

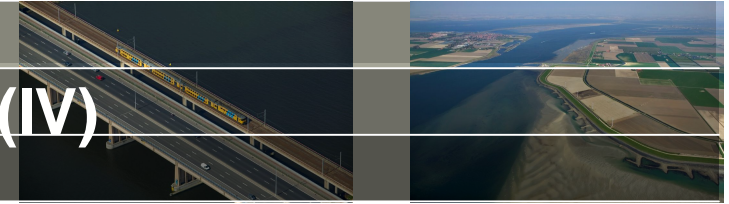


# Zoetzout grondwater cursus, deel 1

Gualbert Oude Essink  
Bodem en Grondwater Systemen  
Deltares

**[zoetzout.deltares.nl](http://zoetzout.deltares.nl)**

## Afleiding numerieke dispersie: 1D (IV)



$$\frac{\partial^2 C}{\partial t^2}$$

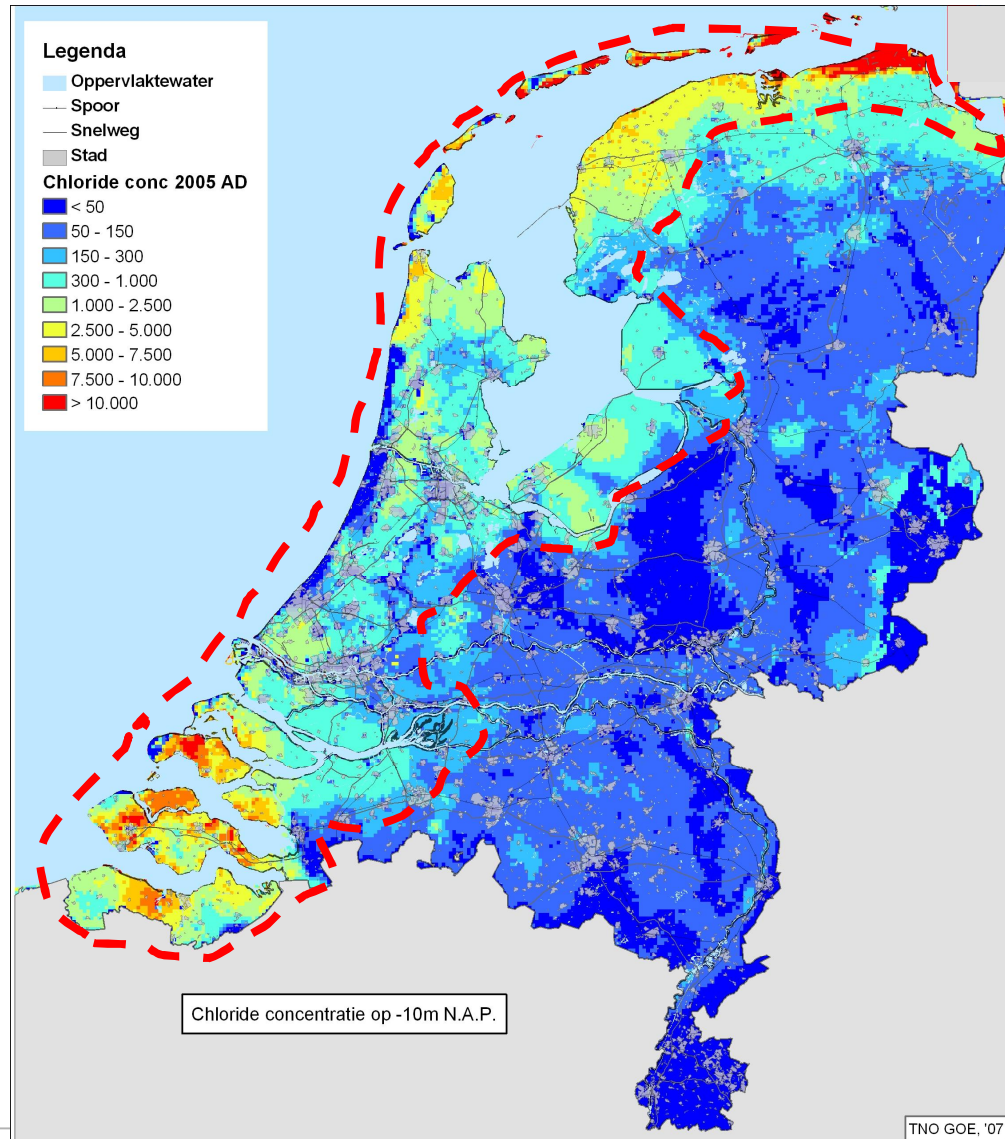
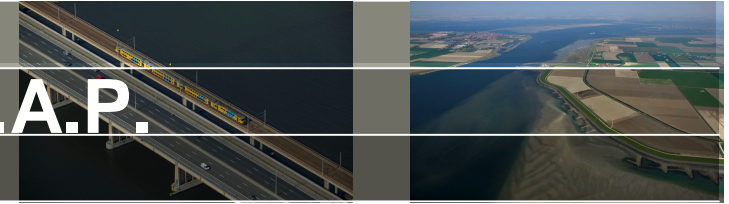
$$\frac{\partial^2 C}{\partial t^2} = \frac{\partial}{\partial t} \left( \frac{\partial C}{\partial t} \right) = \frac{\partial}{\partial t} \left( D \frac{\partial^2 C}{\partial x^2} - V \frac{\partial C}{\partial x} \right) = D \frac{\partial^2}{\partial x^2} \left( \frac{\partial C}{\partial t} \right) - V \frac{\partial}{\partial x} \left( \frac{\partial C}{\partial t} \right)$$

$$\frac{\partial^2 C}{\partial t^2} = D \frac{\partial^2}{\partial x^2} \left( D \frac{\partial^2 C}{\partial x^2} - V \frac{\partial C}{\partial x} \right) - V \frac{\partial}{\partial x} \left( D \frac{\partial^2 C}{\partial x^2} - V \frac{\partial C}{\partial x} \right)$$

$$\frac{\partial^2 C}{\partial t^2} = D^2 \frac{\partial^4 C}{\partial x^4} - VD \frac{\partial^3 C}{\partial x^3} - VD \frac{\partial^3 C}{\partial x^3} + V^2 \frac{\partial^2 C}{\partial x^2}$$

$$\frac{\partial^2 C}{\partial t^2} \approx V^2 \frac{\partial^2 C}{\partial x^2}$$

# Chloride verdeling op -10m N.A.P.

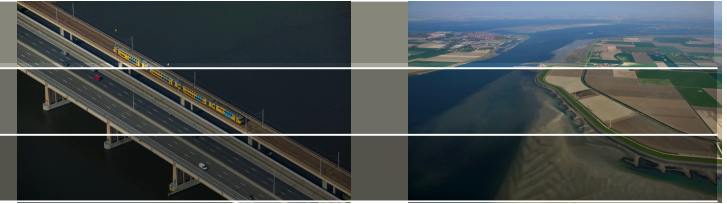


gebied waar de chloride concentratie  
In de ondergrond zodanig hoog is  
dat dichtheidseffecten op de  
grondwaterstroming waarschijnlijk  
een rol spelen

Drinkwater norm: <math>< 150</math> mg Cl-/l  
Oppervlaktewater: 200 mg Cl-/l  
Noordzee: ~ 16800 mg Cl-/l  
Stille Oceaan: ~19000 mg Cl-/l  
Middellandse Zee: ~22000 mg Cl-/l

**Deltares**

# Verzoet of verzilt Nederland?

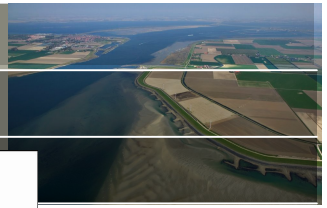
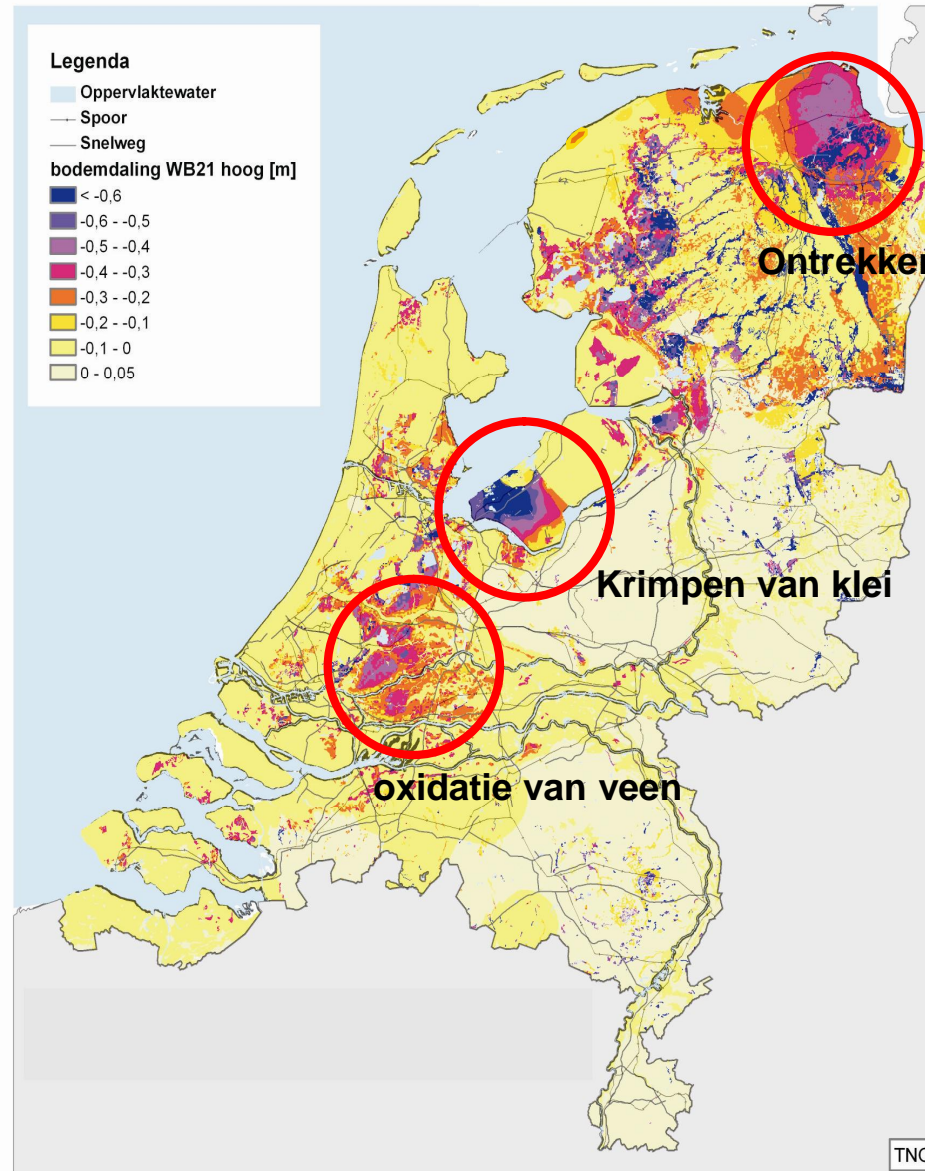


## Oorzaken:

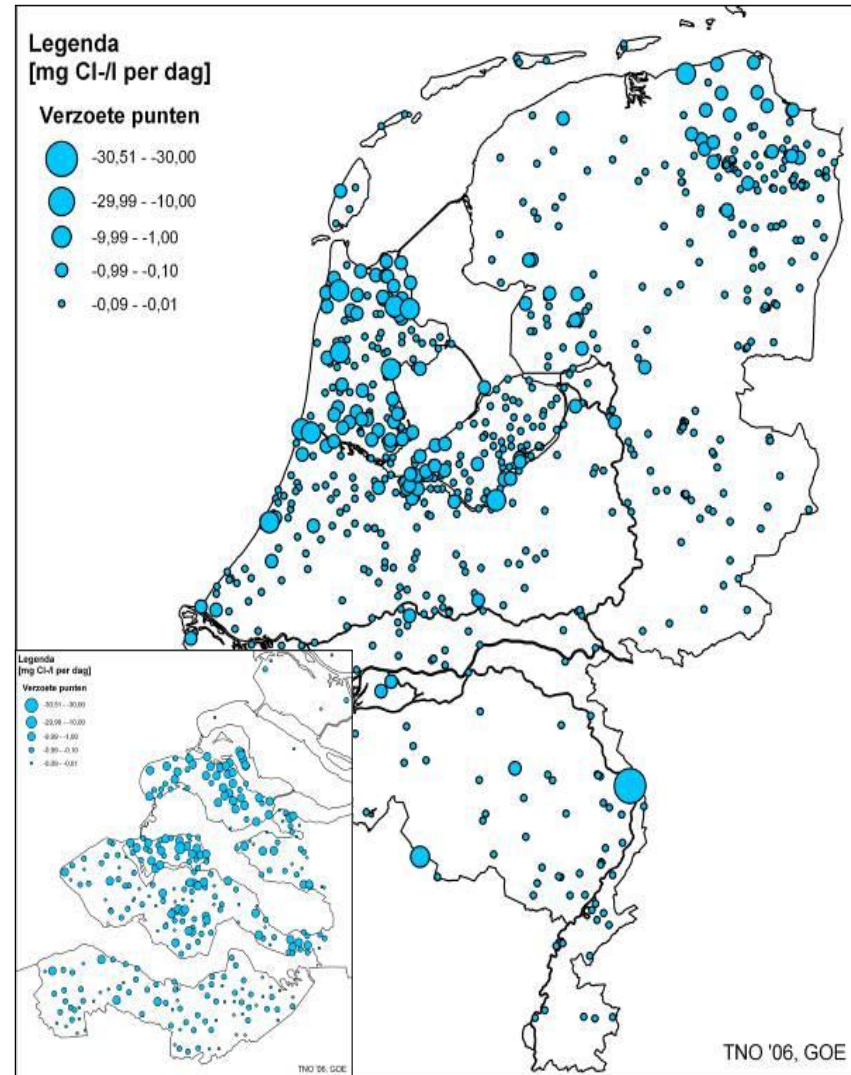
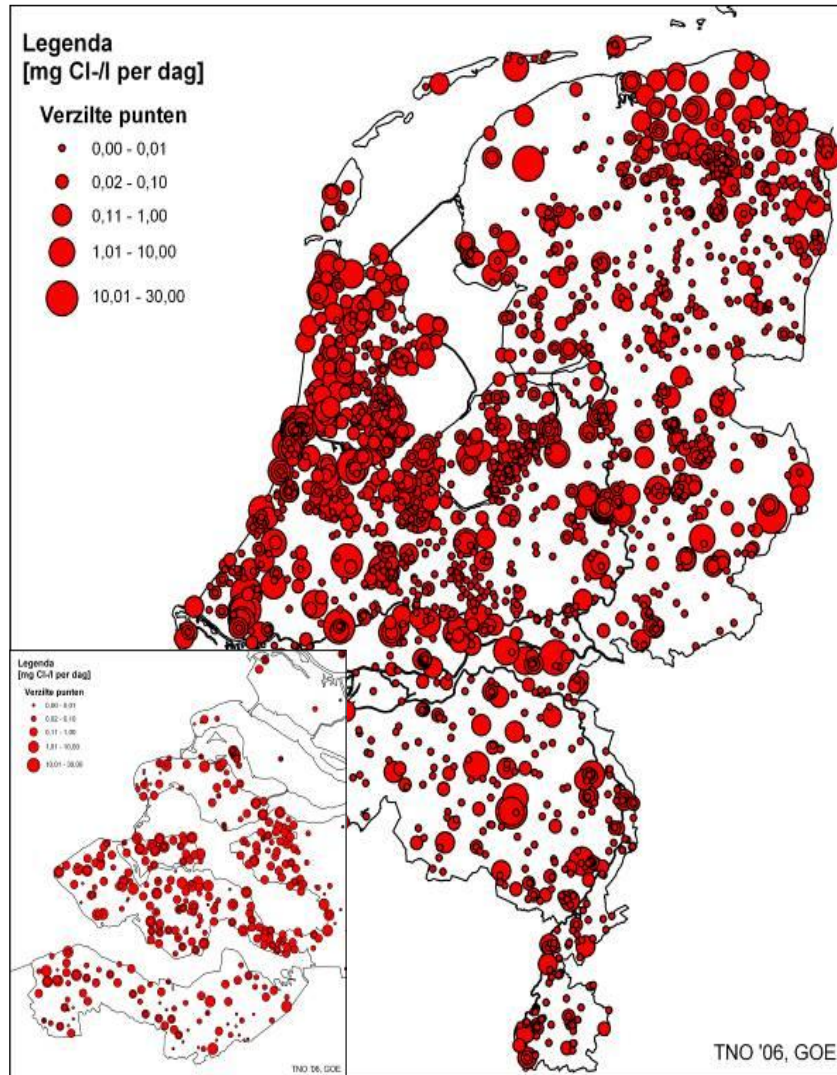
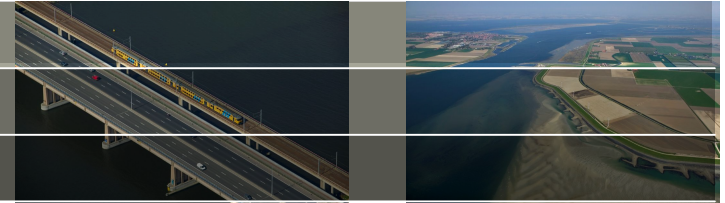
1. Autonome effecten
  - bodemdaling
  - inpolderingen
  - land onder NAP
  
2. Menselijk ingrijpen
  - peilbeheer
  - grondwateronttrekkingen
  - van zoet naar zout (VZM, Waterdunen)
  
3. Klimaatverandering
  - zeespiegelstijging
  - grondwateraanvulling



# Toekomstige bodemdaling, 2050 AD

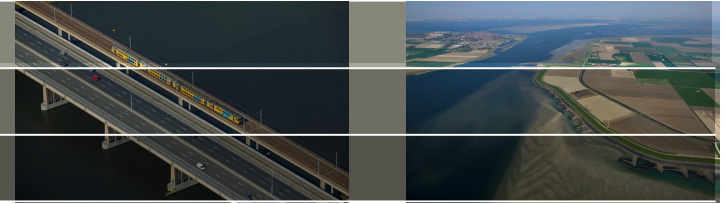


# Verzoet of verzilt Nederland?





# Concept van een kustgebied



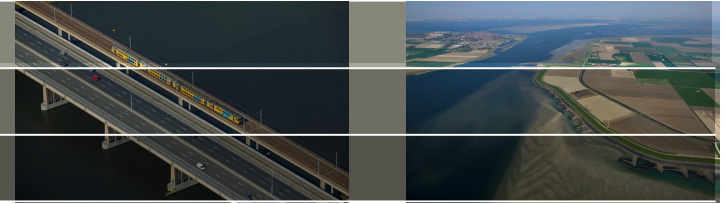




# Poreuze media



# De Wet van Darcy



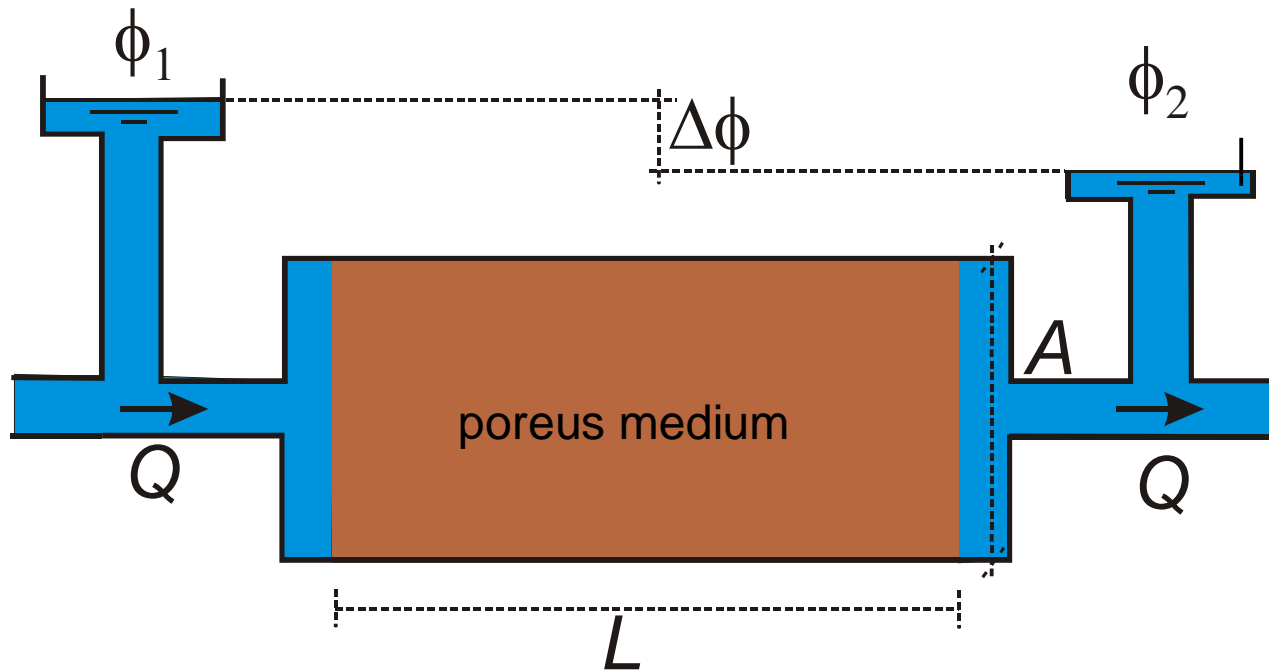
De beweging van water door een poreus medium is grondwaterstroming

Het kan worden beschreven door de bewegingsvergelijking:

De Wet Van Darcy

# I. Wet van Darcy (1856)

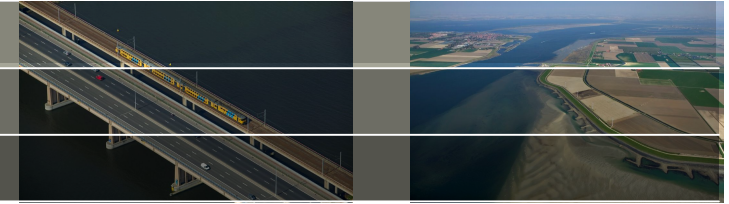
$$Re = \frac{\rho g R}{\mu} < 1-10$$



$$Q \propto \phi_1 - \phi_2 \quad Q \propto \frac{1}{L} \quad Q \propto A \quad \text{geeft} \quad Q \propto A \frac{\phi_1 - \phi_2}{L}$$

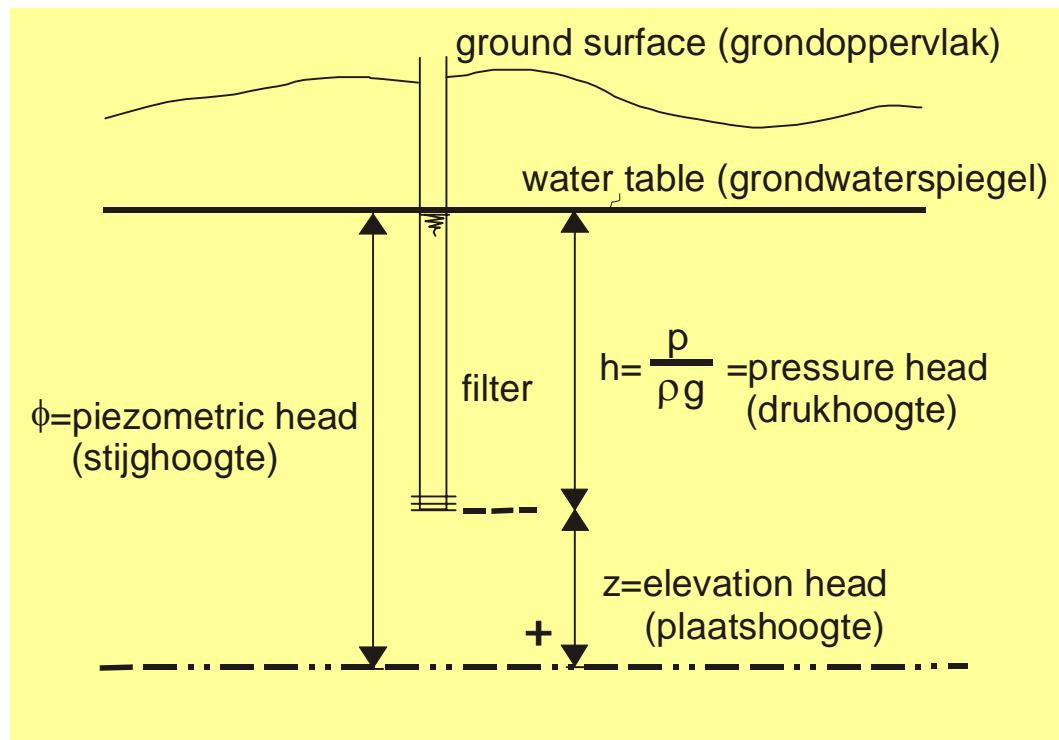
$$Q = KA \frac{\phi_1 - \phi_2}{L} \quad \text{waarin } K = \text{doorlatendheid [L/T]}$$

## II. Wet van Darcy



In drukken:  $q_x = -\frac{\kappa_x}{\mu} \frac{\partial p}{\partial x}$      $q_y = -\frac{\kappa_y}{\mu} \frac{\partial p}{\partial y}$      $q_z = -\frac{\kappa_z}{\mu} \left( \frac{\partial p}{\partial z} + \rho g \right)$

Definitie stijghoogte:  $\phi = \frac{p}{\rho g} + z \Leftrightarrow p = \rho g(\phi - z)$



$p$ =druk [ $M/LT^2$ ]

$\kappa$ =intrinsieke permeabiliteit [ $L^2$ ]

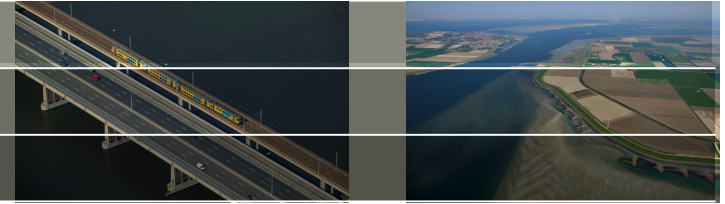
$\mu$ =dynamische viscositeit [ $M/LT$ ]

$\rho$ =dichtheid [ $M/L^3$ ]

$g$ =versnelling [ $L/T^2$ ]

$k_f$ =hydraulische conductiviteit [ $L/T$ ]

### III. Wet van Darcy

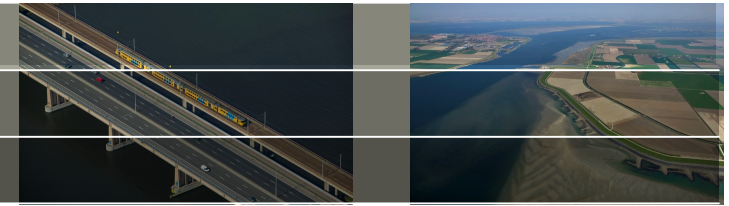


geeft  $q_x = -\frac{\kappa_x \rho g}{\mu} \frac{\partial \phi}{\partial x}$      $q_y = -\frac{\kappa_y \rho g}{\mu} \frac{\partial \phi}{\partial y}$      $q_z = -\frac{\kappa_z \rho g}{\mu} \frac{\partial \phi}{\partial z}$

als  $\frac{\kappa_i \rho g}{\mu} = k_i$  dan:

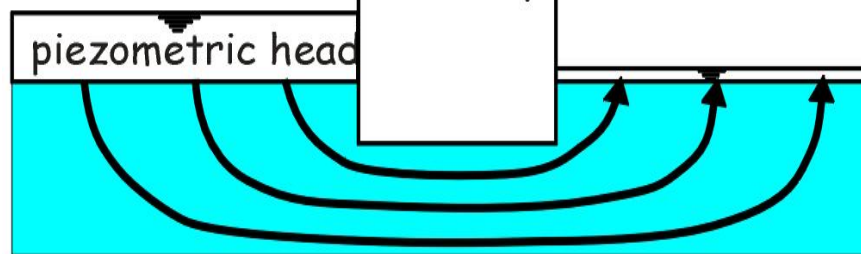
geeft  $q_x = -k_x \frac{\partial \phi}{\partial x}$      $q_y = -k_y \frac{\partial \phi}{\partial y}$      $q_z = -k_z \frac{\partial \phi}{\partial z}$

# Analogy physical processes

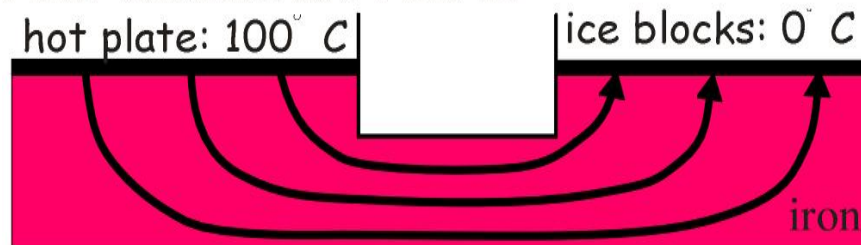


Heat transport (analogy with solute transport)

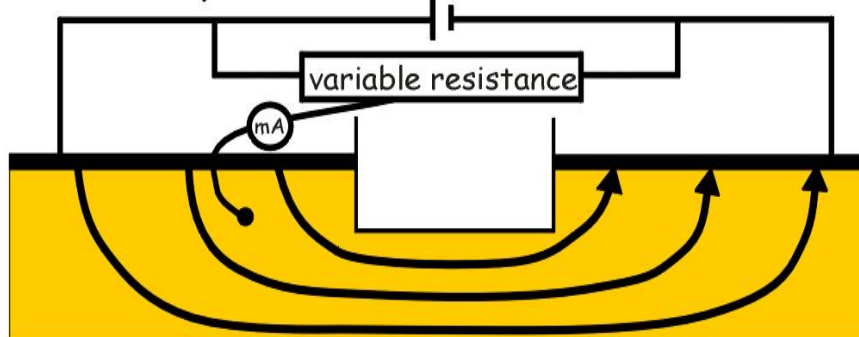
Groundwater flow: Darcy



Heat conduction: Fourier



Electrodynamics: Ohm



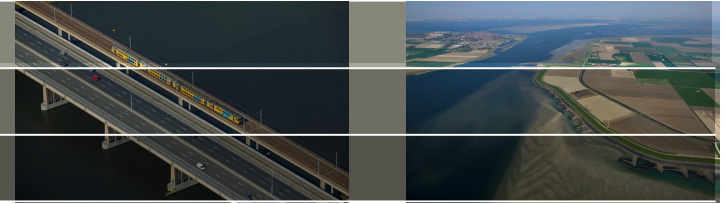
$$q = -k \frac{\partial \phi}{\partial x}$$

$$h = -\lambda \frac{\partial T}{\partial x}$$

$$i = -\sigma \frac{\partial V}{\partial x}$$



# Darcy's law



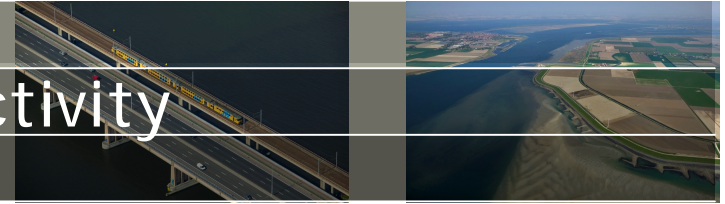
Darcy's law is only valid in the case of laminar flow: viz. at relatively low velocities when water particles move more or less parallel to each other.

In other words, Darcy's law is valid as long as the so-called Reynolds number  $Re$  (-) does not exceed some value between 1 and 10:

$$Re = \rho q R / \mu < 1 - 10$$

where  $R$  = hydraulic radius of the pore

# Permeability and hydraulic conductivity



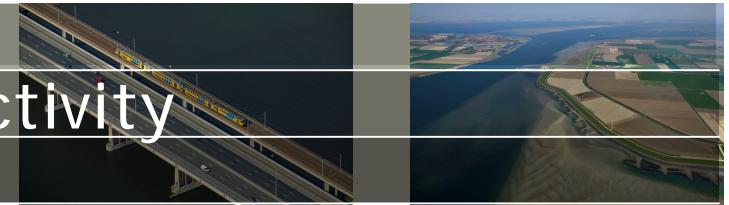
The intrinsic permeability  $\kappa$  largely depends on the size of the pores through the effective porosity  $n_e$ .

