IHE 2019

Density dependent groundwater flow in the coastal zone

Gualbert Oude Essink, PhD

Lecture set-up:

- · PowerPoint sheets
- Lecture Notes
- Practicals numerical modelling

http://freshsalt.deltares.nl





Deltares
Unit Subsurface and Groundwater Systems
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20-24-25-26 June 2019

Introduction

Curriculum Vitae

- Delft University of Technology, Civil Engineering: till 1997
 Ph.D.-thesis: Impact of sea level rise on groundwater flow regimes
- Utrecht University, Earth Sciences: till 2002
- Free University of Amsterdam, Earth Sciences: till 2004
- Deltares
- · Utrecht University (Associate Professor): from 2014

Qualifications:

- Groundwater resources management
- Density-dependent groundwater flow and coupled solute transport
- Salt water intrusion in coastal aquifers
- Assessment of climate change on groundwater resources
- Numerical Modeling
- Teaching and training

http://freshsalt.deltares.nl Deltares: qualbert.oudeessink@deltares.nl

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Colleagues at Deltares Groundwater in the Coastal Zone

http://zoetzout.deltares.nl http://freshsalt.deltares.nl







Gualbert Oude Essink Joost Delsman

Pieter Pauw













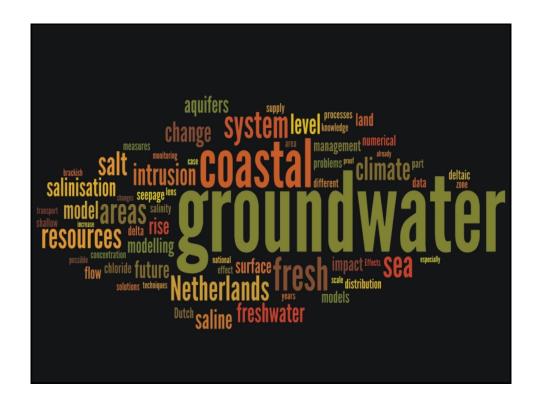
Sandra Galvis

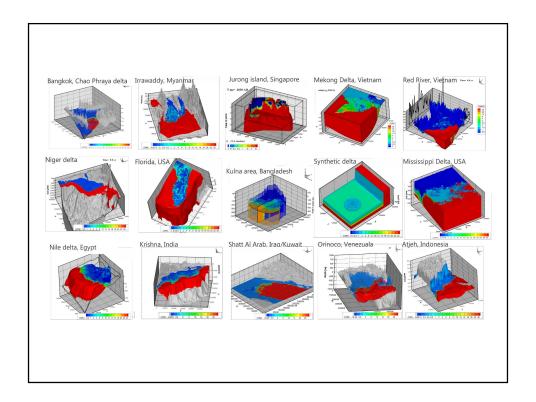
Perry de Louw

Esther van Baaren

Jarno Verkaik

Marta Faneca





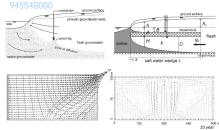
Research on groundwater in the coastal zone

- 20 years experience in modelling variable-density dependent groundwater flow and coupled solute transport in the coastal zone
- Incorporating monitoring campaigns results in numerical modeling tools
- Research on new fresh-saline phenomenae: salty seepage boils and shallow freshwater lenses in saline environments
- Knowledge on creating 3D initial chloride distribution, based on geostatistics and geophysical data (analyses, VES, borehole measures, AEM)
- Quantifying effects of climate change and sea level rise on fresh groundwater resources
- Developing adaptive and mitigative measures to stop salinization in the coastal groundwater system (e.g. ASR, MAR: fresh keeper, coastal collectors, freshwater storage underground)

Lecture notes, practicals and ppt on freshsalt.deltares.nl

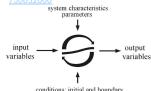
1. Density dependent groundwater flow

http://publicwiki.deltares.nl/download/attachments/22183944/gwm2.pdf?version=1&modificationDate=126



2. Groundwater modelling

http://publicwiki.deltares.nl/download/attachments/22183944/gwm1.pdf?version=1&modificationDate=1268



conditions: initial and boundary
http://publicwiki.deltares.nl/display/FRESHSALT/Upload

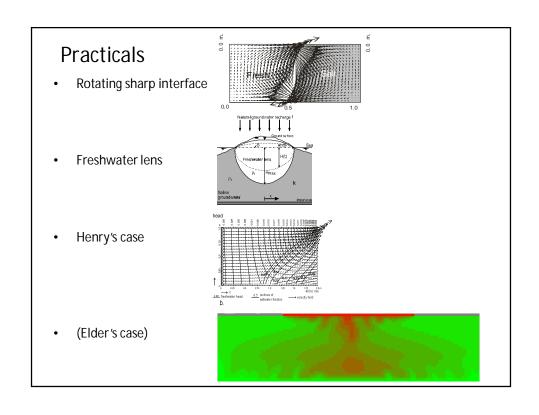
Introduction

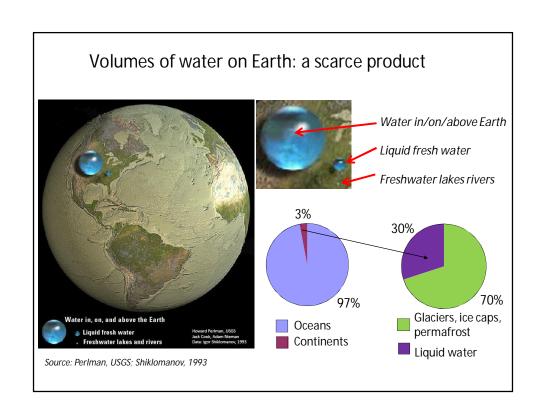
Practicals numeriacl modelling

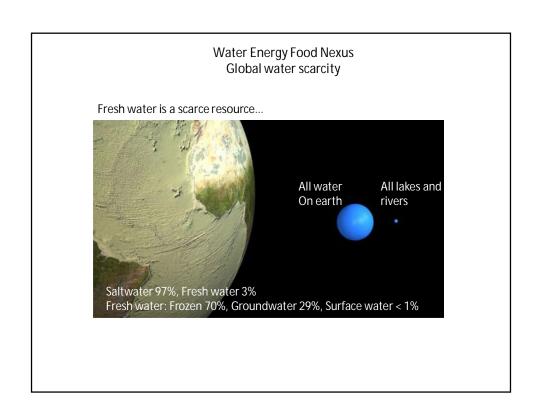
- PMWIN
- SEAWAT
- Cases:
 - Rotating sharp interface
 - Freshwater lens
 - Henry's case
 - (Elder's case)
- Setup practicals:
 - work in small groups of two persons
 - short report of findings (make screenshots)
 - deliver within one week after finish last lectures

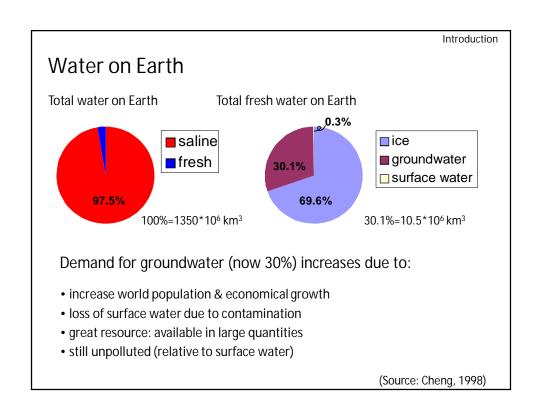
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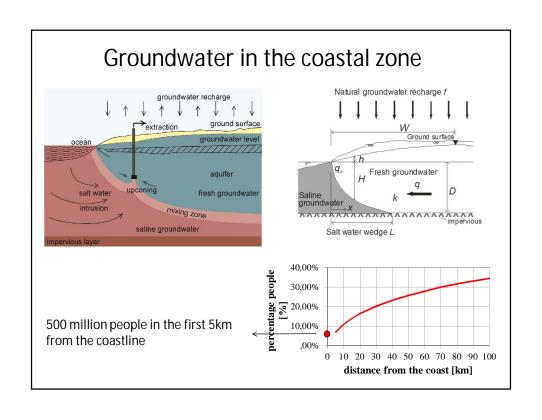
Deltares: gualbert.oudeessink@deltares.nl







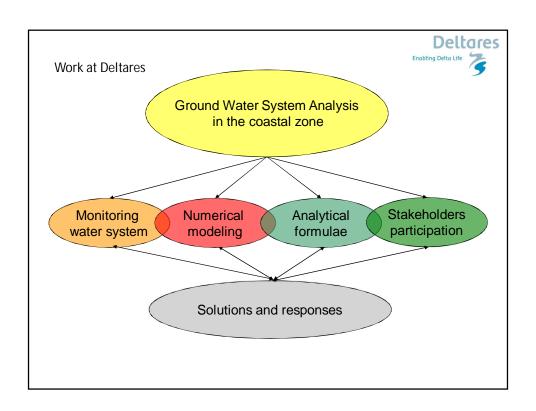


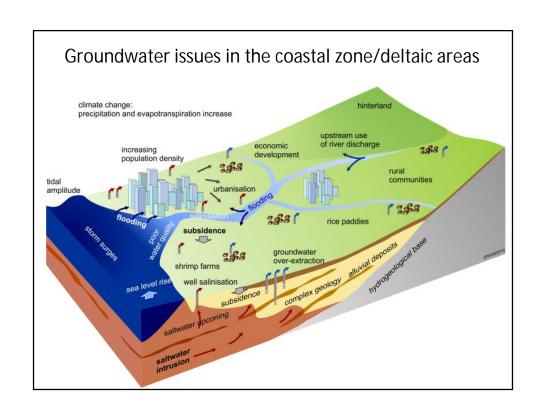


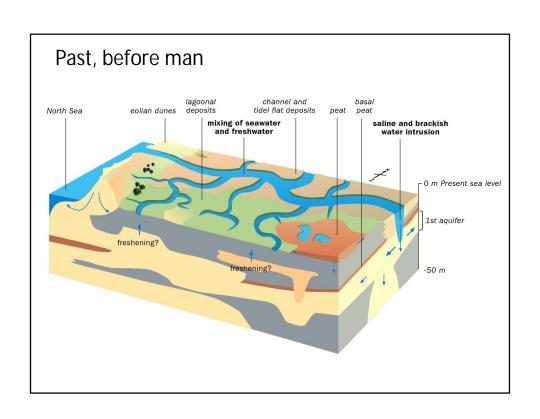
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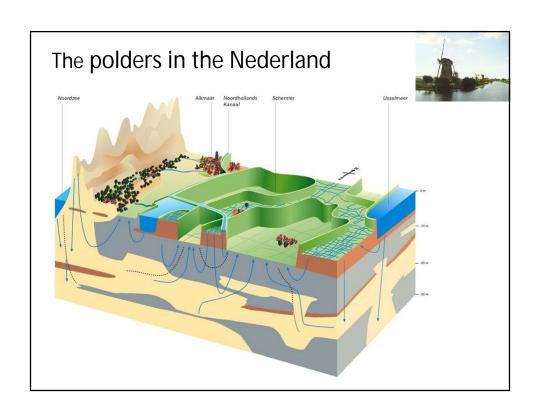
Topics of density driven groundwater flow

- 1. Introduction
 - water on earth
 - salt water intrusion
 - freshwater head
- 2. Interface between fresh and saline groundwater
 - analytical formulae (Badon Ghyben-Herzberg)
 - upconing example
- 3. Numerical modelling
 - mathematical background
 - Benchmark problems: Henry, Elder, Hydrocoin, etc.
- 4. Case-studies
 - hypothetical cases
 - 2D, 3D cases
 - real cases (Dutch coastal zone)









Groundwater in the future

We have to cope which...:

- We have to cope which...:
- Groundwater extractions
- Development energy use/production (heat-cold)
- Climate change
- Land subsidence
- Development spatial land use
- Politics, Policy & Watermanagement

Direct anthopogenic influence on groundwater is more important than climate effect

Salt Water Intrusion Meeting, since 1968

Salt Water Intrusion Meeting, since 1968



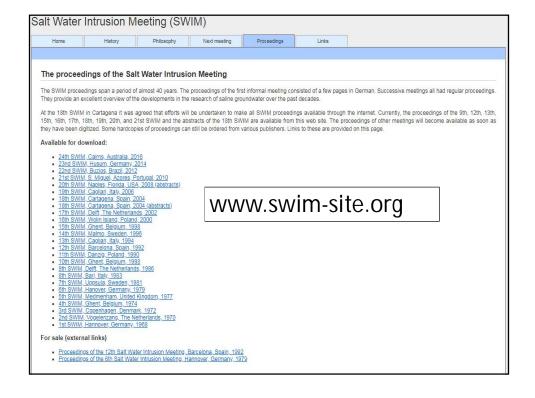
http://www.swim-site.org/

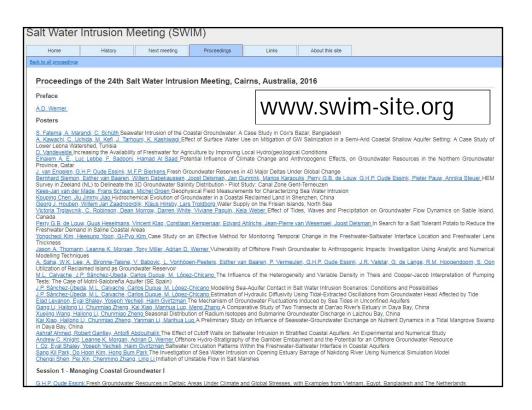
Themes

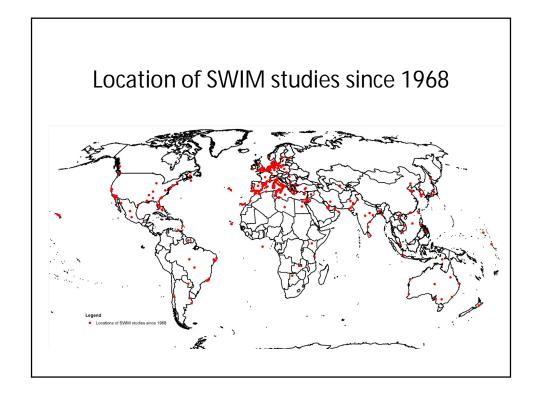
- Water systeem analysis
- Monitoring
- Modelling
- Effects
- Solutions

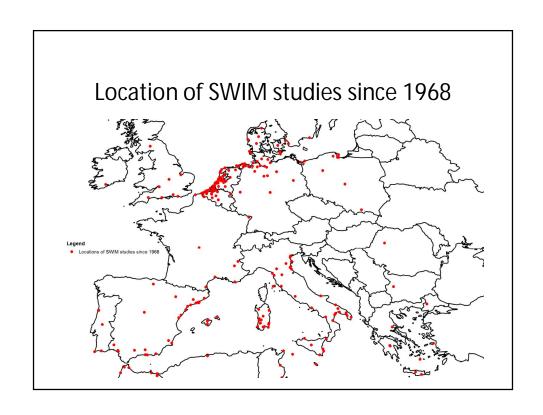












Introduction SWI

Definition of salt water intrusion

Inflow of saline water into an aquifer which contains fresh water

submarine ground surface discharge phreatische groundwater level sea

fresh groundwater

salt water intrusion impervious base

Introduction

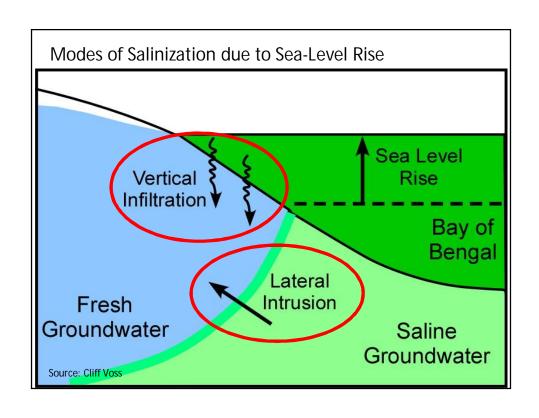
Origin of saline groundwater in the subsoil

