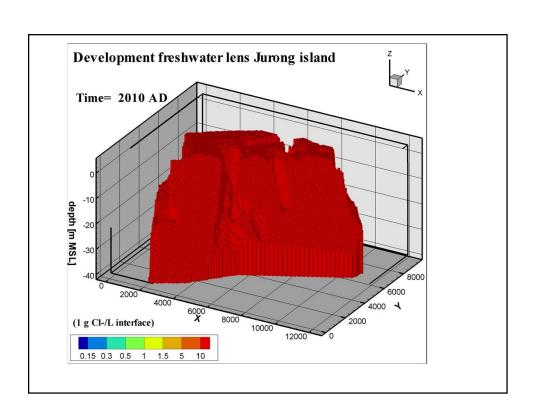


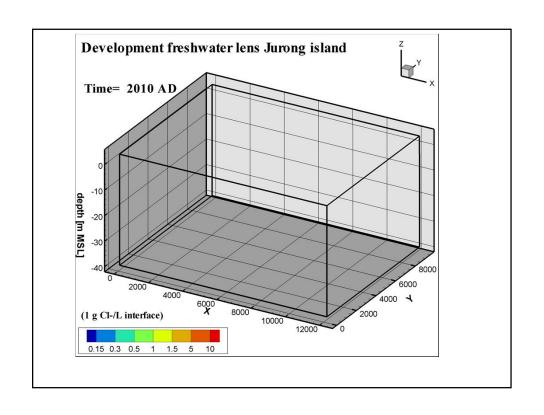


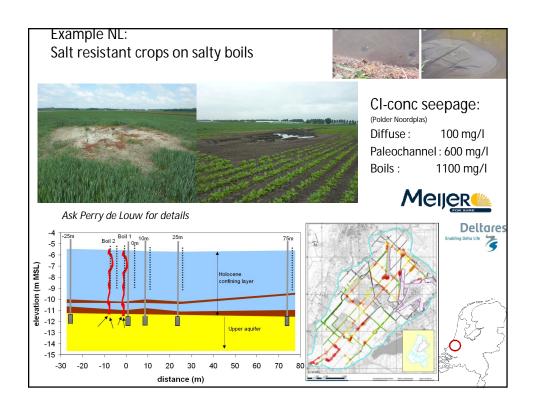


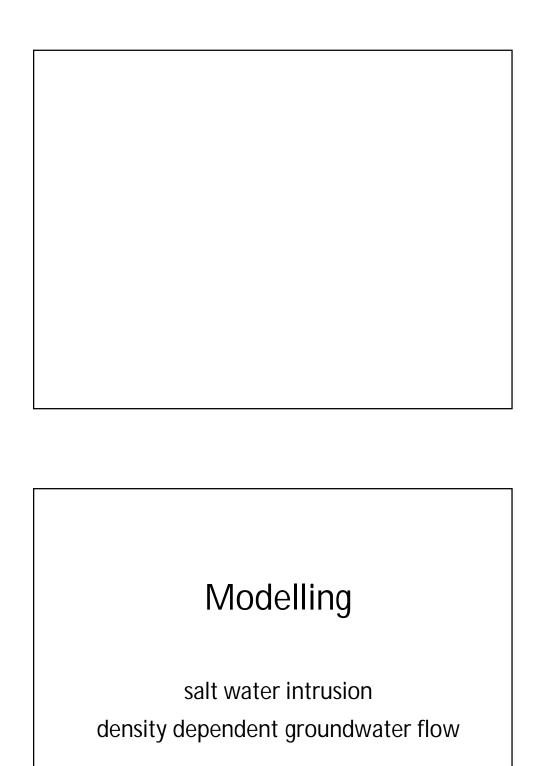












Why mathematical modelling anyway?

A model is only a schematisation of the reality!

modelling

Why mathematical modelling anyway?