A commonly-used equation is that of Kozeny-Carmen to clearly demonstrate the relation between  $\kappa$  and  $n_e$ :

$$\text{Kozeny-Carmen: } \kappa = c d^2 n_e^3 / (1 - n_e)^2 \quad [m^2]$$

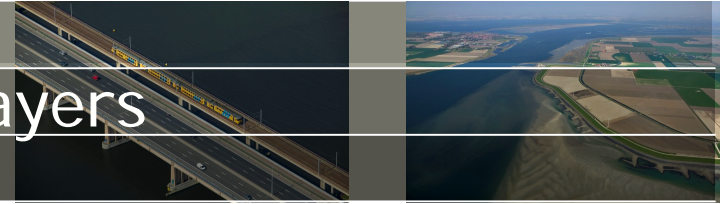
where  $c$ =coefficient depending on pore structure [-]  
 $d$ =main grain size [L]

# Permeability and hydraulic conductivity



Soil	$k$ (m/s)	$K$ (m <sup>2</sup> )
<b>Unconsolidated deposits</b>		
Clay	$<10^{-9}$	$<10^{-17}$
Sandy clay	$10^{-9}$ - $10^{-8}$	$10^{-16}$ - $10^{-15}$
Silt	$10^{-8}$ - $10^{-7}$	$10^{-15}$ - $10^{-14}$
Peat	$10^{-9}$ - $10^{-7}$	$10^{-16}$ - $10^{-14}$
Very fine sand	$10^{-6}$ - $10^{-5}$	$10^{-13}$ - $10^{-12}$
Fine sand	$10^{-5}$ - $10^{-4}$	$10^{-12}$ - $10^{-11}$
Coarse sand	$10^{-4}$ - $10^{-3}$	$10^{-11}$ - $10^{-10}$
Sand with gravel	$10^{-3}$ - $10^{-2}$	$10^{-10}$ - $10^{-9}$
Gravel	$>10^{-2}$	$>10^{-9}$
<b>Rocks</b>		
Unfractured rocks	$<10^{-9}$	$<10^{-17}$
Sandstone	$10^{-10}$ - $10^{-6}$	$10^{-17}$ - $10^{-13}$
Limestone & dolomite	$10^{-9}$ - $10^{-6}$	$10^{-16}$ - $10^{-13}$
Fractured rocks	$10^{-8}$ - $10^{-4}$	$10^{-15}$ - $10^{-11}$
Permeable basalt	$10^{-7}$ - $10^{-2}$	$10^{-14}$ - $10^{-9}$
Karst limestone	$10^{-6}$ - $10^{-2}$	$10^{-13}$ - $10^{-9}$

# Hydraulic conductivity of more layers



Horizontal hydraulic conductivity



$$k_h = \sum k_i \left( \frac{b_i}{b_t} \right)$$

Arithmetic mean (parallel)

Vertical hydraulic conductivity



$$k_h = \frac{b_t}{\sum \left( \frac{b_i}{k_i} \right)}$$

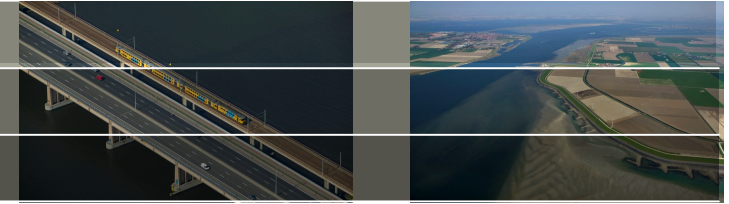
Harmonic mean (serial)



# Stoftransport



# Solute transport equation



Partial differential equation (PDE):

$$R_d \frac{\partial C}{\partial t} = \frac{\partial}{\partial x_i} \left( D_{ij} \frac{\partial C}{\partial x_j} \right) - \frac{\partial}{\partial x_i} (C V_i) + \frac{(C - C') W}{n_e} - R_d \lambda C$$

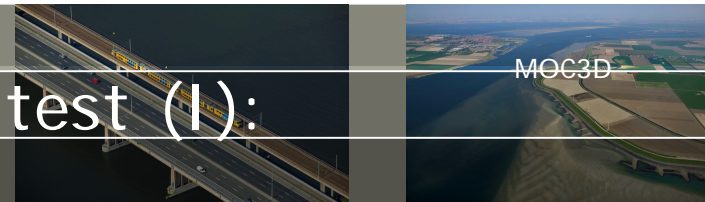
change in concentration      dispersion diffusion      advection      source/sink      decay

$D_{ij}$ =hydrodynamic dispersion [ $L^2 T^{-1}$ ]

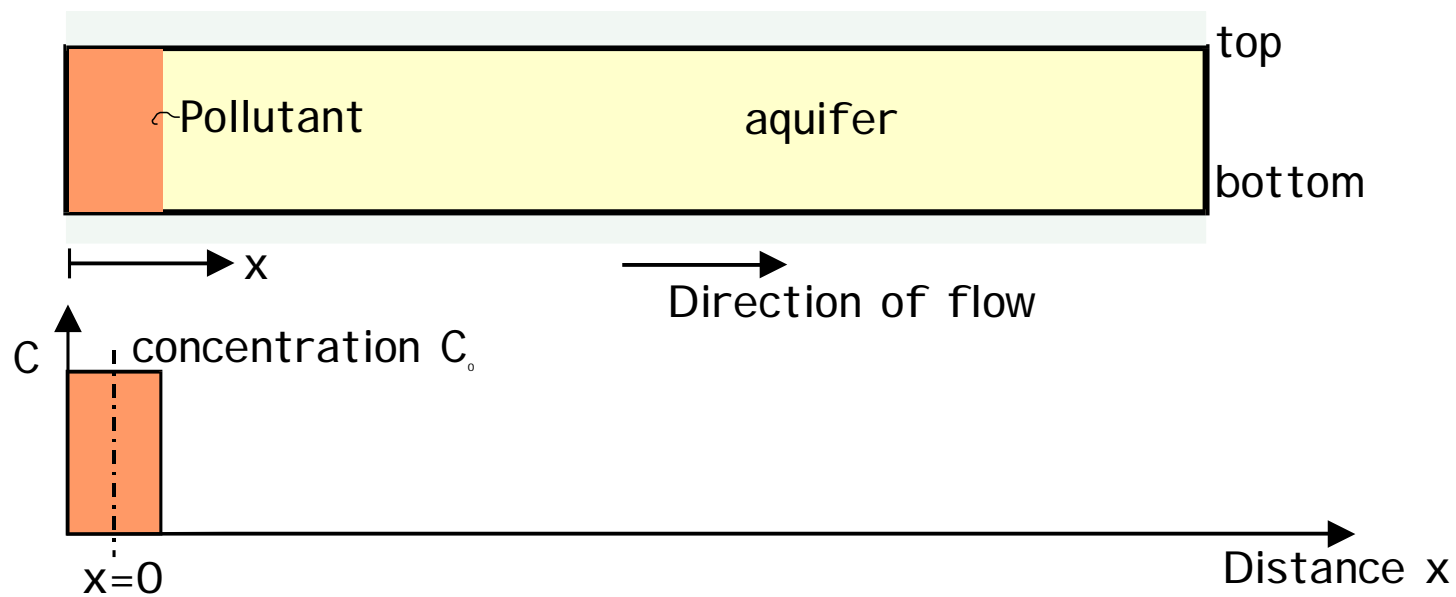
$R_d$ =retardation factor [-]

$\lambda$ =decay-term [ $T^{-1}$ ]

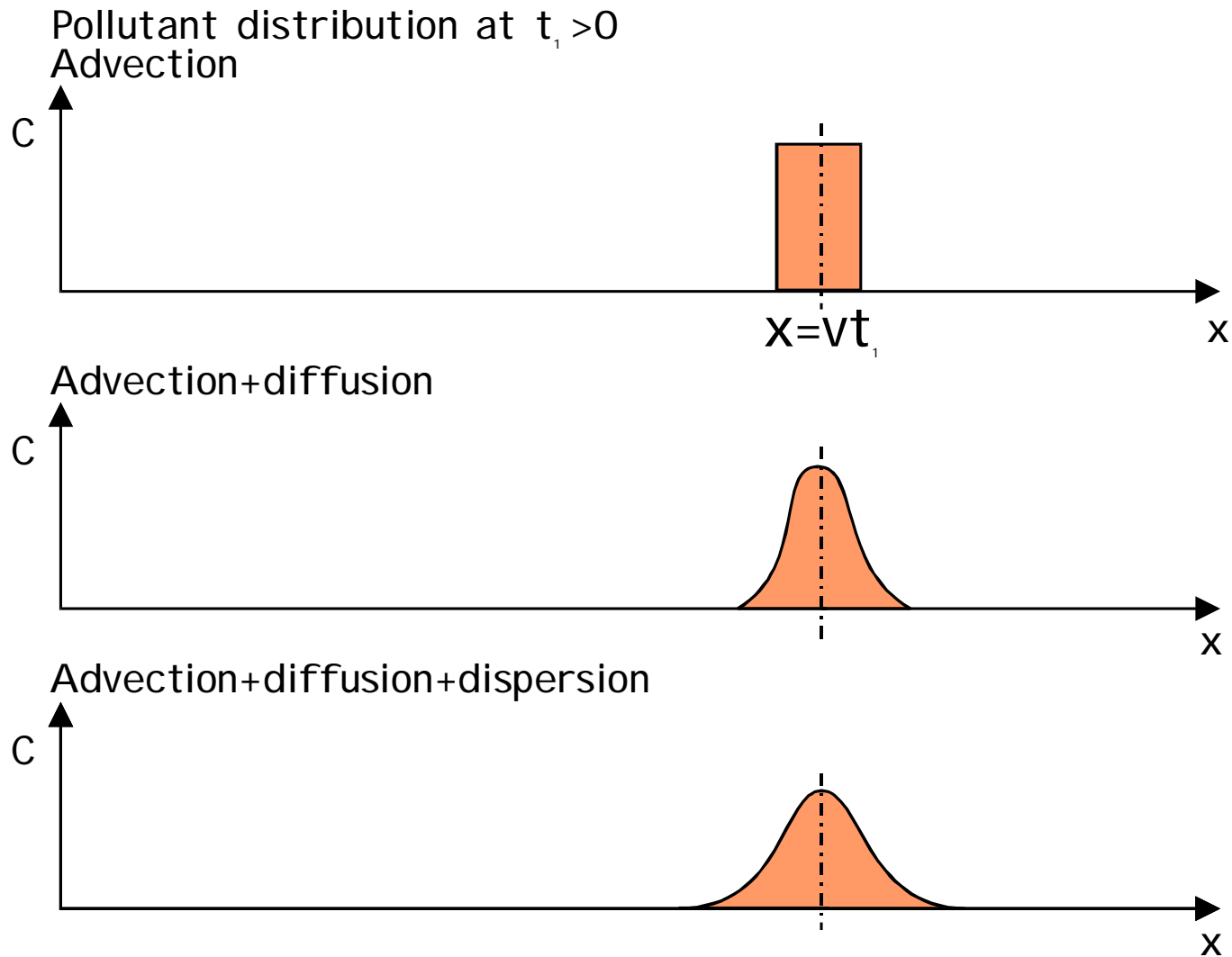
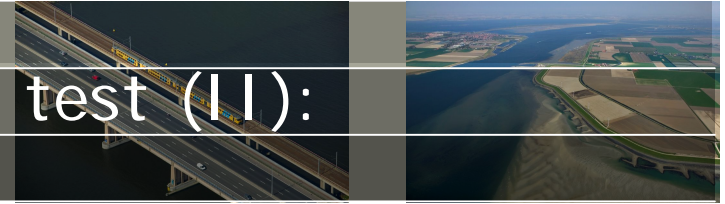
# Solute transport equation: column test (I):



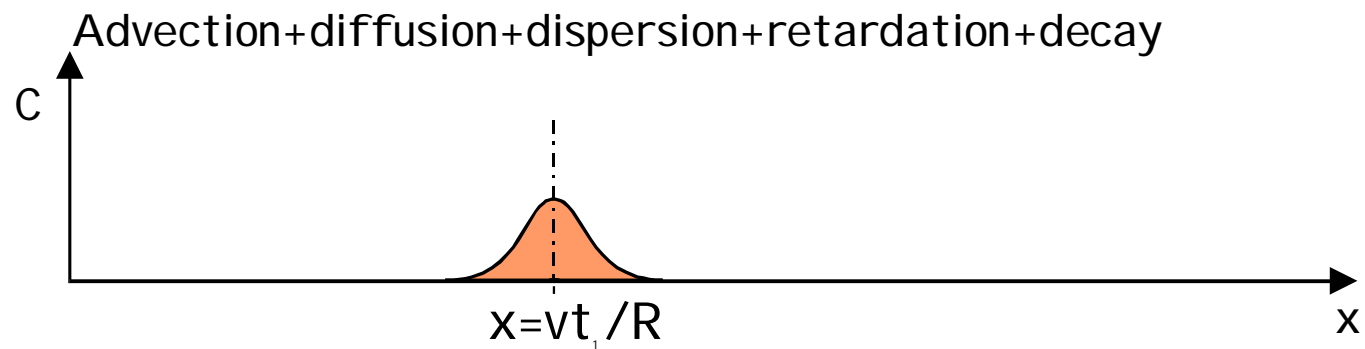
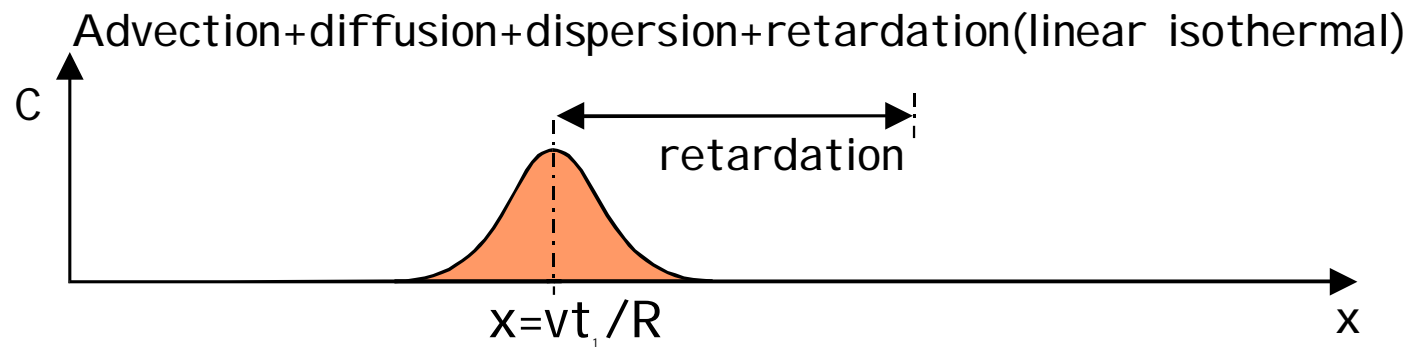
Pollutant distribution at  $t=0$



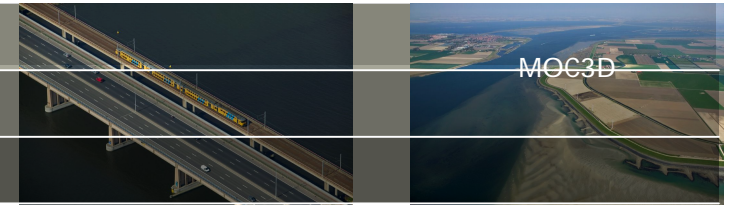
# Solute transport equation: column test (II):



# Solute transport equation: column test (III):



# Hydrodynamic dispersion



$$\text{hydrodynamic dispersion} \\ = \\ \text{mechanical dispersion} + \text{diffusion}$$

mechanical dispersion:

tensor (2D, 3D)

snelheidsafhankelijk

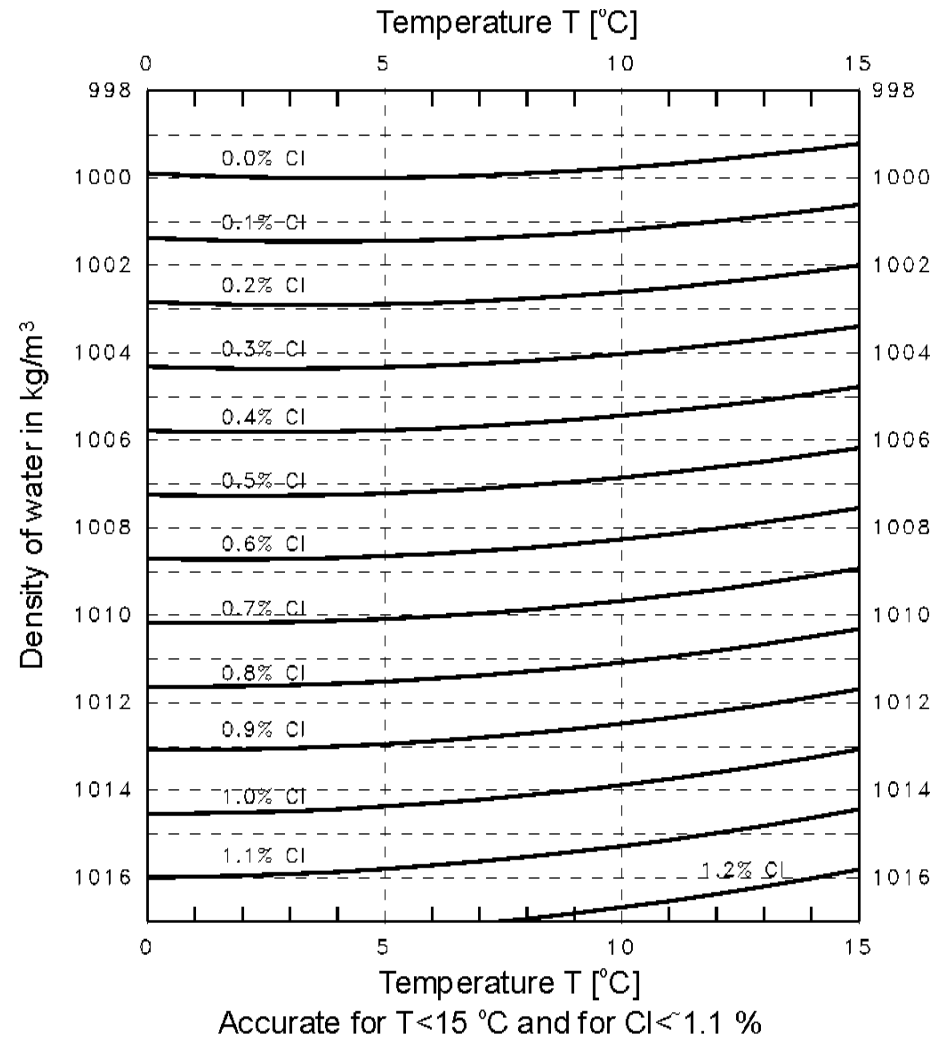
diffusion:

moleculair proces

stoffen verspreiden door concentratieverschillen



# Density depends on salinity and temperature



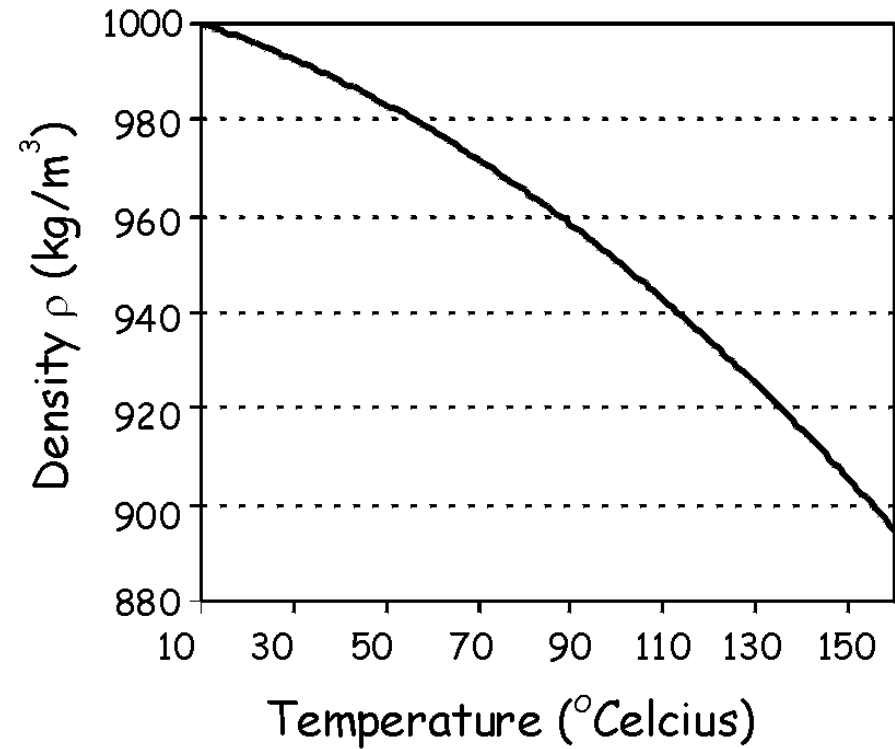
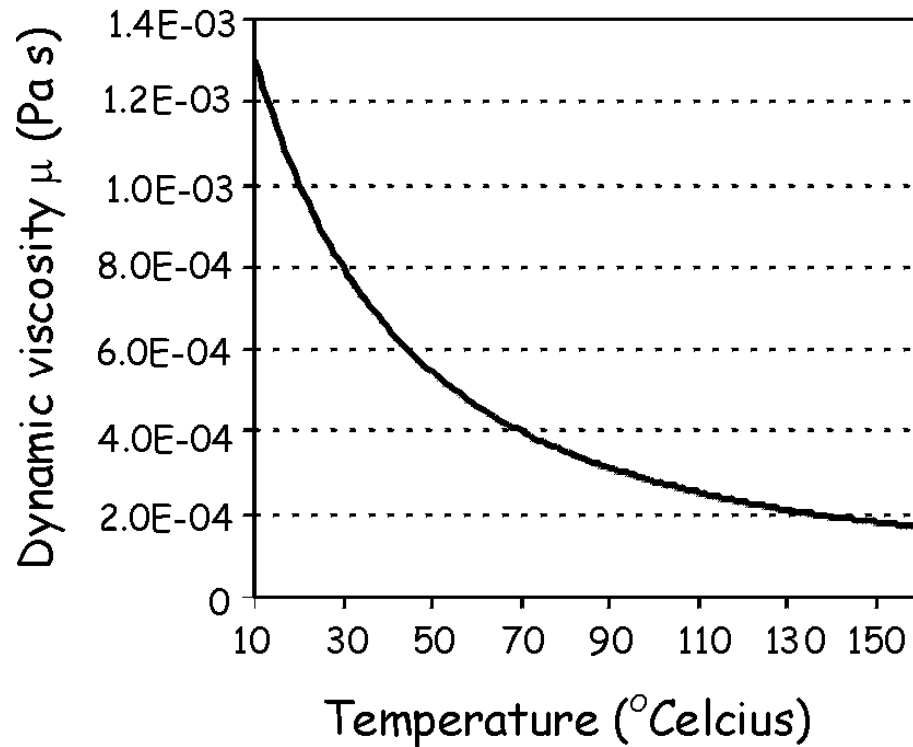
$$\rho_{(S,T)} = 1000 + 0.8054S - 0.0065(T - 4 + 0.2214S)^2 \quad \text{Knudsen (1902)}$$

**Deltares**

# Density and viscosity depend on temperature



(10°C-160 °C)



# Introductie grondwater in het kustgebied

## Concepten van zoutwater intrusie



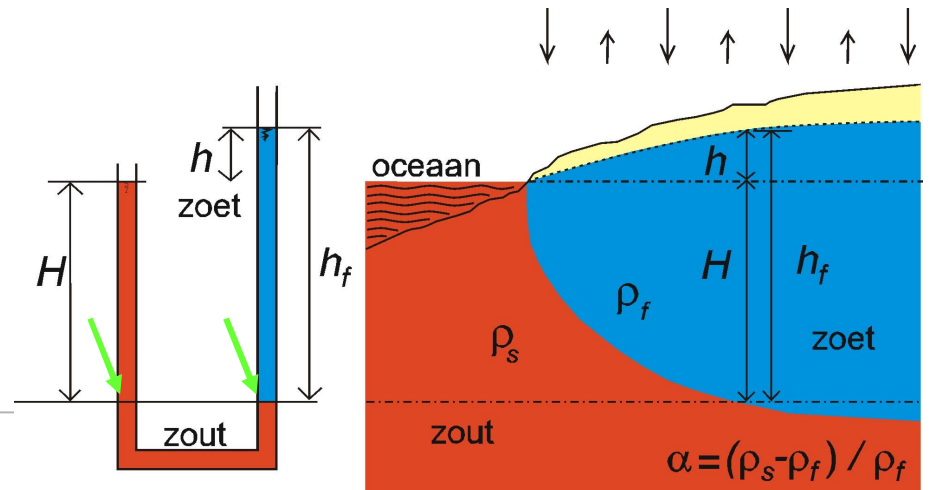


druk zout grondwater = druk zoet grondwater

$$\rho_s H g = \rho_f (H + h) g$$

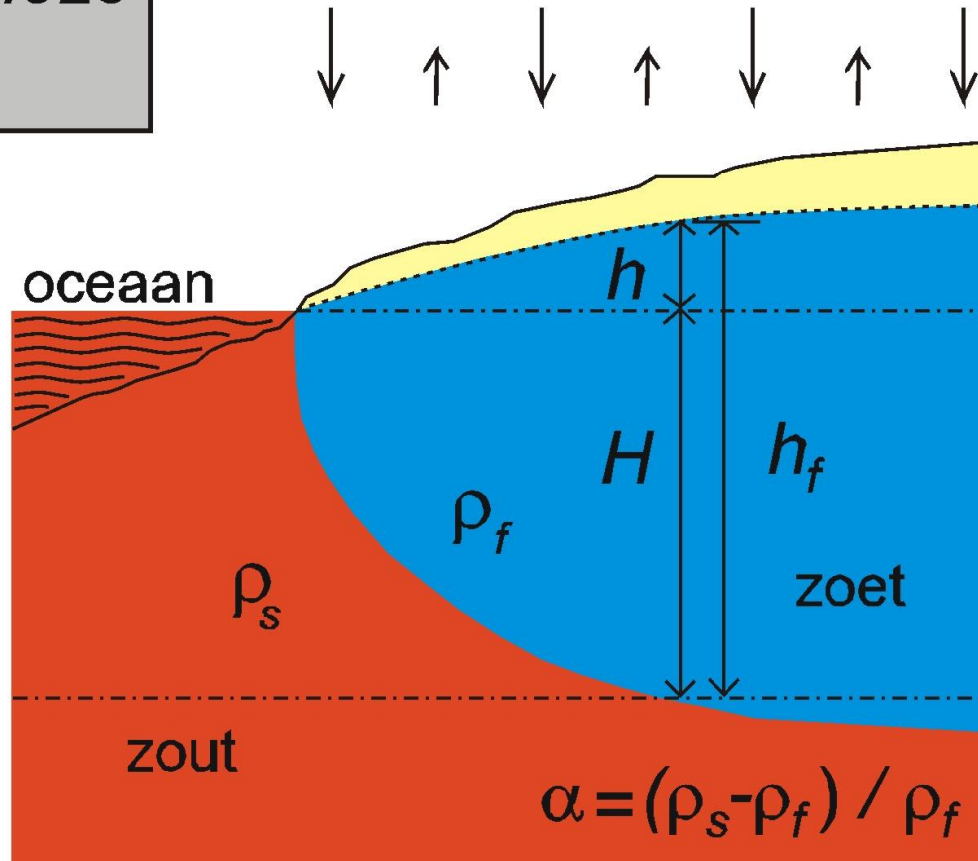
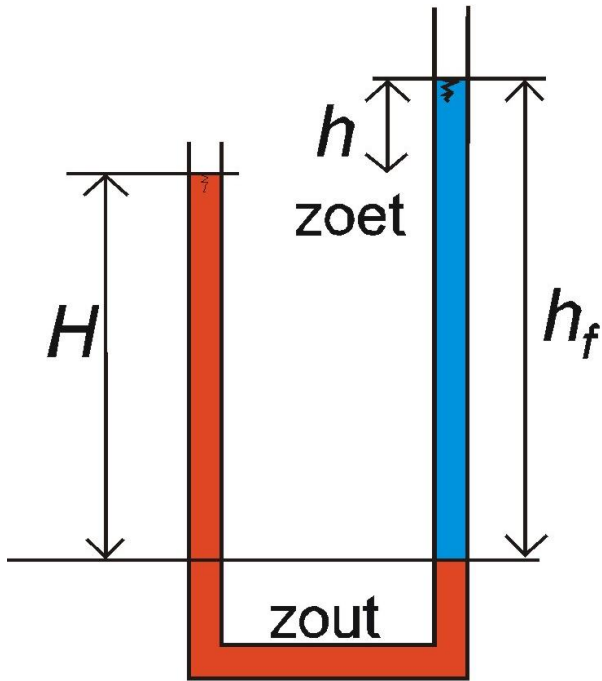
$$h = \frac{\rho_s - \rho_f}{\rho_f} H$$

$$h = \alpha H$$



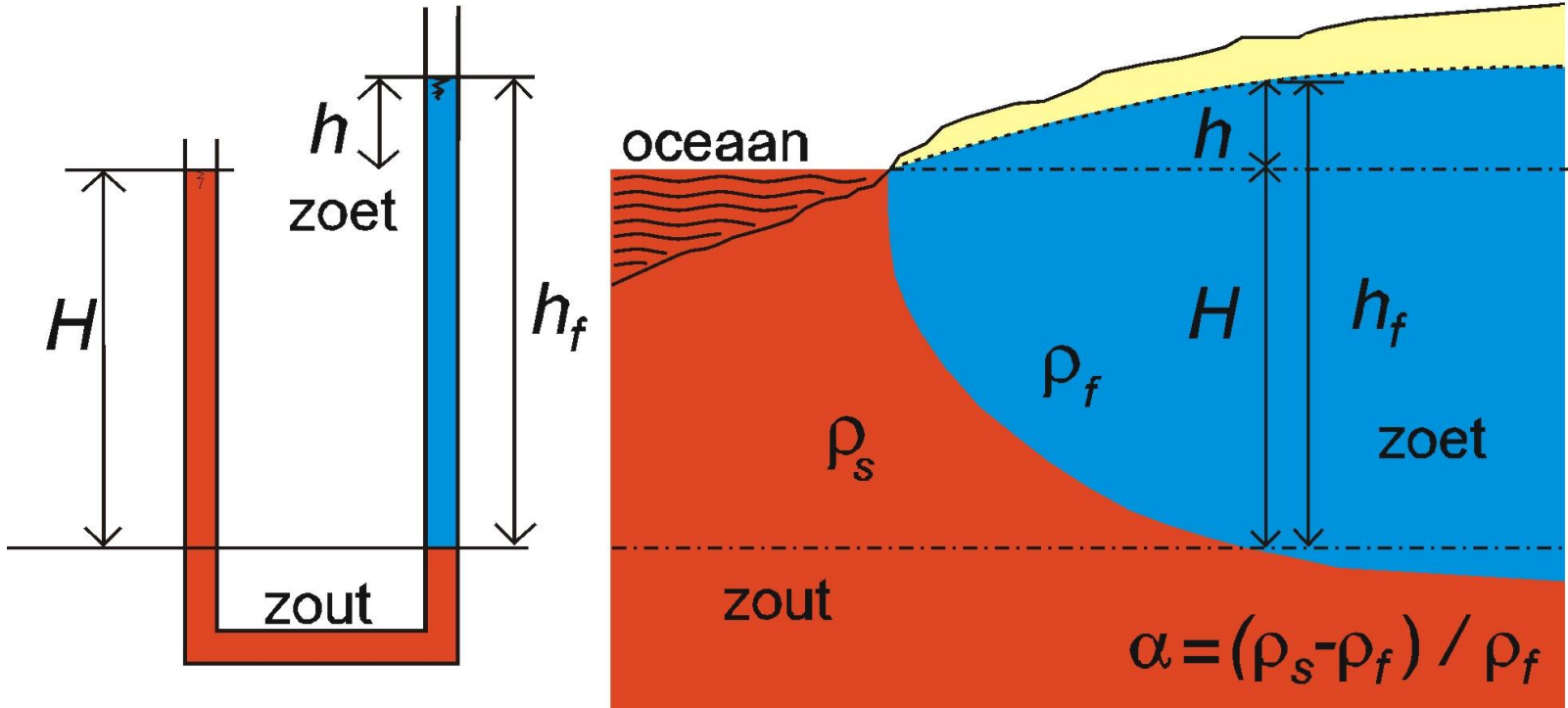


$h = \alpha H$   
 in ocean water  $\alpha = 0.025$   
 $h = 1 \text{ m}, H = 40 \text{ m}$



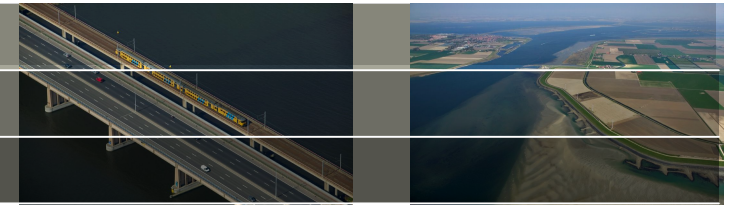


$h = \alpha H$   
 Mediterranean Sea  $\alpha = 0.028$   
 $h = 1 \text{ m}, H = 35.7 \text{ m}$