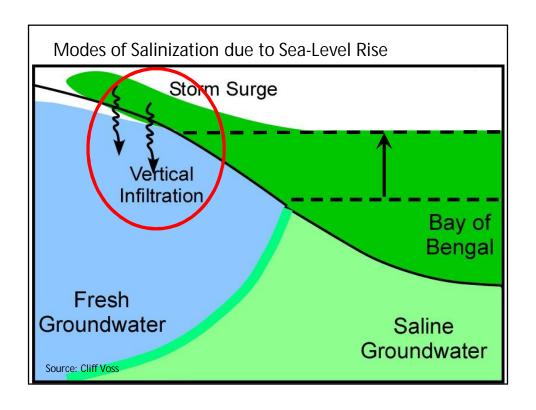
Geological causes:

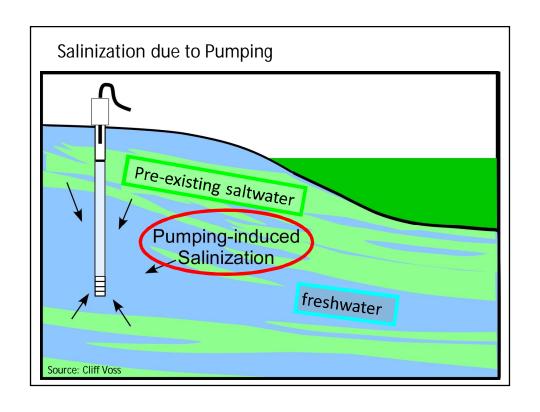
- -marine deposits during geological times
- -trans- and regressions in coastal areas (deltas)
- -salt/brine dome

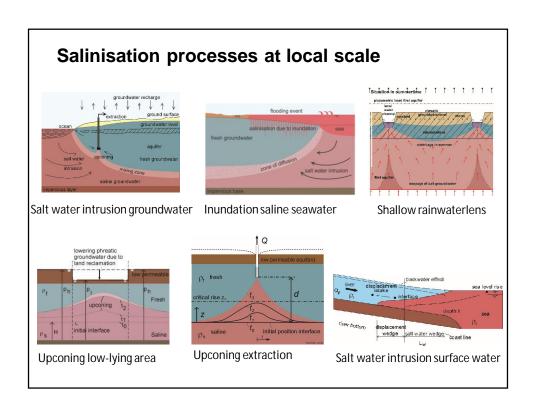
Anthropogenic causes:

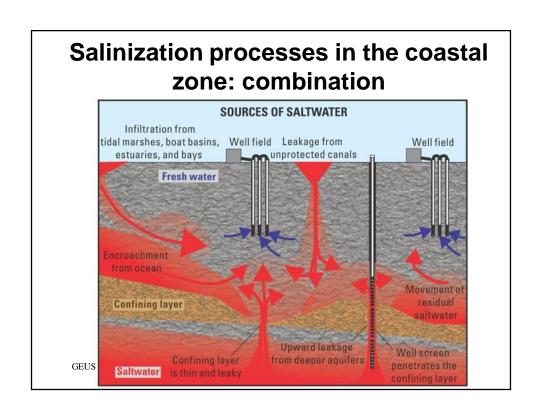
- -agriculture/irrigation (salt damage Middle East & Australia)
- -upconing under extraction wells throughout the world
- -upconing under low-lying areas (e.g. Dutch polders)

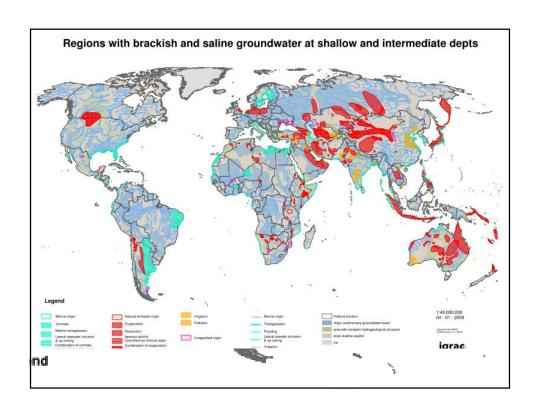


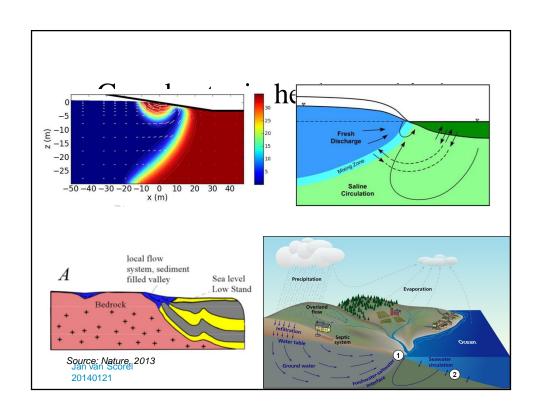


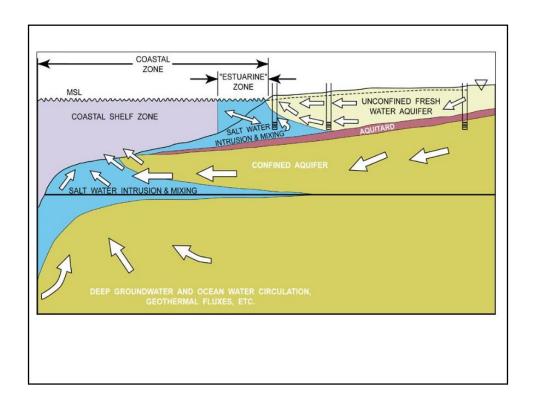


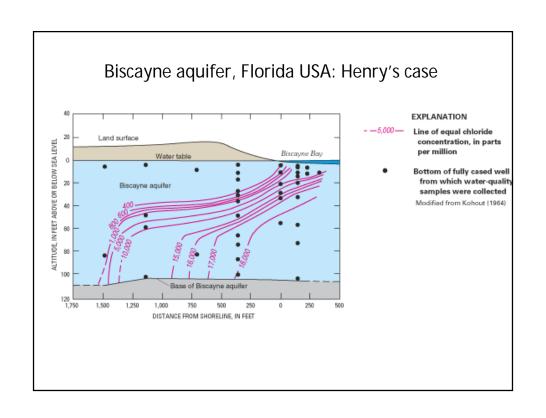


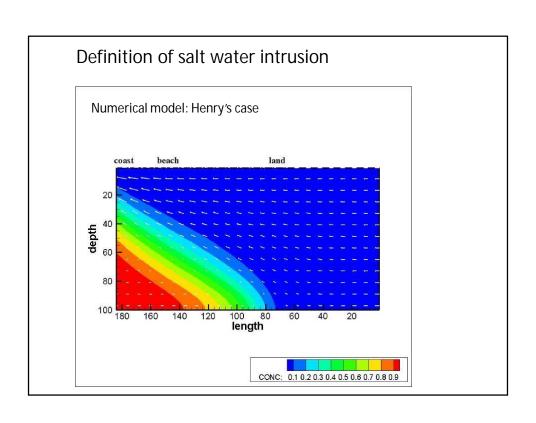


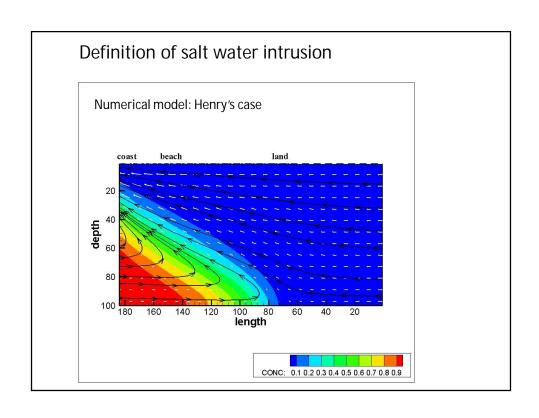


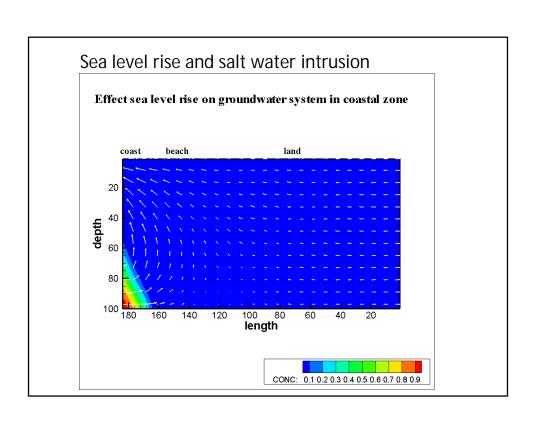


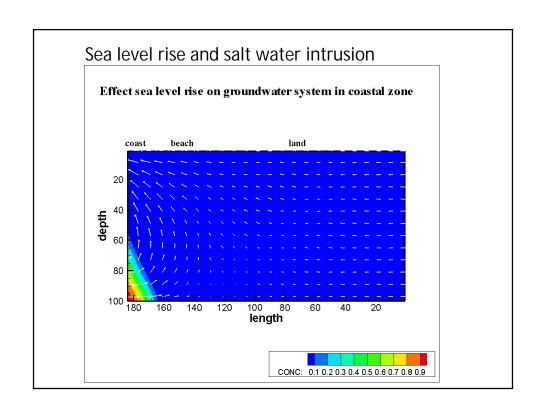


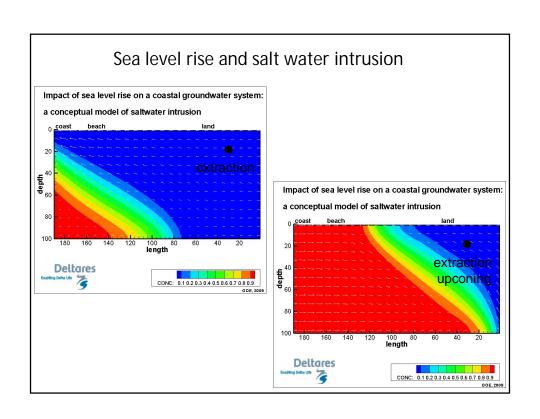


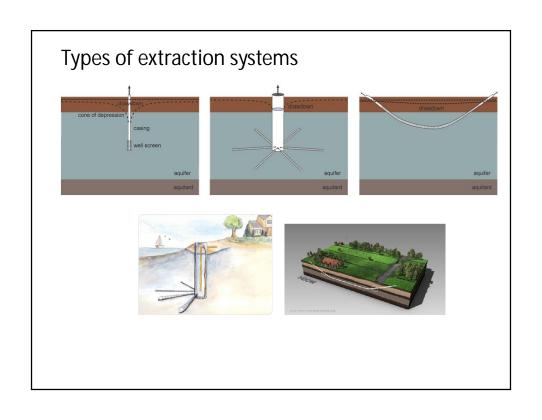


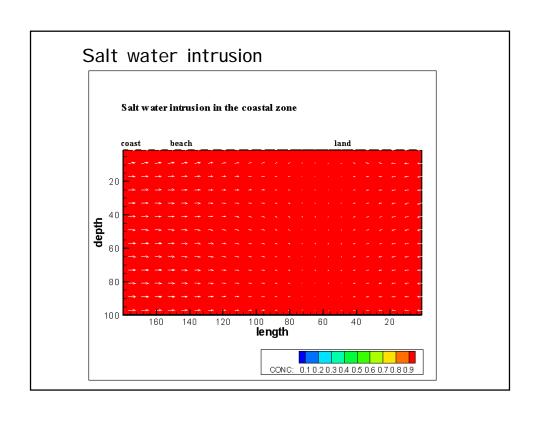


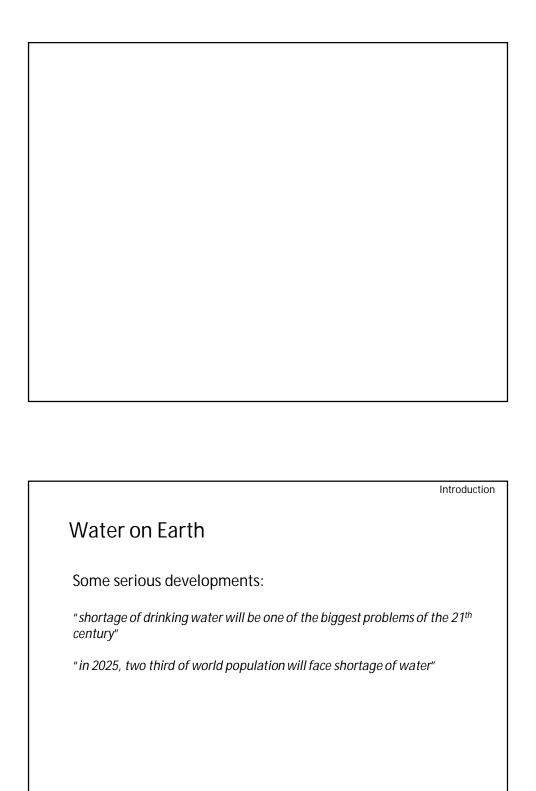


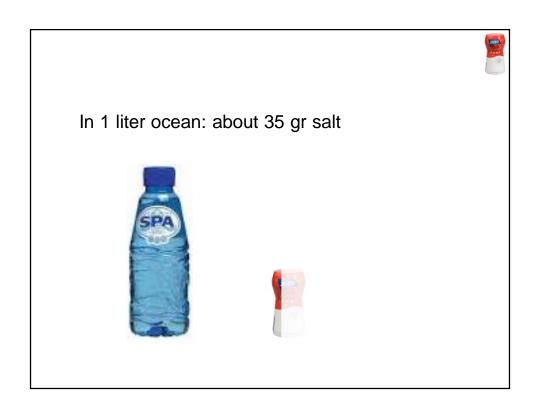


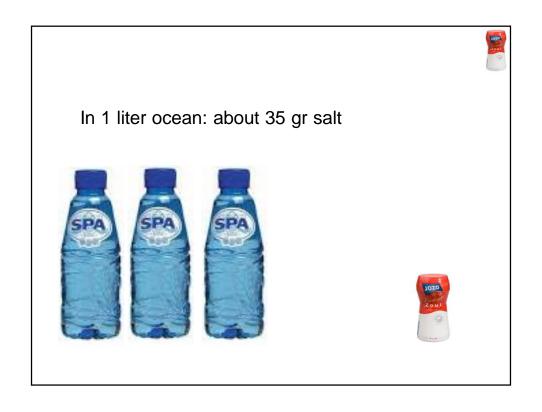




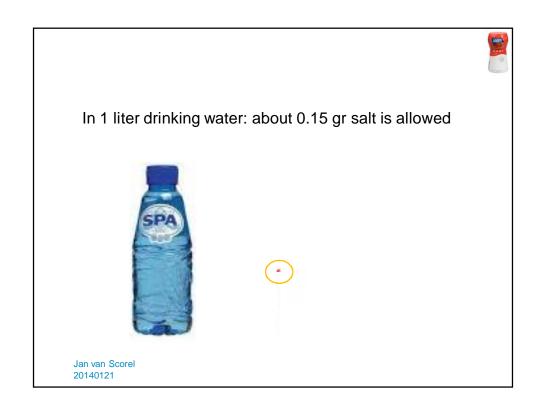


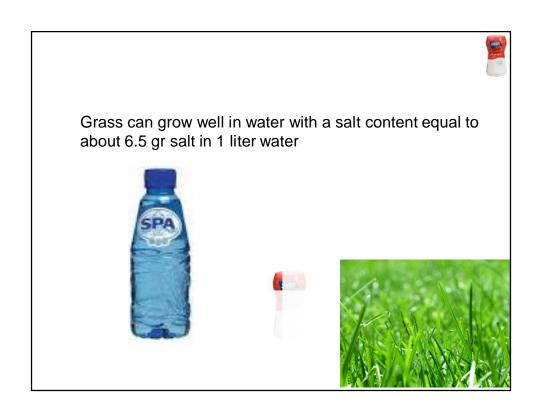












Fresh-brackish-saline groundwater [mg/L] Negative ions 19000 SO₄-2 2700 HCO 3 140 Br⁻ 65 Total negative ions 21905 **Positive ions** 10600 Na⁺ Mg +2 1270 Ca +2 400 380 Total positive ions 12650 Total Disssolved Solids (TDS) 34555

Definition fresh-brackish-saline groundwater

Main type of groundwater	Chloride concentration [mg Cl ⁻ /L]
oligohaline	0 - 5
oligohaline-fresh	5 - 30
fresh	30 - 150
fresh-brackish	150 -300
brackish	300 - 1000
brackish-saline	1000 - 10.000
saline	10.000 - 20.000
hyperhaline or brine	≥20.000

Type	[mS/cm]	[mg TDS/L]	Drinking- or irrigation water	
Non-saline or fresh water	<0.7	<500	Drinking and irrigation water	
Slightly saline	0.7 - 2	500-1.500	Irrigation water	
Moderately saline	2 - 10	1.500-7.000	Primary drainage water and groundwater	
Highly saline	10 - 25	7.000-15.000	Secondary drainage water and groundwater	
Very highly saline	25 - 45	15.000-35.000	Seawater is about 35000 TDS mg/L	
Brine	>45	>35.000	n.a.	