+:

- · cheaper than scale models
- analysis of very complex systems is possible
- a model can be used as a database
- to increase knowledge about a system (water balances)

-:

- simplification of the reality
- · only a tool, no purpose on itself
- garbage in=garbage out: (field)data important
- perfect fit measurement and simulation is suspicious

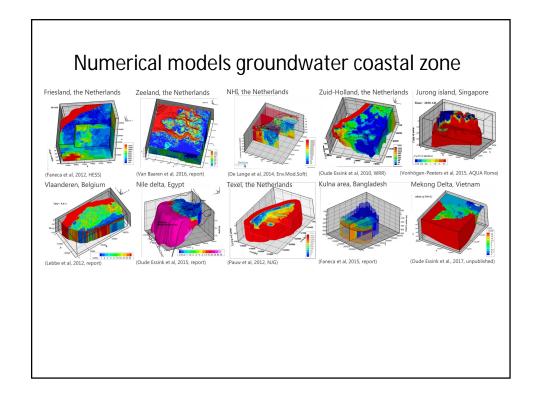
Numerical modelling variable density flow

Type:

- sharp interface models
- solute transport models

State of the art:

- three-dimensional
- solute transport
- transient



Some existing 3D codes which simulate variable density groundwater flow in porous media:

SEAWAT (Guo & Bennett, 98) METROPOL (Sauter, '87) FEFLOW (Diersch, '94) MVAEM (Strack, '95) D3F (Wittumet al., '98) MOCDENS3D (Oude Essink, '98) HydroGeoSphere (Therrien, '92) SWICHA (Huyakorn et al., '87) SWIFT (Ward, '91) FAST-C 3D (Holzbecher, 98) MODFLOW+MT3D96 (Gerven, '98) HST3D (Kipp, '86) SUTRA (beta-version, Voss, '02)

Restrictions 3D salt water intrusion modelling

- the data problem:
 - -not enough hydrogeological data available
 - -e.g. the initial density distribution
 - -especially important issue in data-poor countries
- the computer problem:
 - -modelling transient 3D systems: computer only good enough at high costs
- the numerical dispersion problem:
 - -numerical dispersion is large in case of coarse grid

Restrictions 3D salt water intrusion modelling now

- the data problem:
 - -not enough hydrogeological data available
 - -e.g. the initial density distribution
 - -especially important issue in data-poor countries
- the computer problem:
 - modelling transient of bits computer only good solution is
- the nume ical dispersion problemsolvers
 -numerical dispersion problemsolvers
 -numerical dispersion problemsolvers
 -numerical dispersion problemsolvers

variable density

Stability criteria for solute transport equation (I)

1. Neumann criterion:

$$\frac{D_{xx}\Delta t_s}{\Delta x^2} + \frac{D_{yy}\Delta t_s}{\Delta y^2} + \frac{D_{zz}\Delta t_s}{\Delta z^2} \le 0.5$$

$$\Delta t_s \le \frac{0.5}{\frac{D_{xx}}{\Delta x^2} + \frac{D_{yy}}{\Delta y^2} + \frac{D_{zz}}{\Delta z^2}}$$

variable density

Stability criteria for solute transport equation (II)

2. Mixing criterion:

$$\Delta t_s \le \frac{n_e b_{i,j,k}^k}{Q_{i,j,k}'}$$

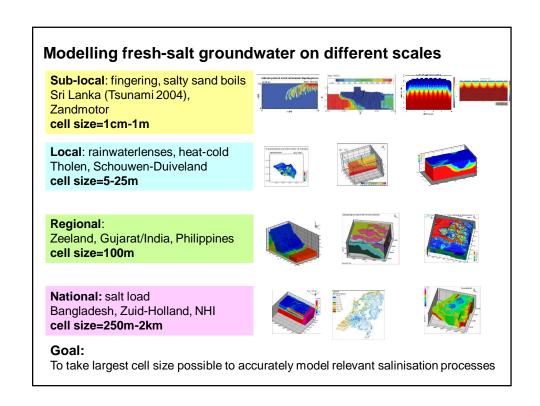
Change in concentration in element is not allowed to be larger than the difference between the present concentration in the element and the concentration in the source

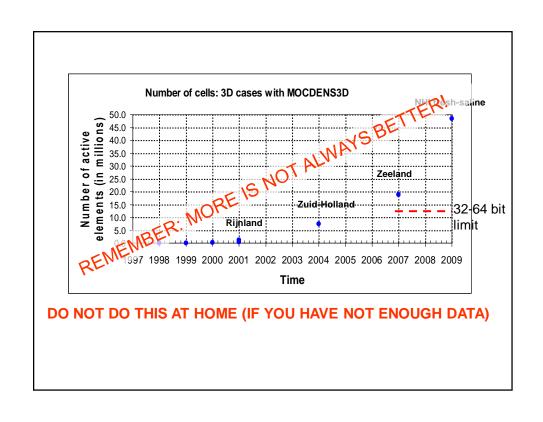
variable density

Stability criteria for solute transport equation (III)