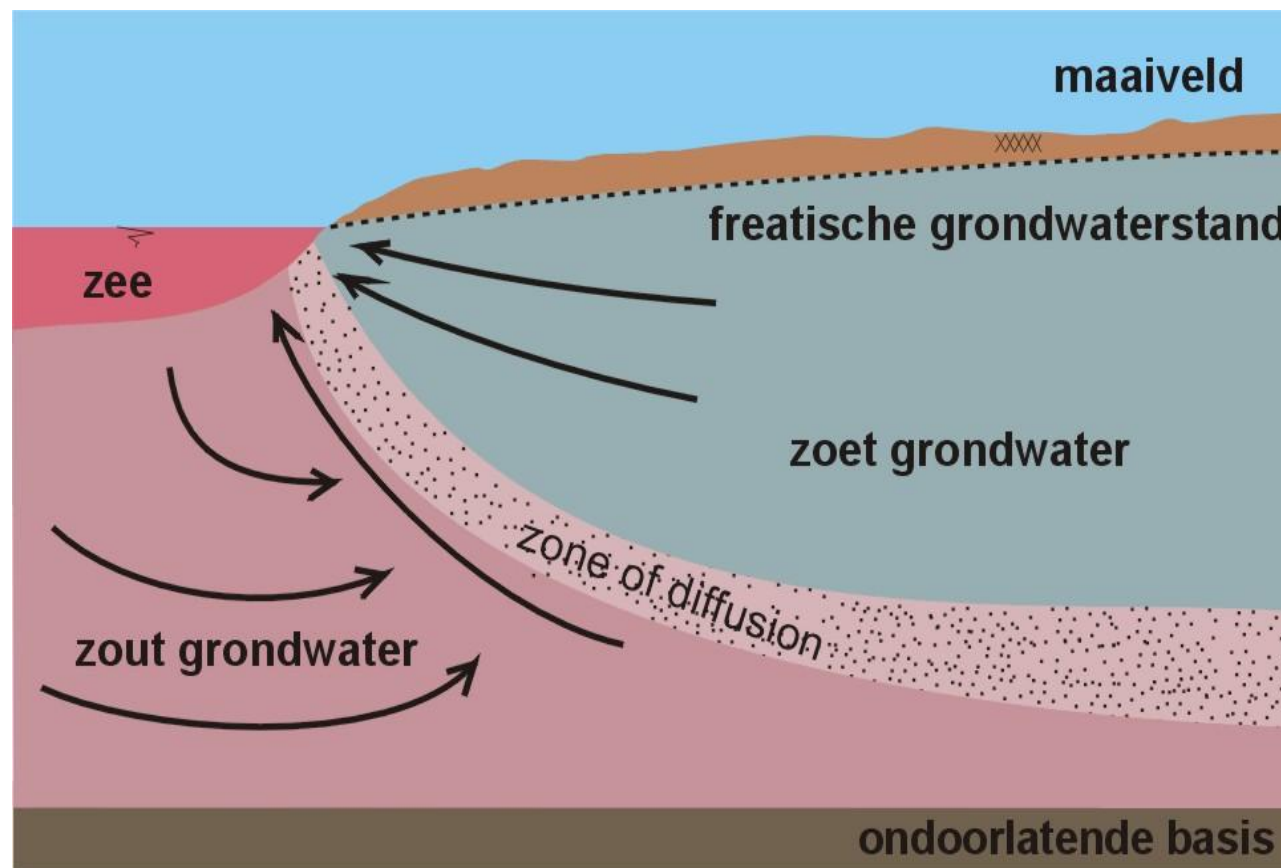




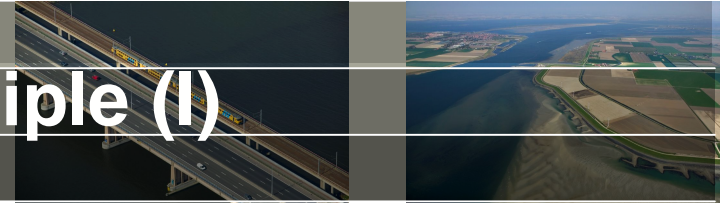
# Klassieke definitie zoutwaterintrusie



***Horizontale instroming van zout grondwater in een watervoerend pakket dat zoet grondwater bevat***

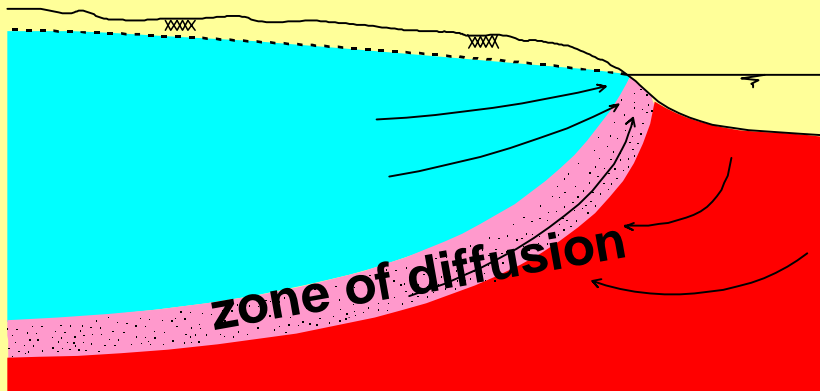


# Badon Ghyben-Herzberg principle (I)

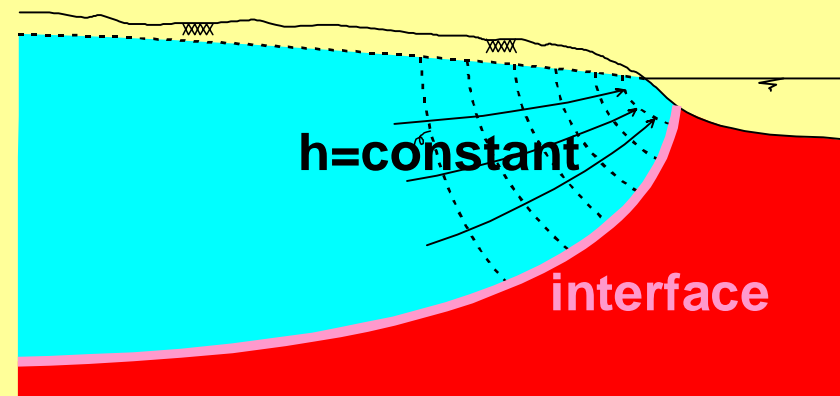


## Difference between reality and Badon Ghyben-Herzberg approximation

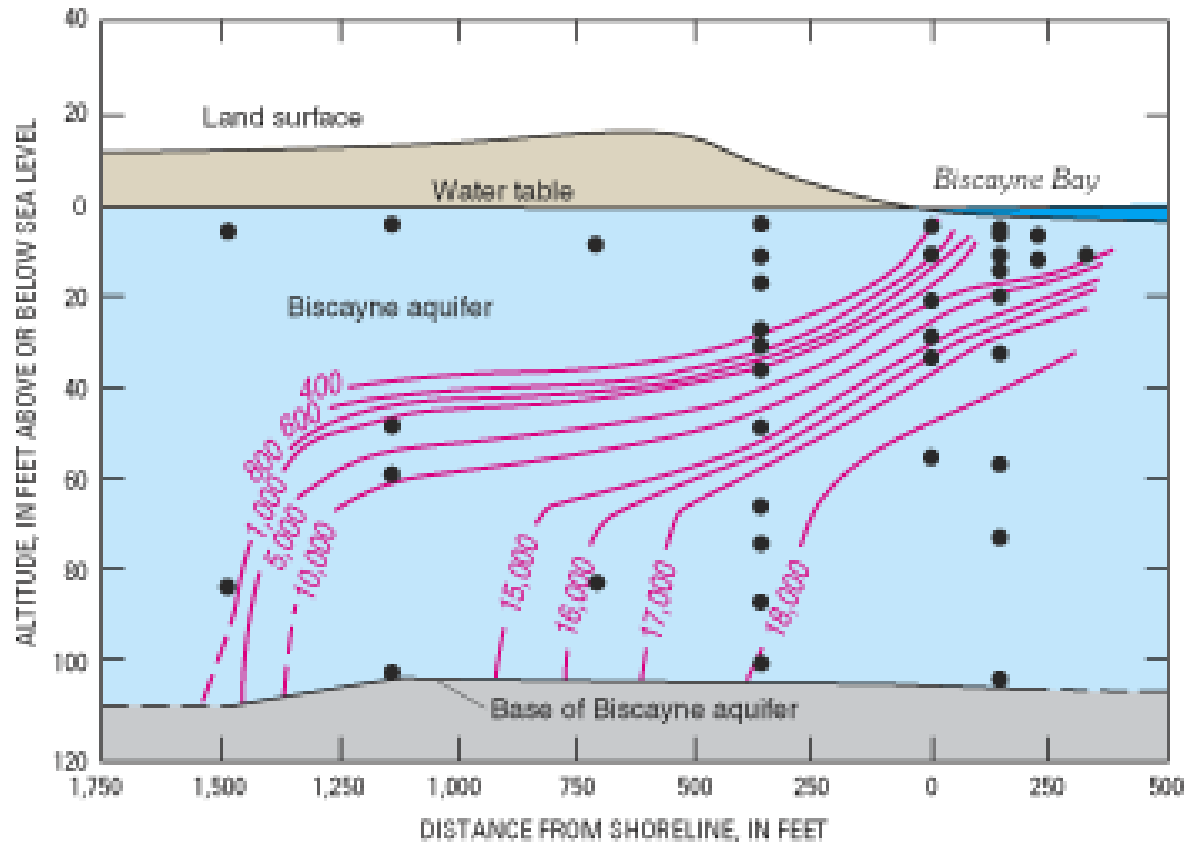
concept: mixing zone in reality





concept: interface between fresh and saline groundwater



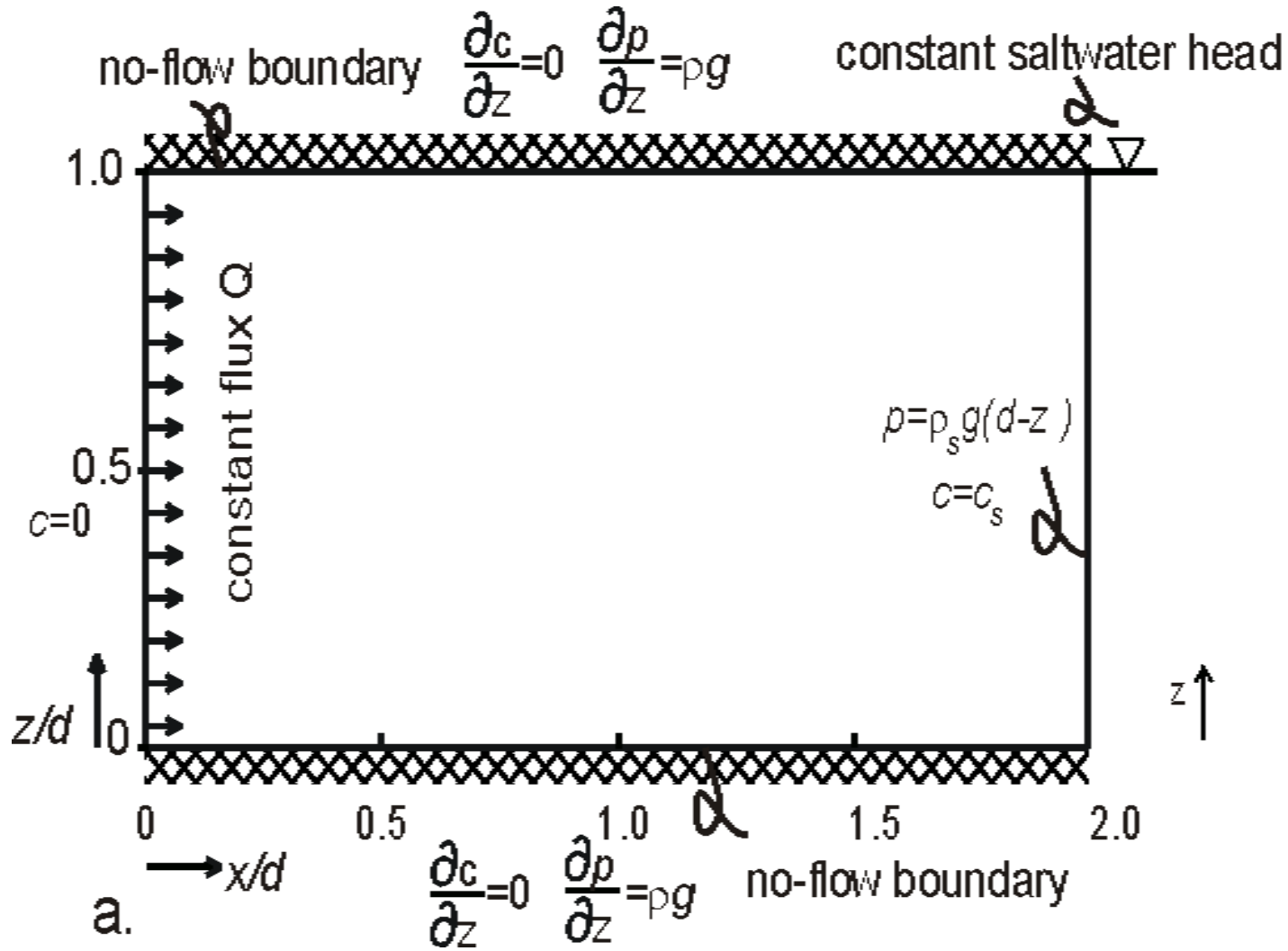
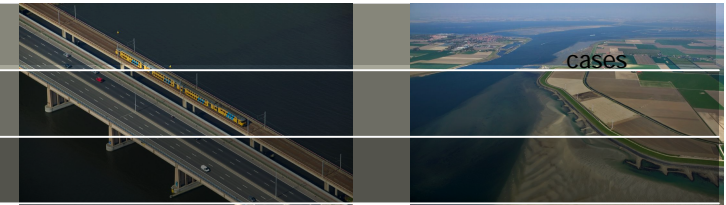
# Biscayne aquifer, Florida: Henry's case



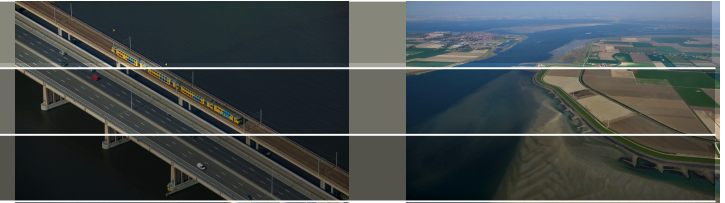
## EXPLANATION

-  5,000 Line of equal chloride concentration, in parts per million
  -  Bottom of fully cased well from which water-quality samples were collected
- Modified from Kohout (1964)

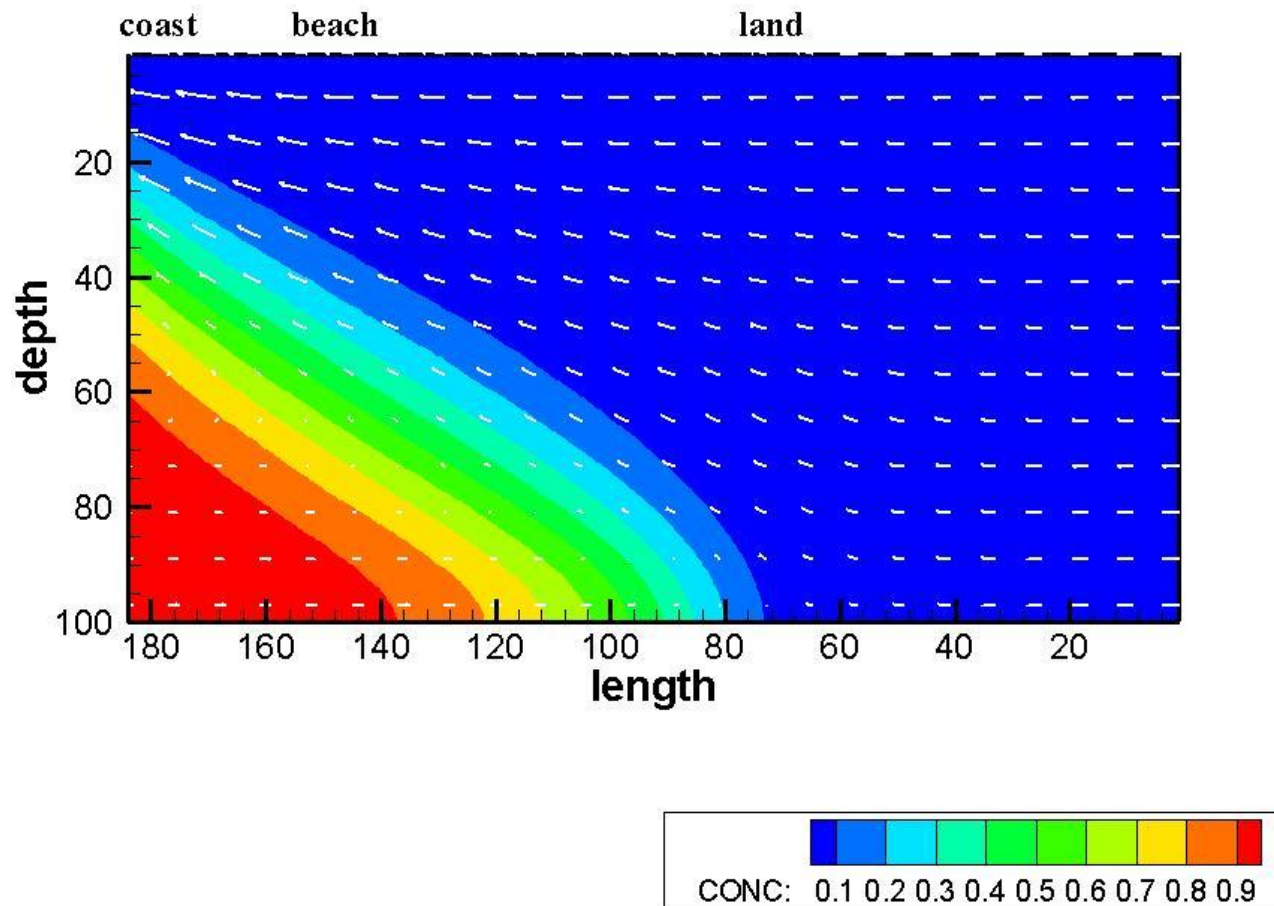
# Henry's problem (1964)



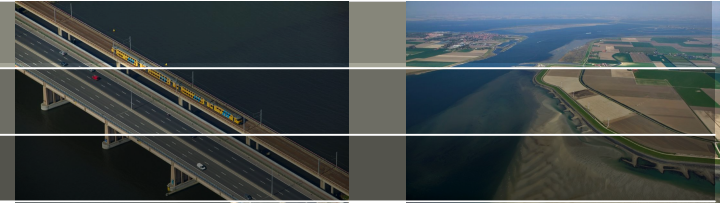
# Definitie zoutwaterintrusie



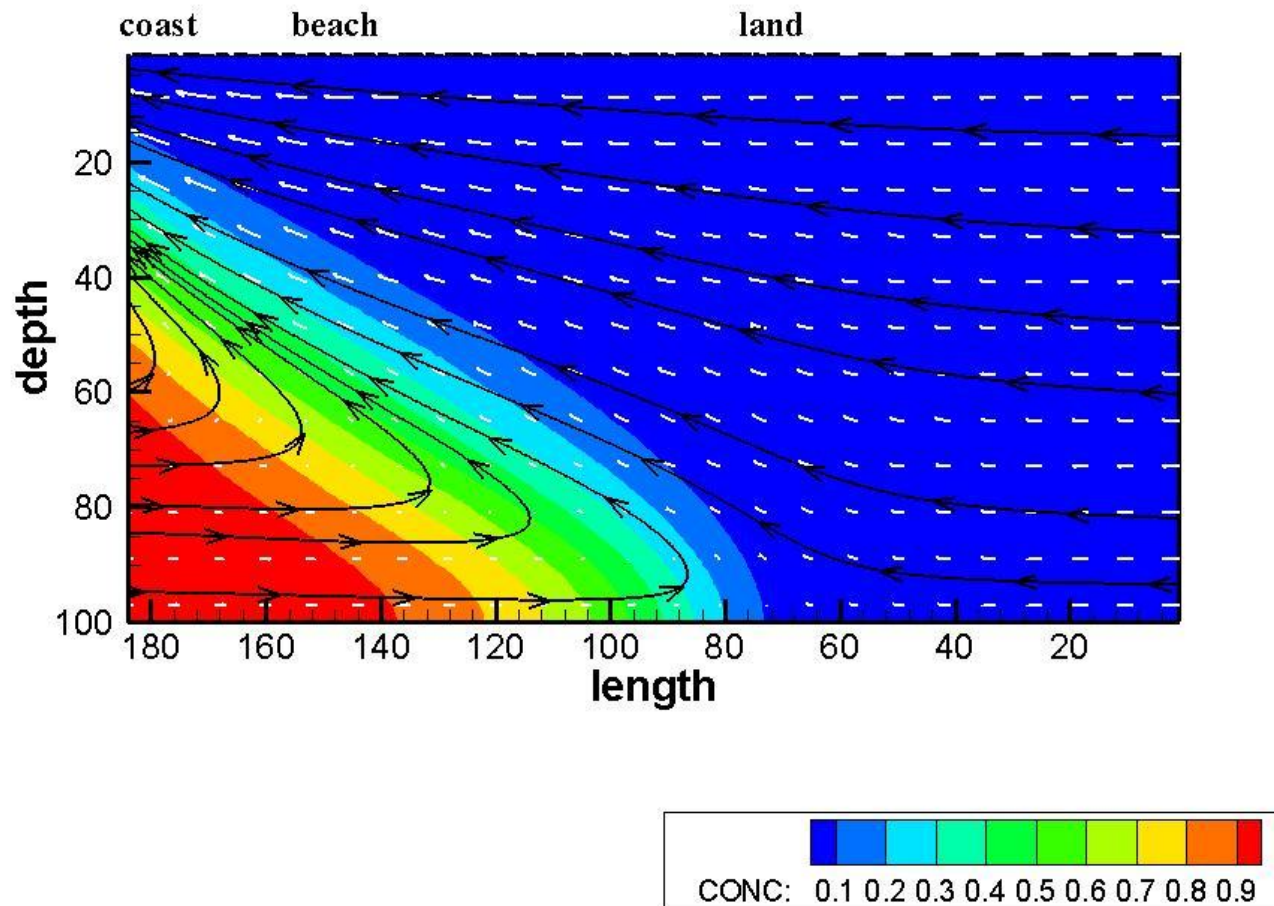
Numerical model: Henry's case



# Definitie zoutwaterintrusie



Numerical model: Henry's case

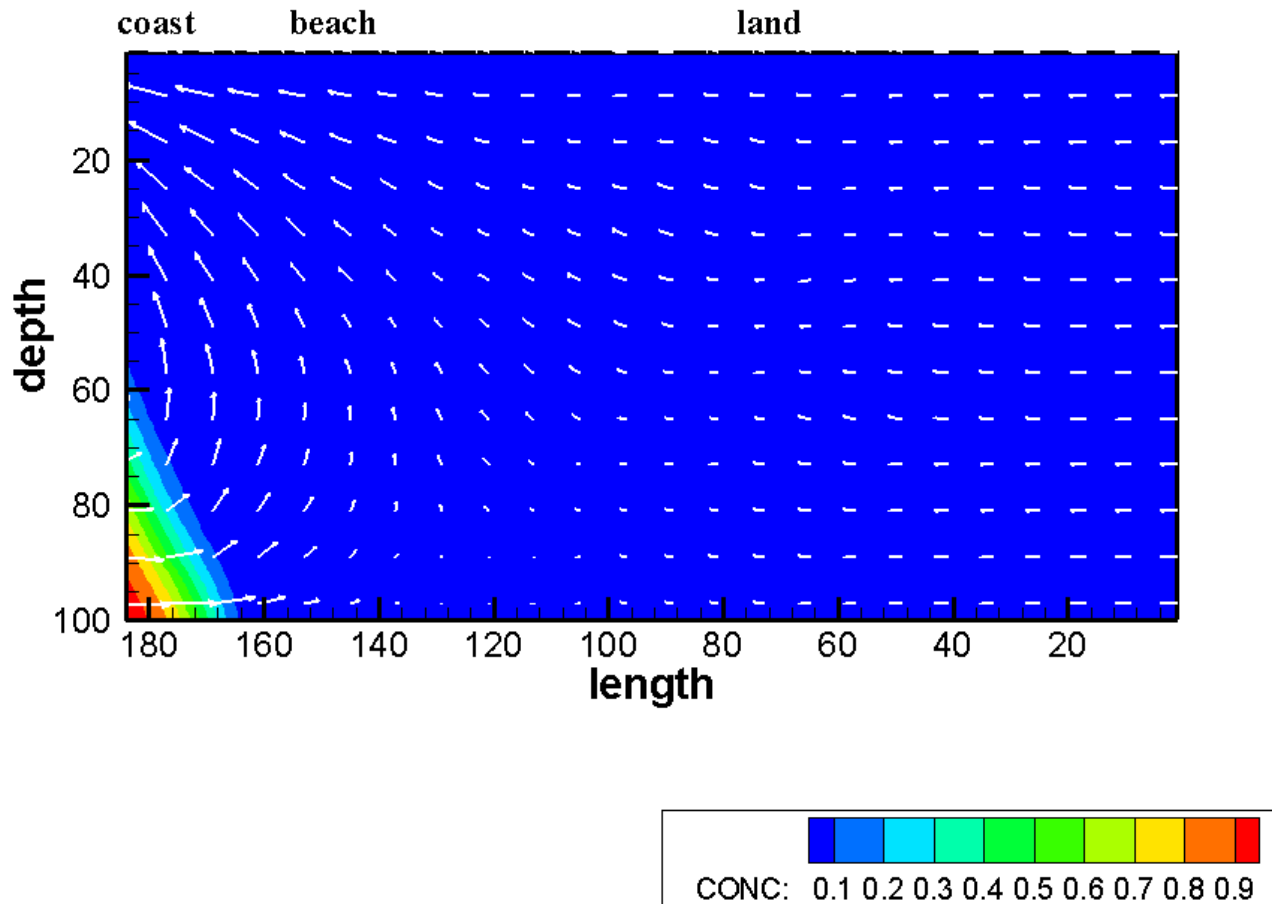




# Zeespiegelstijging en zoutwaterintrusie



## Effect sea level rise on groundwater system in coastal zone

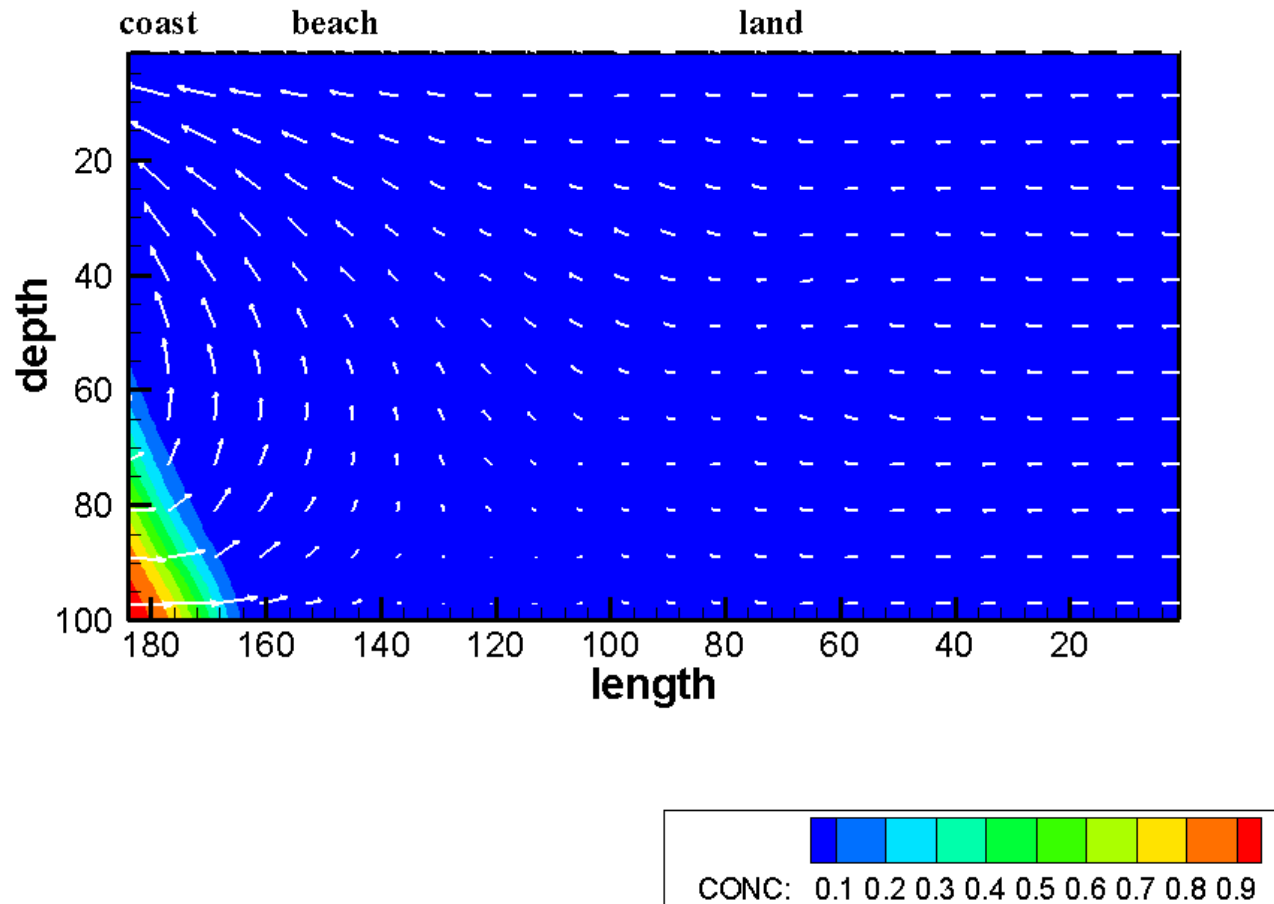




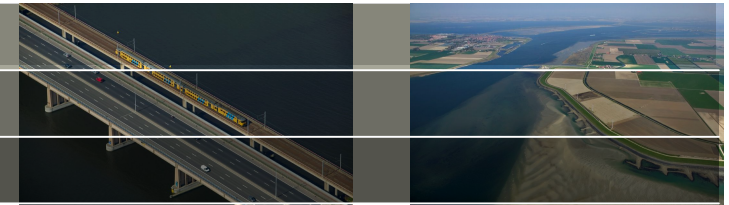
# Zeespiegelstijging en zoutwaterintrusie



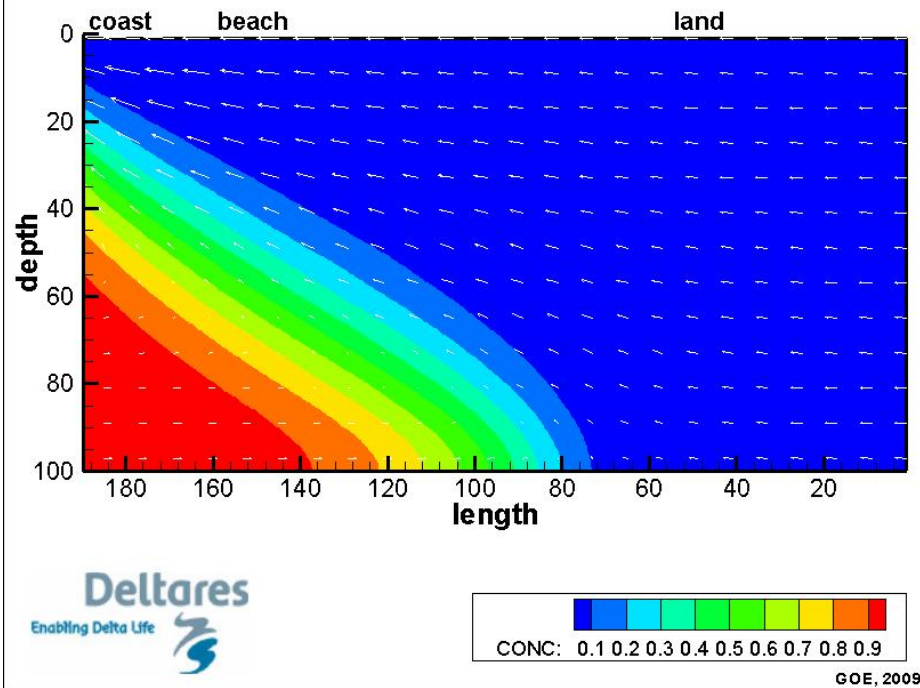
## Effect sea level rise on groundwater system in coastal zone



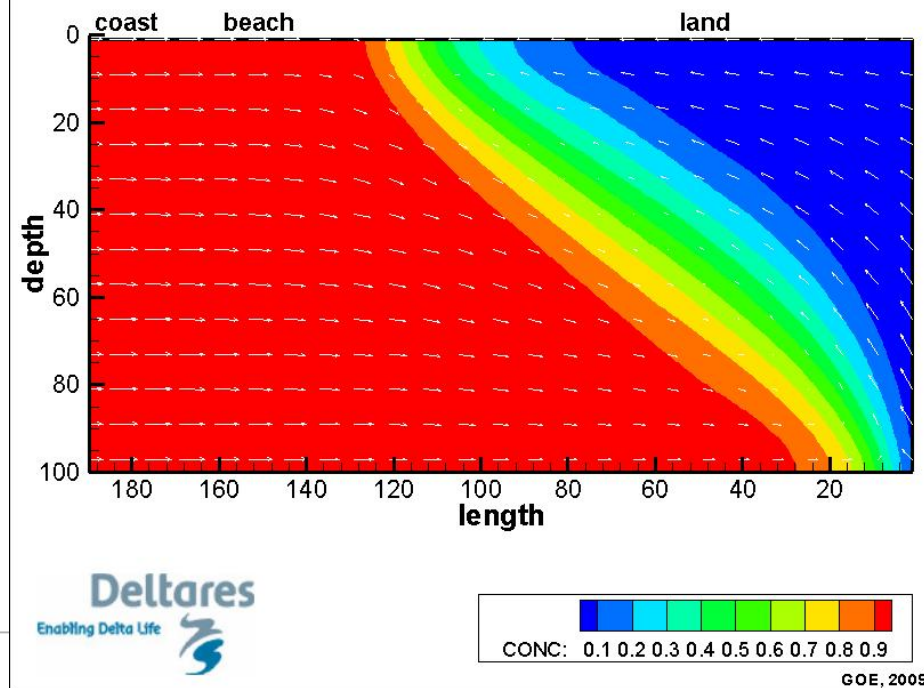
# Sea level rise and salt water intrusion



Impact of sea level rise on a coastal groundwater system:  
a conceptual model of saltwater intrusion



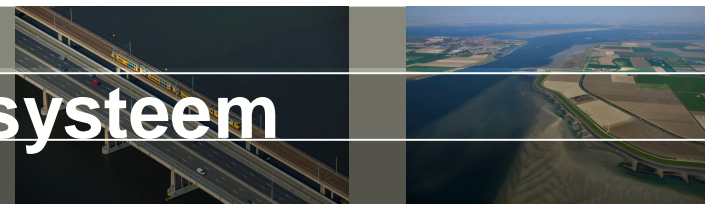
Impact of sea level rise on a coastal groundwater system:  
a conceptual model of saltwater intrusion





# Klimaatverandering en zeespiegelstijging

# Ons toekomstige grondwatersysteem

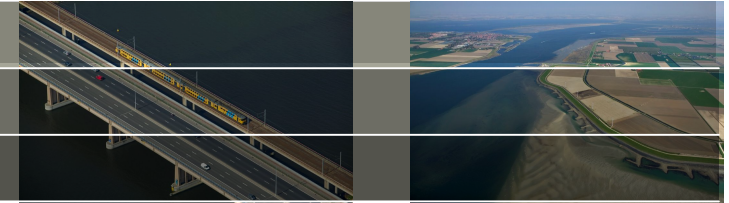


Wat komt er allemaal op ons af?

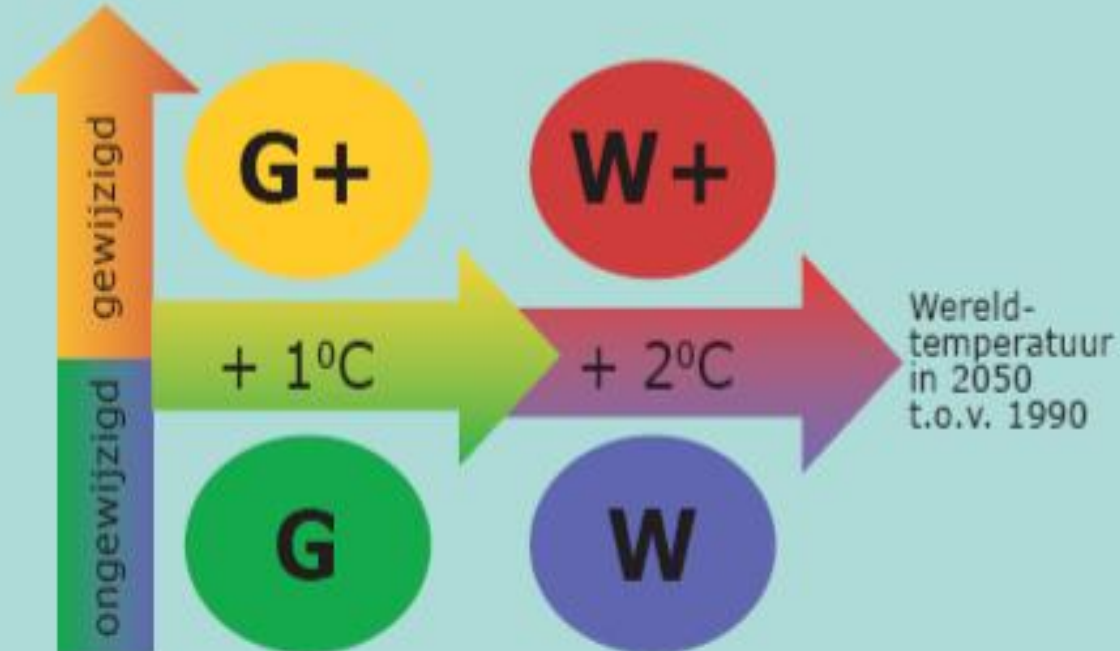
- Klimaatverandering
- Grondwaterwinning
- Ontwikkeling energiegebruik/productie
- Bodemdaling
- Ontwikkeling ruimtegebruik
- **Politiek, Beleid & Waterbeheer**

*Directe invloed door de mens op het grondwater is belangrijker dan klimaateffect*

# KNMI'06 scenario's



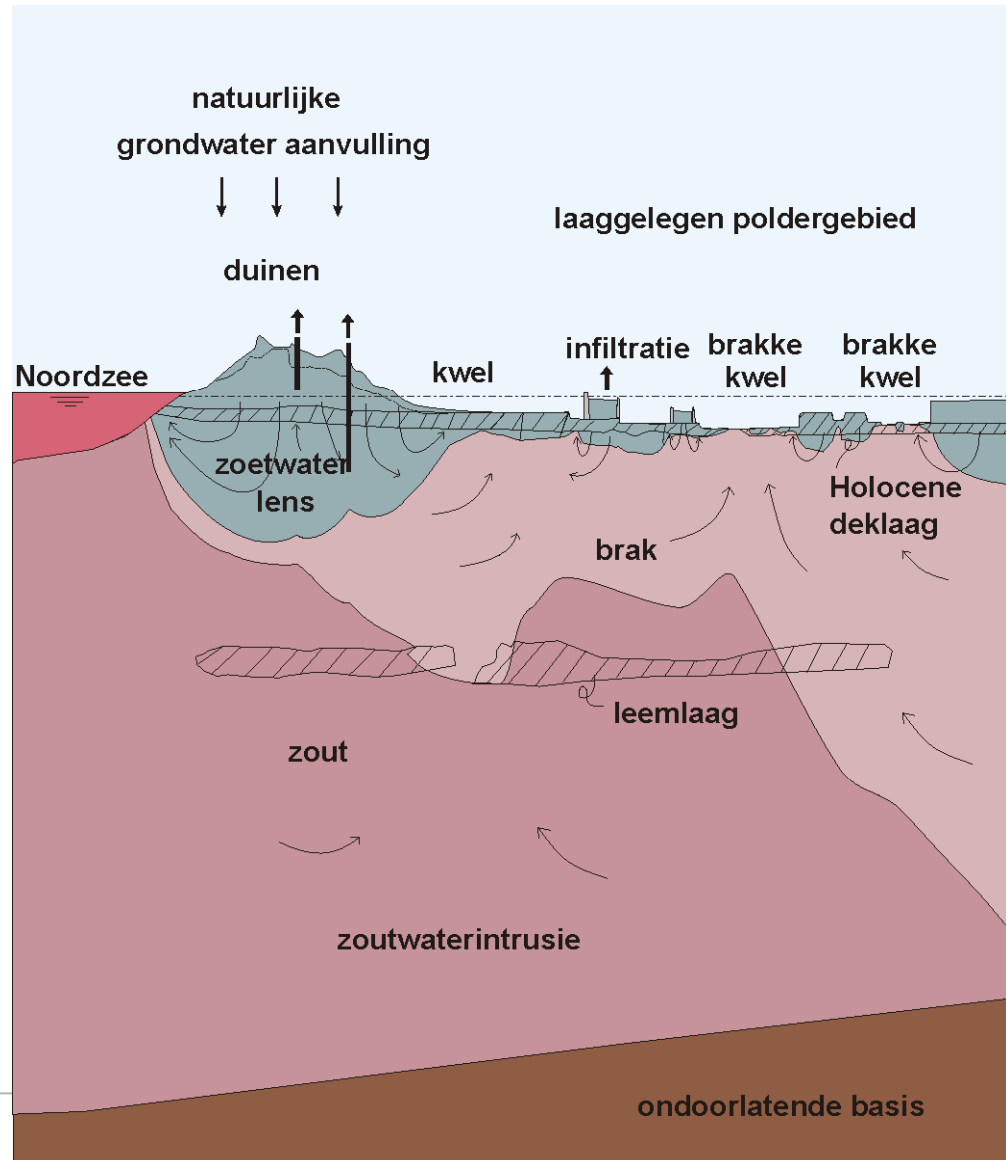
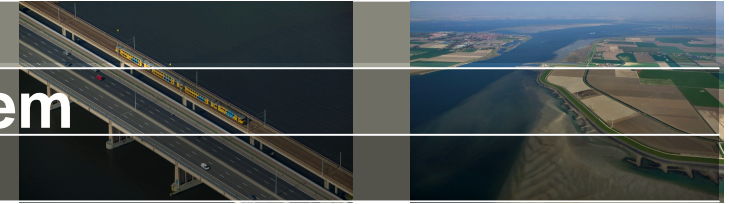
Luchtstromingspatronen