EOS

Examples of equations of state

Knudsen (1902)

$$\rho_{(S,T)} = 1000 + 0.8054S - 0.0065(T - 4 + 0.2214S)^{2}$$

$$\rho_{(S,T)} = 1000 + 0.8054S - 0.0065(T - 4 + 0.2214S)^2$$
 Linear (concentration)
$$\rho_{(C)} = \rho_f \left[1 + \alpha \frac{C_i}{C_s}\right] \quad \text{where a=relative density difference}$$
 Linear (temperature)

Linear (temperature)

$$\rho_{\scriptscriptstyle (T)} = \rho_f [1 - \beta (T - T')]$$

Exponential (temperature, pressure, salt)

$$\rho_{(T,p,\omega)} = \rho_f e^{-\alpha(T-T_0)+\beta(p-p_0)+\gamma\omega}$$

Equation of state (SEAWAT)

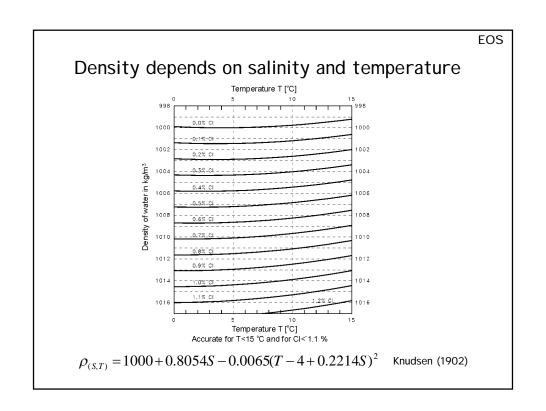
$$\rho_{i,j,k} = \rho_f + \frac{\partial \rho}{\partial C} C_{i,j,k}$$

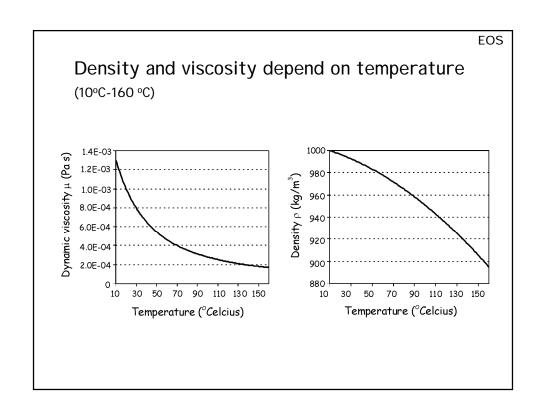
e.g.:

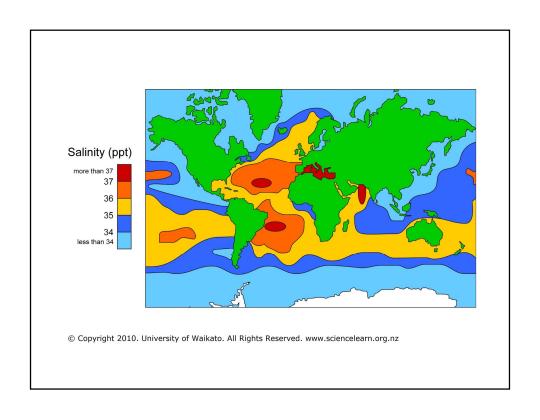
1. conc=35 TDS g/I: DRHODC=0.7143

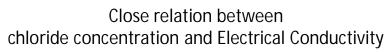
2. conc=19000 mg CI-/I: DRHODC=0.001316 (as 1025=1000+0.001316*19000)

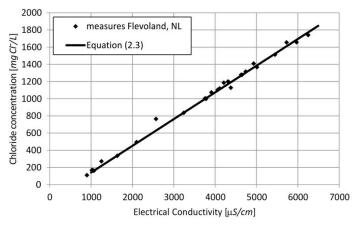
3. conc=1: DRHODC=25 (example practicals)











$$Cl^{-}(mg/L) = EC_{w}(\mu S/cm) \cdot 0.305 - 137$$

Close relation between chloride concentration and Electrical Conductivity

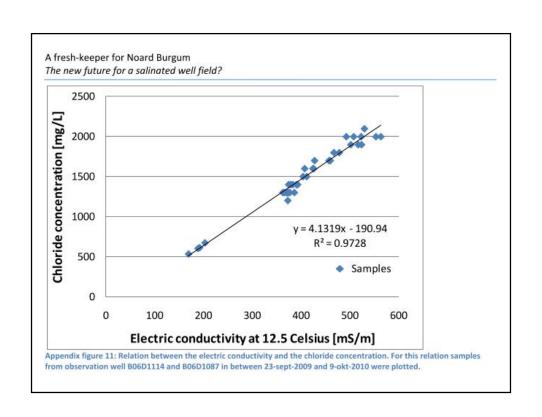
$$10^6 \mu S/cm = 10^3 mS/cm = 1 S/cm$$

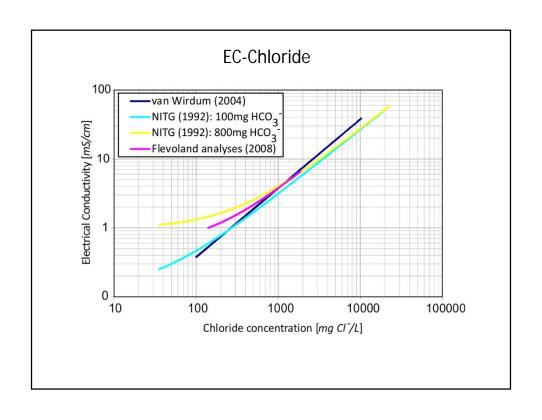
1 $\mu S/cm = 100 \mu S/m$

ocean water:

- ~19000 mg Cl-/L or ~34555 mg TDS/L
- ~5 S/m or ~48 mS/cm

the ratio Cl⁻ over TDS equal to ~0.554, under stable normal seawater environments



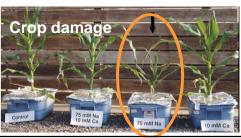


Salt in water is a problem









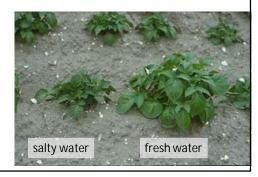


Introduction

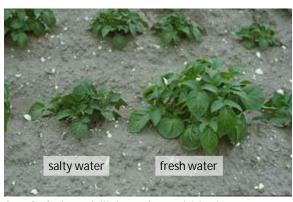
Salt in water is a problem for different water management sectors:

- -drinking water:
 - •taste (100-300 mg Cl⁻/l)
 - •long term health effect
 - •norm: EC& WHO=150 mg Cl⁻/I (live stock=1500 mg Cl⁻/I)
- -industry:
 - •corrosion pipes
 - preparation food
- -irrigation/agriculture:•production crops

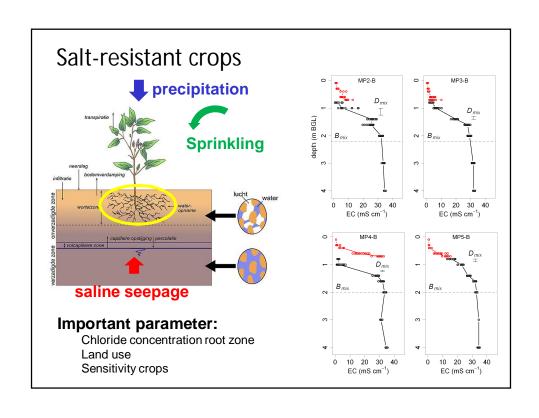
 - •salt damage



Effects salinisation: salt damage



Source: Proefstation voor de Akkerbouw en Groenteteelt, Lelystad



Salt damage to crops

Important parameters:

- Chloride concentration in the root zone
- Land use
- Sensitivity crops

Land use	Threshold root zone (mg CI-/I)	value	Gradient root zone (-)
Grass		3606	0.0078
Potatoes		756	0.0163
Beet		4831	0.0057
Grains		4831	0.0058
Horticulture		1337	0.0141
Orchard (trees)		642	0.0264
Bulb		153	0.0182

Relatio trusen zoutgehalte en opbrengstschade landbouwgewessen

100

Schiede bij overschrijding geheel groeiseizonn (%)

60

40

20

1000

2000

3000

4000

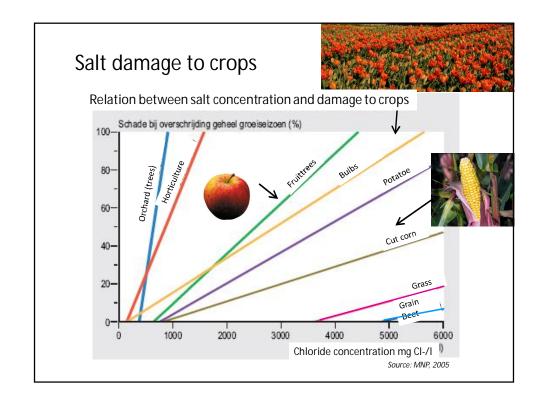
5000

5000

5000

Source: MNP, 2005

Source: Roest et al., 2003 en Haskoning

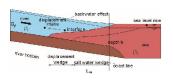


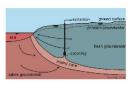
	Soil moisture Irrigation wat		Irrigation water	
	Limi	Gradient	Limit	Gradient
Crop	mg/I CI	%/mg/l Cl	mg/I CI	%/mg/I CI
Potatoe	756	0.0163	202	0.0610
Grass	3606	0.0078	962	0.0294
Sugar beat	4831	0.0057	1288	0.0212
Cut Corn	815	0.0091	217	0.0343
Grains	4831	0.0058	1288	0.0218
Fruit trees	642	0.0264	171	0.0991
Orchard (trees)	378	0.1890	101	0.7086
Vegetables	917	0.0158	245	0.0591
Horticulture	1337	0.0141	356	0.0527
Bulbs	153	0.0182	41	0.0683

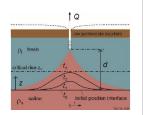
Introduction

Why is salinisation a pressing problem?

- 30% of world population lives < 100 km from coastline
- economic and tourist activities increase
- enormous increase in extraction
- irreversible process
- increase saltwater intrusion problem world-wide:
 - upconing
 - salt water wedge
 - decrease outflow q₀
- climate change:
 - sea level rise
 - natural groundwater recharge







Introduction

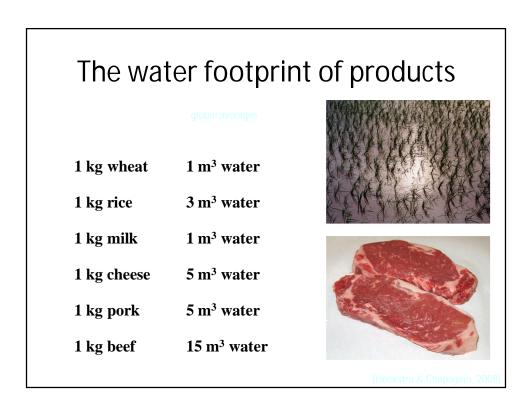
Processes that accelerate salt water intrusion:

- Sea level rise
- Land subsidence
- Human activities

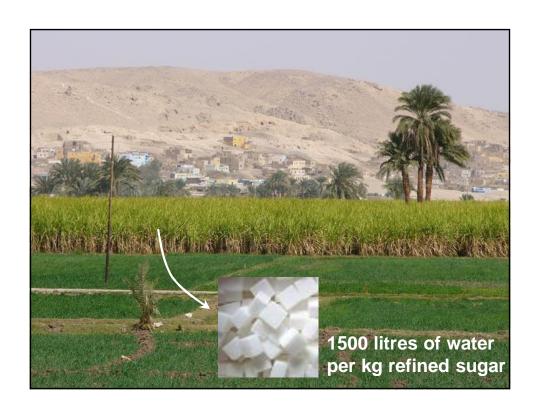
Threats for:

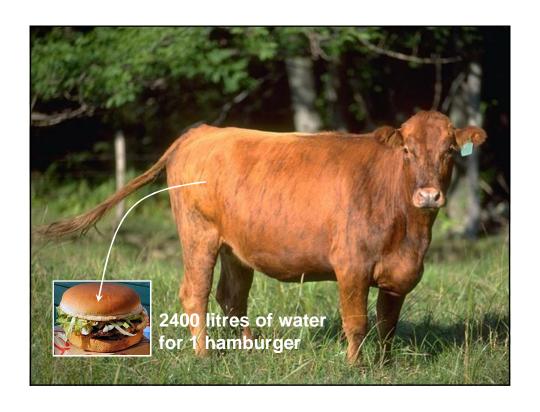
- •drinking water supply in dunes: upconing of saline groundwater decrease of fresh groundwater resources recharge areas reduction
- agriculture: salt damage to crops: salt load and seepage
- water management low-lying areas: flushing water channels
- ecology













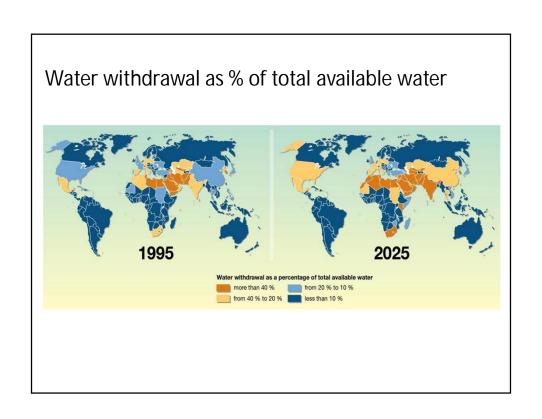


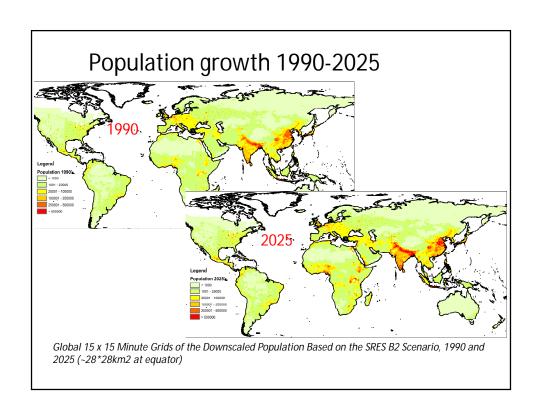
Introduction

Question:

Demand fresh water per capita per day?:

- a. 10 litre/day b. 25 litre/day
- c. 100 litre/day
- d. 200 litre/day





Introduction

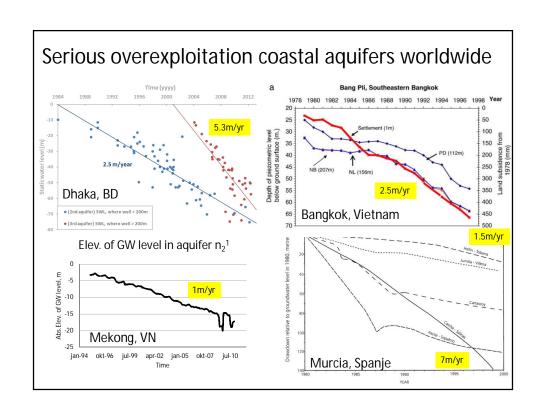
Reasons and drawbacks of using groundwater

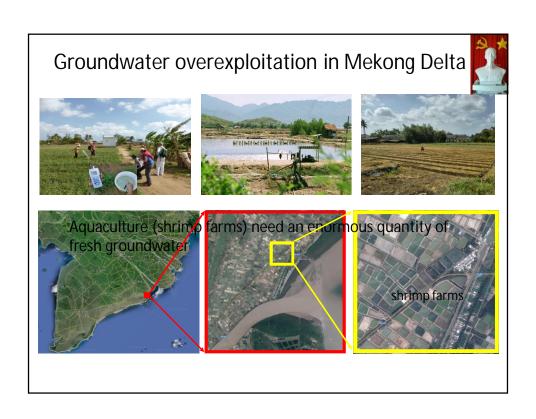
Advantage:

- -no seasonal effects
- -high quality
- -low storage costs
- -large quantities
- -no spatial limitations

Disadvantage:

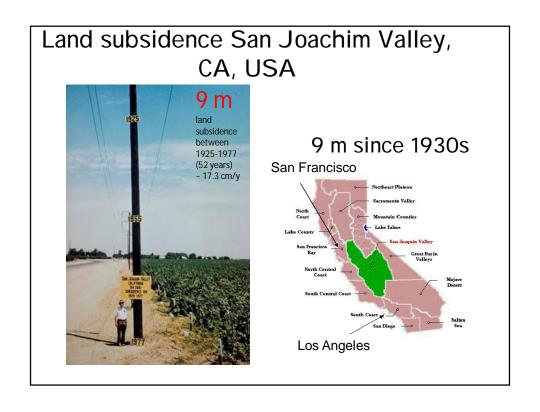
- -high extraction costs
- -local droughts
- -high mineral content
- -land subsidence....
- -salt water intrusion!