3. Courant criterion:

$$\Delta t_{s} \leq \frac{\xi \Delta x}{V_{x,\text{max}}} \qquad \Delta t_{s} \leq \frac{\xi \Delta y}{V_{y,\text{max}}} \qquad \Delta t_{s} \leq \frac{\xi \Delta z}{V_{z,\text{max}}}$$





Modelling effect climate change on fresh-salt groundwater

Modelling:

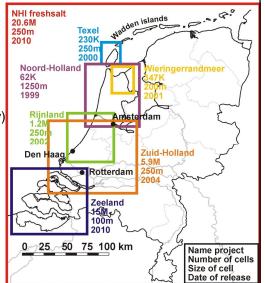
- variable-density
- 3D, non-steady
- groundwater flow
- coupled solute transport

Code:

MOCDENS3D (MODFLOW family) similar to SEAWAT

Assessing effects:

- autonomous salinisation
- sea level rise
- · changing recharge pattern
- land subsidence
- changing extraction rates
- · adaption measures



Fields of application of fresh-saline groundwater models

- Water system analysis in brackish-saline environments (salt loads, salt boils, freshwater lenses)
- Quantifying effects of climate change & sea level rise
- Drinking water issues: upconing saline groundwater under extraction wells
- Developing measurements to stop salinization groundwater systems (e.g. fresh keeper, coastal collectors, freshwater storage underground)
- Impact of the disasters as tsunamis on fresh groundwater resources
- Submarine Groundwater Discharge (marine water pollution, Harmful Algae)

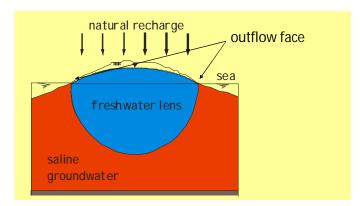
variable density

Difficulties with variable density groundwater flow

- Initial density distribution (effects on velocity field)!
- Velocities freshwater lens at the outflow face near the sea
- Boundary conditions (especially concentration boundaries)
- Choice of element size
- Length of flow time step to recalculate groundwater flow

problems

Outflow face at the coast is difficult to model



Flow converges and thus velocities are very high at the outflow face

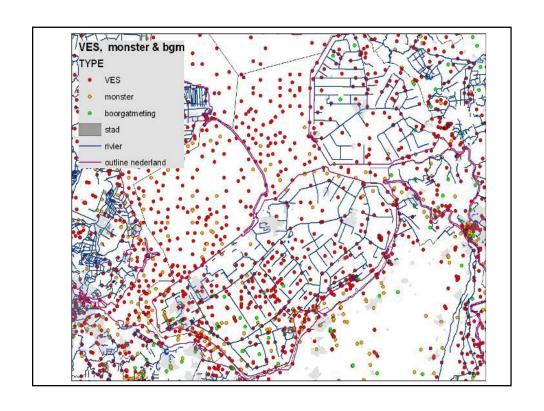
This is numerically difficult to handle

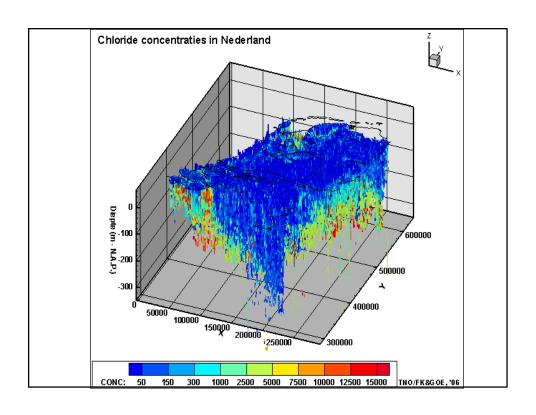
A good initial density distribution is essential

- Because groundwater and solute transport are coupled, the density influences grondwater velocities
- Numerous density measurements are necessary to get a reliable 3D density matrix