G	Gematigd	1°C temperatuurstijging op aarde in 2050 ten opzichte van 1990 geen verandering in luchtstromingspatronen in West Europa
G+	Gematigd +	1°C temperatuurstijging op aarde in 2050 ten opzichte van 1990 + winters zachter en natter door meer westenwind + zomers warmer en droger door meer oostenwind
W	Warm	2°C temperatuurstijging op aarde in 2050 ten opzichte van 1990 geen verandering in luchtstromingspatronen in West Europa
W+	Warm +	2°C temperatuurstijging op aarde in 2050 ten opzichte van 1990 + winters zachter en natter door meer westenwind + zomers warmer en droger door meer oostenwind

Bron: Van der Hurk **Deltares**  
et al. (2006)

# Het Nederlandse grondwatersysteem

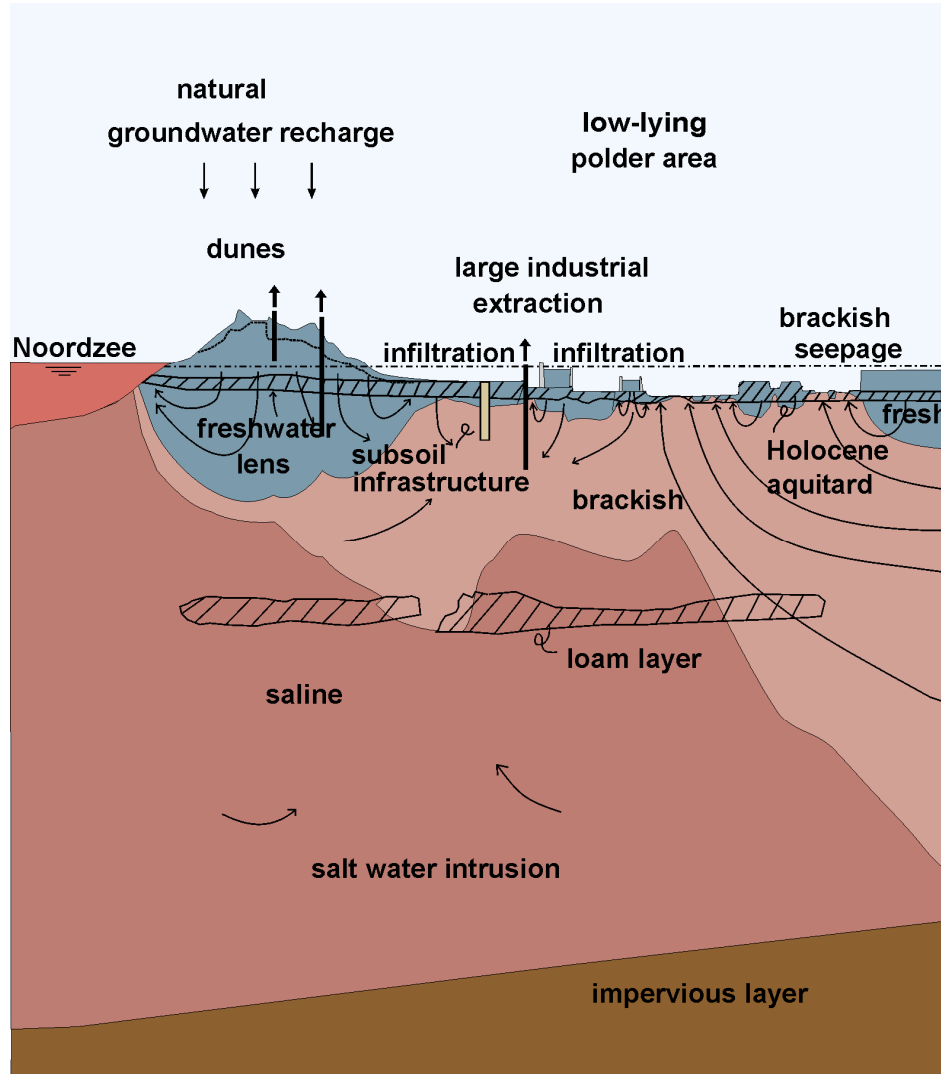


Peilverlagingen in het verleden hebben grondwaterstromingen naar het binnenland veroorzaakt

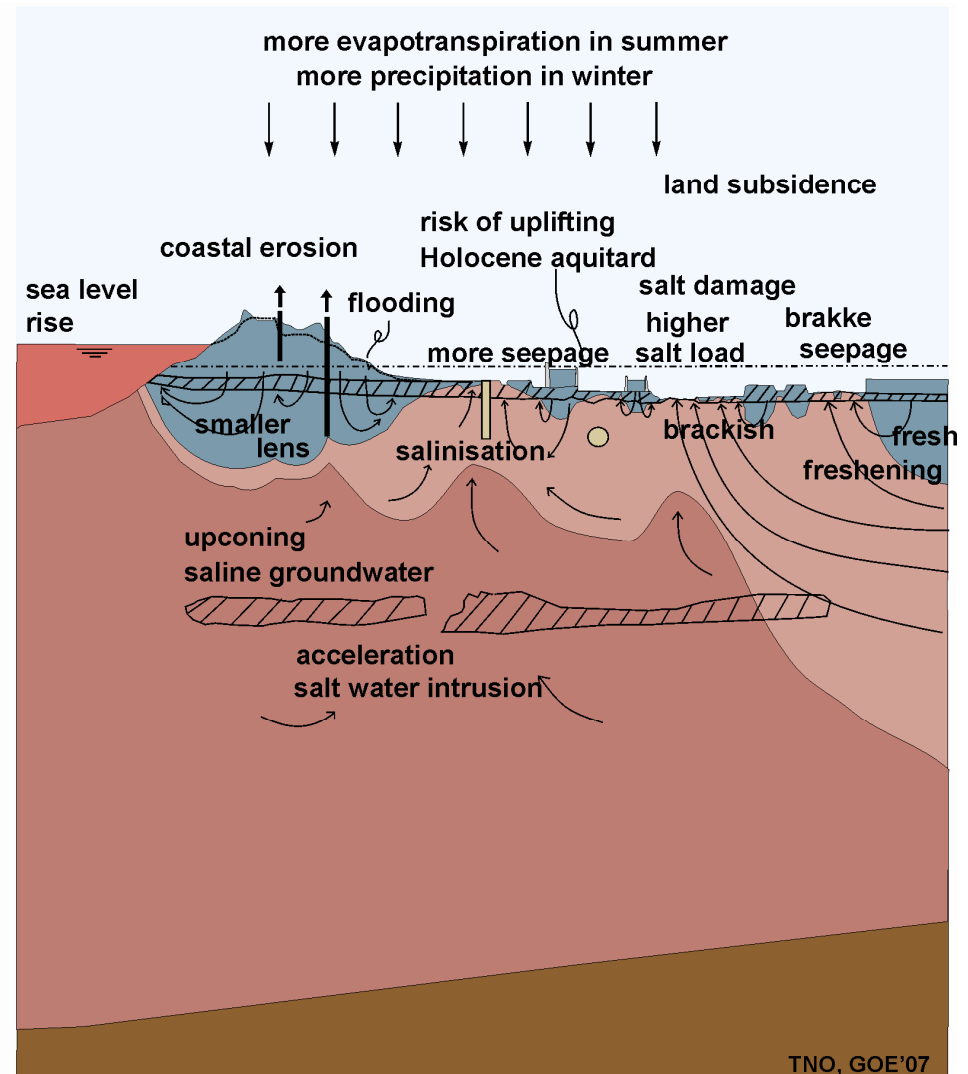


# Het Nederlandse grondwatersysteem staat onder druk

## Present processes

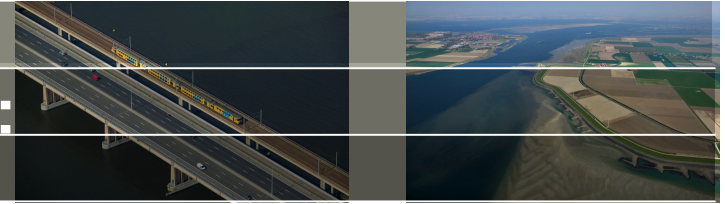


## Future changes

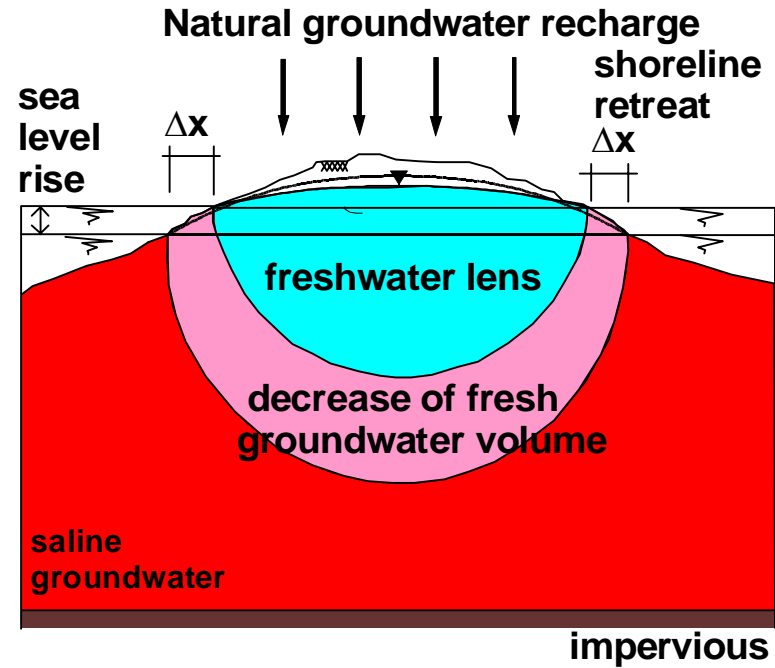
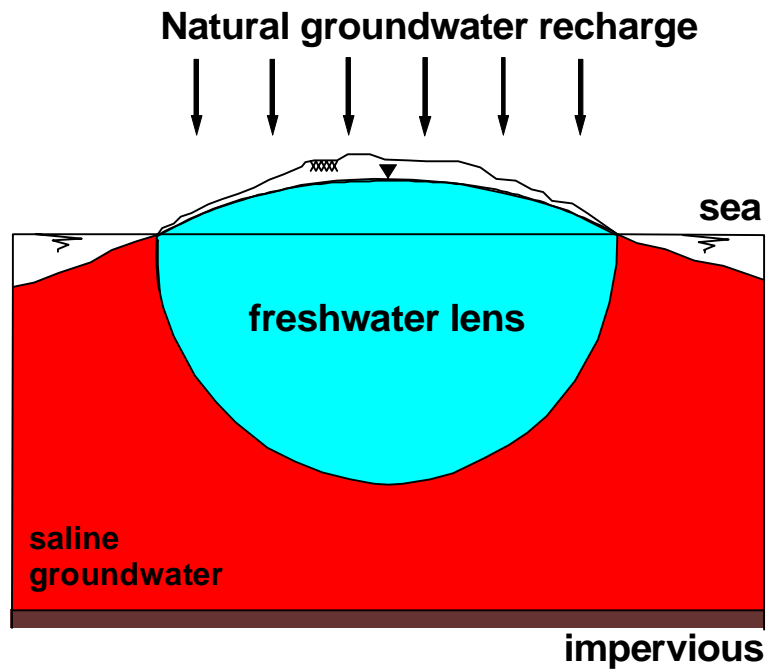




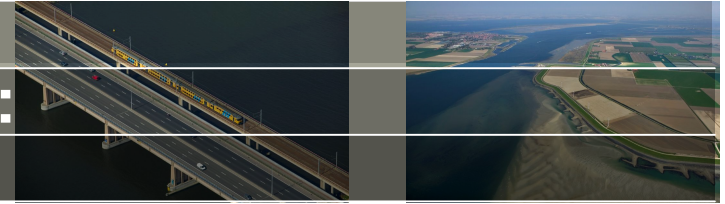
# Effect of a relative sea level rise (1):



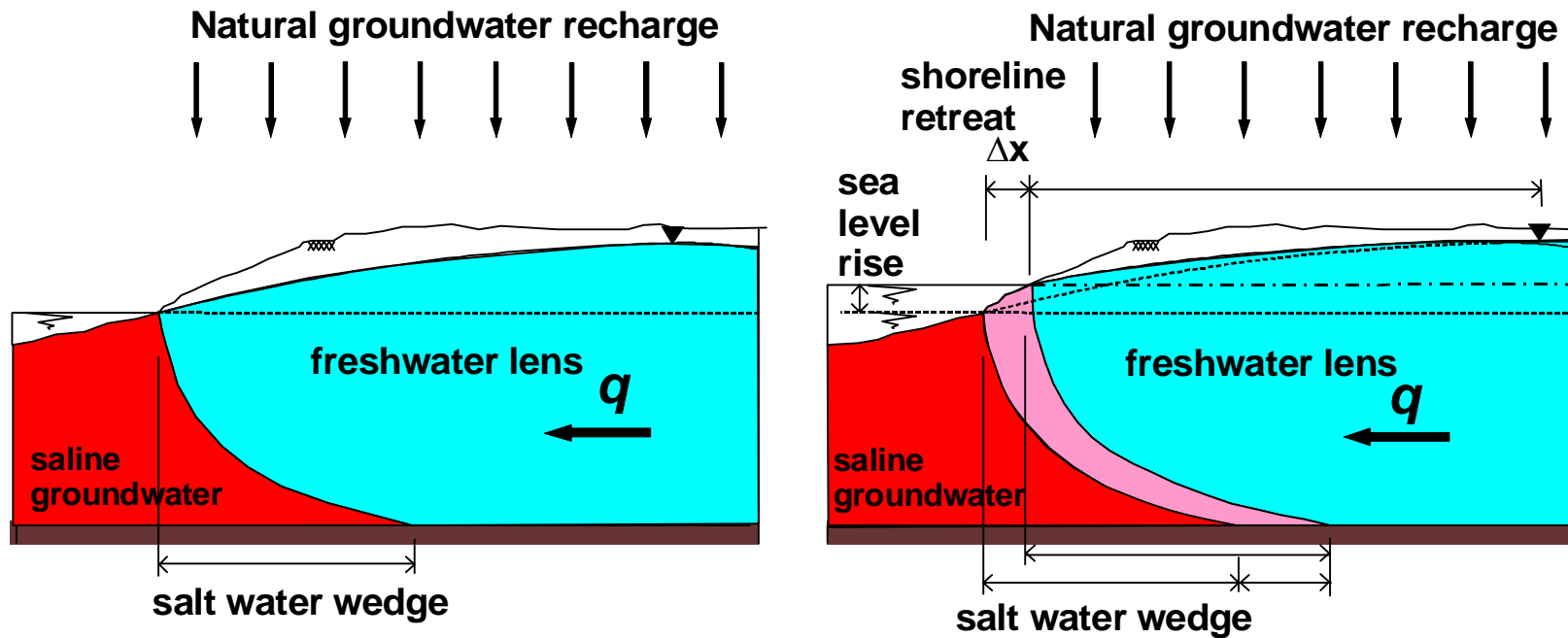
## Deep aquifer



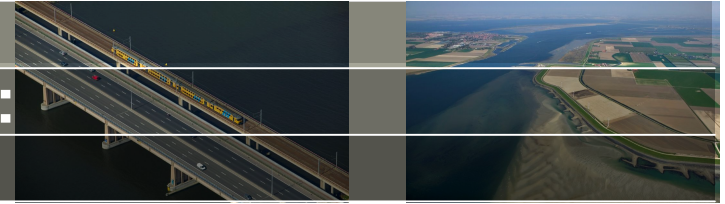
# Effect of a relative sea level rise (2):



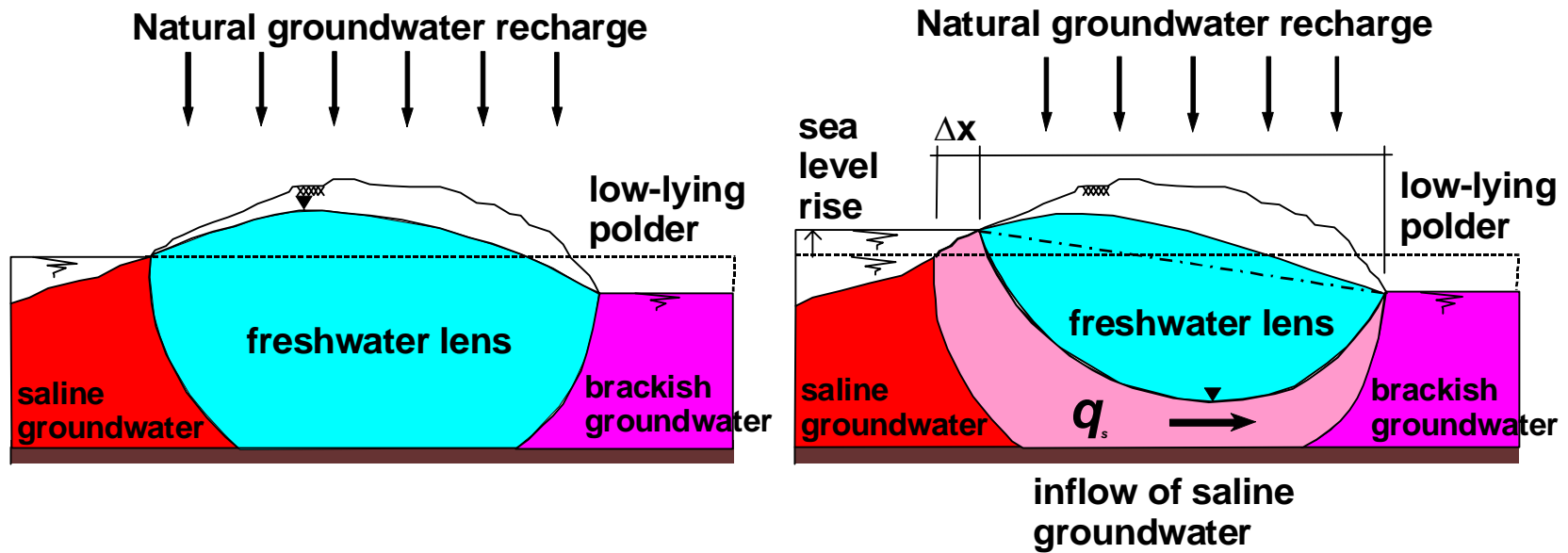
## Shallow aquifer



# Effect of a relative sea level rise (3):



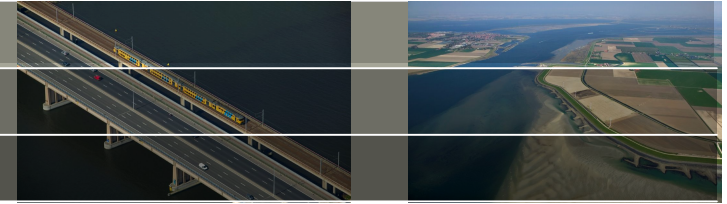
## Shallow aquifer





# Effecten voor het oppervlaktewater systeem en de landbouw

# Gevolgen verzilting



## Gevolgen verzilting grondwater:

- Zoutschade gewassen
- Zoutbelasting vanuit het grondwater naar het oppervlaktewater
- [Verzilting onttrekkingputten]



*Bron: Proefstation voor de Akkerbouw en Groenteteelt, Lelystad*

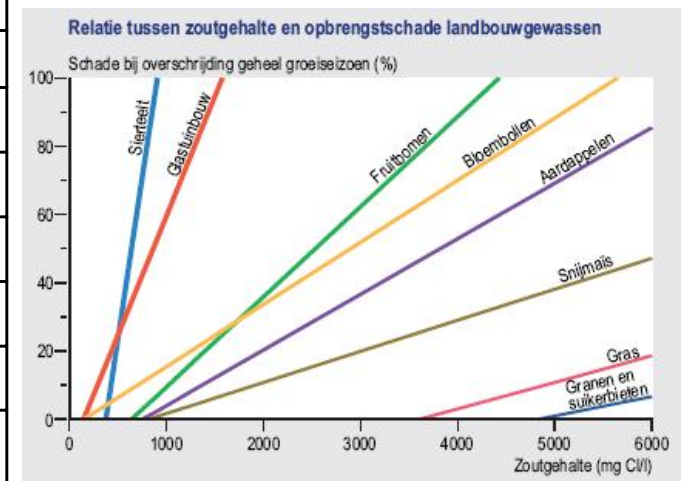
# Gevolgen verzilting: zoutschade aan gewassen bij wortelzone

Belangrijke parameters:

- Chloride concentratie bij wortelzone
- Landgebruik
- Gevoeligheid gewassen

Landgebruik	Drempelwaarde wortelzone (mg Cl-/l)	Helling wortelzone (-)	Gewasopbrengst prijspeil 2050 (Euro/ha)
gras	3606	0.0078	1080
aardappelen	756	0.0163	4240
<b>bieten</b>	<b>4831</b>	0.0057	2640
granen	4831	0.0058	880
glastuinbouw	1337	0.0141	17850
boomgaard	642	0.0264	11900
<b>bollen</b>	<b>153</b>	0.0182	26000

Bron: Roest et al., 2003 en Haskoning



Bron: MNP, 2005

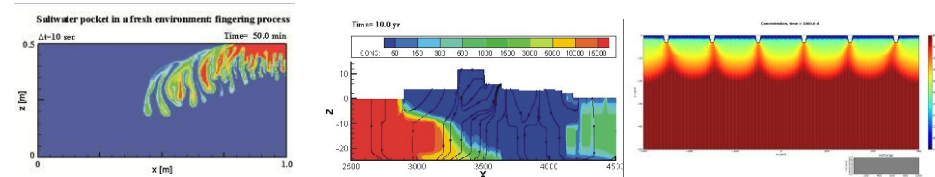


# Modelleren zoet-zout grondwater

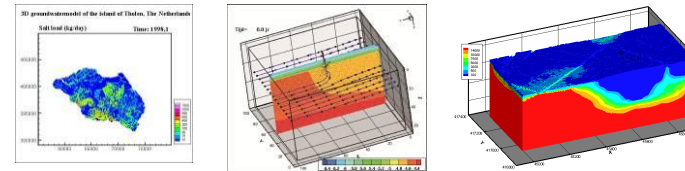


# Verschillende modelschalen om verschijnselen te analyseren

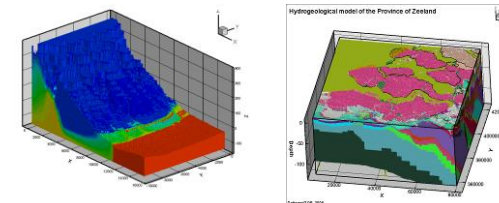
**Sub-local:** fingering, zoute wellen  
 Sri Lanka (Tsunami 2004), De Zandmotor  
**modelschaal=1cm-1m**



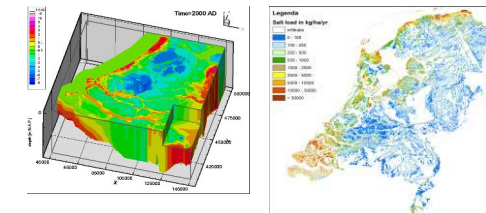
**Local:** regenwaterlenzen, WKO,  
 Tholen, Schouwen-Duiveland  
**modelschaal=5-25m**



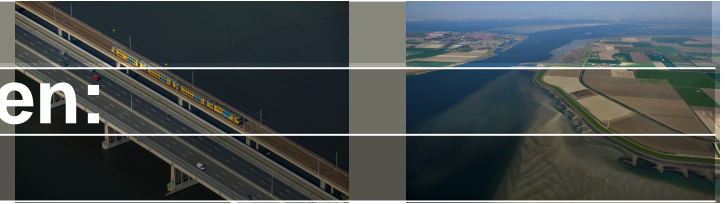
**Regionaal:**  
 Zeeland, Gujarat/India, Philippines  
**modelschaal=100m**



**Nationaal:** salt load  
 Zuid-Holland, NHI  
**modelschaal=250m-1 km**



# Numerieke computer modellen:



## ROTZOOI IN = ROTZOOI UIT

echter...

- steeds betere 3D geologische opbouw (REGIS II, Geotop)
- steeds betere 3D initiële dichtheidsverdeling
- steeds betere concepten
- steeds betere ijkings technieken
- steeds nauwkeuriger (snellere computers, meer geheugen)

# Modelling effect climate change on fresh-salt groundwater

## Modelling:

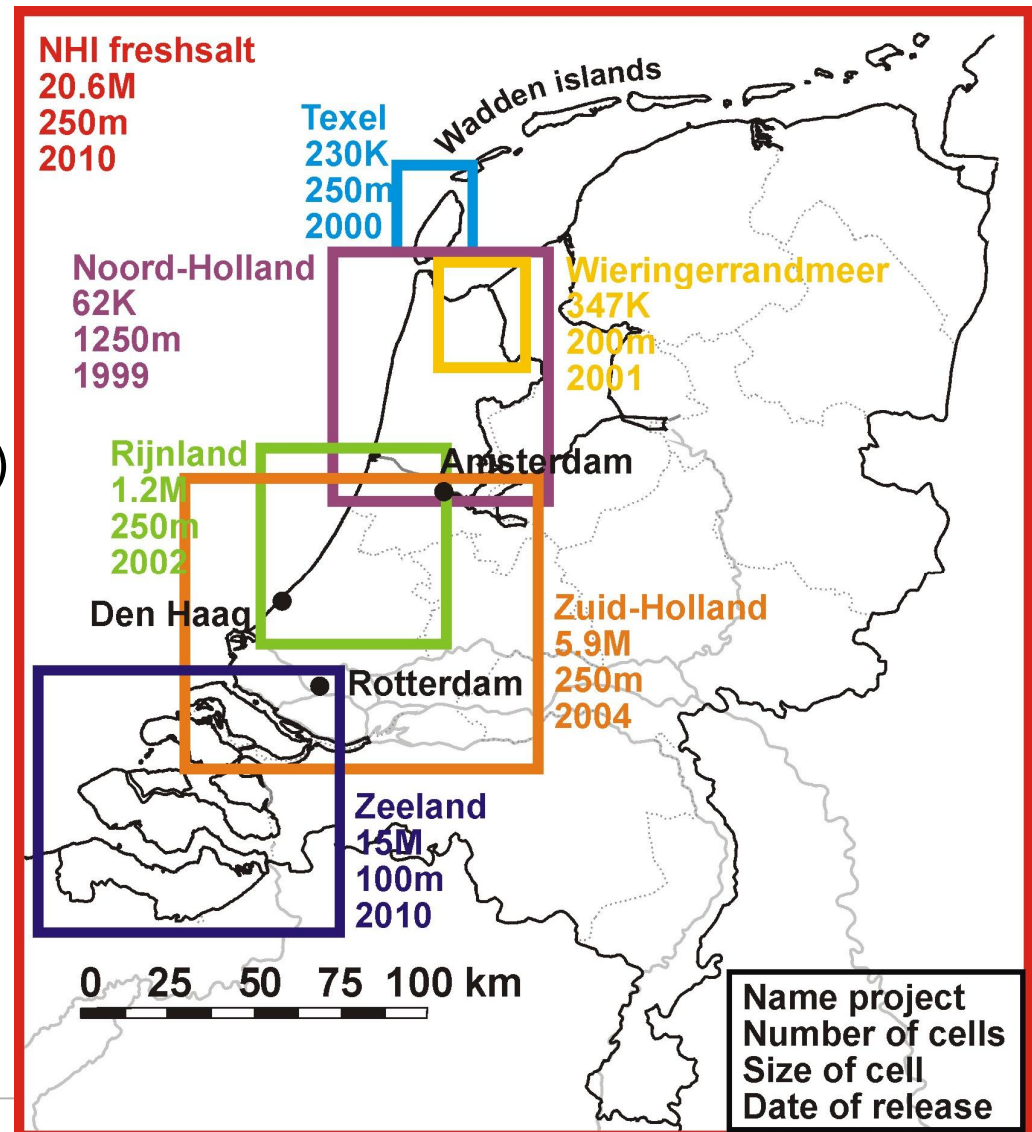
- variable-density
- 3D, non-steady
- groundwater flow
- coupled solute transport

## Code:

MOCDENS3D (MODFLOW family)  
similar to SEAWAT

## Assessing effects:

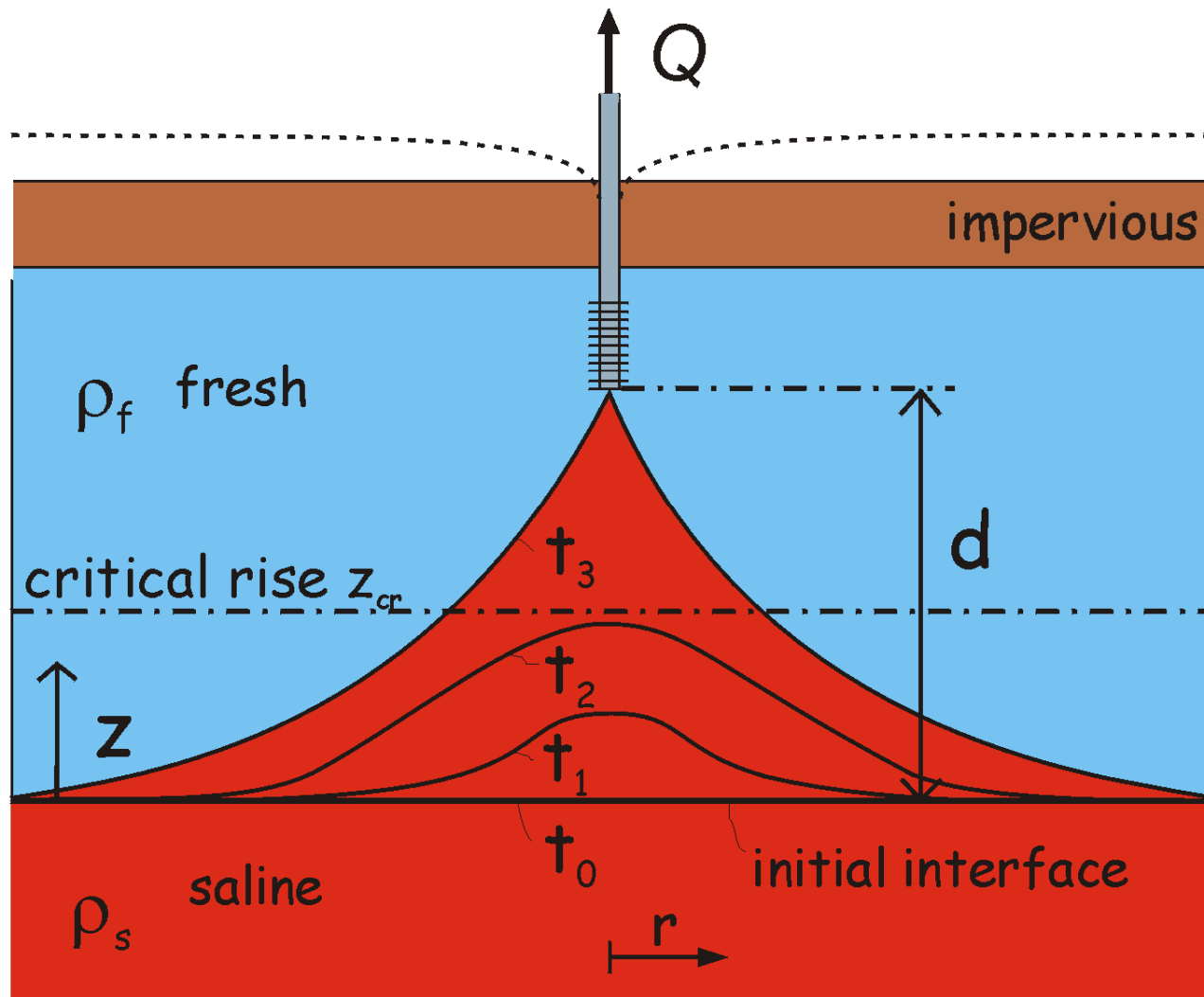
- autonomous salinisation
- sea level rise
- changing recharge pattern
- land subsidence
- changing extraction rates
- adaption measures



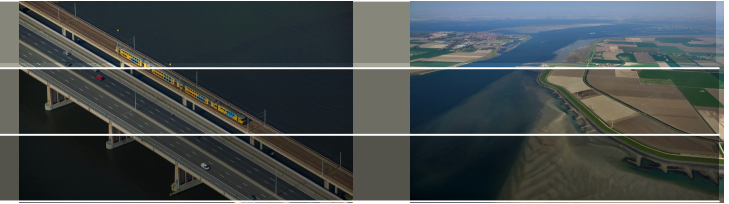


# Opkegeling van zout grondwater onder een onttrekkingsput

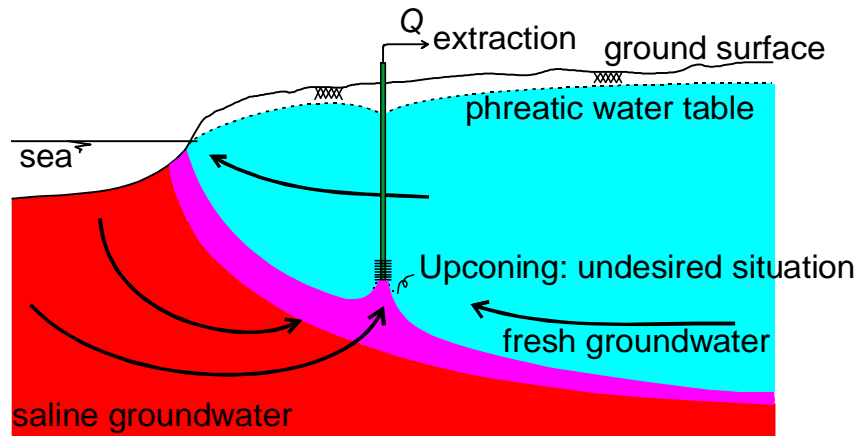
# Opkegeling van zout grondwater onder een onttrekkingsput