Land	\sim 1	hai	ia	$l \sim r$	
Land	SU	N2	IU	lei.	ICE

Megacity	Maximum subsidence [m]	Date commenced
Shanghai	2.80	1921
Tokyo	5.00	1930's
Osaka	2.80	1935
Bangkok	1.60	1950's
Tianjin	2.60	1959
Jakarta	0.90	1978
Manila	0.40	1960
Los Angeles	9.00	1930's



What causes the land to subside?

Natural causes (geological processes):

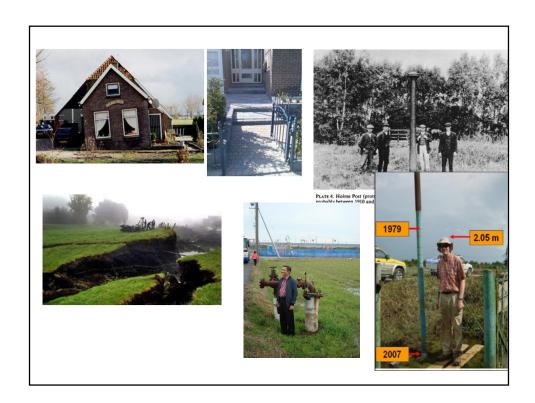
- Loading of the earth's crust by ice sheets, sediment (delta's), the ocean/sea
- Compaction of older sediments after sedimentation

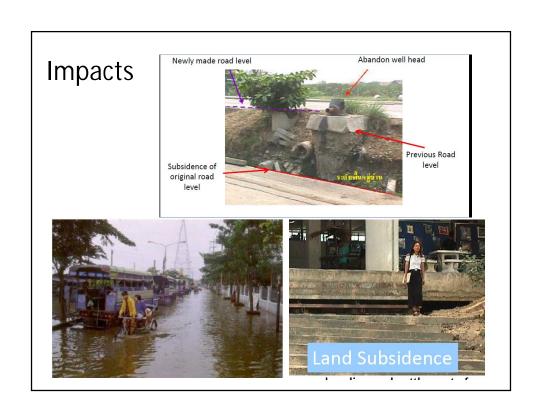
Anthropogenic causes (human-induced processes):

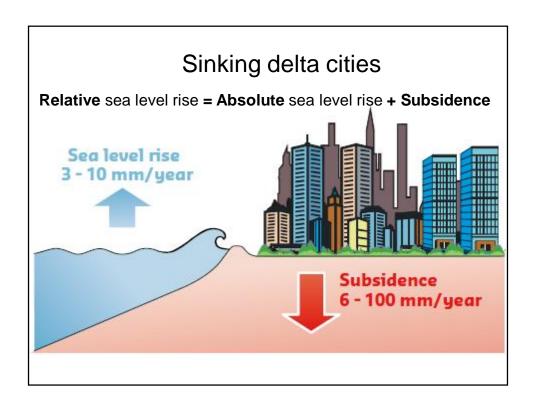
- Oil/gas extraction (usually relatively deep)
- Groundwater extraction (usually moderately deep)
- □ Drainage of soils ⇒ oxidation of peat, soil compaction

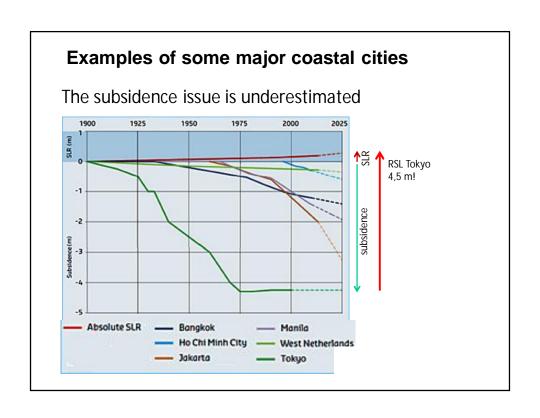
Why discriminating between human-induced and natural processes?

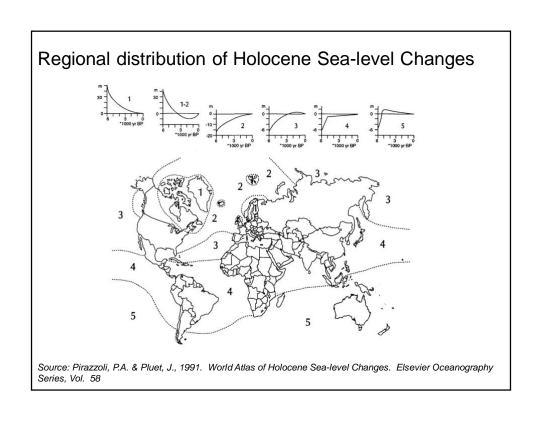
- Magnitude
- Cooping strategy (mitigation versus adaptation)

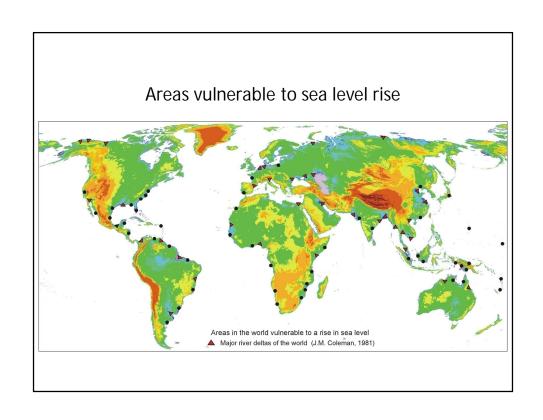






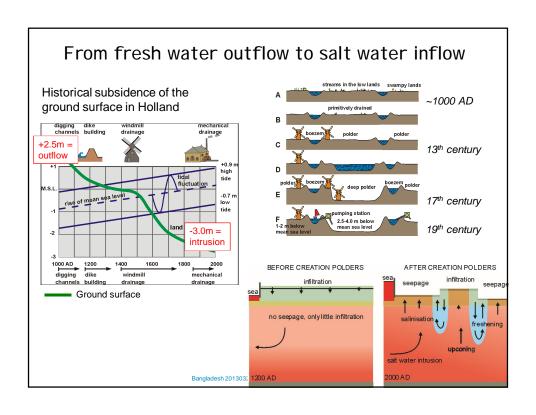


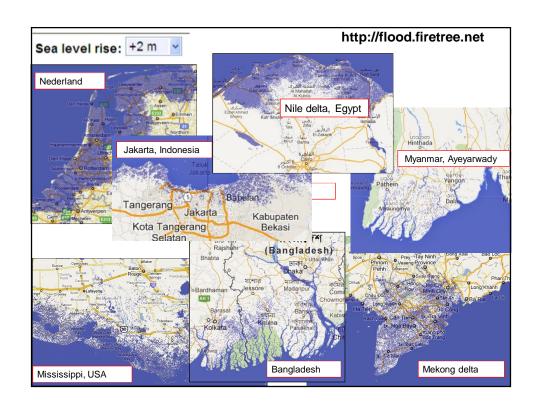


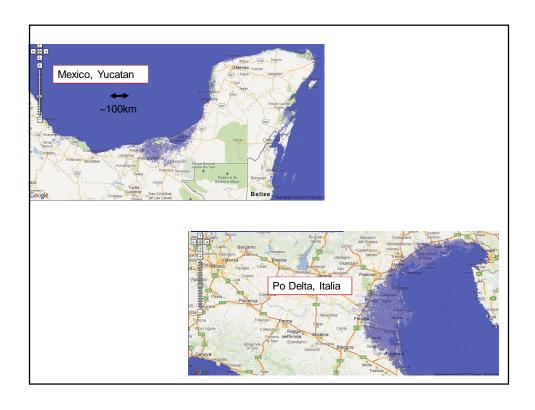


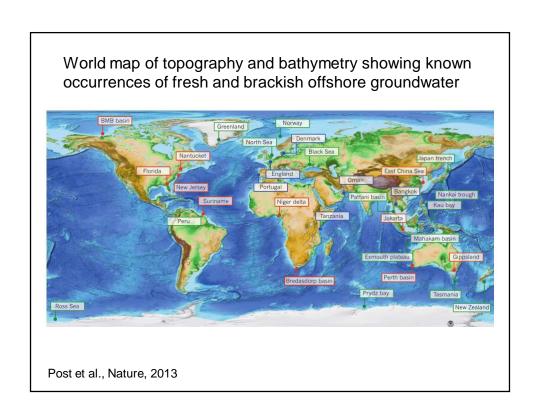


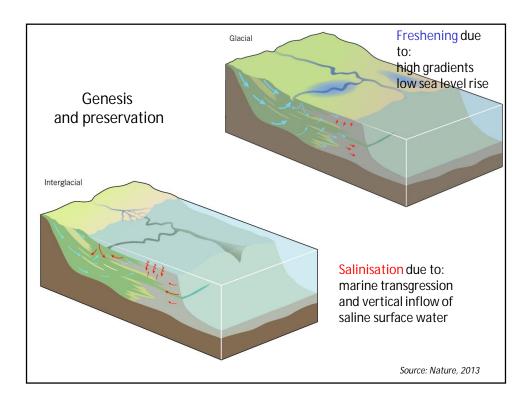
Saline seepage leads to: Salinization and eutrophication of surface waters Salinization of shallow groundwater Salinization of root zone (crop damage) Deep seepage polders dunes Peat lands dike Lake IJssel fresh/saline interface

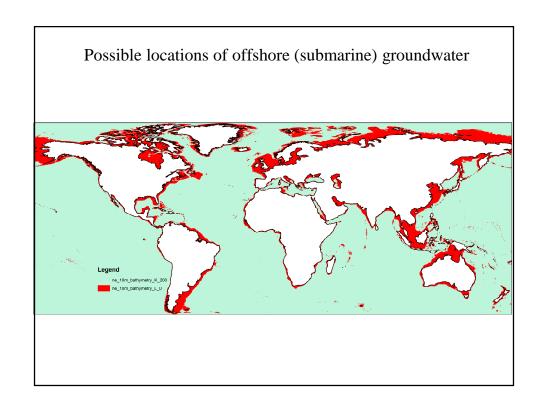


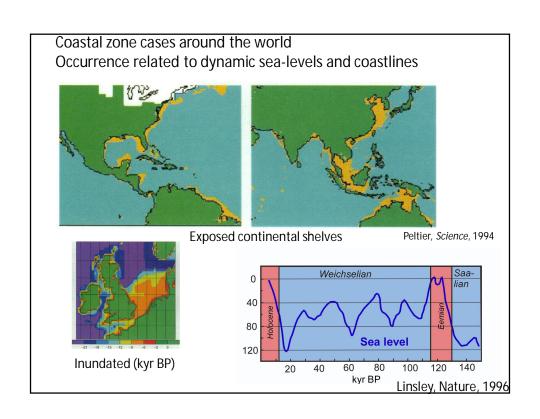


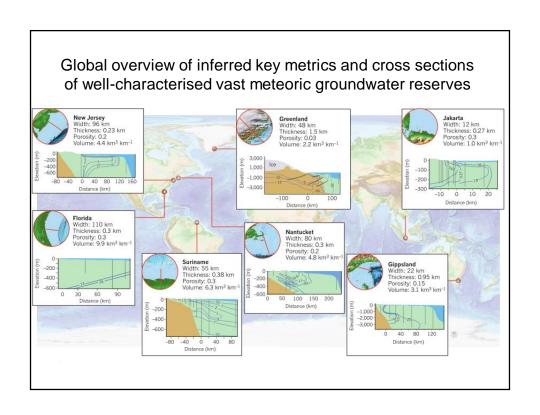




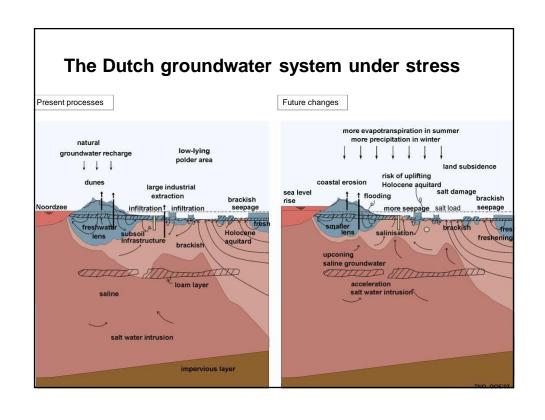


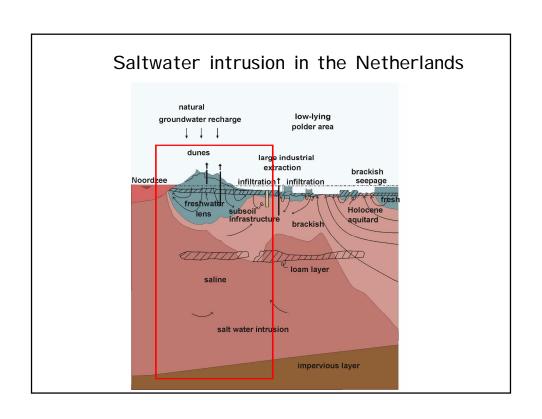


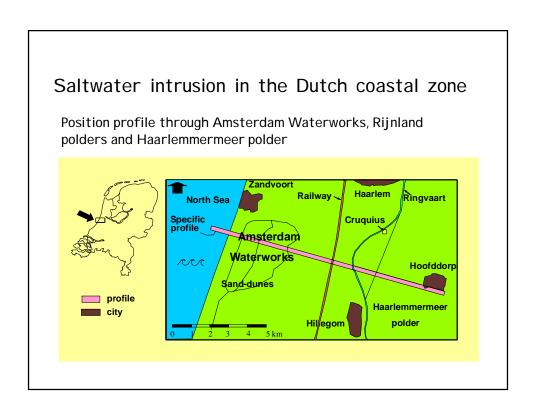


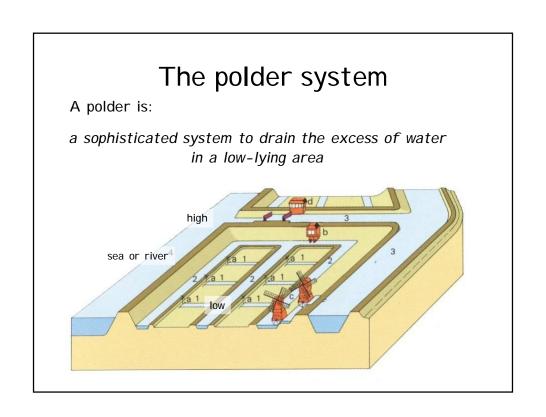


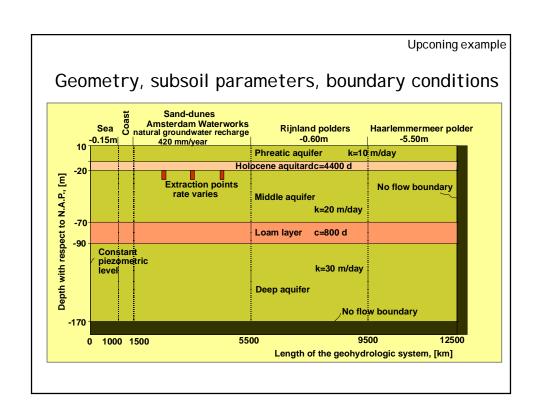
My first density dependent groundwater flow and solute transport model in 1990!