'Procedure' to improve initial density distribution

- · Implement all chloride data
 - Analyses, Borehole, VES, Airborne techniques (HEM, SkyTem)
 - Better old then nothing
 - Better VES then nothing
- Interpolate and extrapolate
 - Sea = easy (salt)
 - Inland = fresh?
- Start with simulation (10/20/30 years) with mol.diffusion*1000 to smooth out artificial densities

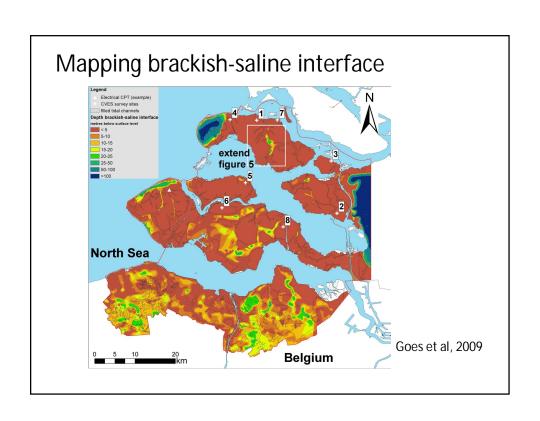


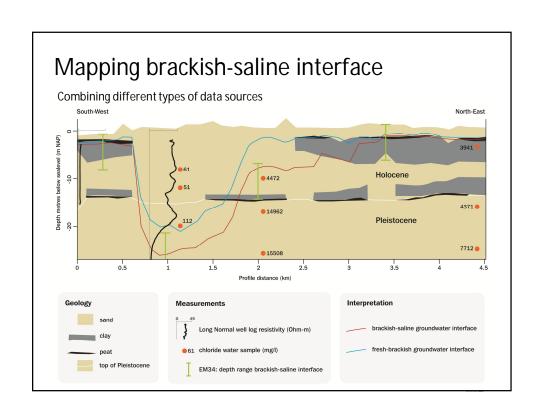


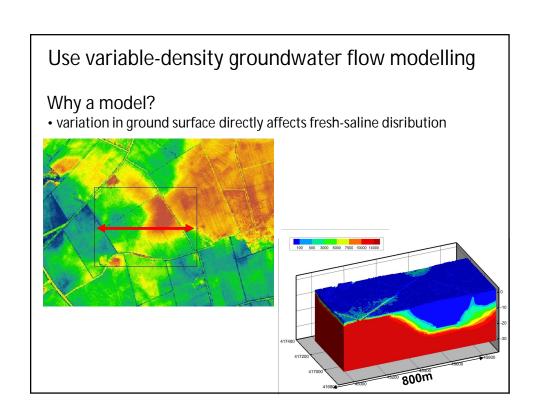
Mapping brackish-saline interface Zeeland

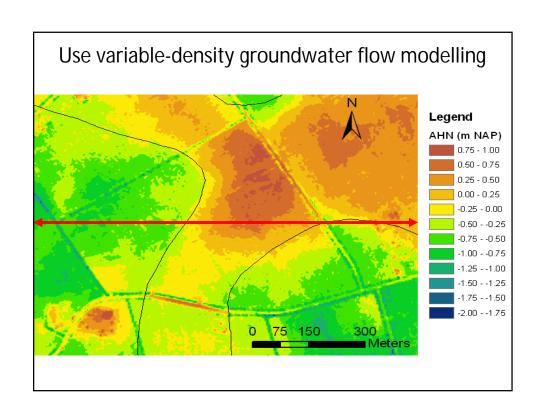
Combining different types of data sources:

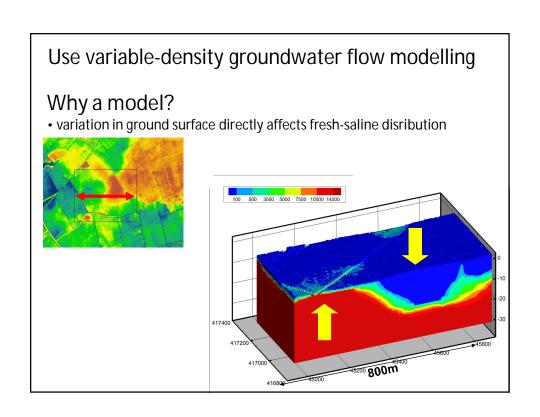
Data type	Characteristics of measurement	# Data	Determined	Accuracy depth of interfaces
Groundwater Samples	0D in situ	721	Chloride concentration	Depends on positions of screens
Geo-electrical borehole logs	1D in situ	149	1D chloride profile, Depth fresh-brackish and brackish-saline interface, Inversions.	±1 m
Electrical CPT	1D in situ (max. depth 50 m)	71	Borehole log	±1 m
VES	1D from surface	1113	Depth brackish-saline interface, Major inversions, (1D chloride profile).	±20% of depth
EM34	1D from surface	3251	Depth brackish-saline interface	ranges of 7.5, 15 or 30 m (accuracy decreases with depth)
Groundwater Abstractions	0D in situ	716	Depth brackish-saline interface	a range depending on screen depth
Unique locations		6021		

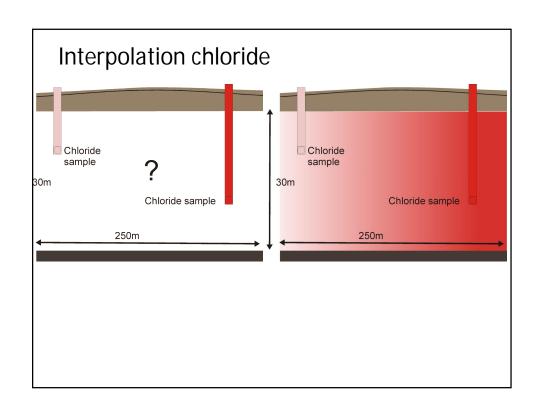


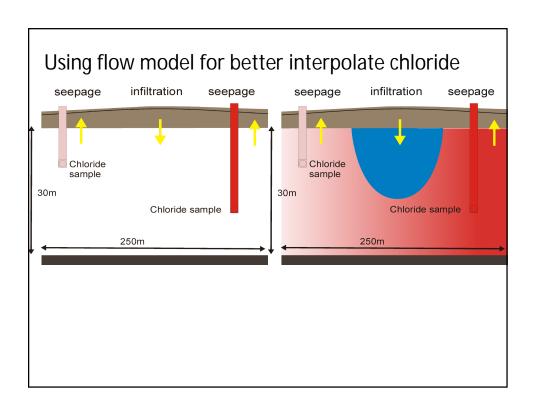


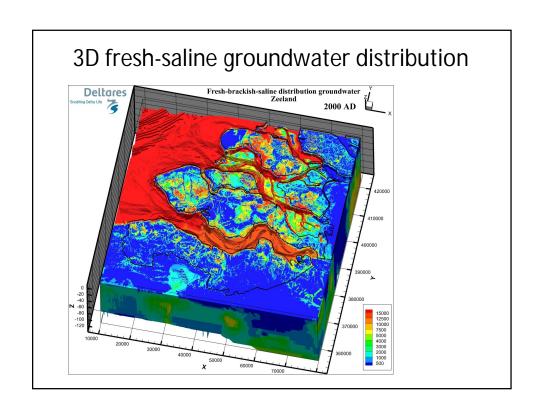


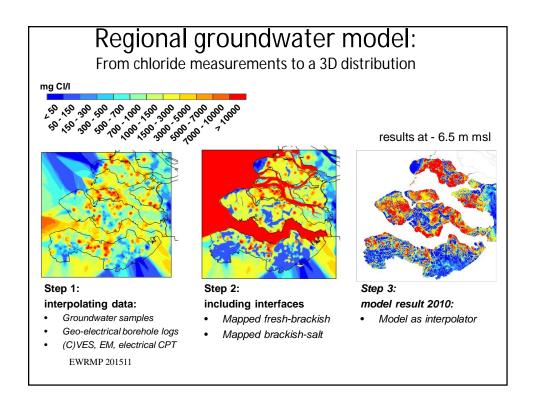












Examples of variable-density groundwater flow

- Rotating immiscible interface
- Henry's problem
- Evolution freshwater lens
- Hydrocoin
- Salt water pocket
- Broad 14 Basin, North Sea
- Heat transport: Elder and Rayleigh=4000
- 5 Dutch 3D cases
- Freshwater lenses
- Effect of Tsunami on groundwater resources