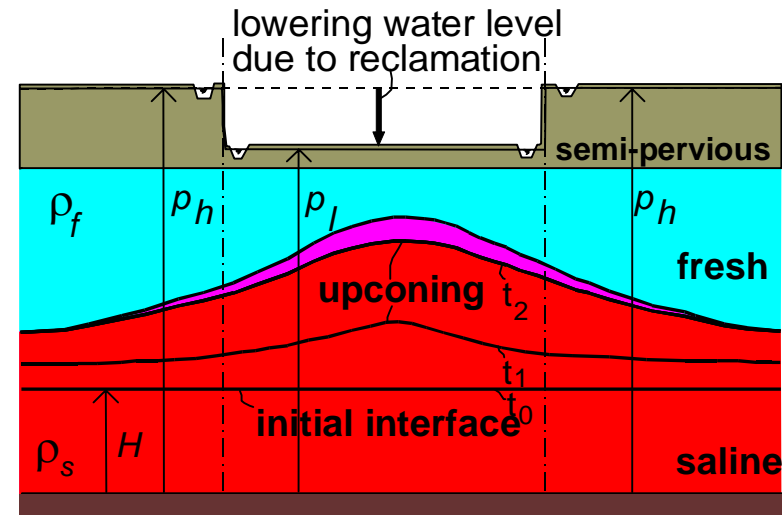
# Upconing of saline groundwater



Under an extraction well



Under a low-lying area



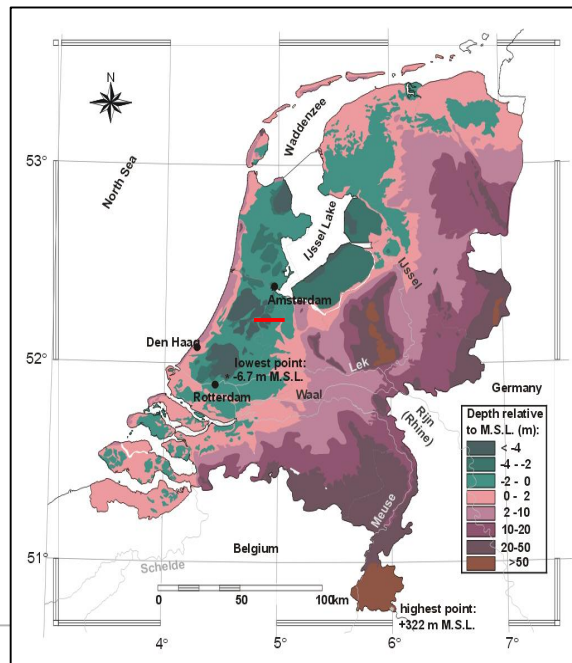
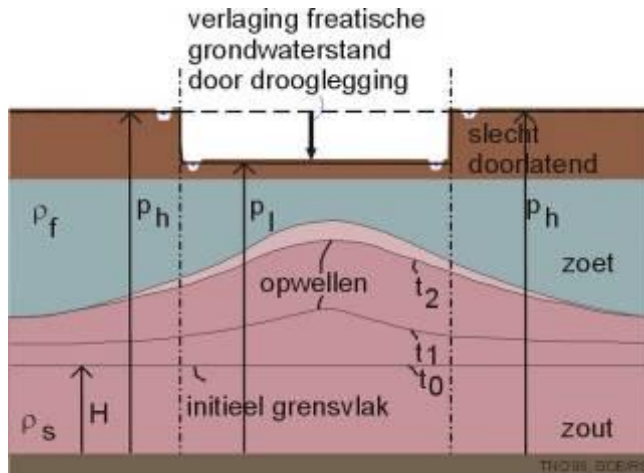
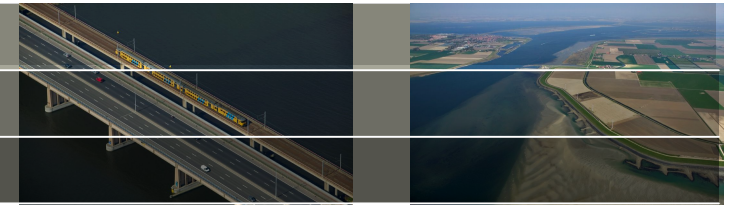
- movement of saline groundwater to extraction wells
- increase in salinity (>150-200 mg Cl-/l)
- lowering of the piezometric head (leads to land subsidence:  
e.g. Los Angeles: 9 m in the 1930's)

'Solutions': reduce extraction rate, abandon well, inundate polder

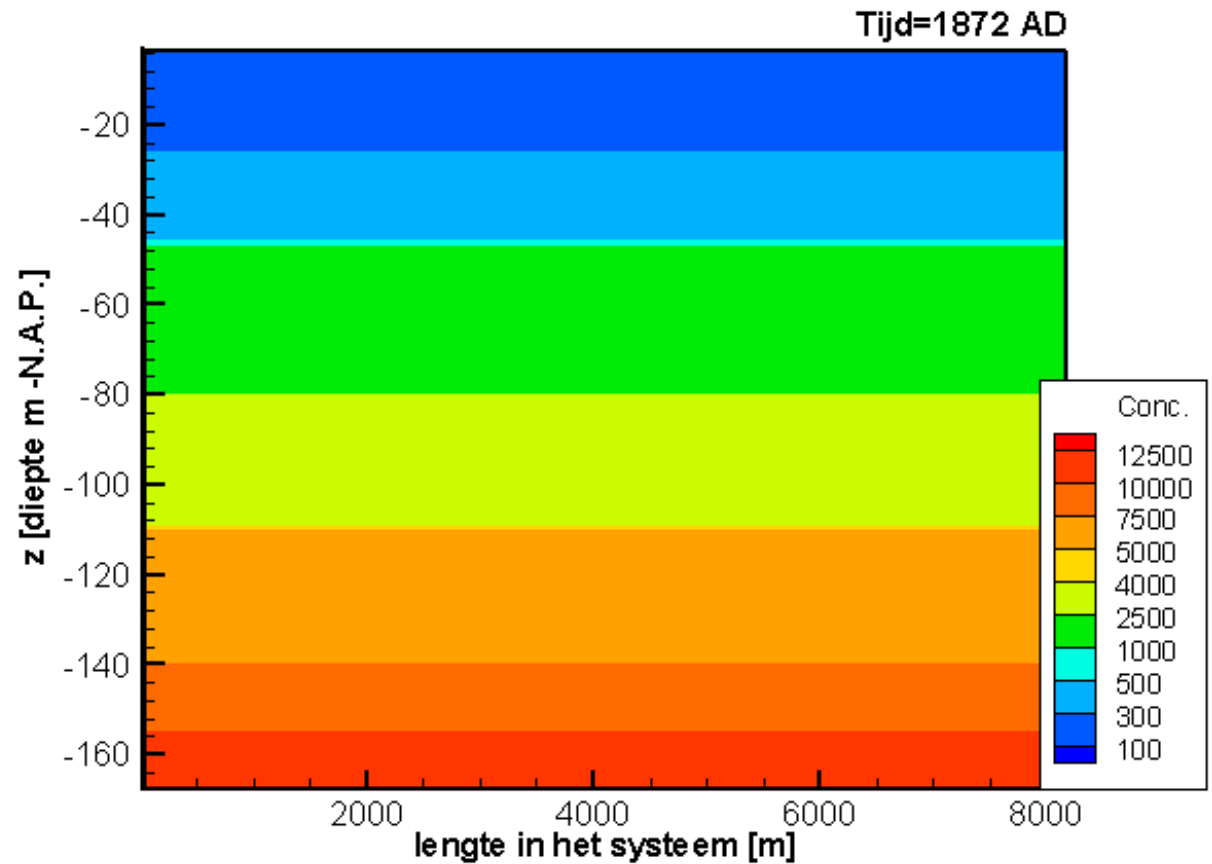
Deltares



# Upconing under a low-lying polder



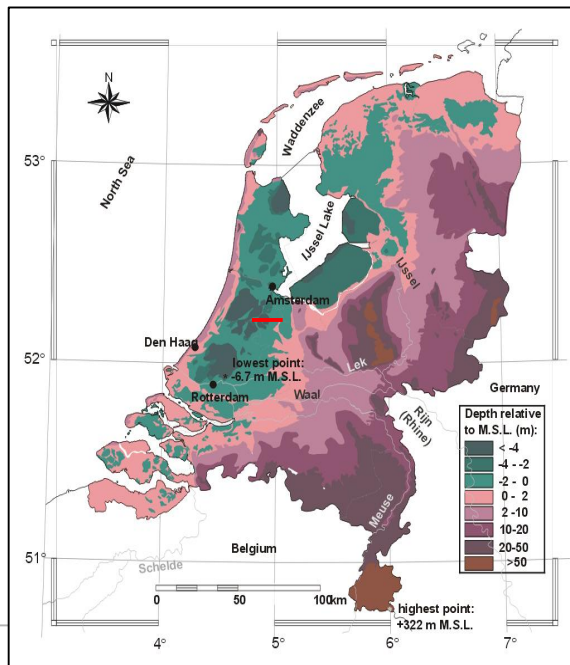
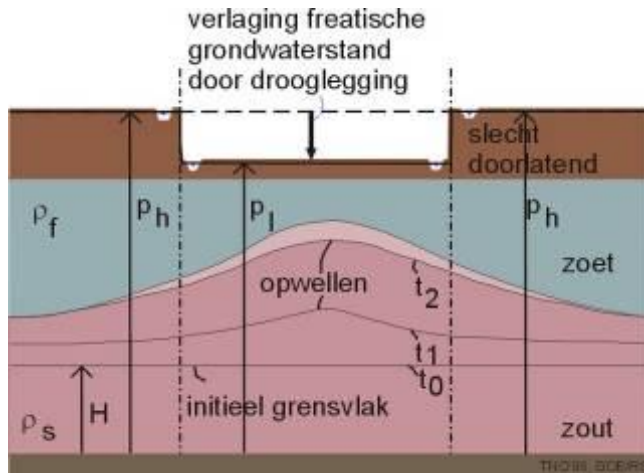
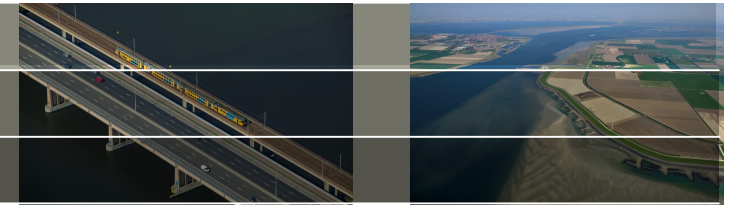
Verzilting onder de polder Groot-Mijndrecht: opwellen zout grondwater



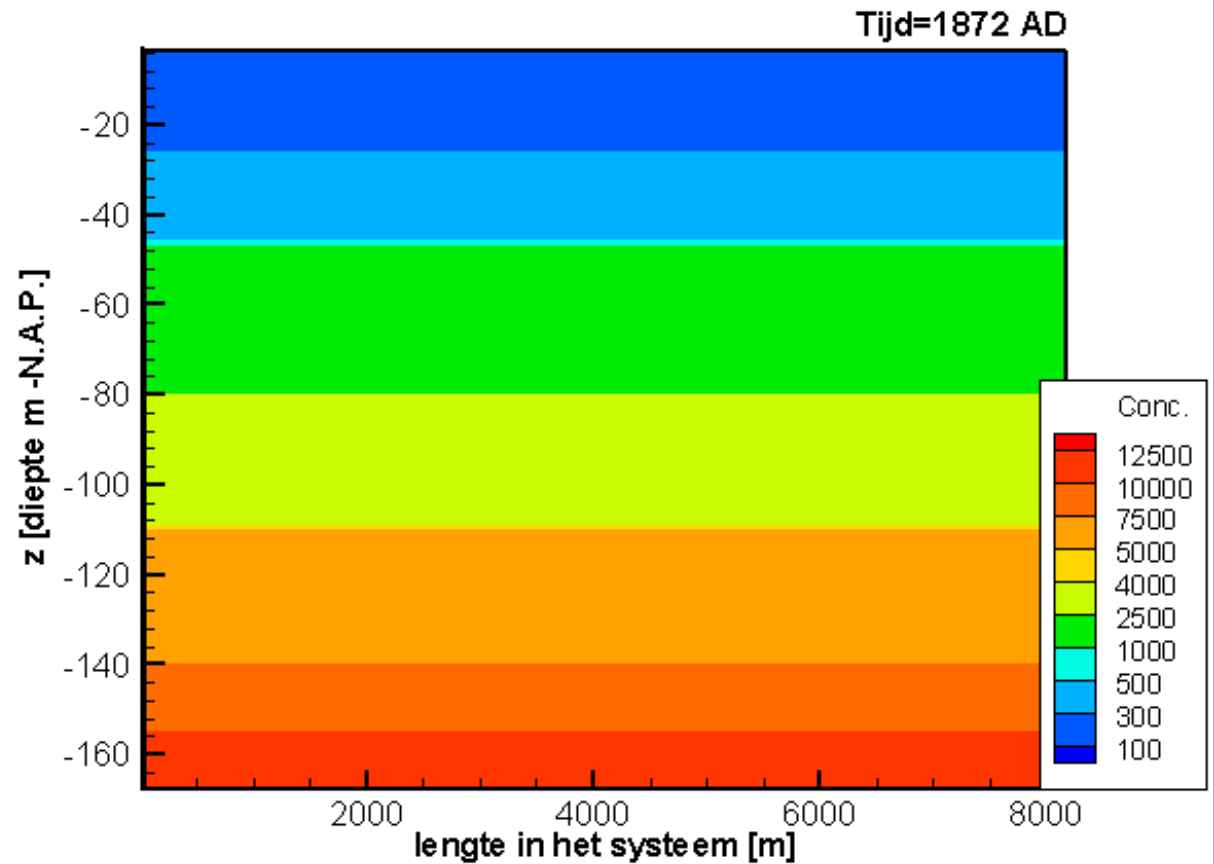
TNO-BO, GOE, '06



# Upconing under a low-lying polder

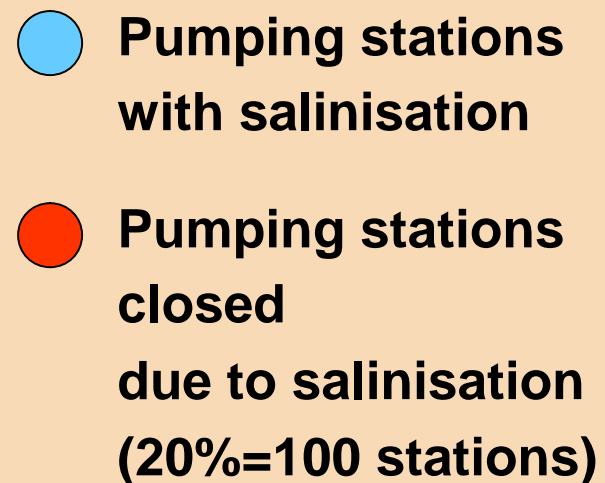
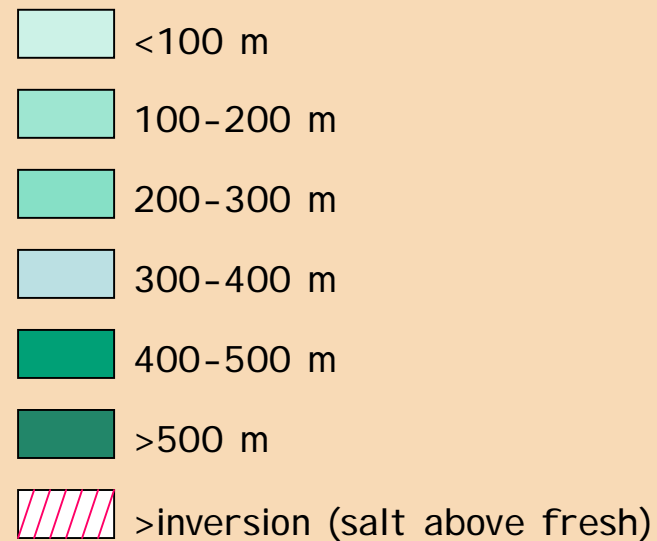
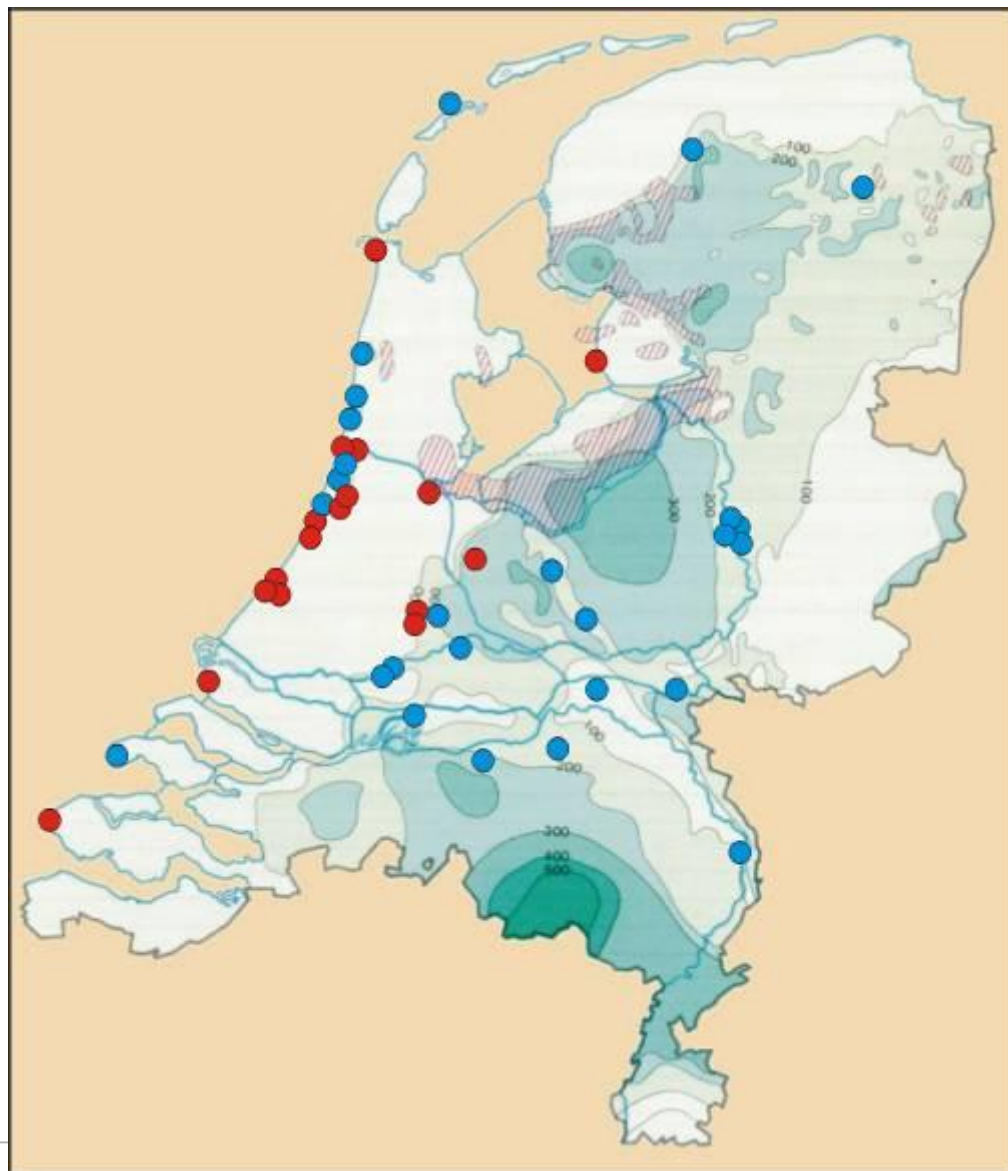


Verzilting onder de polder Groot-Mijndrecht: opwellen zout grondwater

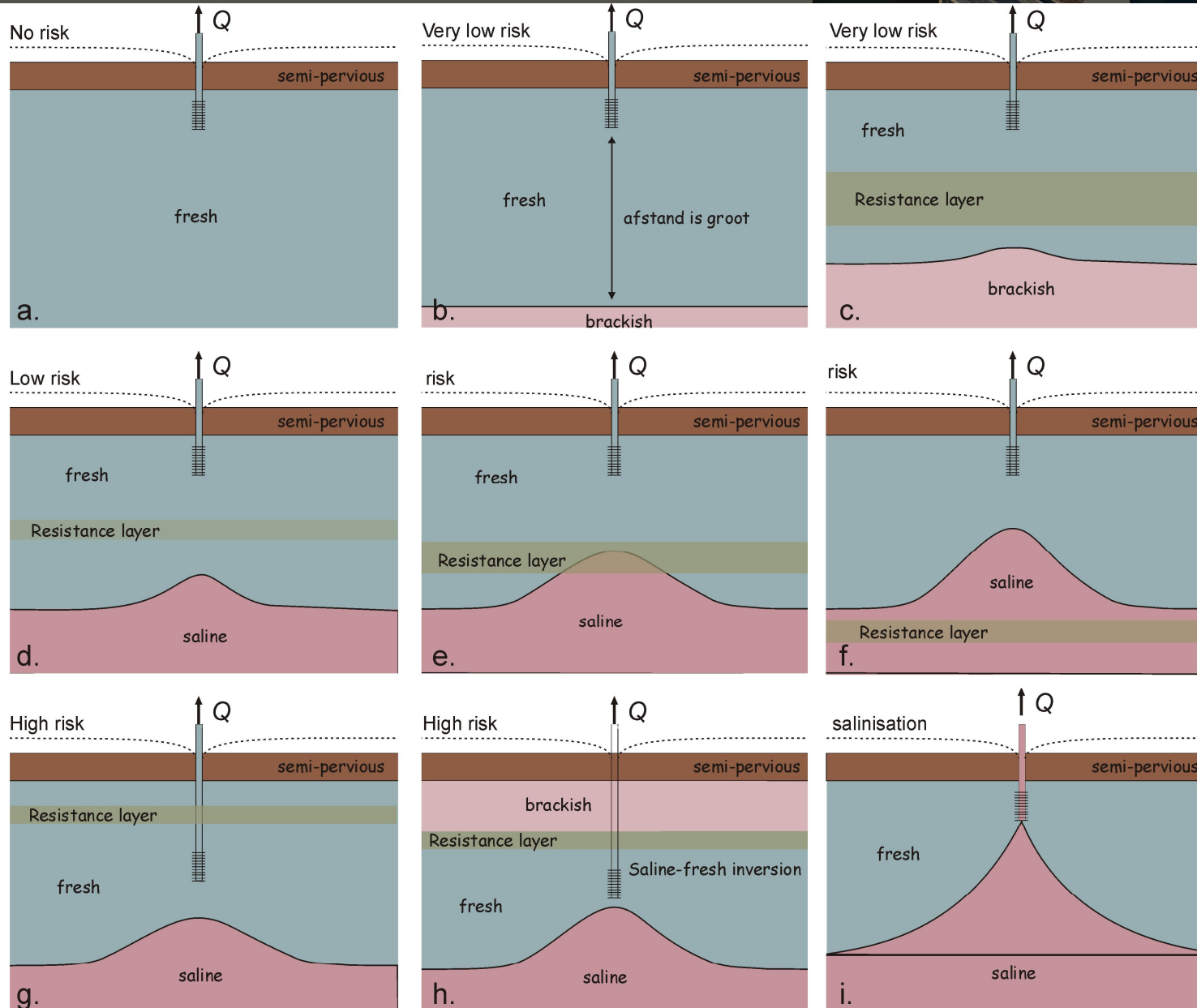


TNO-BO, GOE, '06

# Availability of fresh groundwater for drinking water



# Different risks of upconing saline groundwater



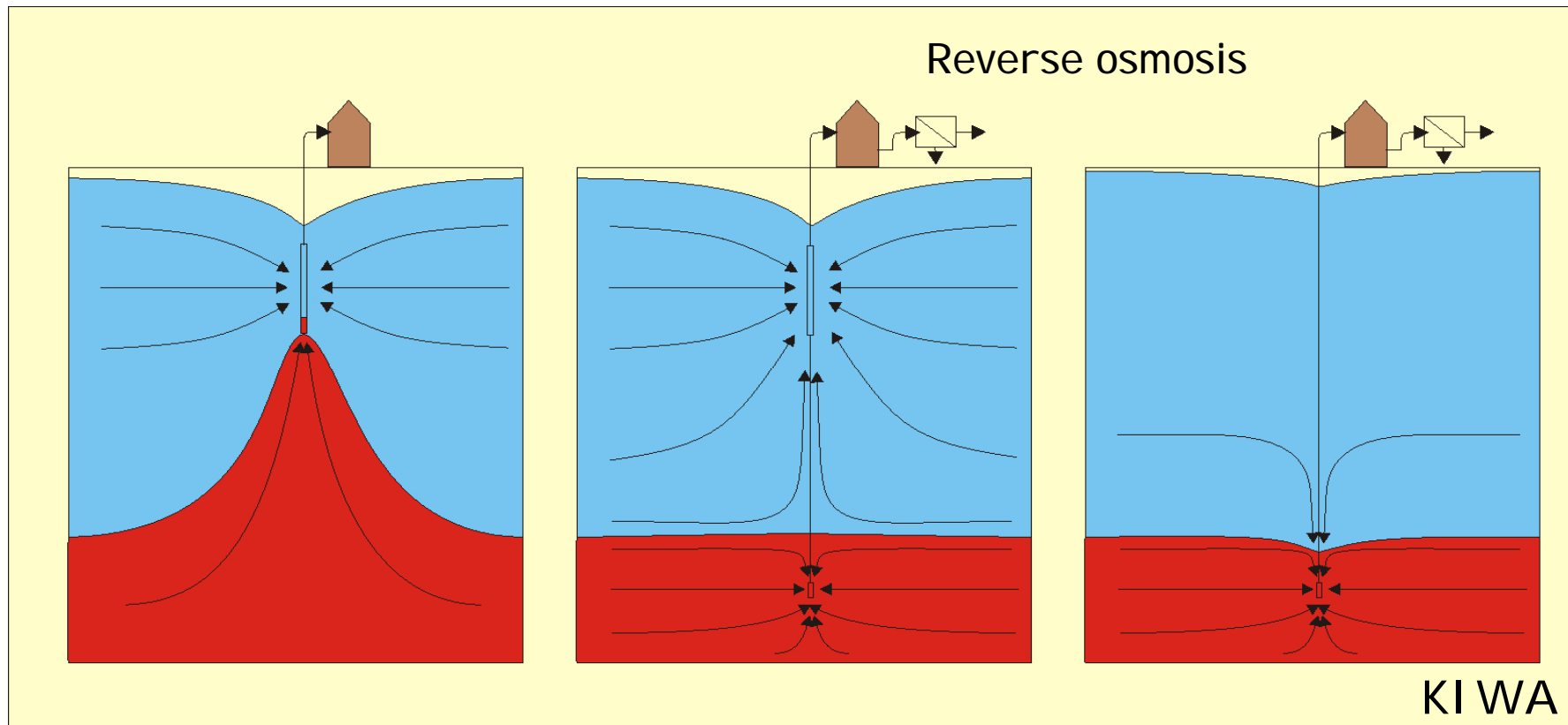
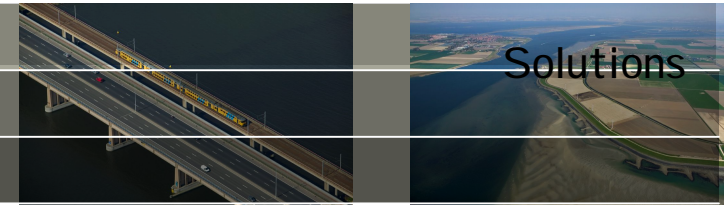


# Technische maatregelen

# Measures to compensate salt water intrusion

- 'The Fresh Holder'
- Extraction of saline/brackish groundwater
- Infiltration of fresh surface water
- Modifying pumping rates
- Land reclamation in front of the coast
- Creating physical barriers (crystallisation or biosealing)

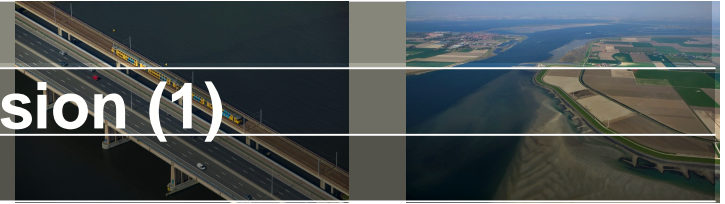
# Solution: The Fresh Holder



Upconing can be prevented by the extraction of brackish groundwater

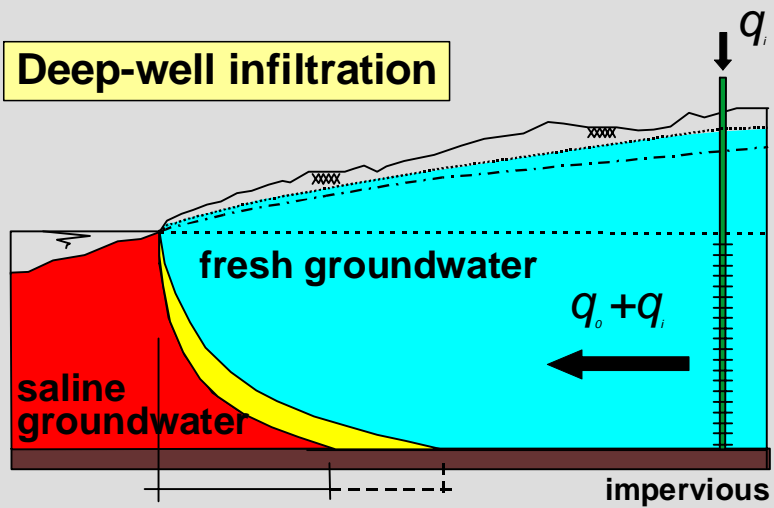
This brackish groundwater can be transformed to water of agricultural water quality by using the membrane filtration technique

# Countermeasures of salt water intrusion (1)

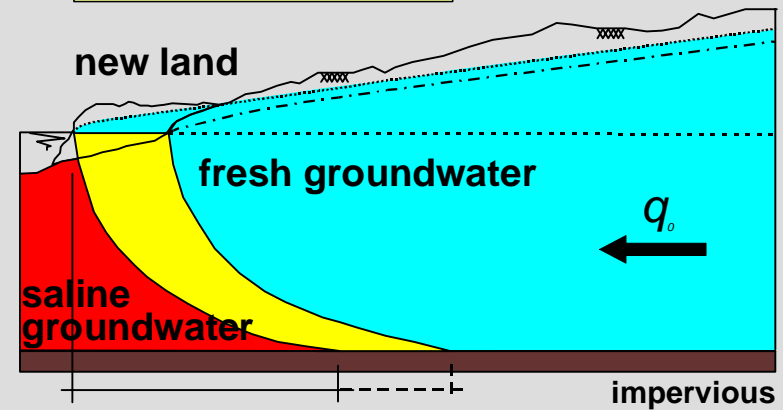


 increase of fresh groundwater volume due to countermeasure

Deep-well infiltration

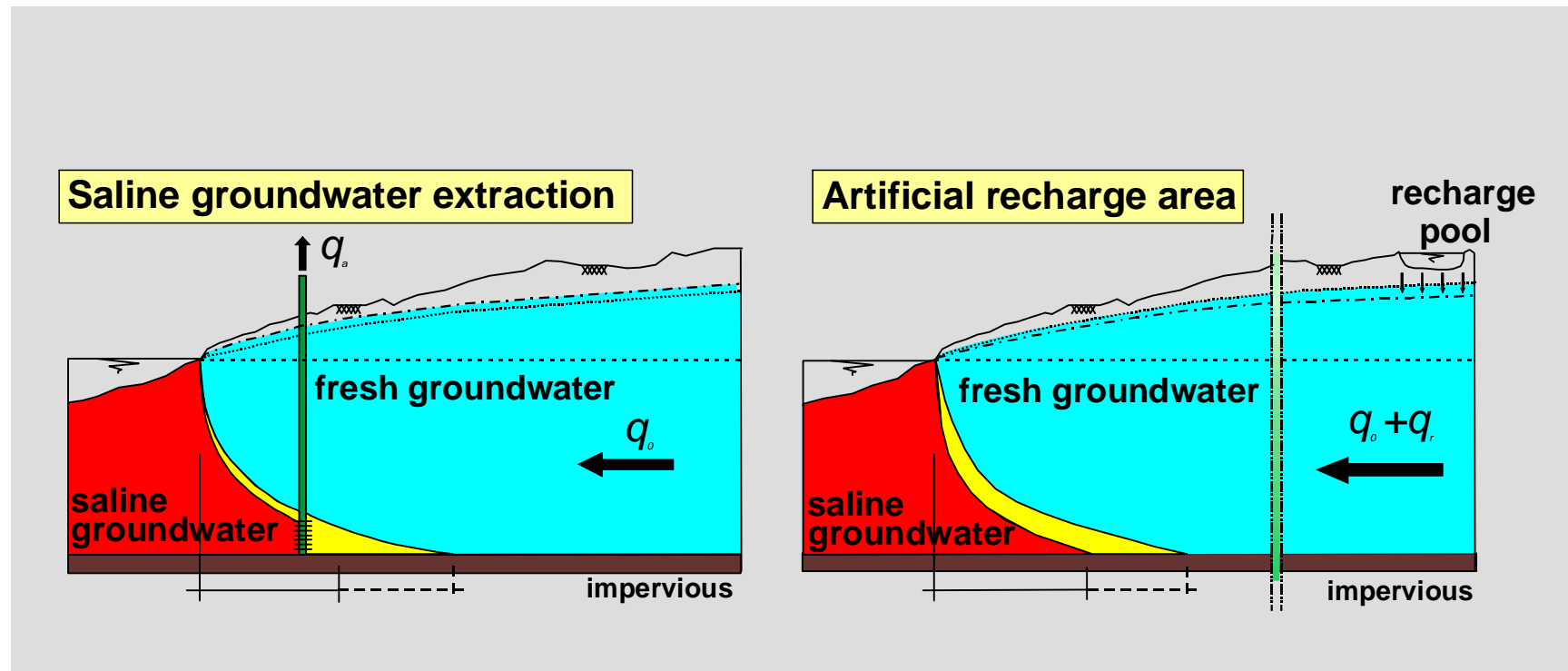


Land reclamation

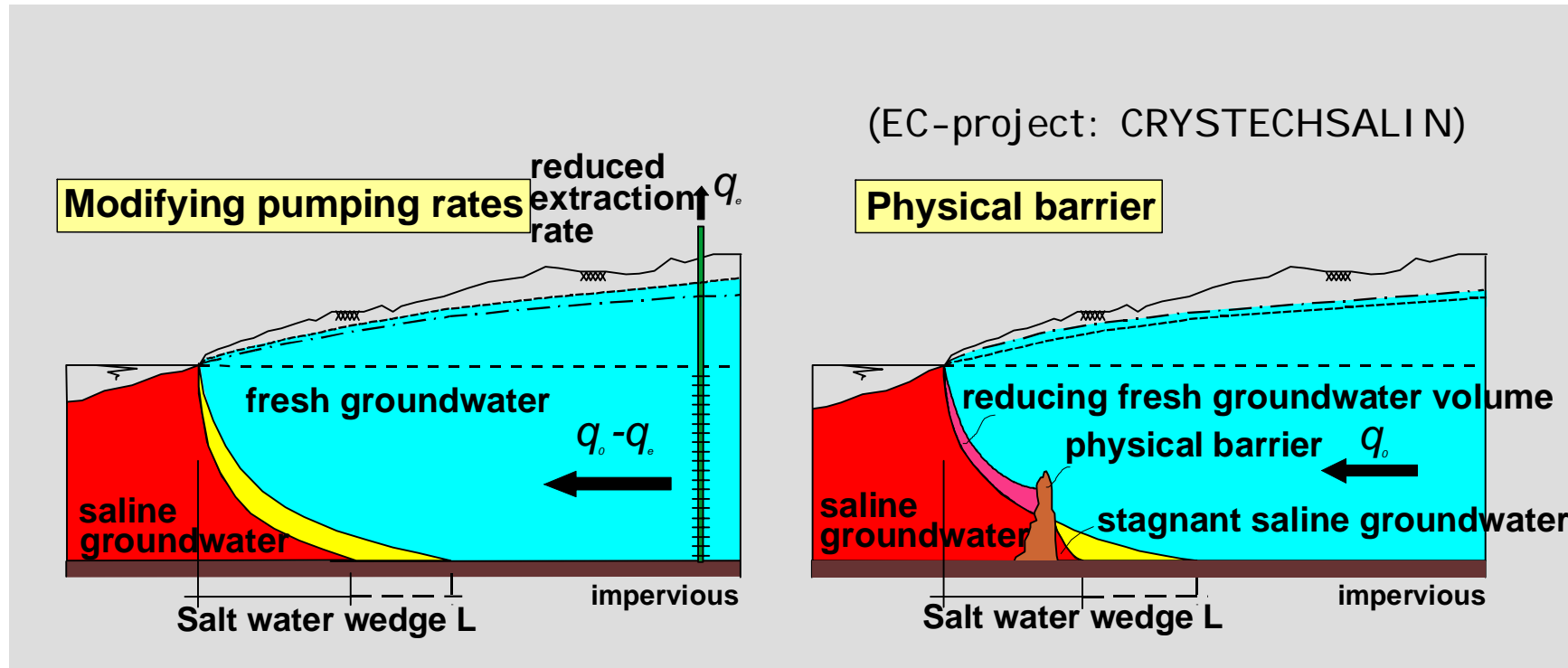




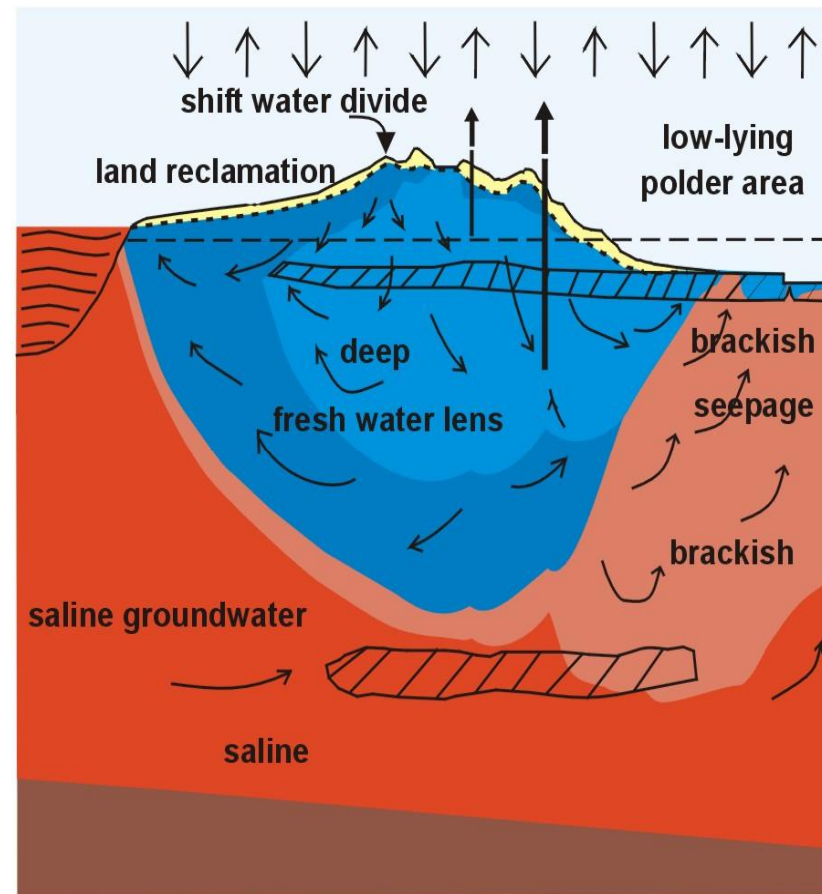
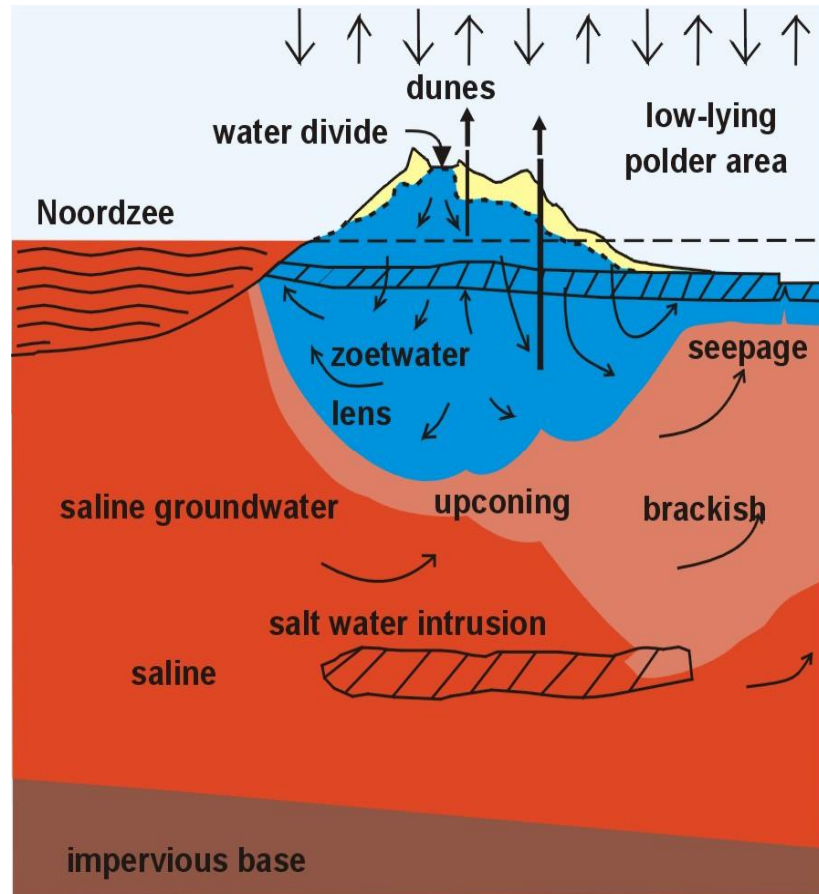
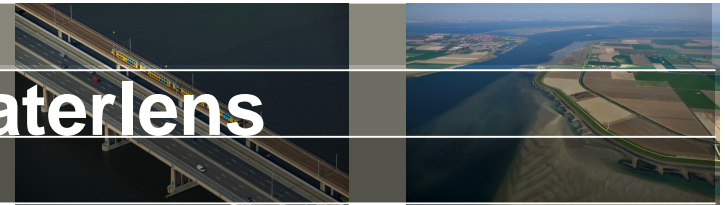
# Countermeasures of salt water intrusion (2)