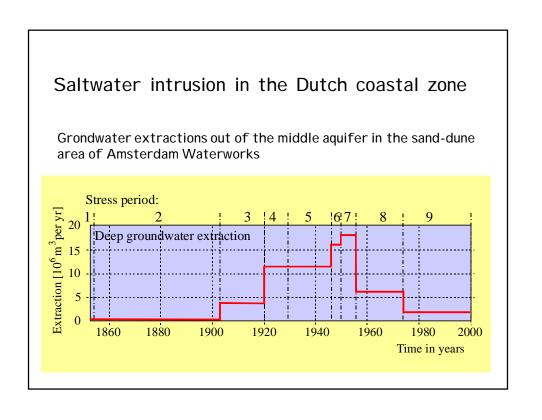


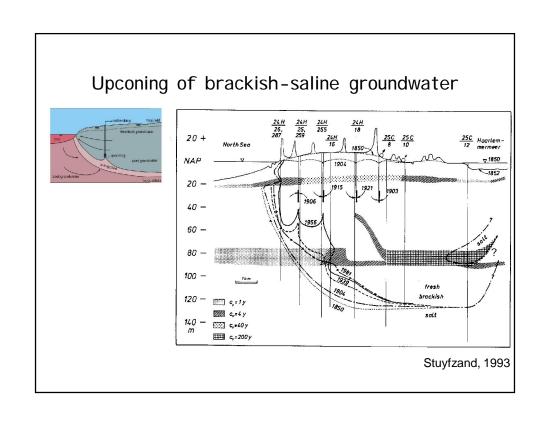


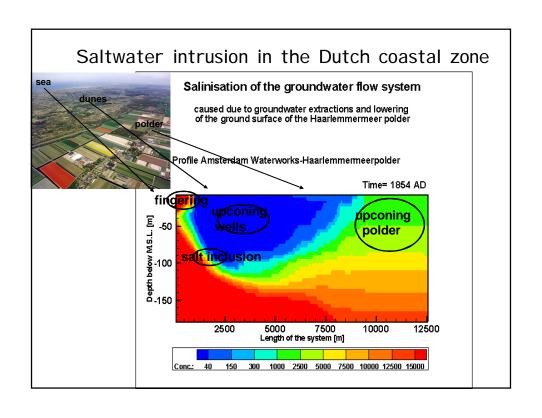


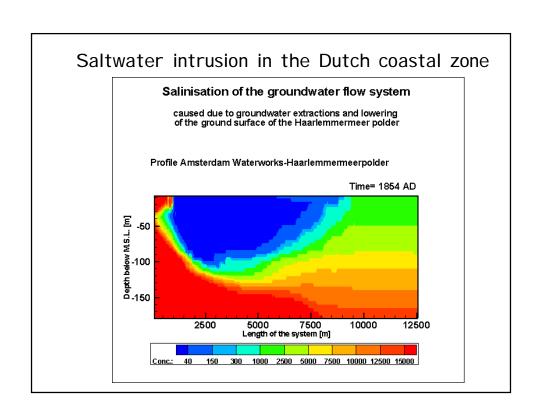








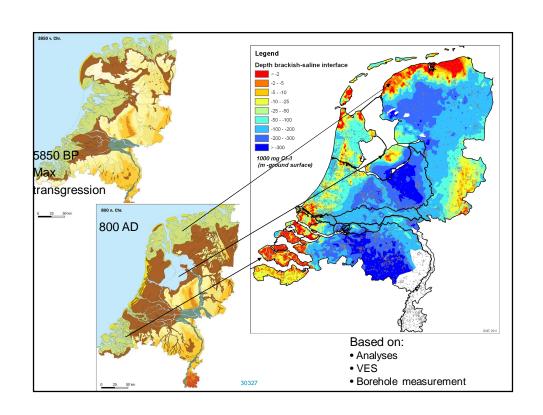


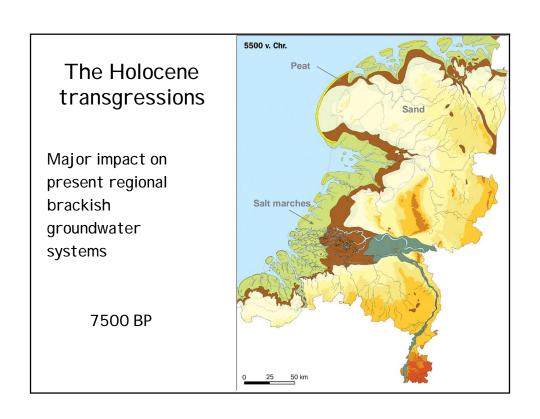


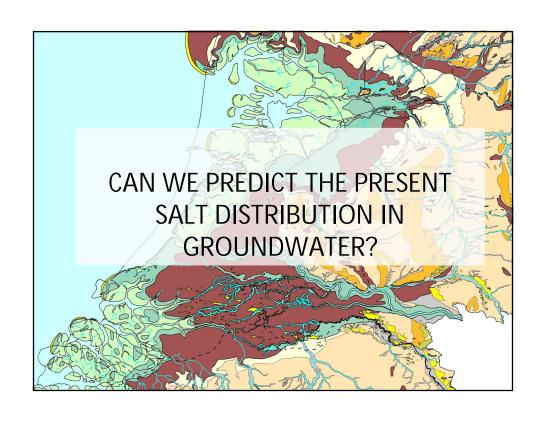
Palaeo hydrogeological modelling

Palaeo-modeling salt water intrusion during the Holocene: an application to the Netherlands

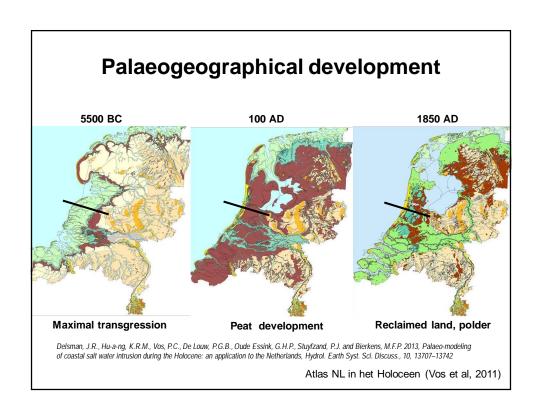
J.R. Delsman, K. Hu-a-ng, P.C. Vos, P.G.B. de Louw, G.H.P. Oude Essink and M.F.P. Bierkens









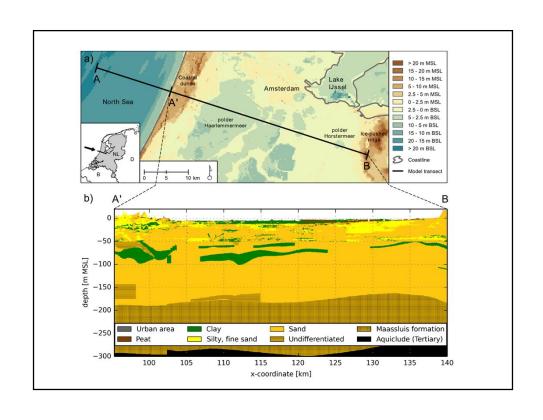


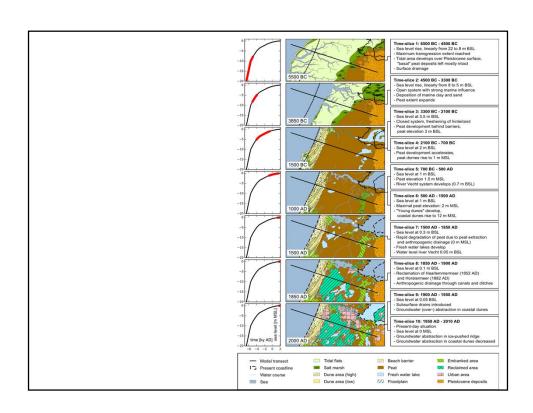
Occurrence of salt under the polder Haarlemmermeer

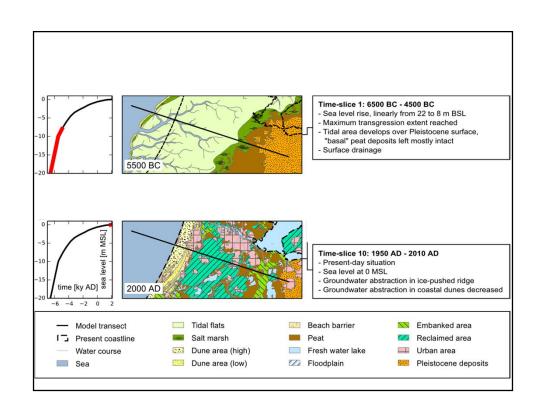
- Model profile Zandvoort Hoofddorp Hilversum
- Palaeogeographical development (Vos et al, 2011)
- 6500 BC 2010 AD
- marine transgression
- · Peat development, peat degradation, drainage,

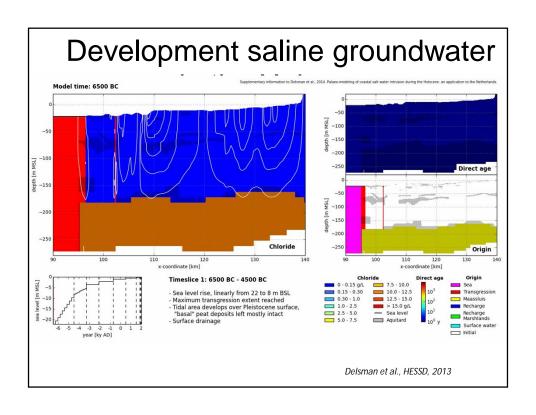


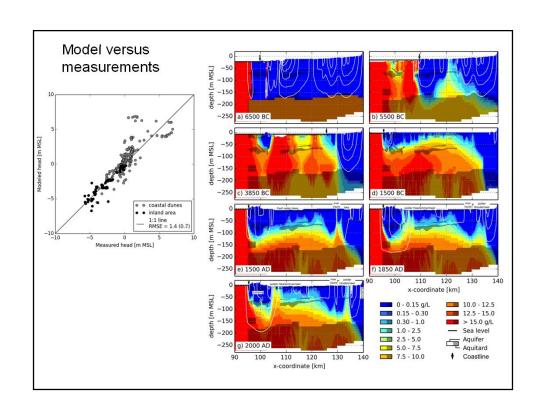
Delsman et al., HESS, 2013

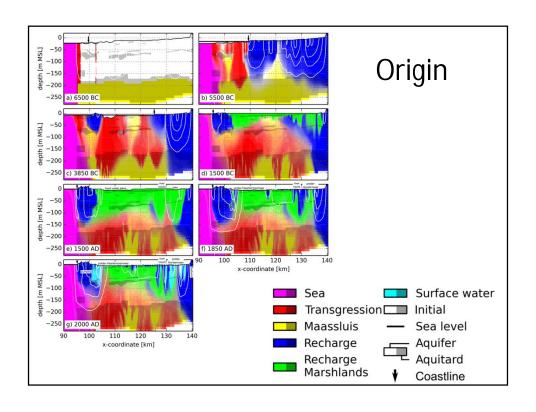


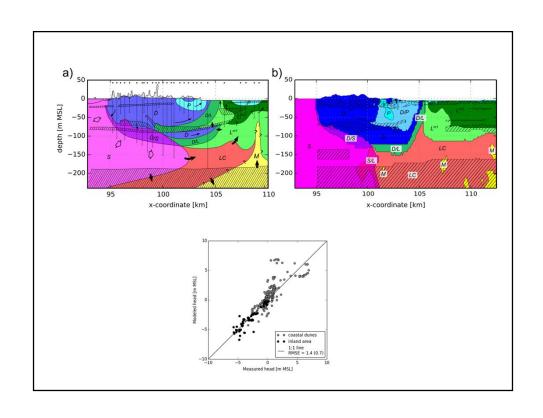


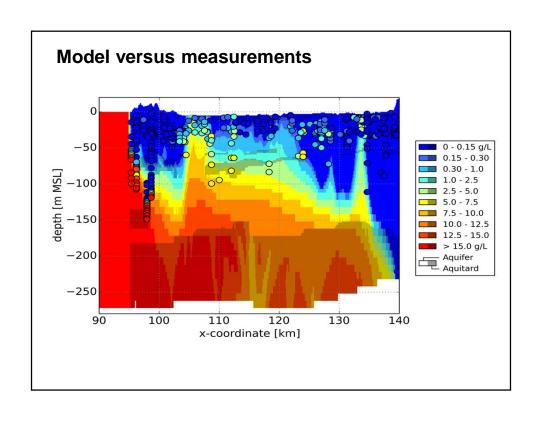




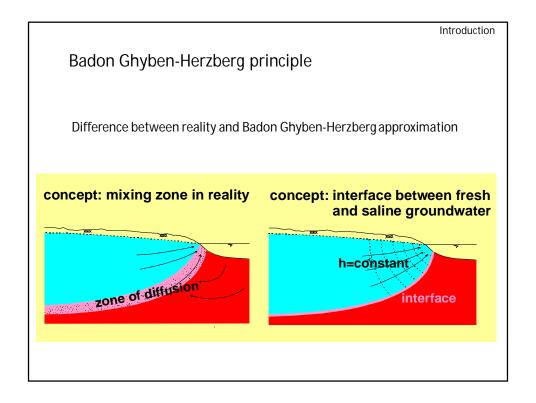


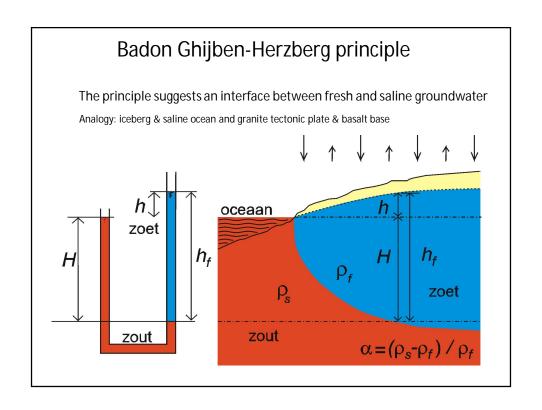


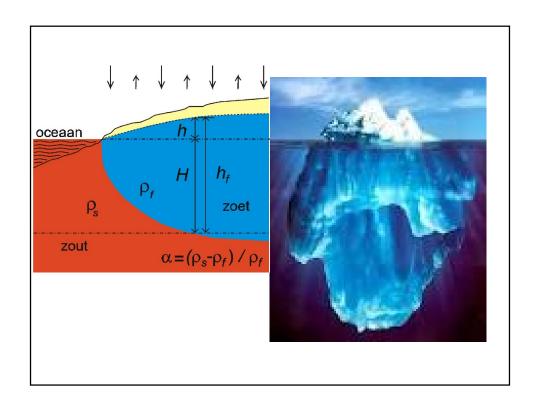




Sharp interface between fresh and saline groundwater







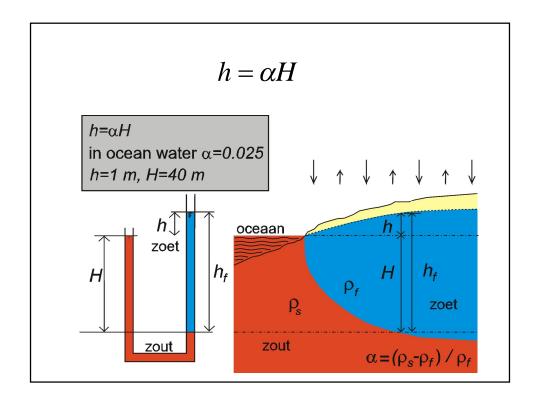
pressure saline groundwater=pressure fresh groundwater
$$\rho_s Hg = \rho_f (H+h)g$$

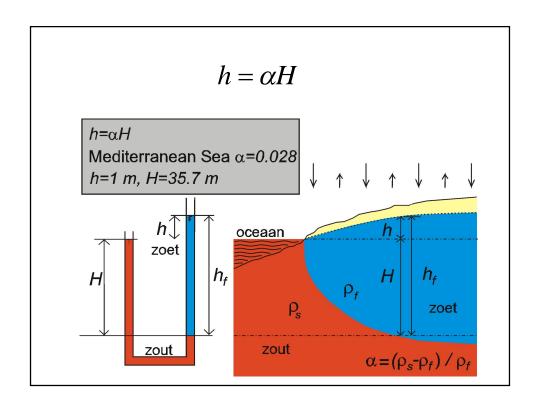
$$h = \frac{\rho_s - \rho_f}{\rho_f} H$$

$$h = \alpha H$$

$$h = \alpha H$$

$$h_f = \frac{\rho_s - \rho_f}{\rho_f} H$$





Badon Ghyben-Herzberg principle

- gives analytical solutions (see later and lectures)
- educational
- interface is a simple approximation
- dispersion zone <10m
- relative simple geometries

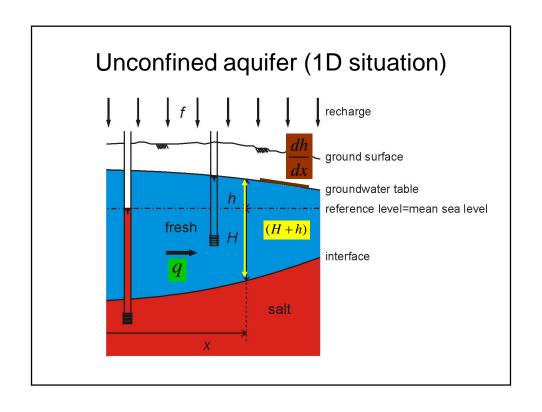
Badon Ghyben-Herzberg principle

What is the case then $h\neq\alpha H$?