# Countermeasures of salt water intrusion (3)



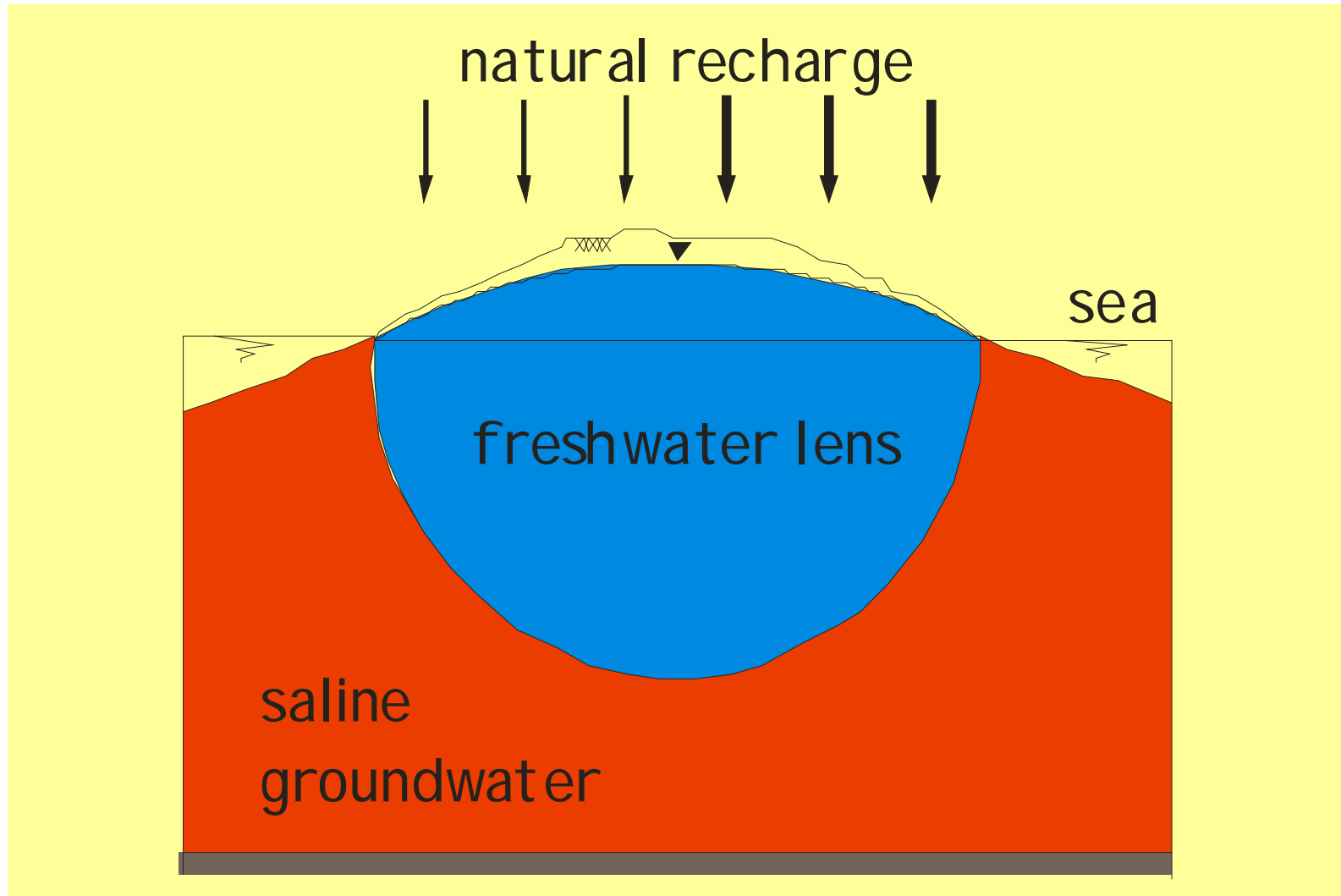
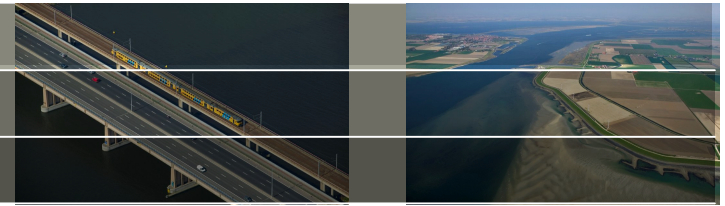
# Landaanwinning en de zoetwaterlens



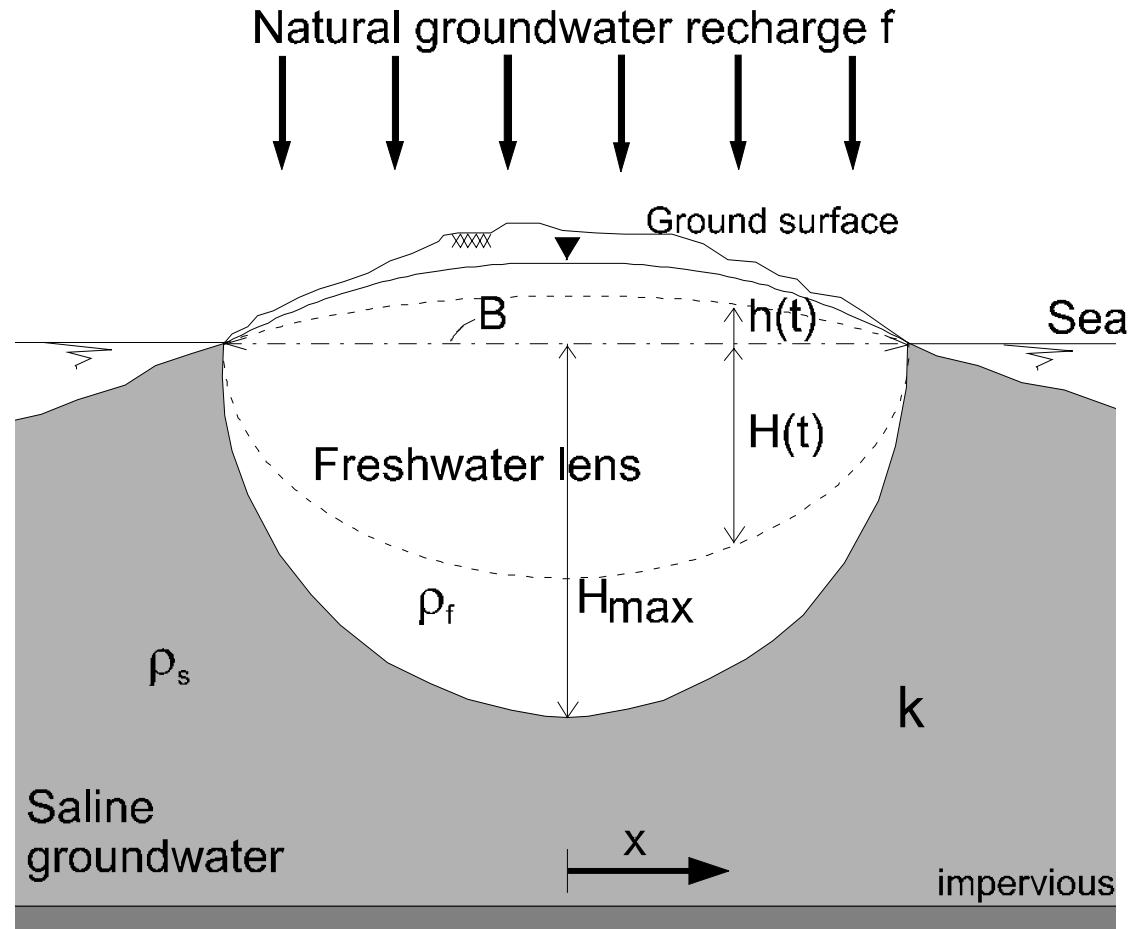
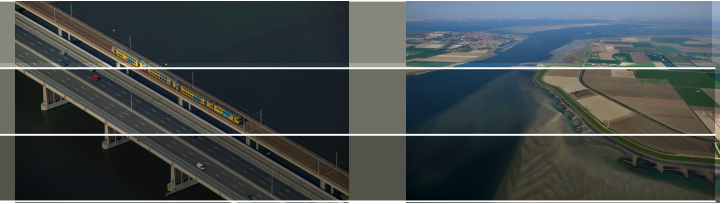


# Voorbeelden zoet-zout grondwater

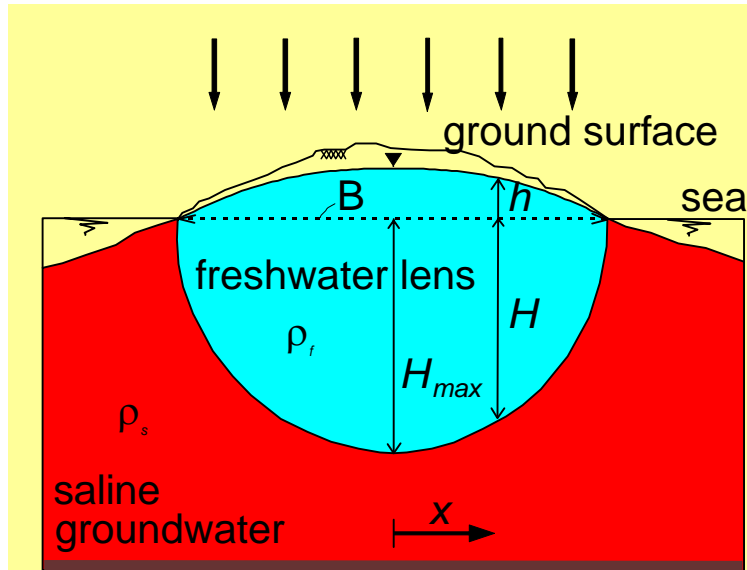
# Evolution of a freshwater lens



# Development of a freshwater lens



# Zoetwaterlens duingebied: analytische vergelijking



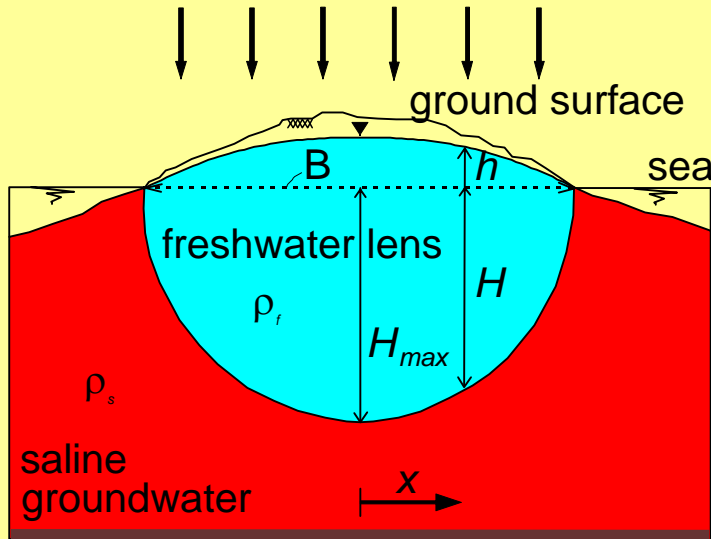
Depth of fresh-saline interface  $H$

$$H = \sqrt{\frac{f(0.25B^2 - x^2)}{k\alpha(1 + \alpha)}}$$

$$h = \alpha H$$



# Zoetwaterlens duingebied: analytische vergelijking



## Depth of fresh-saline interface $H$

$$B = 2000\text{m}, f = 0.001\text{m/day}$$

$$k = 10\text{m/day}, \alpha = 0.025$$

$$n_e = 0.35$$

Maximal thickness lens

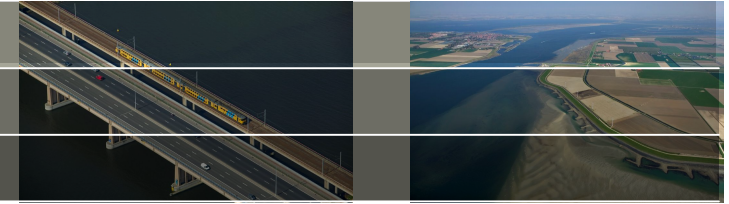
$$H_{\max} = 62.5\text{m}, h_{\max} = 1.56\text{m}$$

Volume lens (wrong in lectures notes)

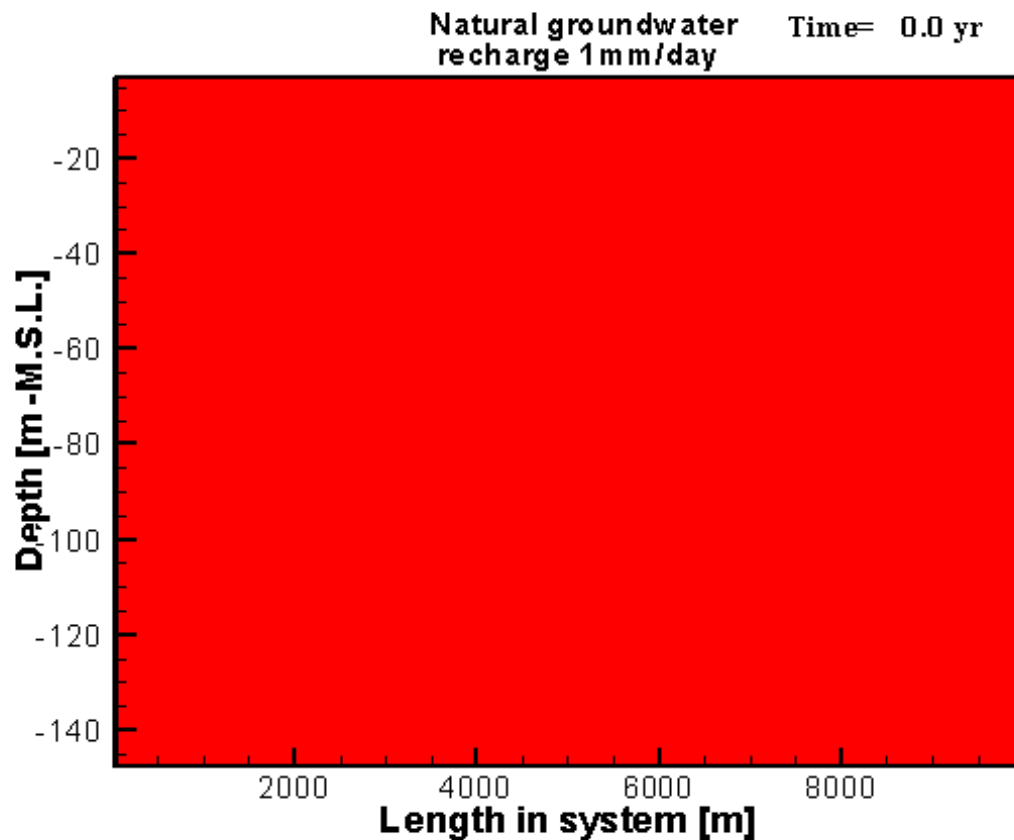
$$V = 35203\text{m}^3/\text{m}'$$

Characteristic time  $T = \frac{35203}{2} \text{days} = 48.2 \text{years}$

# Ontwikkeling van een zoetwaterlens



## Evolution of a freshwater lens



T = karakteristieke tijdschaal

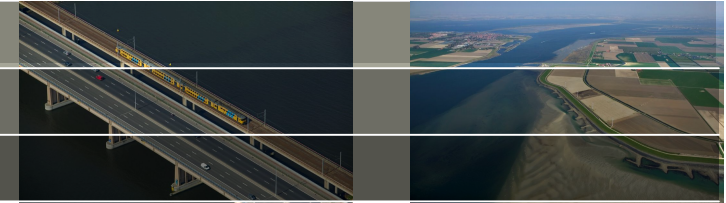
T = tijdsduur voordat lens 95% van zijn uiteindelijke vorm heeft

In Nederland: T = 75-200 jaar,

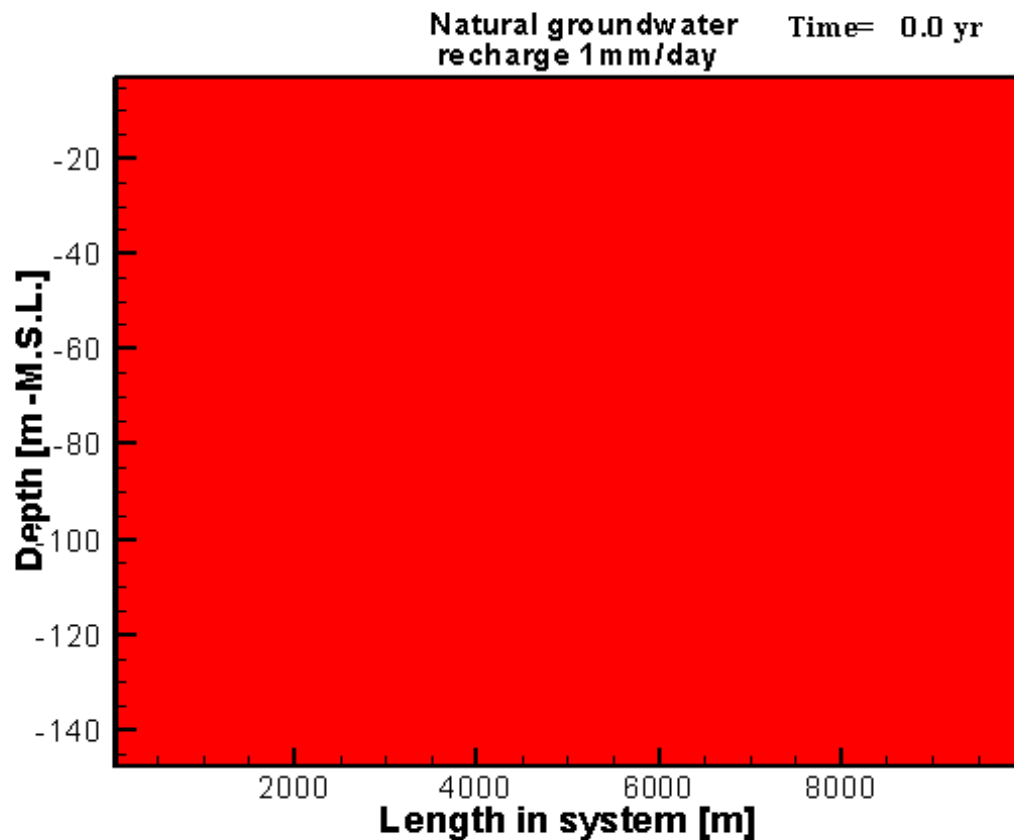
afhankelijk van:

- breedte duingebied
- neerslag
- doorlatendheid bodem

# Ontwikkeling van een zoetwaterlens



## Evolution of a freshwater lens



T = karakteristieke tijdschaal

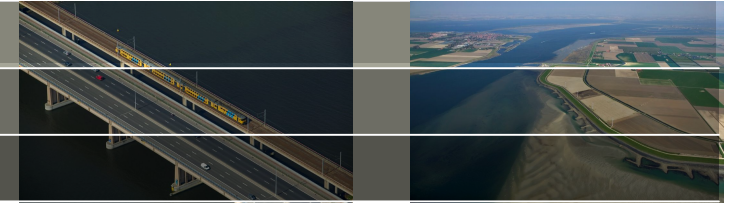
T = tijdsduur voordat lens 95% van zijn uiteindelijke vorm heeft

In Nederland: T = 75-200 jaar,

afhankelijk van:

- breedte duingebied
- neerslag
- doorlatendheid bodem

# Griend

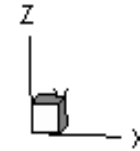


1. Het kleine eiland beweegt ~7.5m per jaar naar het oosten
2. Effect op de zoetwatervoorraad:
  - Kan er een zoetwater lens ontwikkelen?
  - Wat is de dikte van de lens?



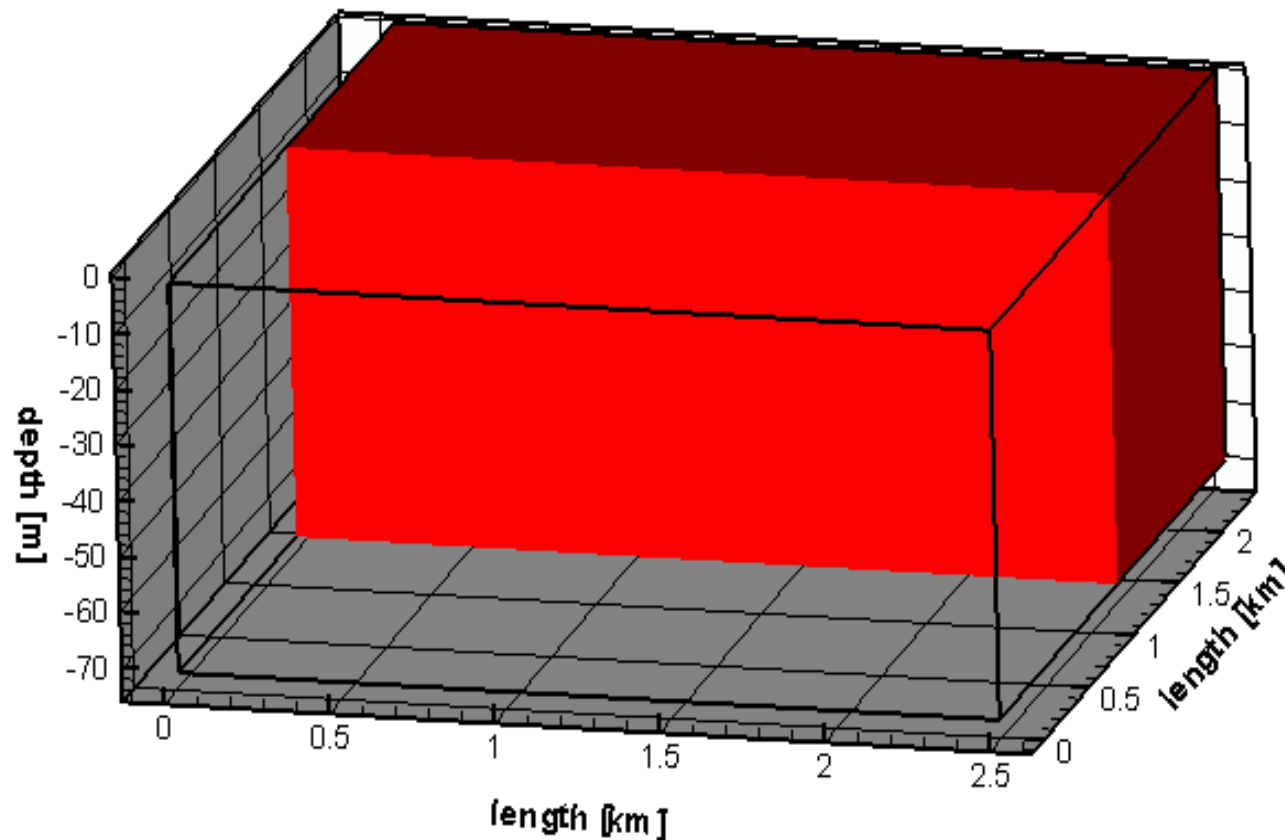
# Beweging van De Griend en de creatie van de lens

Movement of the island of Griend (Friesland) through the Waddenzee



7.5m/jaar to the east

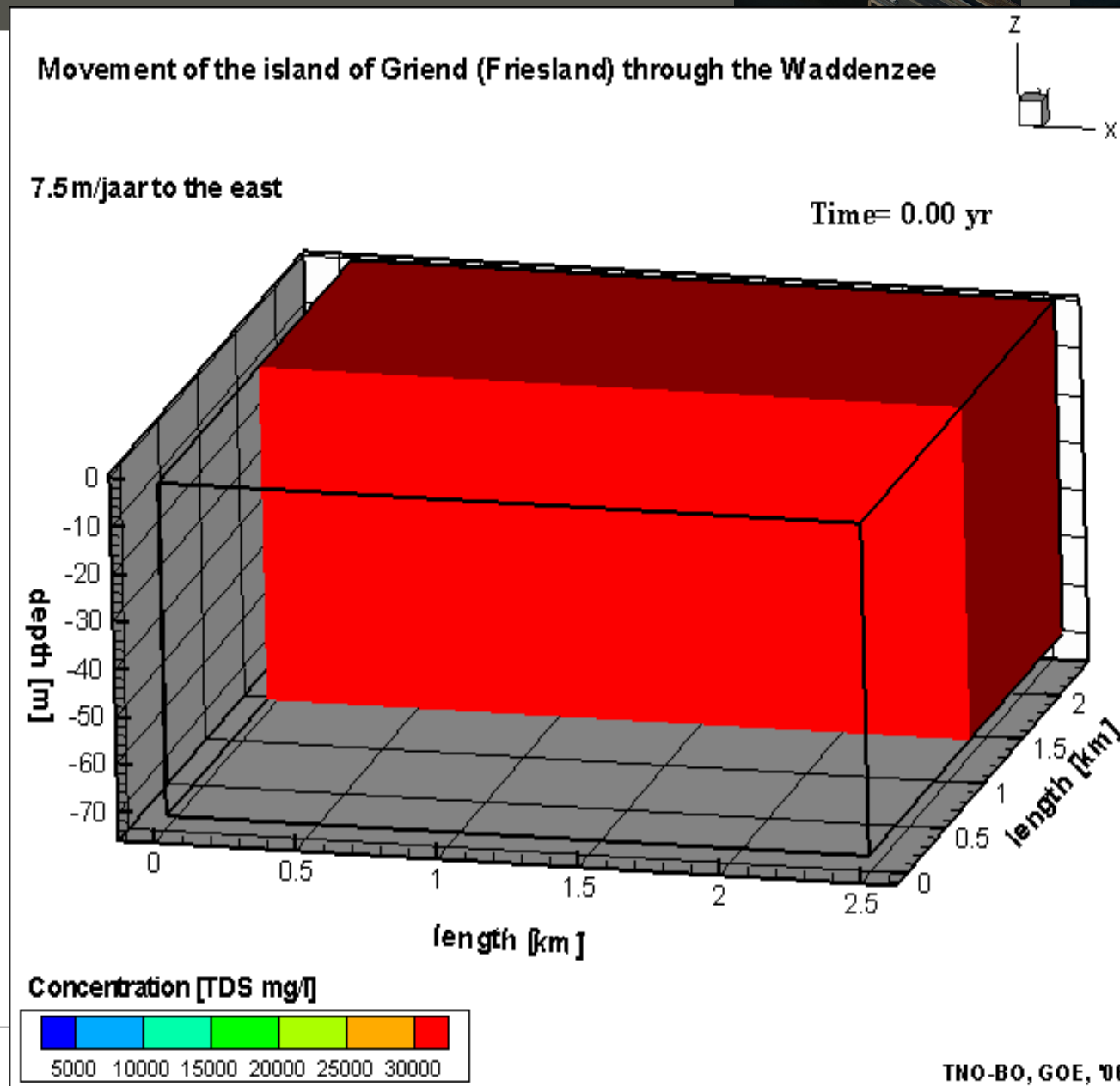
Time= 0.00 yr



Concentration [TDS mg/l]

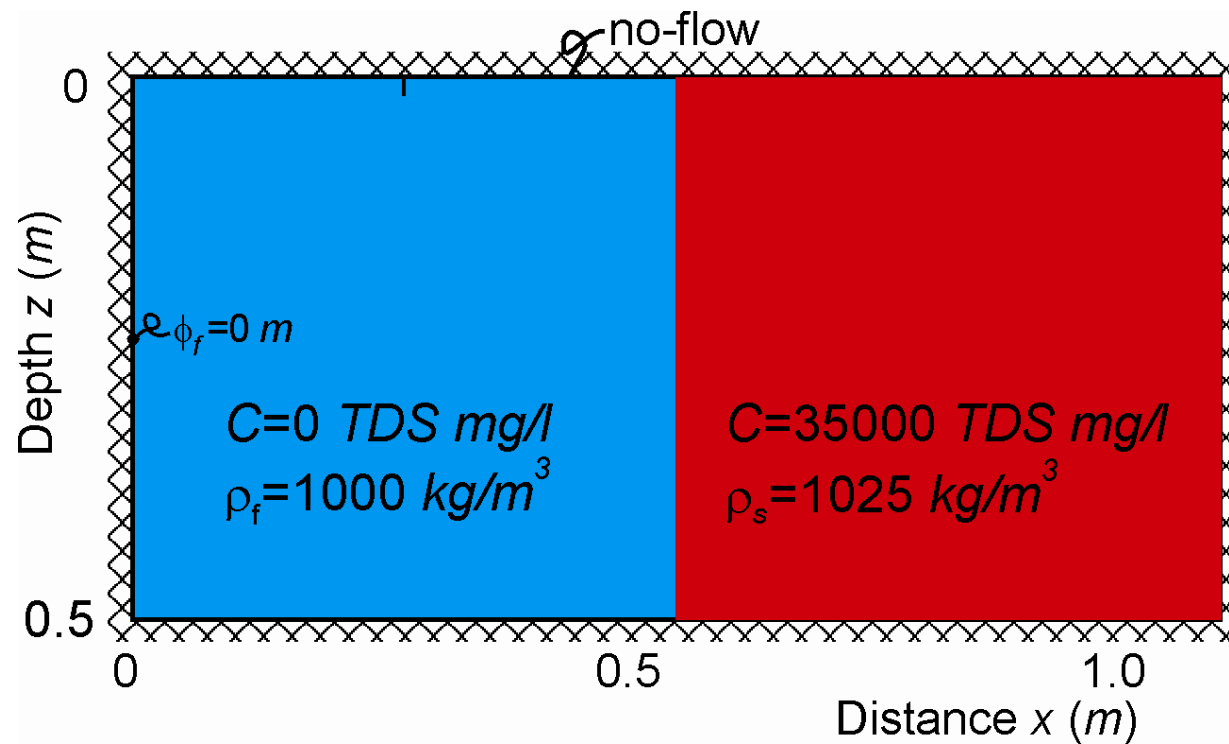


# Beweging van De Griend en de creatie van de lens





# Vertical interface between fresh and saline groundwater

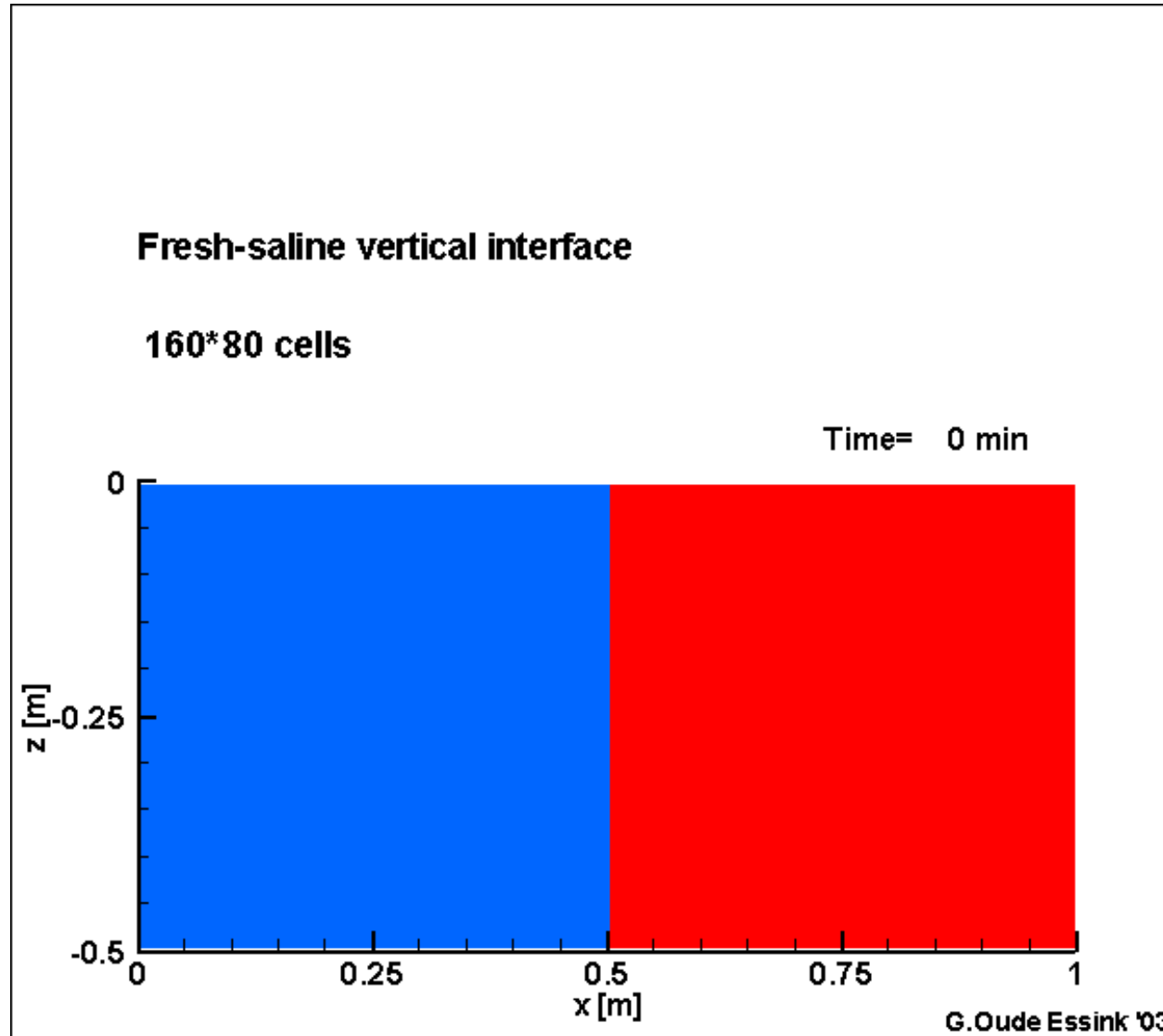


Subsoil parameters:

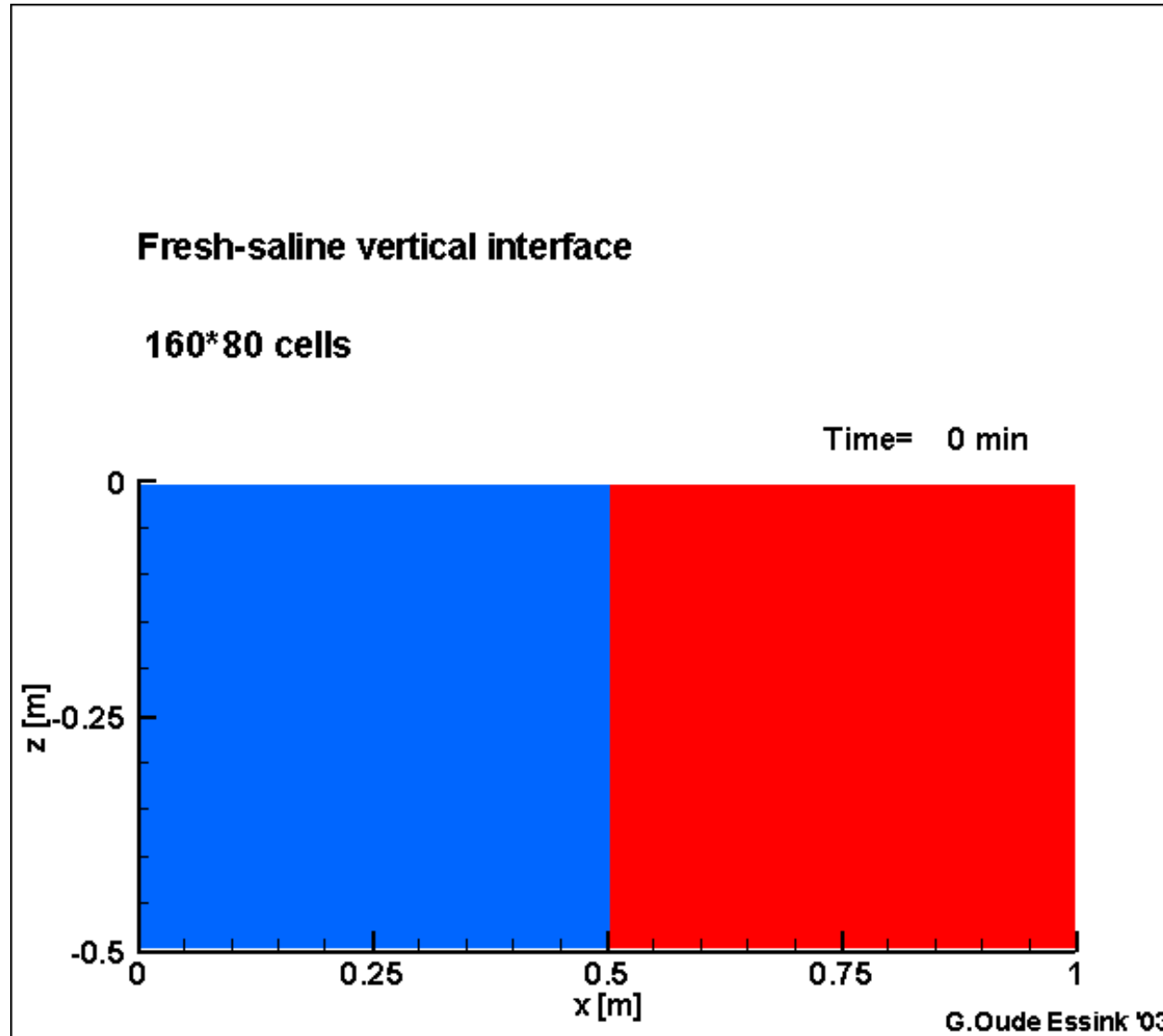
$$k = 10^{-3} \text{ m/s}$$
$$n_e = 0.10$$
$$\alpha_L = \alpha_{TH} = \alpha_{TV} = 0 \text{ m}$$
$$D_m = 0 \text{ m}^2/\text{s}$$



# Dichtheidsafhankelijke grondwaterstroming



# Dichtheidsafhankelijke grondwaterstroming



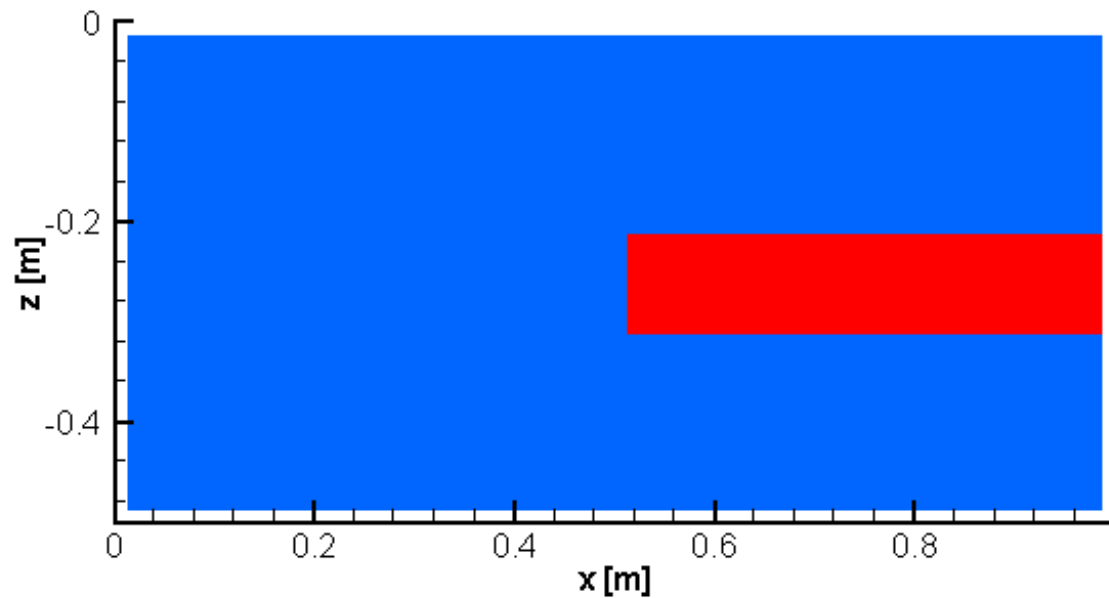
# Dichtheidsafhankelijke grondwaterstroming



**Saline pocket in fresh groundwater: fingering process**

**40\*20 cells**

Time= 0 min



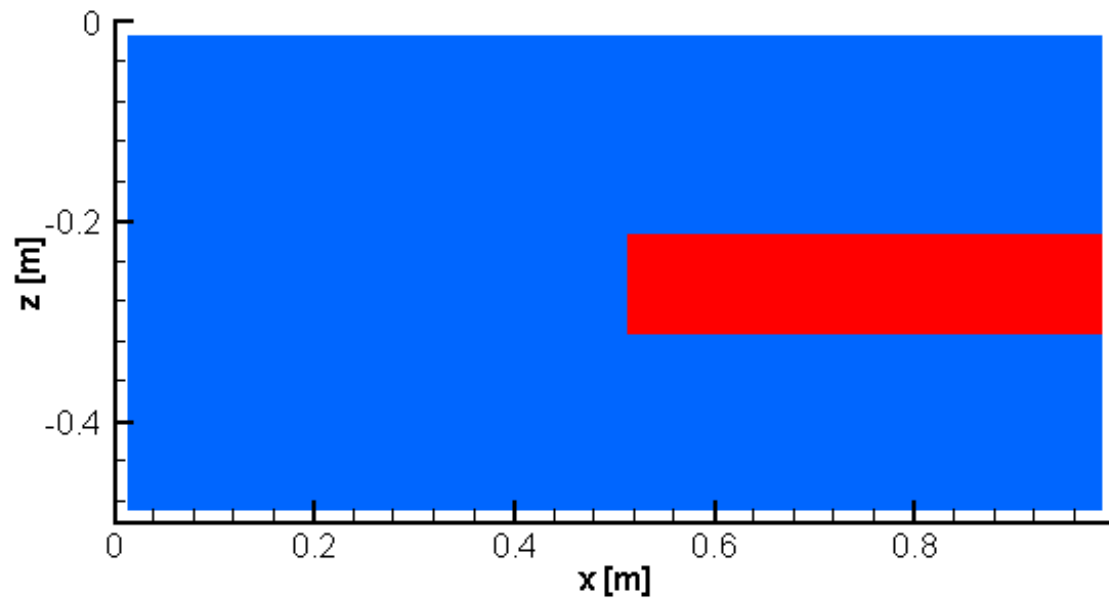
# Dichtheidsafhankelijke grondwaterstroming



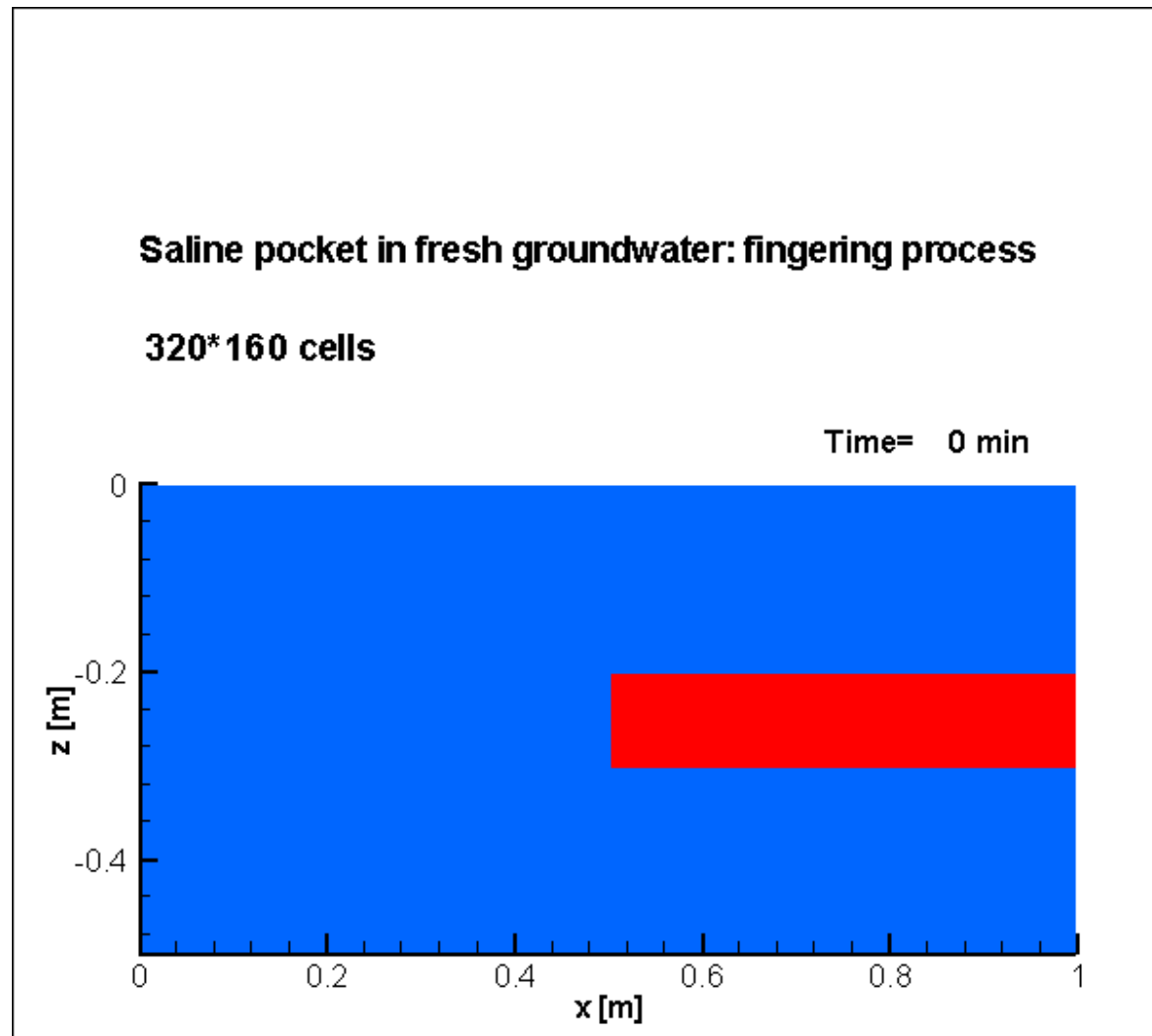
**Saline pocket in fresh groundwater: fingering process**

**40\*20 cells**

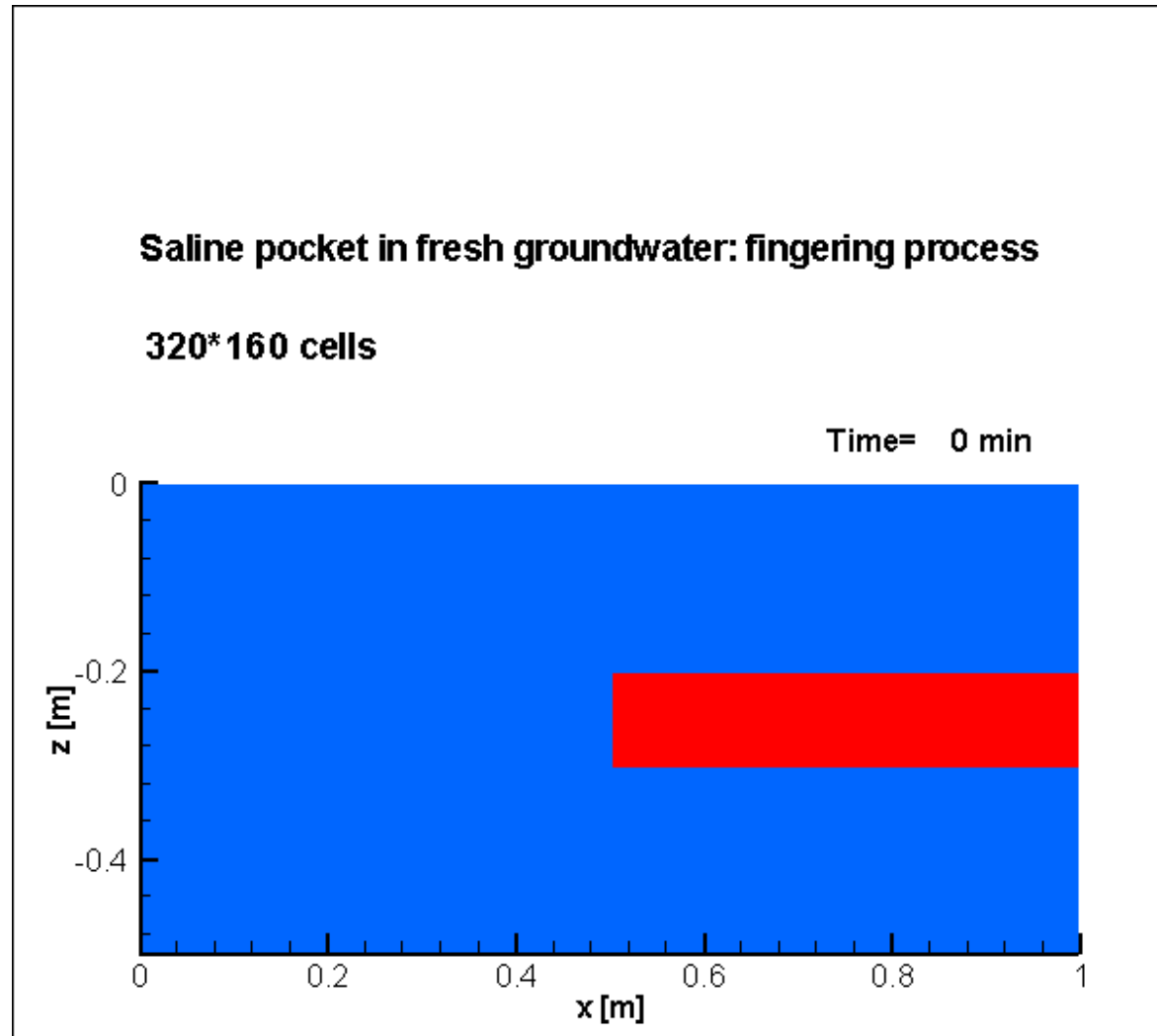
Time= 0 min



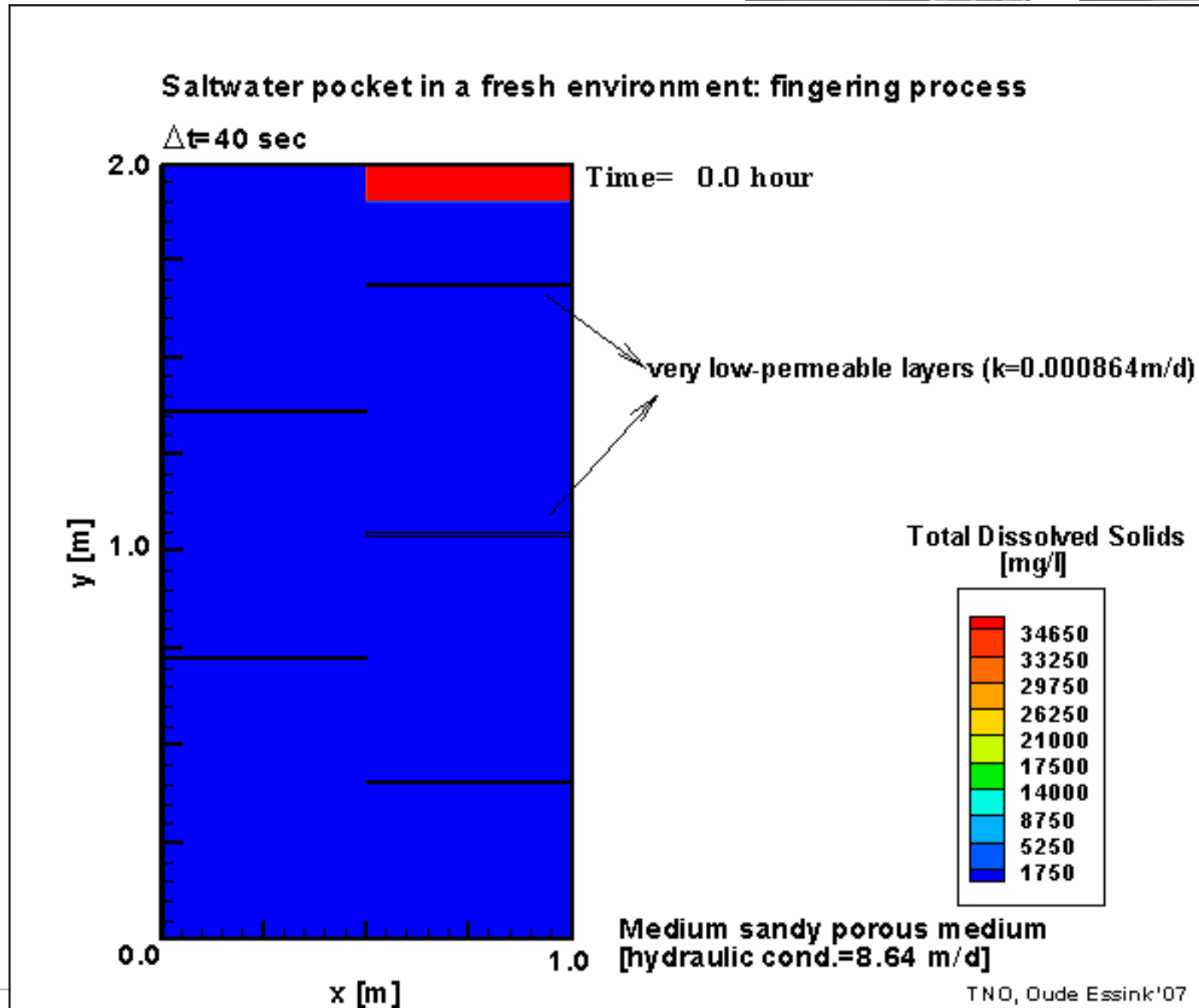
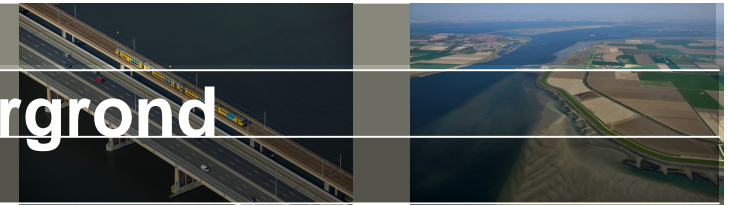
# Dichtheidsafhankelijke grondwaterstroming



# Dichtheidsafhankelijke grondwaterstroming

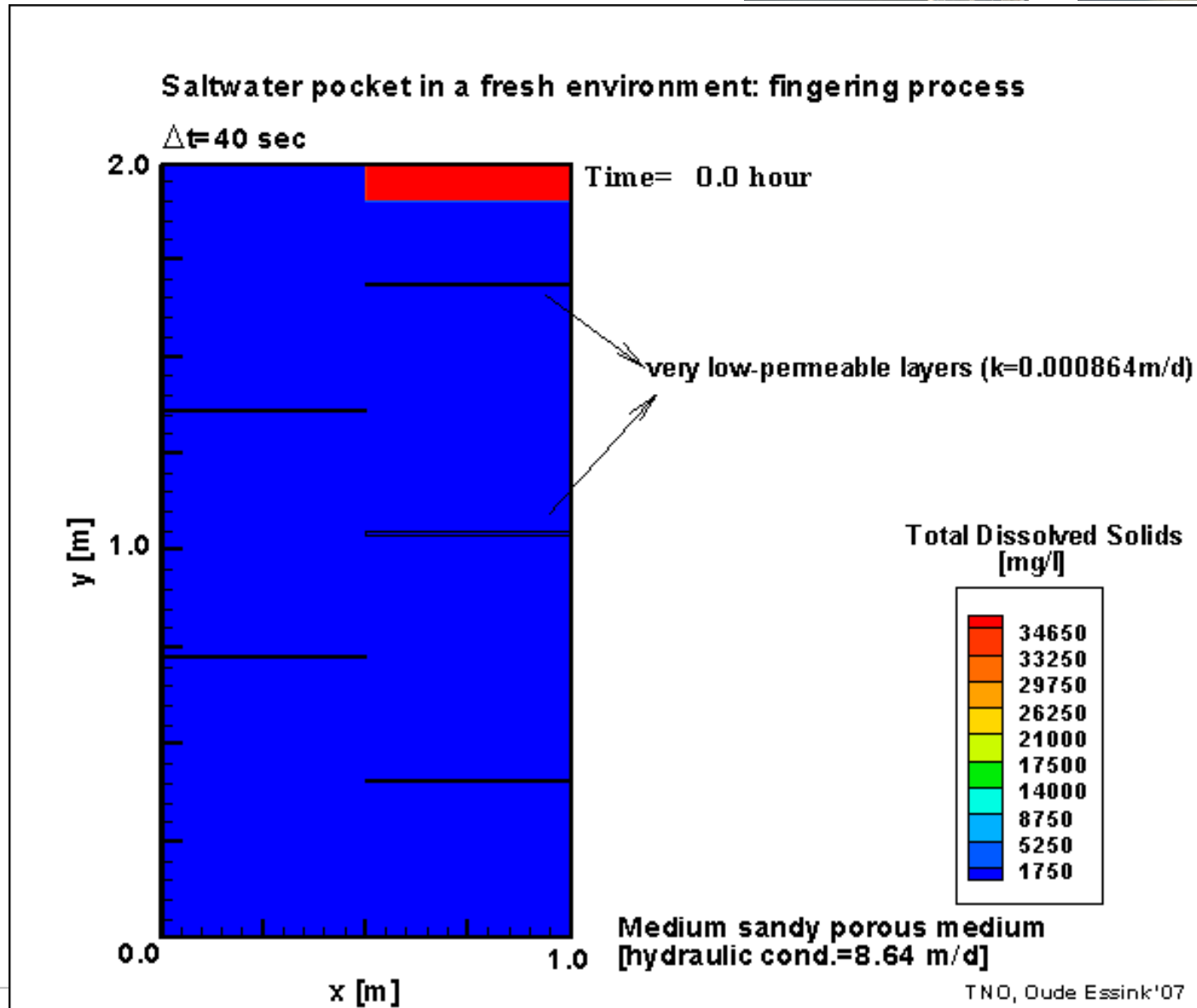
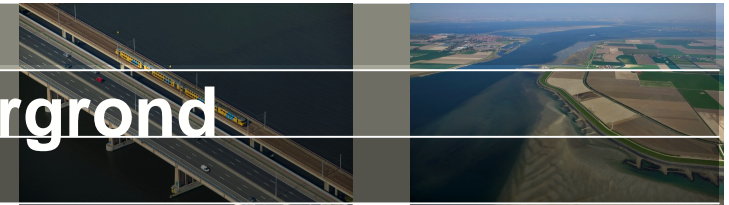


# Fingering processes in de ondergrond





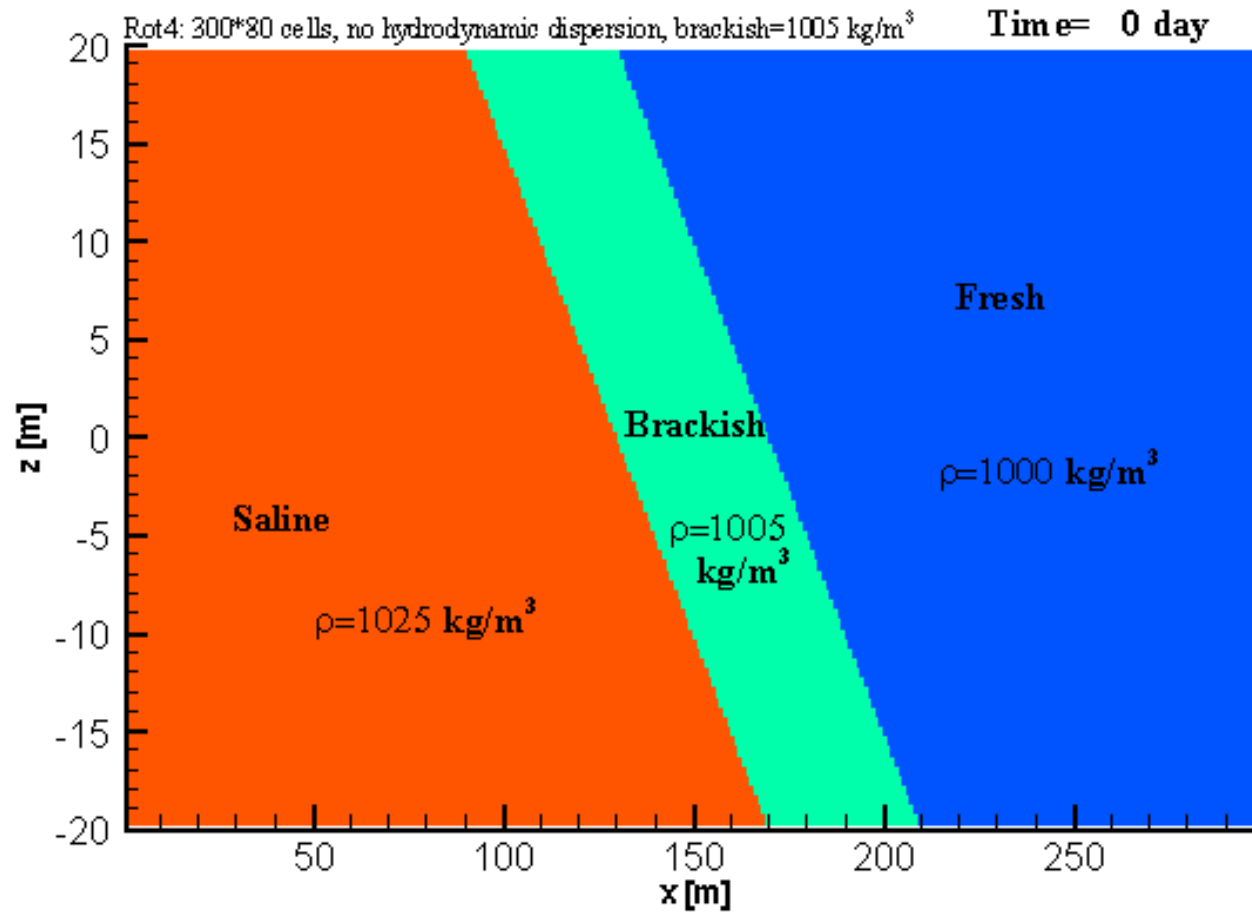
# Fingering processes in de ondergrond



# Roterend asymmetrisch scherp grensvlak



Rotating movement of three immiscible fluids (asymmetric case)



J. Hydrology, 2004  
nr. 287, pp. 270-278

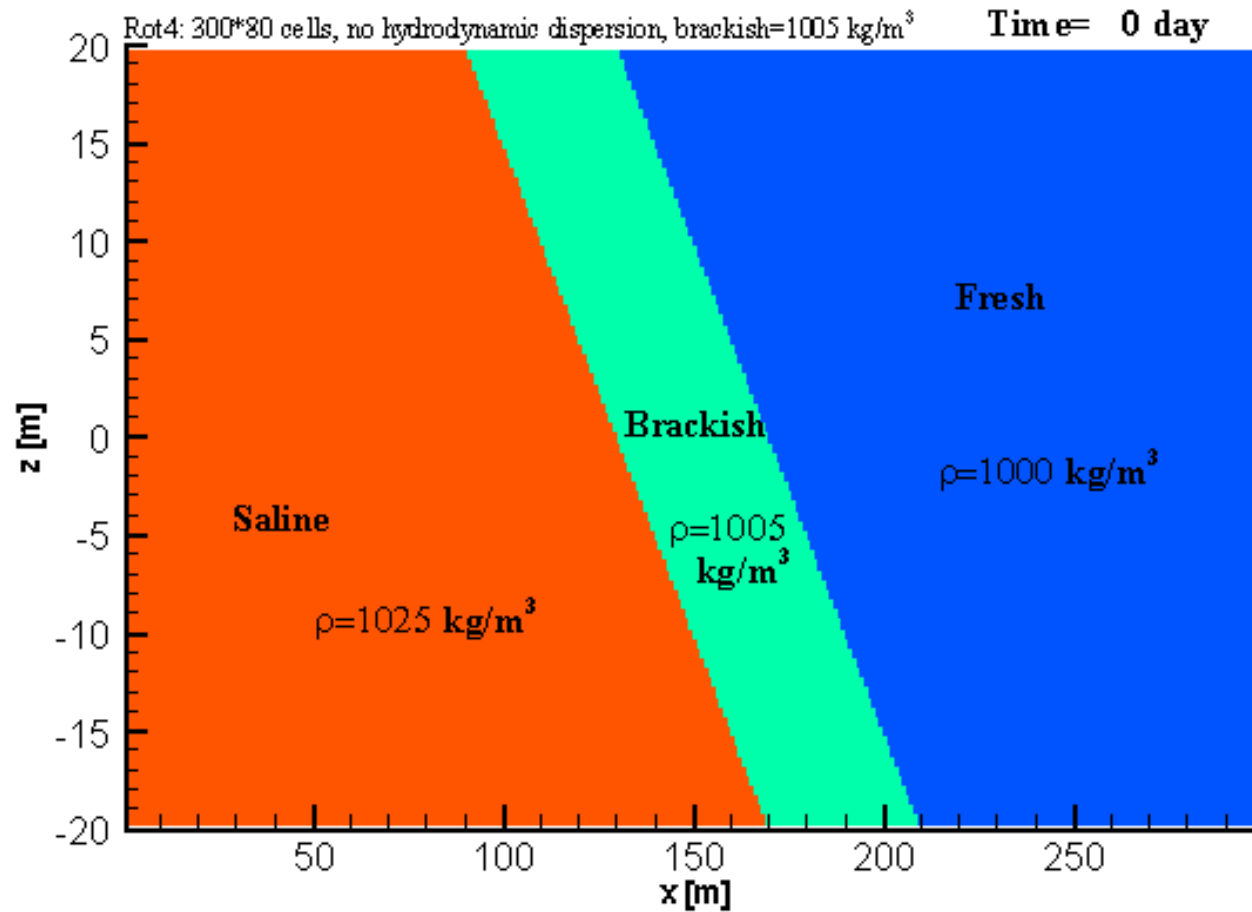
**Deltares**

GOE

# Roterend asymmetrisch scherp grensvlak



Rotating movement of three immiscible fluids (asymmetric case)