- 1. still dynamic situation
- 2. occurrence resistance layer
- 3. natural groundwater recharge not constant
- 4. relative density difference α is not ok
- 5. occurrence shallow bedrock
- 6. groundwater extractions

Analytical solutions

Analytical solutions See lecture notes Density dependent groundwater flow (p. 29-48) **The state of the sta



Unconfined aquifer (1D situation)

- (I) Darcy $q = -k(H+h)\frac{dh}{dx}$
- (II) Continuity dq = fdx
- (III) BGH $h = \alpha H$

Unconfined aquifer (1D situation)

$$dq = fdx$$
 integration $q = fx + C1$ gives

$$-k(H+h)\frac{dh}{dx} = fx + C1$$

$$h = \alpha H \rightarrow -k(H + \alpha H)\alpha \frac{dH}{dx} = fx + C1$$

$$HdH = -\frac{fx + C1}{k\alpha(1+\alpha)}dx$$

Unconfined aquifer (1D situation)

$$HdH = -\frac{fx + C1}{k\alpha(1+\alpha)}dx$$

integration gives

$$\frac{1}{2}H^{2} = \frac{-\frac{1}{2}fx^{2} - C1x + C2}{k\alpha(1+\alpha)}$$

$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}}$$

Unconfined aquifer (1D situation)

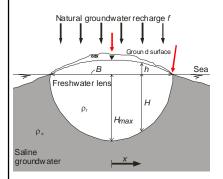
$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}}$$

$$h = \alpha H$$

$$q = fx + C1$$

Example 1: Elongated island

$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}} \qquad q = fx + C1$$

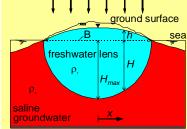


Boundary conditions

$$x = 0: q = 0 \rightarrow C1 = 0$$

$$x = 0.5B : H = 0 \rightarrow C2 = fB^2 / 8$$

Example of analytical solutions (I)



Depth of fresh-saline interface H

$$H = \sqrt{\frac{f(0.25B^2 - x^2)}{k\alpha(1+\alpha)}}$$

$$h = \alpha H$$

Maximal thickness lens

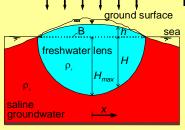
$$H_{\text{max}} = \frac{1}{2} B \sqrt{\frac{f}{k\alpha(1+\alpha)}} \qquad V = \frac{1}{4} \pi (1+\alpha) H_{\text{max}} B n_e$$

$$V = \frac{1}{4}\pi(1+\alpha)H_{\text{max}}Bn_e$$

Characteristic time $T = \frac{\text{volume of water in lens}}{\text{inflow of water}} = \frac{\pi n_e B}{8} \sqrt{\frac{(1+\alpha)}{kf\alpha}}$

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Example of analytical solutions (I)



Depth of fresh-saline interface H

B = 2000 m, f = 0.001 m/day

$$k = 10 \text{m/day}, \alpha = 0.025$$

$$n_e = 0.35$$

Maximal thickness lens

Volume lens (wrong in lectures notes)

$$H_{\text{max}} = 62.5 \text{m}, h_{\text{max}} = 1.56 \text{m}$$
 $V = 35203 \text{m}^3/\text{m}'$

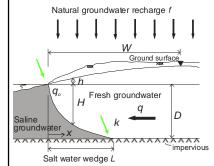
$$V = 35203 \text{m}^3/\text{m}'$$

Characteristic time
$$T = \frac{35203}{2} days = 48.2 years$$

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Example 2: salt water wedge

$$H = \sqrt{\frac{-fx^2 - 2C1x + 2C2}{k\alpha(1+\alpha)}} \qquad q = fx + C1$$



Boundary conditions

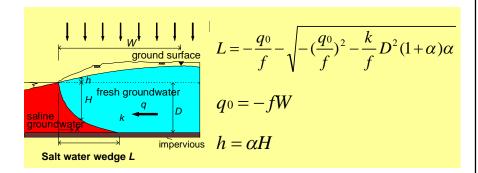
$$x = 0 : q = q_0 \rightarrow q_0 = -fW \rightarrow C1 = q_0$$

$$x = 0: H = 0 \rightarrow C2 = 0$$

Length of salt water wedge

$$x = L : H = D$$

Example of analytical solutions (II)



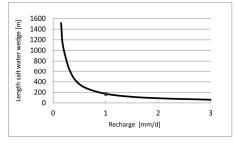
Example:

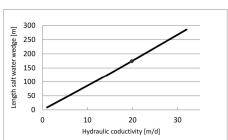
$$W = 3000 \text{m}, f = 0.001 \text{m/day}, \alpha = 0.020, k = 20 \text{m/day}, D = 50 \text{m}$$

$$L = 175.1$$
m

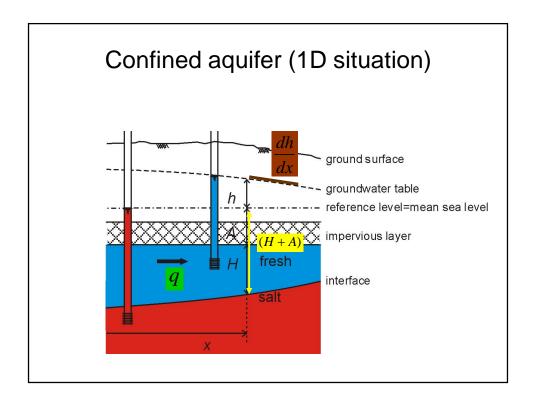
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Length of the salt water wedge as a function of a. recharge and b. hydraulic conductivity





the dots resample with the example mentioned above



Confined aquifer (1D situation)

- (I) Darcy $q = -kH \frac{dh}{dx}$
- (II) Continuity $q=q_0$
- (III) BGH $h = \alpha(H + A)$

Confined aquifer (1D situation)

$$-kH\frac{dh}{dx} = q_0$$

$$HdH = -\frac{q_0}{k\alpha}dx$$

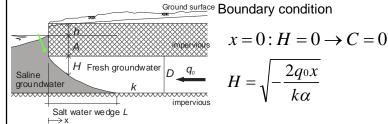
integration gives

$$\frac{1}{2}H^2 = \frac{q_0x}{k\alpha} + C$$

$$H = \sqrt{-\frac{2q_0x}{k\alpha} + 2C}$$

Example 3: salt water wedge confined aquifer

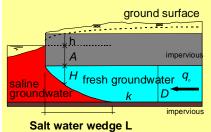
$$H = \sqrt{-\frac{2q_0x}{k\alpha} + 2C}$$



Length of salt water wedge x = L : H = D

$$L = -\frac{kD^2\alpha}{2q_0}$$

Example of analytical solutions (III)



Length of salt water wedge

$$H = \sqrt{-\frac{2q_0x}{k\alpha}}$$

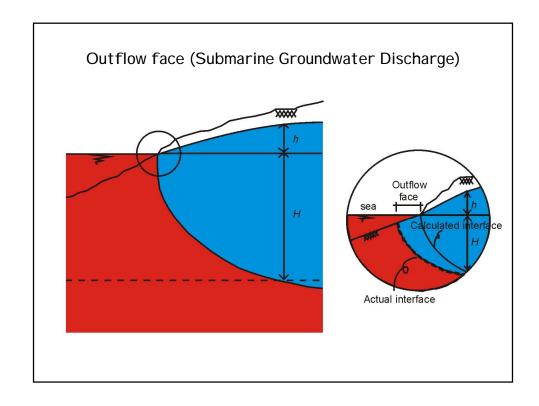
$$L = -\frac{kD^2\alpha}{2q_0}$$

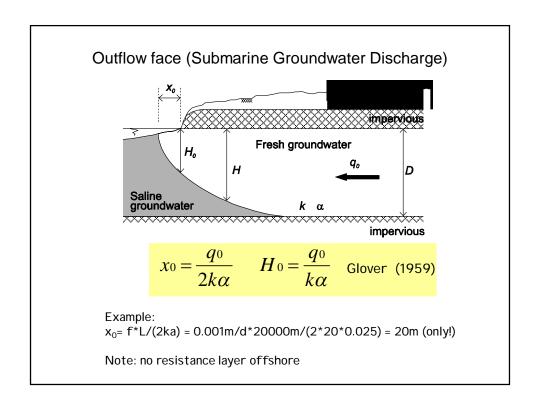
Example:

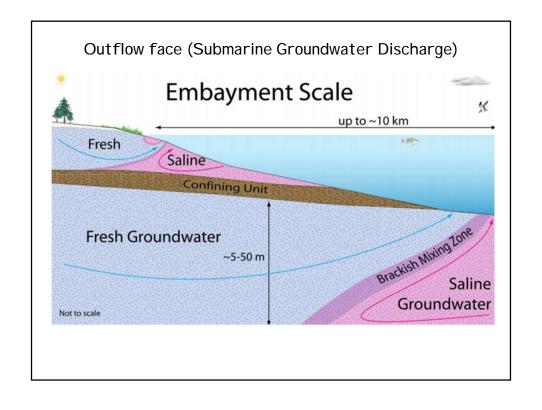
$$W = 2000 \text{m}, f = 0.001 \text{m/day}, \alpha = 0.025, k = 25 \text{m/day}, D = 40 \text{m}$$

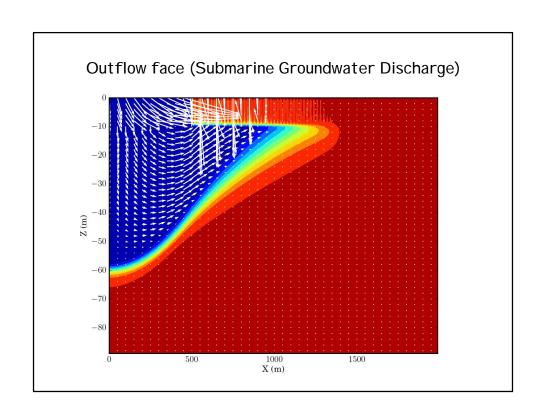
$$L = 250 \text{m}$$

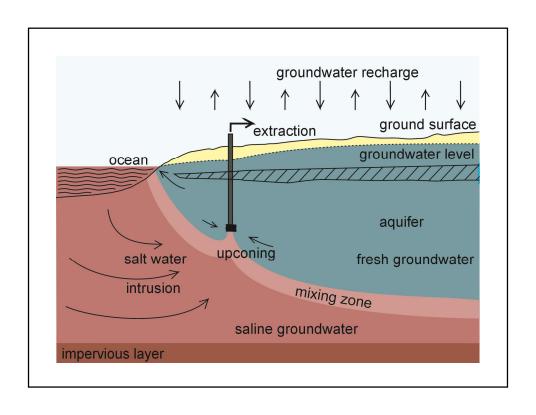
Lecture notes p. 35-36

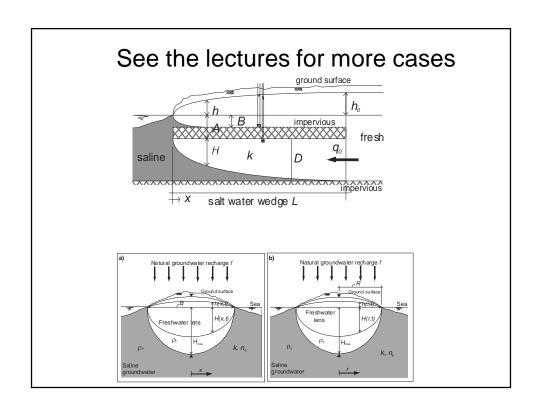




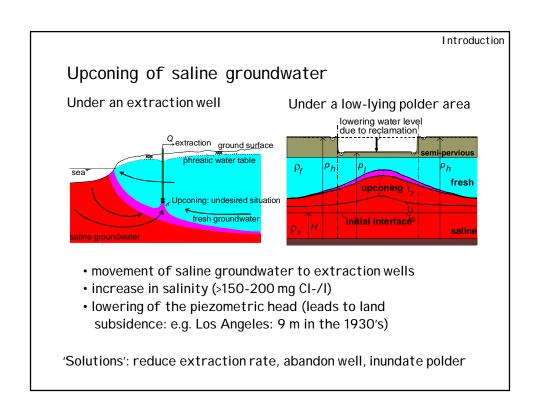


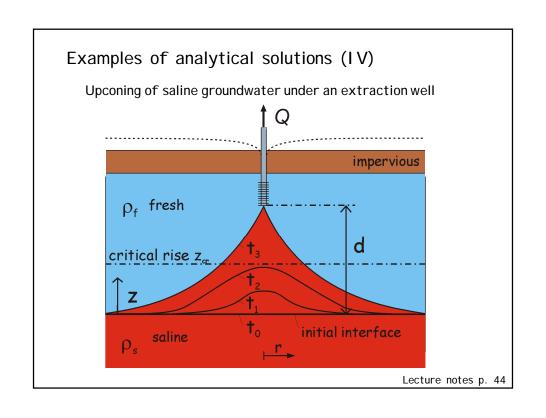


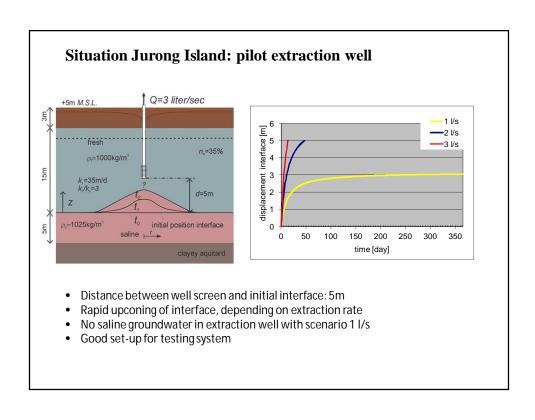


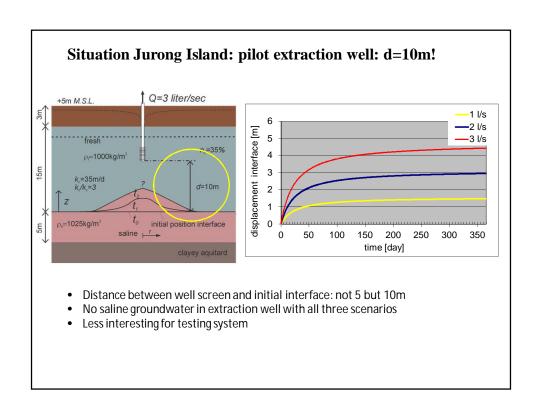


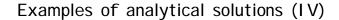
Upconing processes











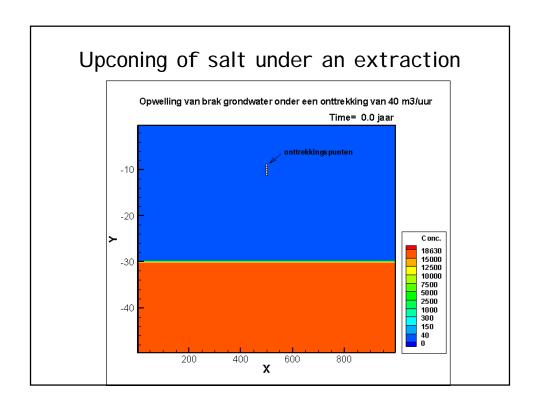


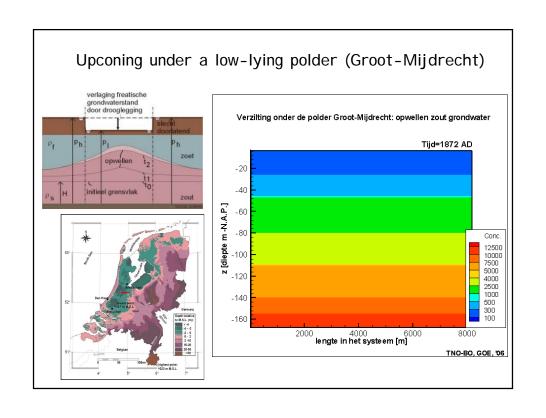
Upconing of saline groundwater under an extraction well

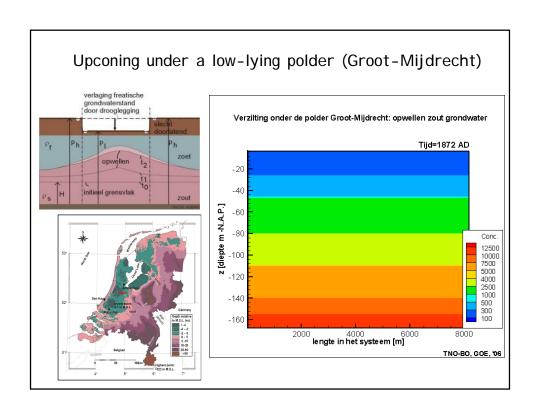
$$\begin{split} z(r,t) &= \frac{Q}{2\pi\alpha k_x d} \left[\frac{1}{(1+R'^2)^{1/2}} - \frac{1}{[(1+\gamma')^2 + R'^2]^{1/2}} \right] \\ R' &= \frac{r}{d} \frac{k_z}{k_x}^{1/2} \qquad \gamma' = \frac{\alpha k_z}{2n_e d} t \end{split}$$

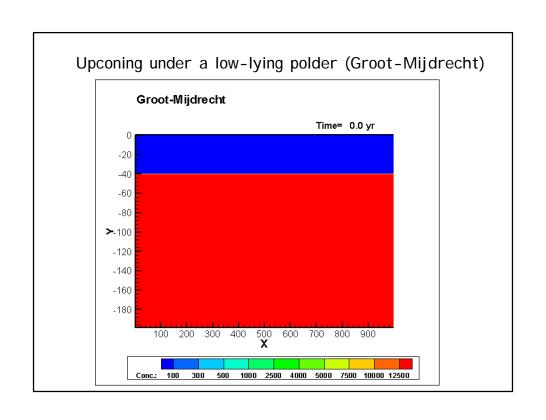
Dagan & Bear, 1968, J. Hydraul. Res 6, 1563-1573

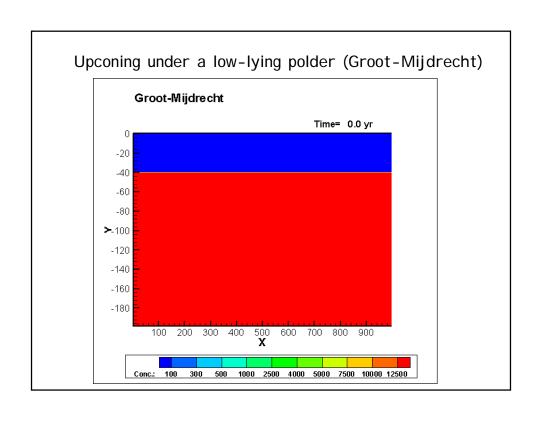
Lecture notes p. 44

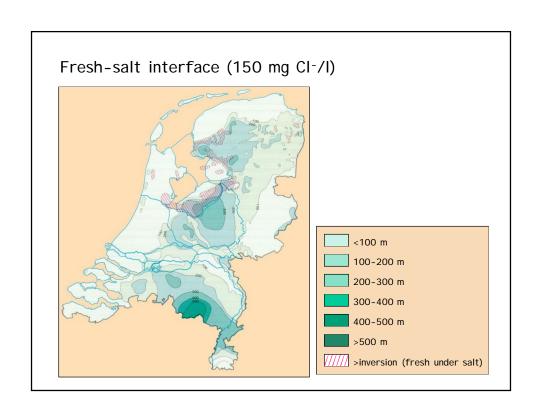


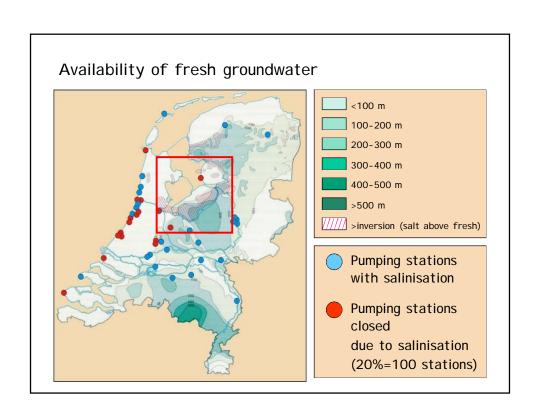


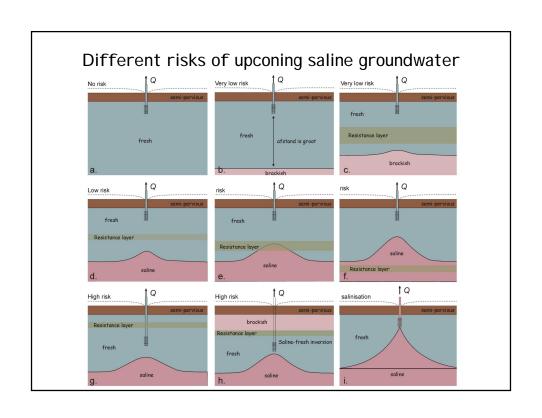


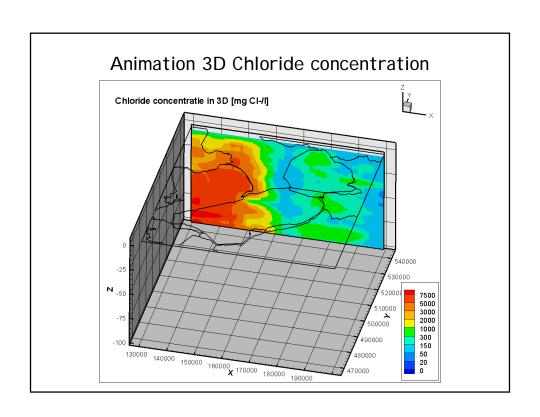


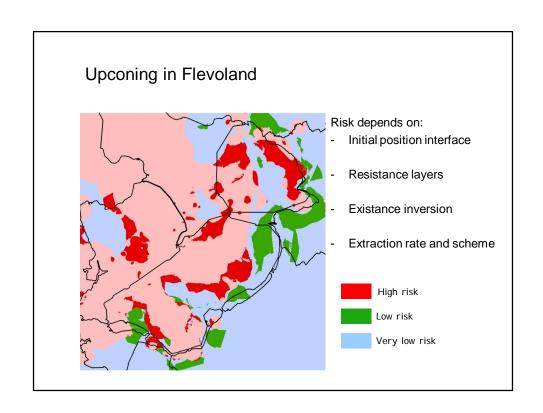












Compensating measures

Possible solutions to stop salt water intrusion:

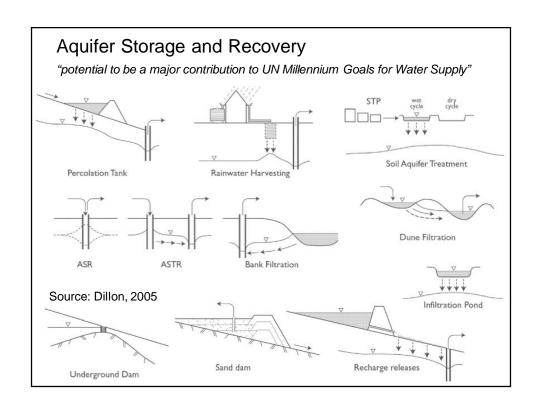
- Restriction of groundwater extractions through permits
- Co-operation between authorities and water users
- Desalinisation of saline water
- Technical countermeasures of salt water intrusion
 - · six examples

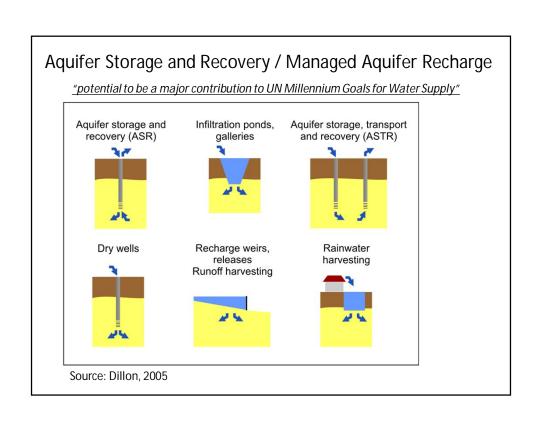
Tools to understand salt water intrusion:

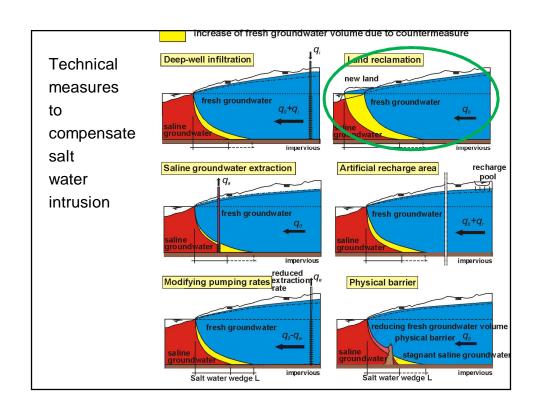
- Monitoring of salinities and piezometric levels
- Numerical modelling of salt water intrusion

Measures to compensate salt water intrusion

- 'The Fresh Holder'
- · Extraction of saline/brackish groundwater
- · Infiltration of fresh surface water
- Modifying pumping rates
- · Land reclamation in front of the coast
- Creating physical barriers (chrystallisation or biosealing)





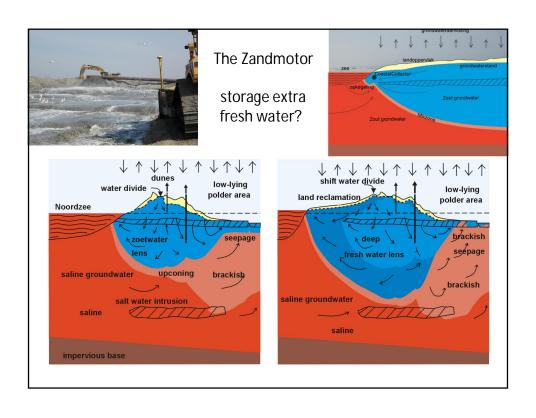




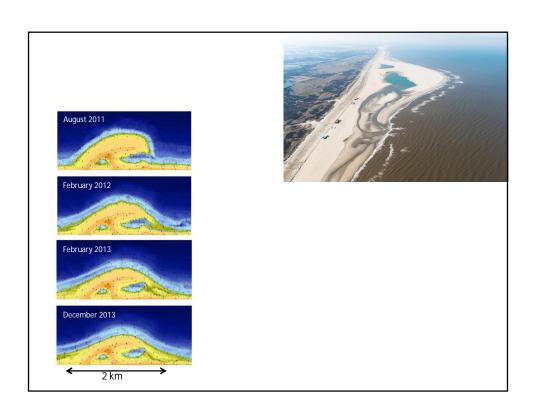


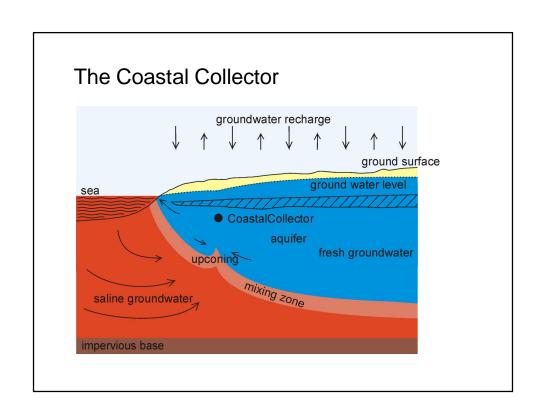


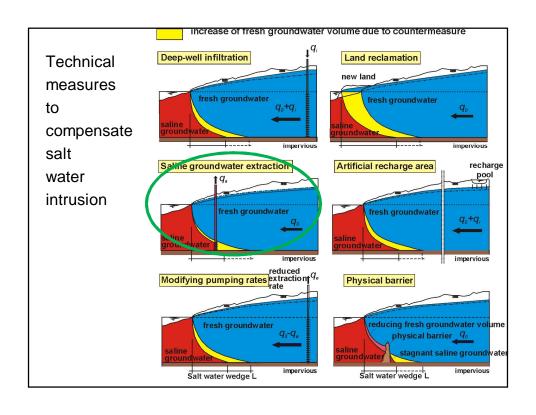


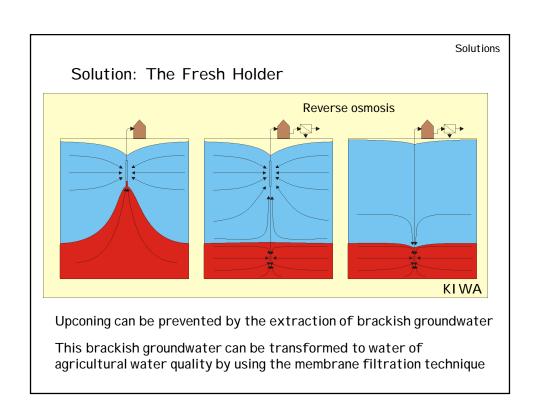


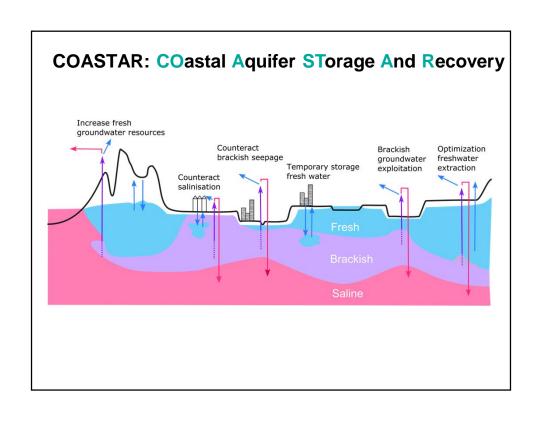


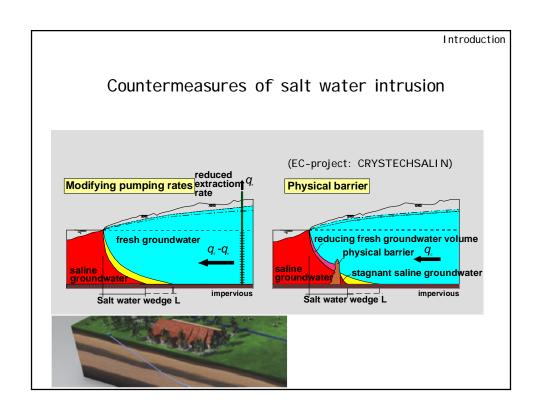


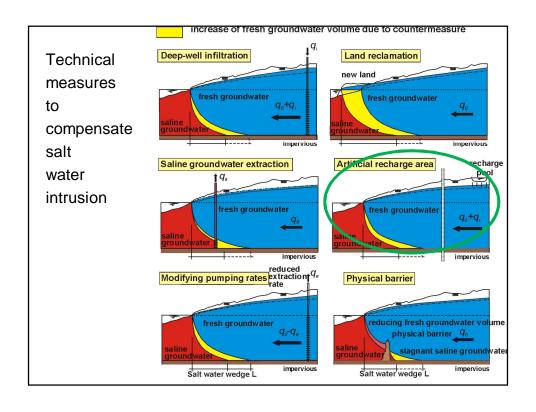






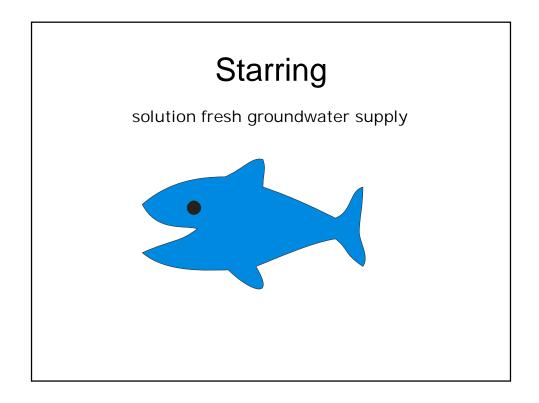






Base idea

Many local solutions for fresh groundwater supply can have regional impact



Starring

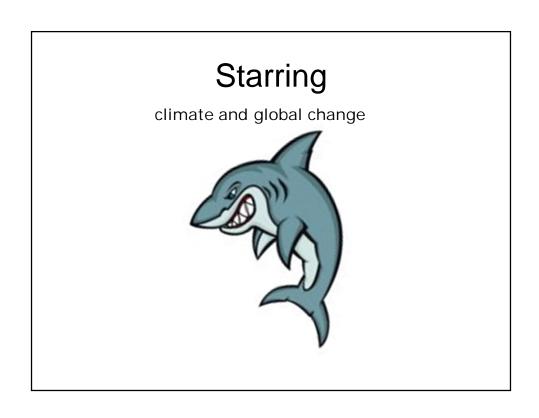
Local solution fresh groundwater supply

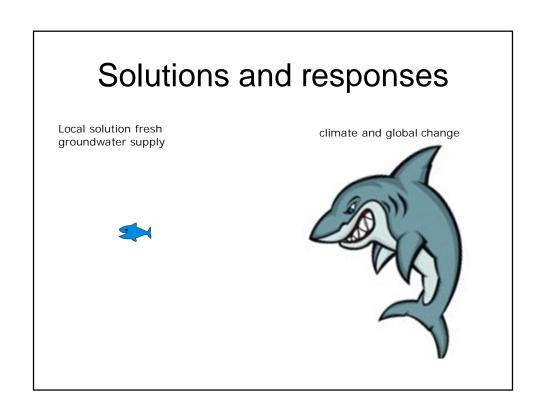


Starring

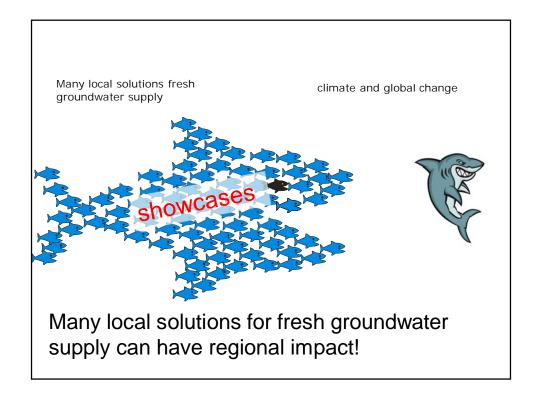
climate and global change

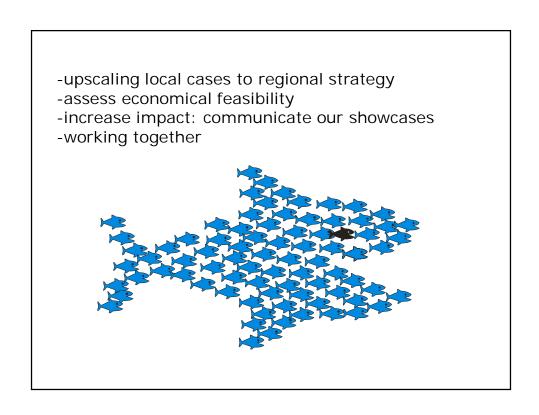


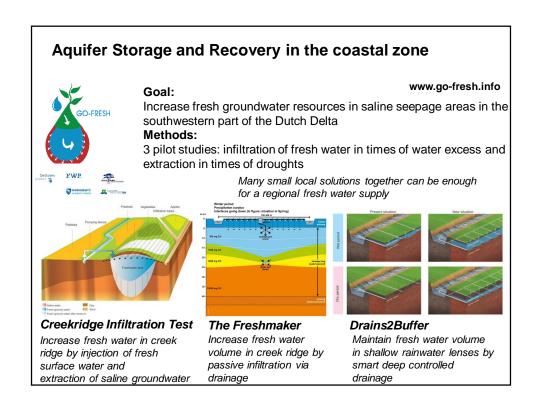




What should be the response?





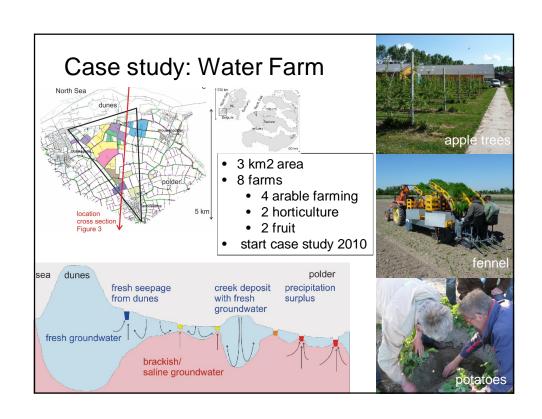


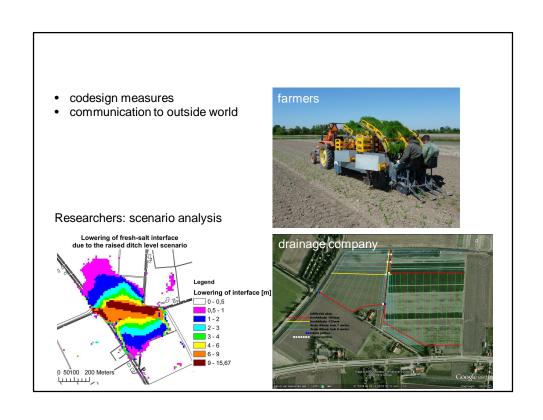
Problem statement

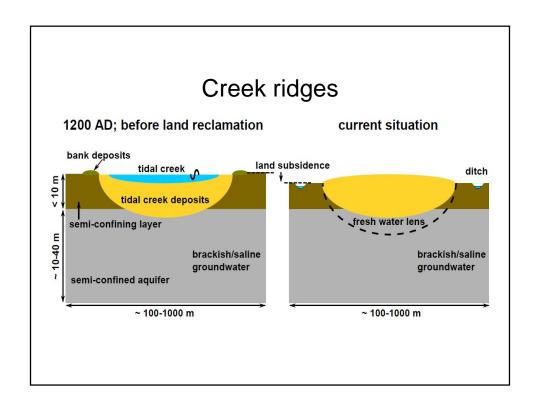
- Crop damage southwestern part of the Netherlands
- Fresh groundwater below creek ridges

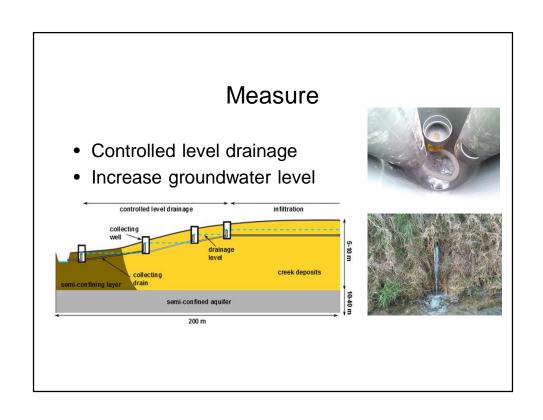


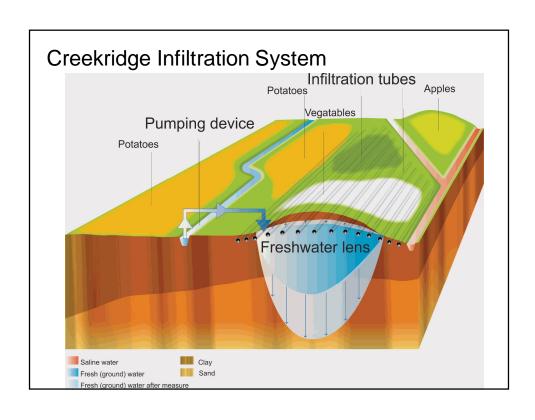


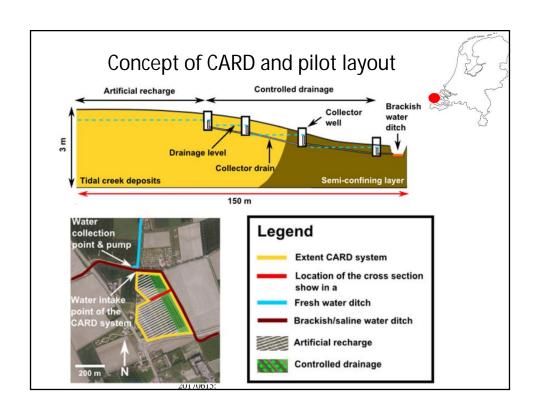


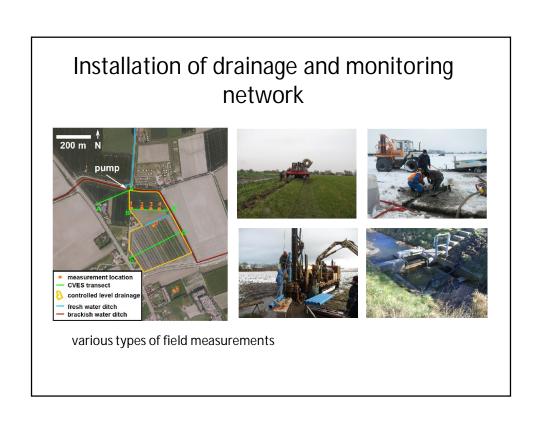












Different types of field measurements applied

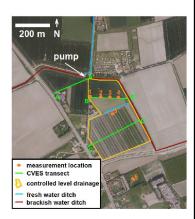
 $\begin{array}{cccc} \hline Pressure \ transducers^a & Groundwater \ levels \\ Sampling \ using & EC_{w20} \\ piezometer \ nest \\ SLIMFLEX^b & EC_{bulk} \\ CPT^c & Lithology \ and \ EC_{bulk} \\ CVES^d & EC_{bulk} \\ SMD^e & EC_{bulk} \end{array}$

Purpose

- a. Schlumberger, The Netherlands (type 'Diver')
- b. Deltares, The Netherlands
- c. Fugro, The Netherlands
- · d. ABEM, Sweden

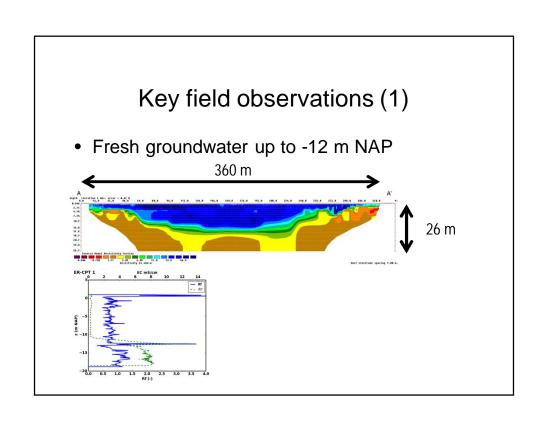
Measurement type

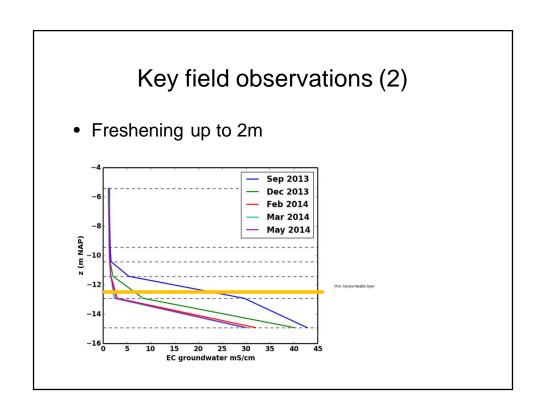
e. Imageau, France

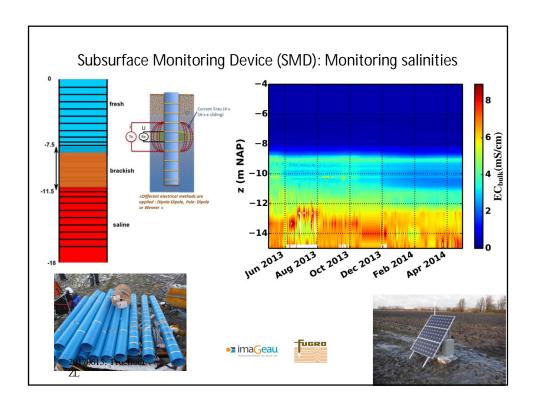


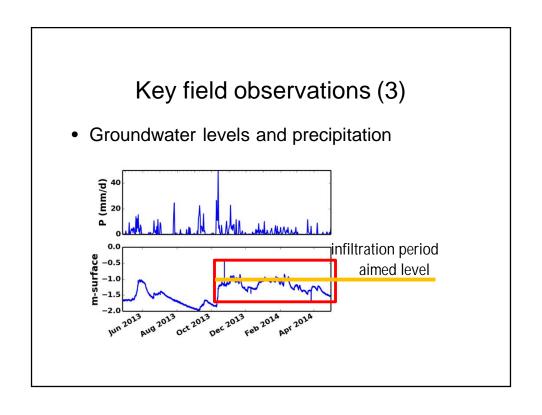


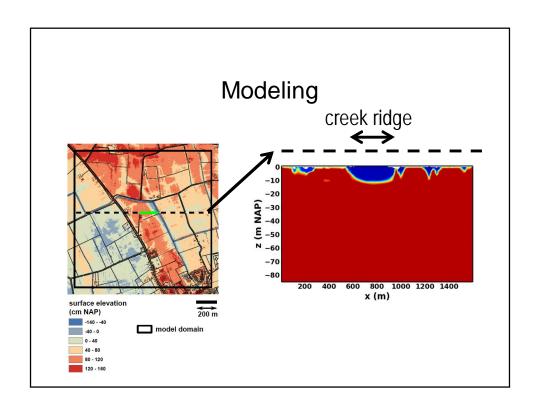


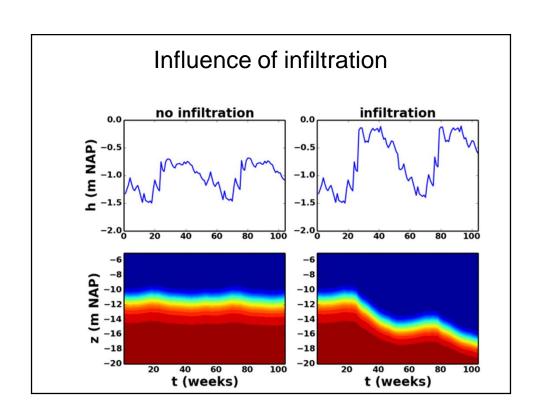












Singapore Jurong Island

Aquifer Storage and Recovery