J. Hydrology, 2004  
nr. 287, pp. 270-278

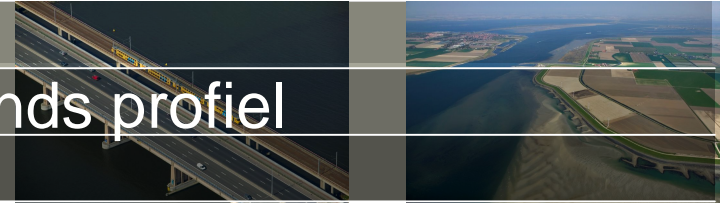
**Deltares**

GOE

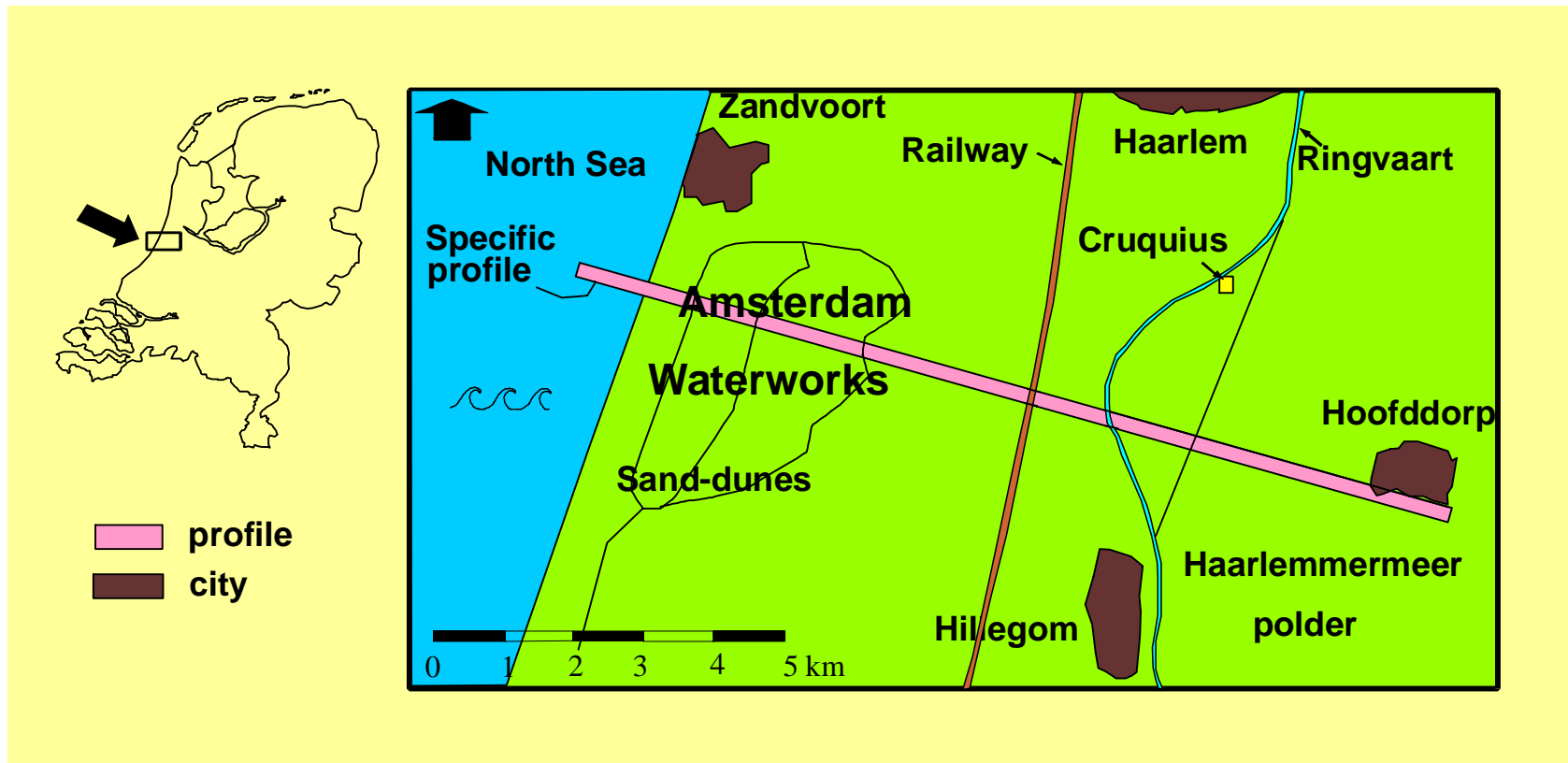


# Mijn eerste zoet-zout model

# Zoutwaterintrusie in een typisch Hollands profiel

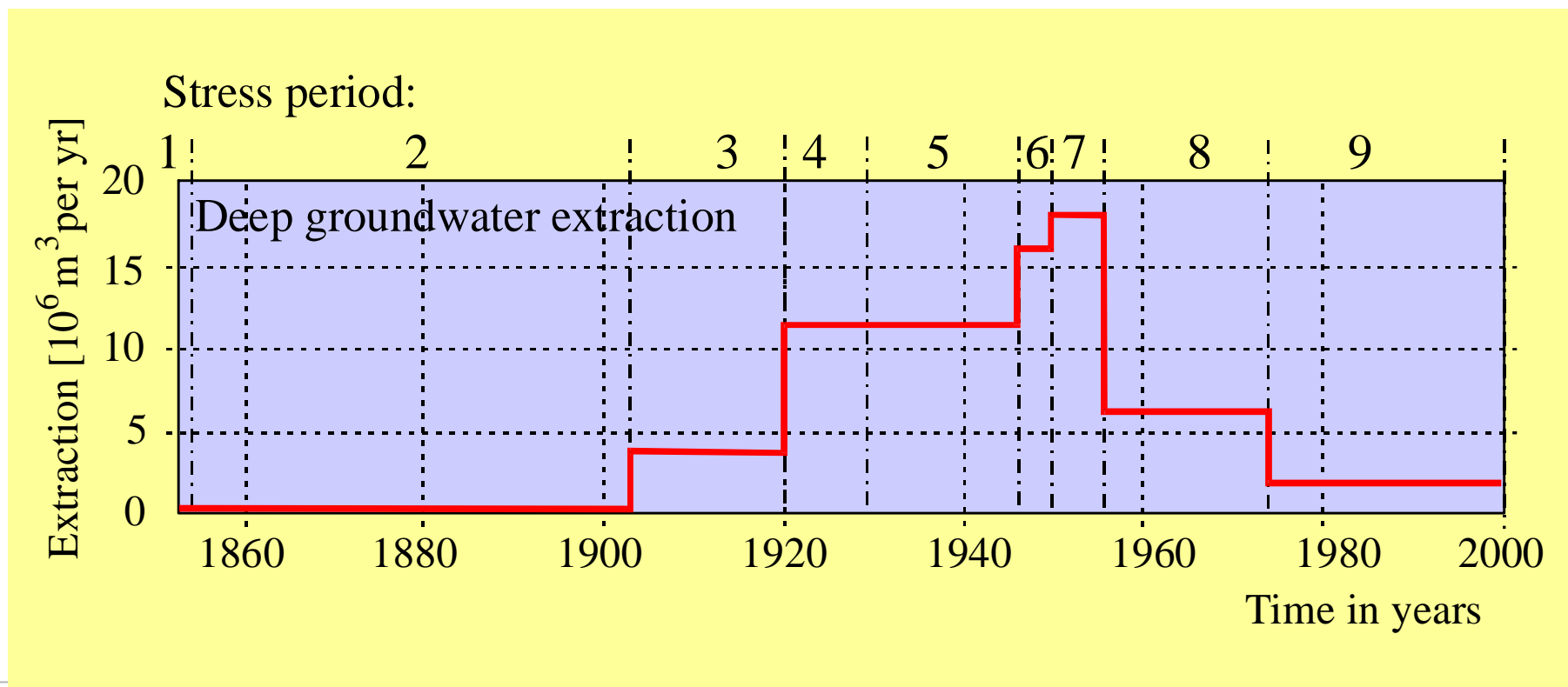


Positie profiel Duingebied Waternet, Rijnland polders en Haarlemmermeer polder

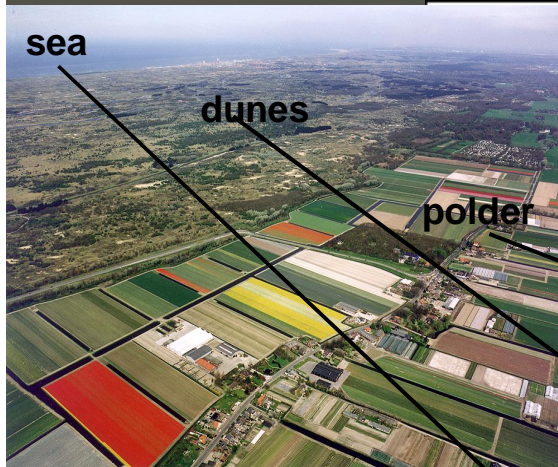


# Zoutwaterintrusie in een typisch Hollands profiel

Grondwater onttrekkingen uit het middelste watervoerend pakket uit het duingebied Waternet



# Zoutwater intrusie in het Nederlandse kustgebied

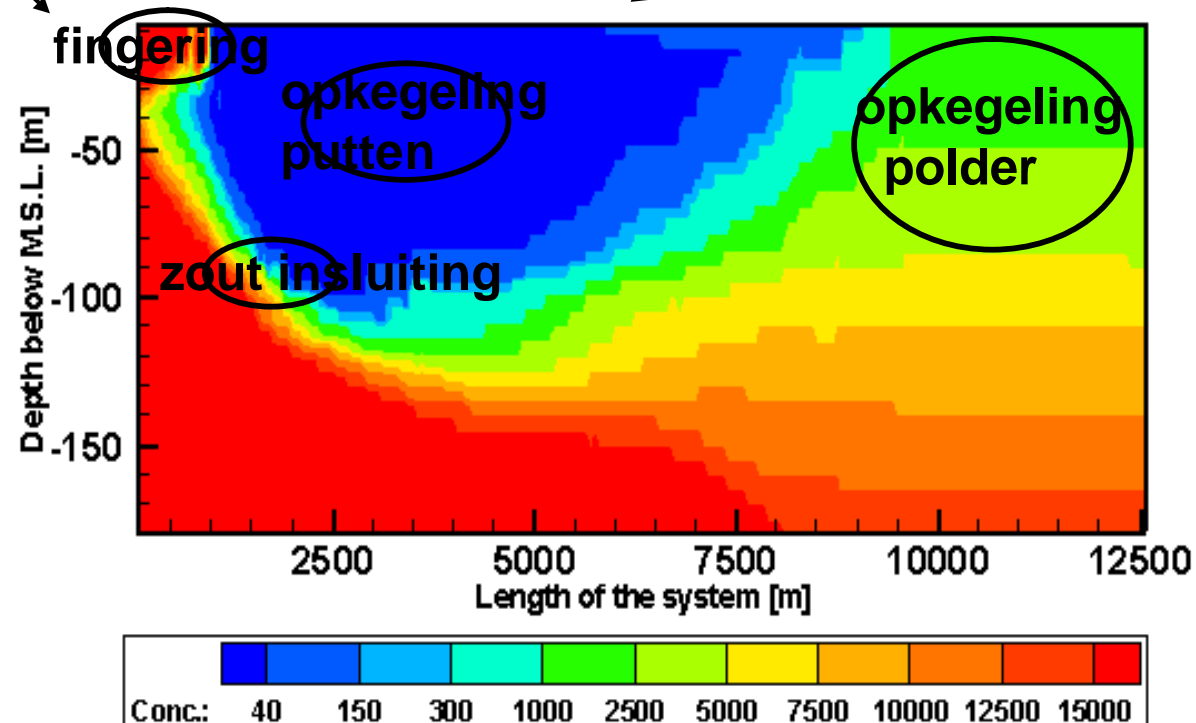


## Salinisation of the groundwater flow system

caused due to groundwater extractions and lowering of the ground surface of the Haarlemmermeer polder

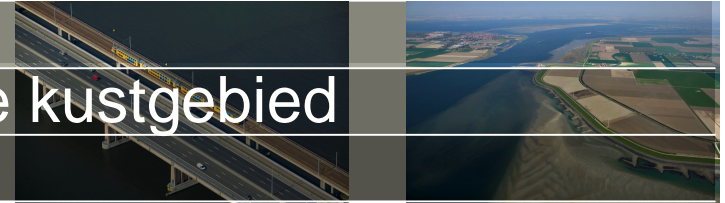
Profile Amsterdam Waterworks-Haarlemmermeerpolder

Time= 1854 AD





# Zoutwater intrusie in het Nederlandse kustgebied

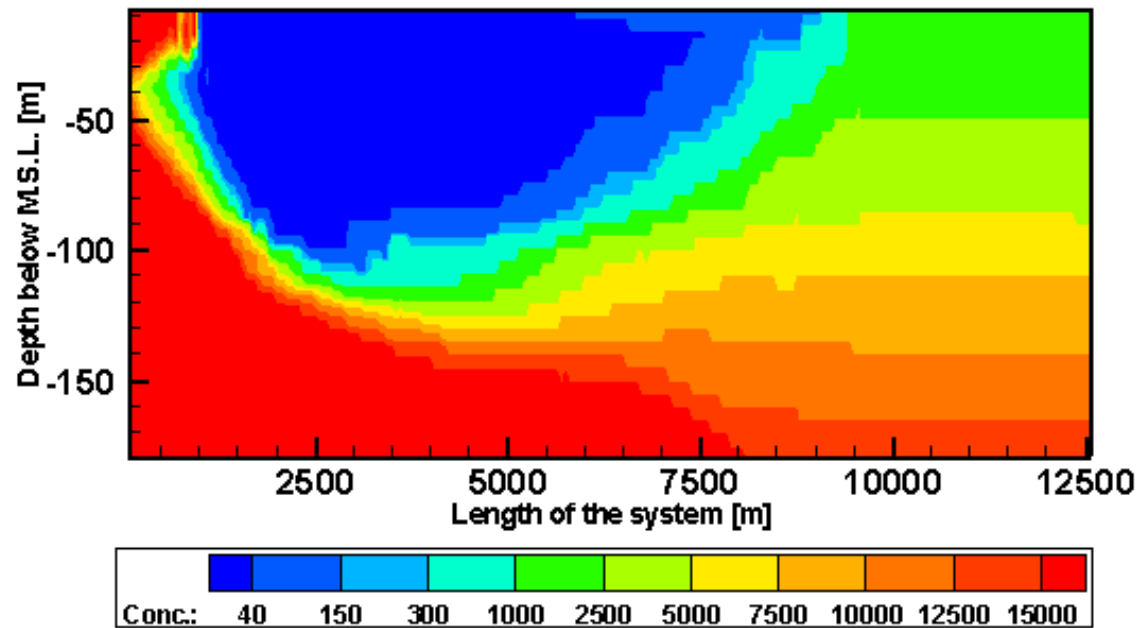


## Salinisation of the groundwater flow system

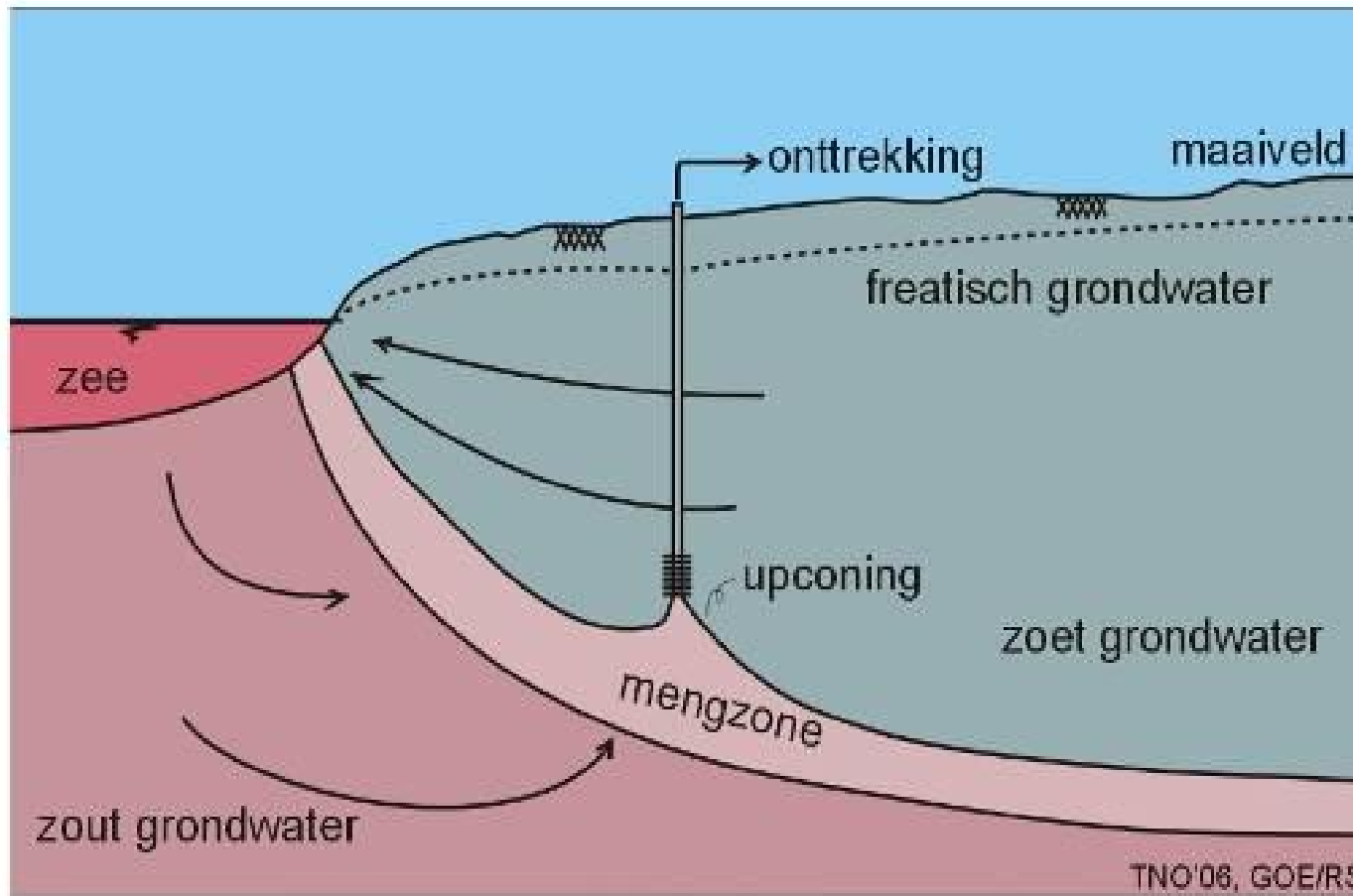
caused due to groundwater extractions and lowering of the ground surface of the Haarlemmermeer polder

Profile Amsterdam Waterworks-Haarlemmermeerpolder

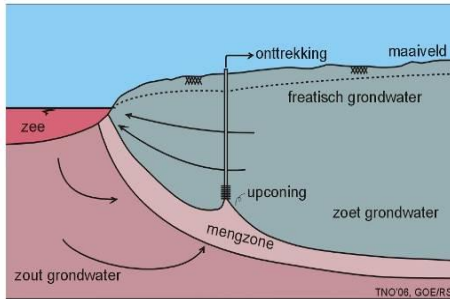
Time= 1854 AD



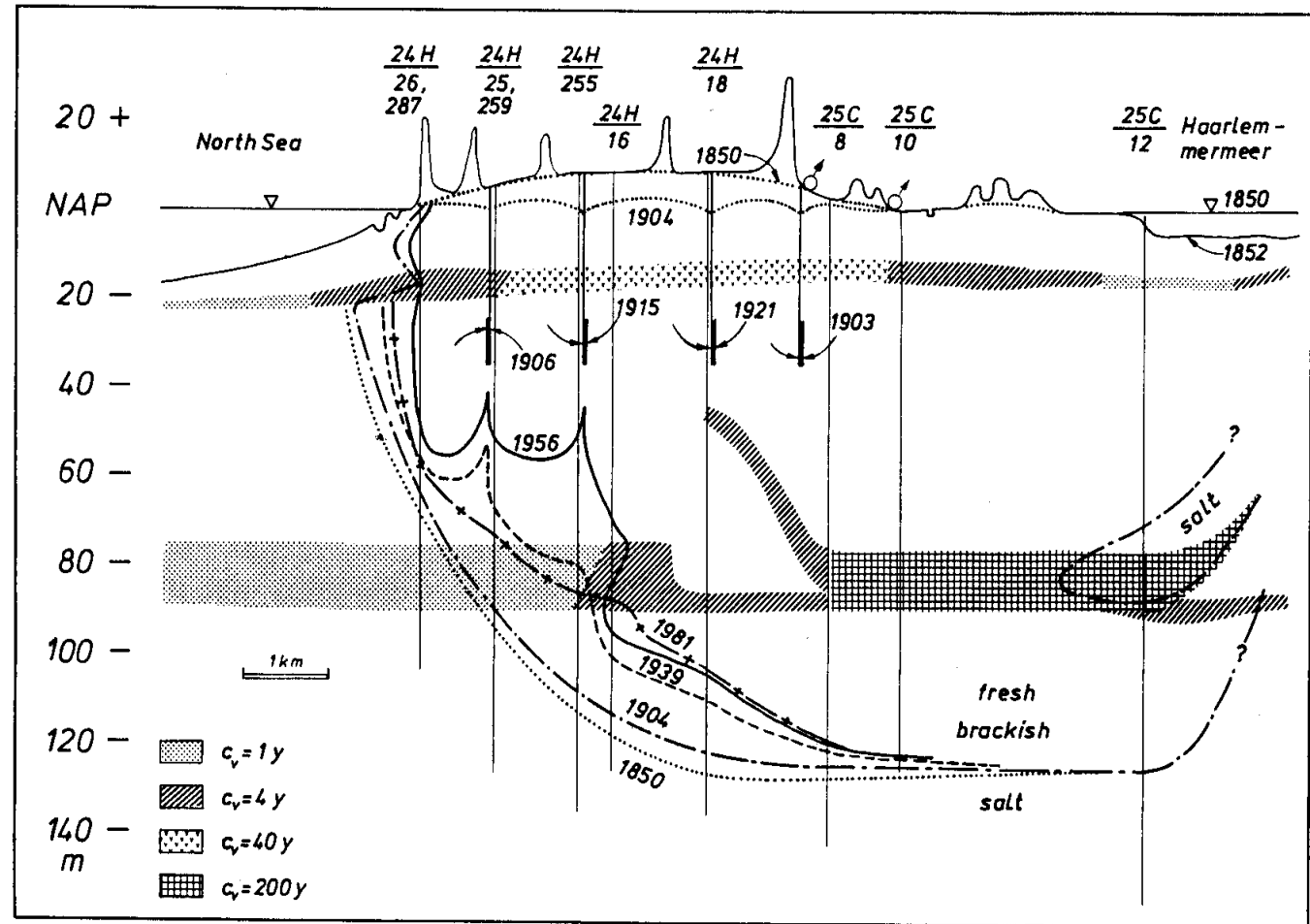
# Opkegeling van brak-zout grondwater, voorbeeld



# Opkegeling van brak-zout grondwater, voorbeeld



Zandvoort -  
Haarlemmermeer



Stuyfzand, '93

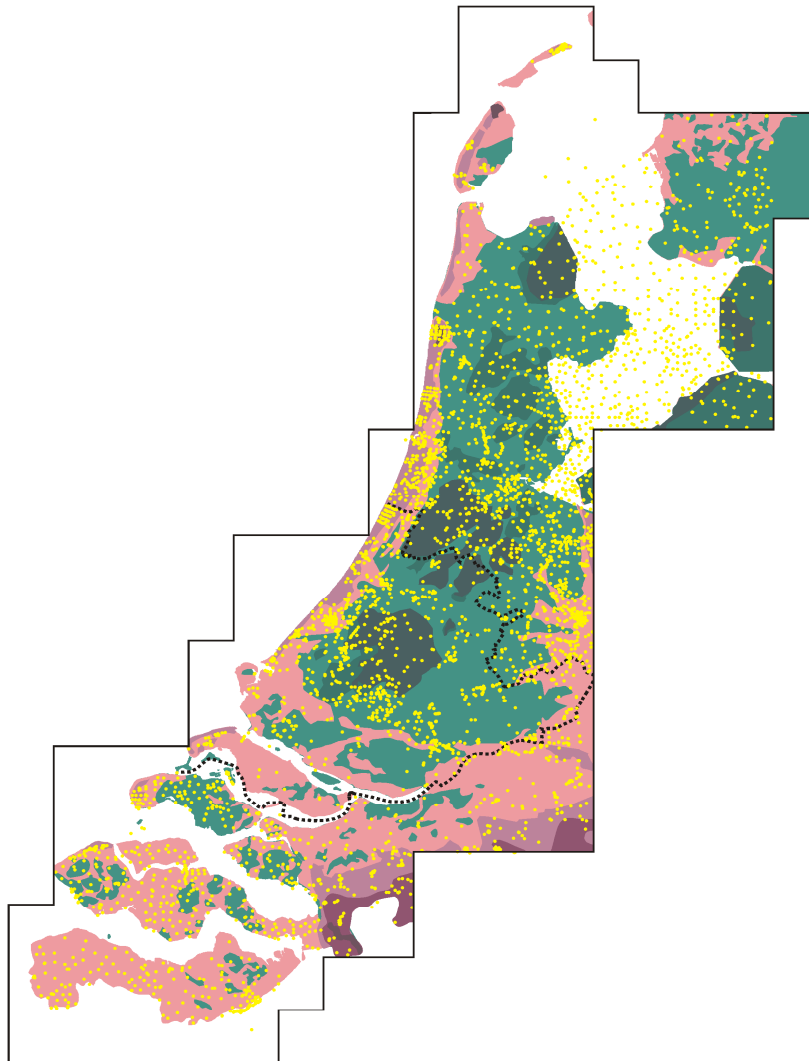


Grondwateronttrekkingen in het kustgebied  
veroorzaken opkegeling van brak-zout  
grondwater



# Initiele zoet-zout verdeling

# Database kustgebied: positie Cl-observatie filters



## Data: DINO-Qua

- 17500 Chloride-metingen

Diepte (m)	Aantal
mv- -5	4135
-5 -10	1095
-10 -25	5287
-25 -50	3678
-50 -75	1438
-75 -100	760
-100 -200	982
-200 -300	110
> -300	16

- Maximaal 64 metingen in 1 put

- Langste tijdreeks in 1 filter: 72 jaar

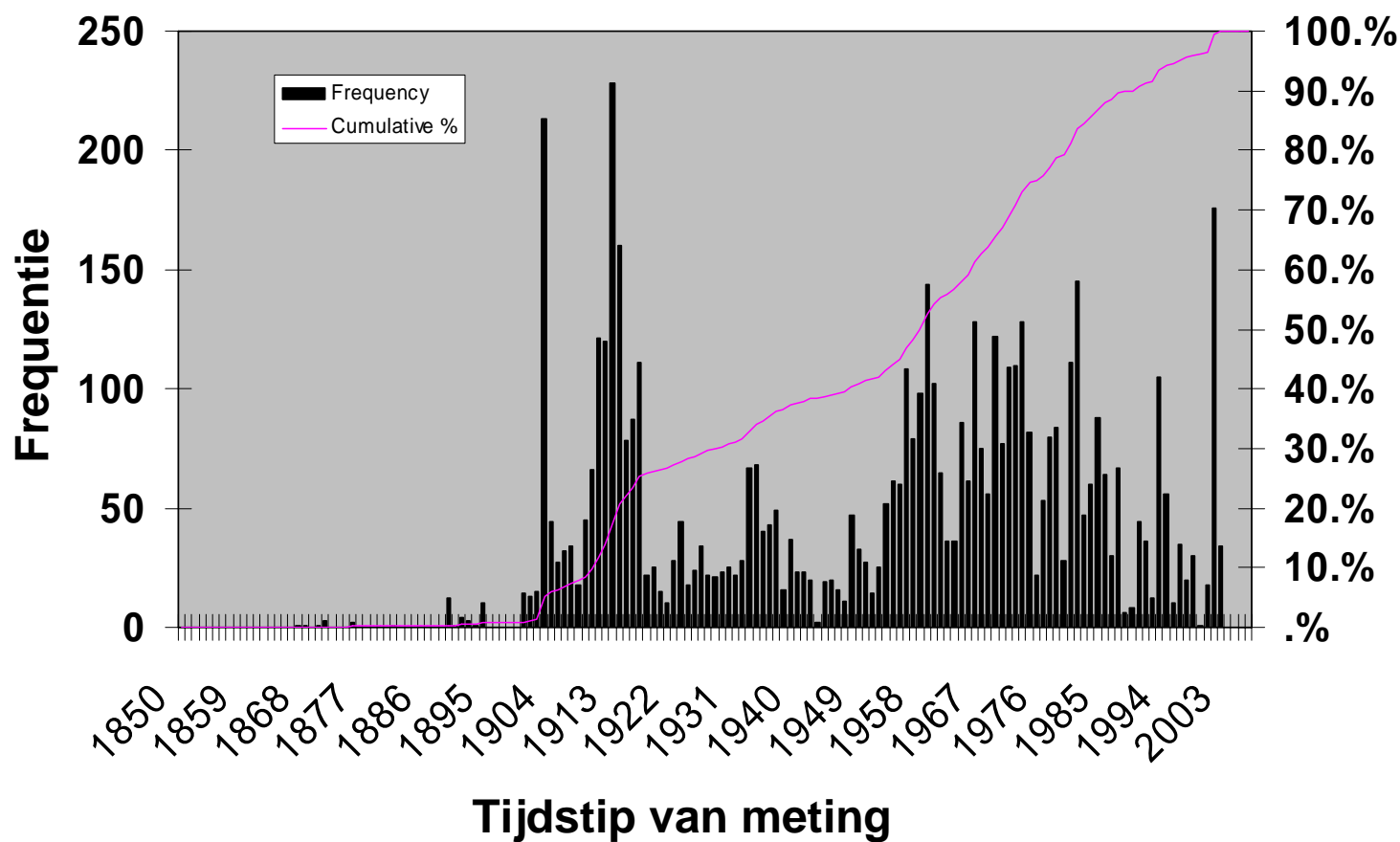
- Eerste meting op 2 jan 1853

- Aantal tijdreeksen

(minimaal twee metingen in een filter)  
2126 (gehele gebied)

# Chloride concentraties in Zuid-Holland: ~6000 metingen

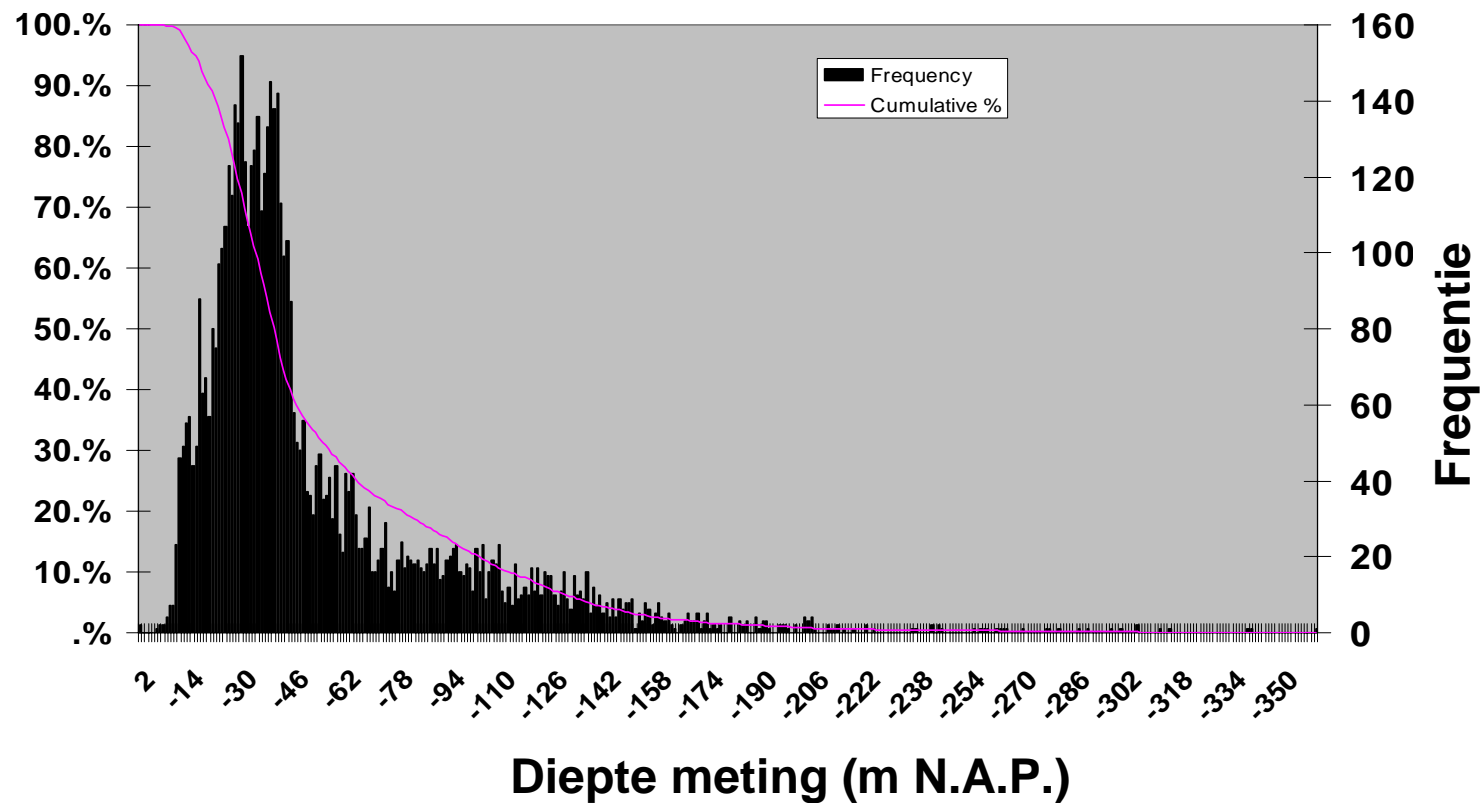
## Histogram



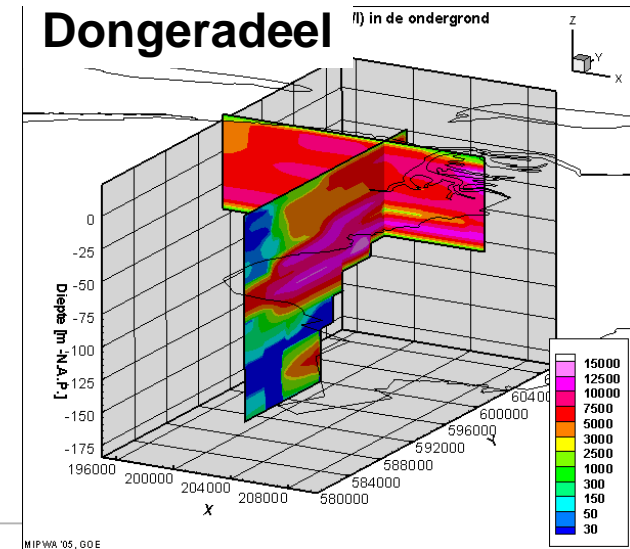
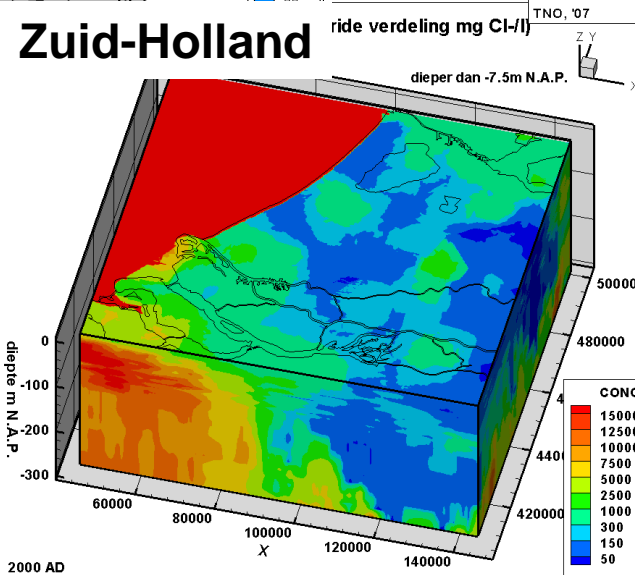
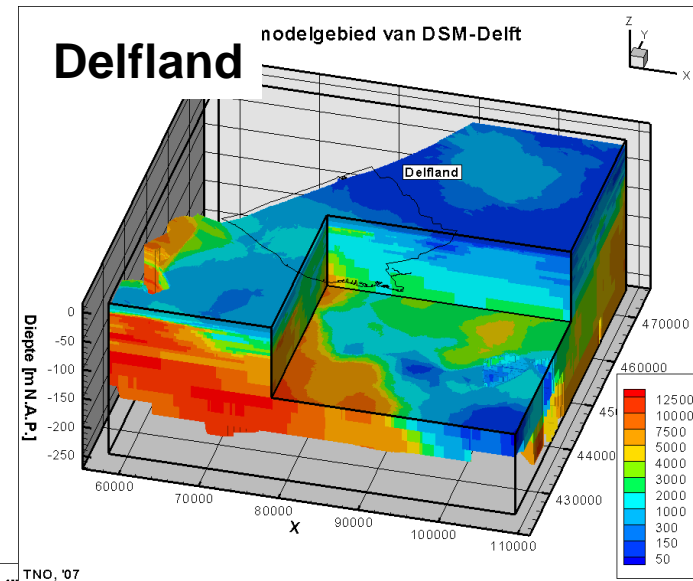
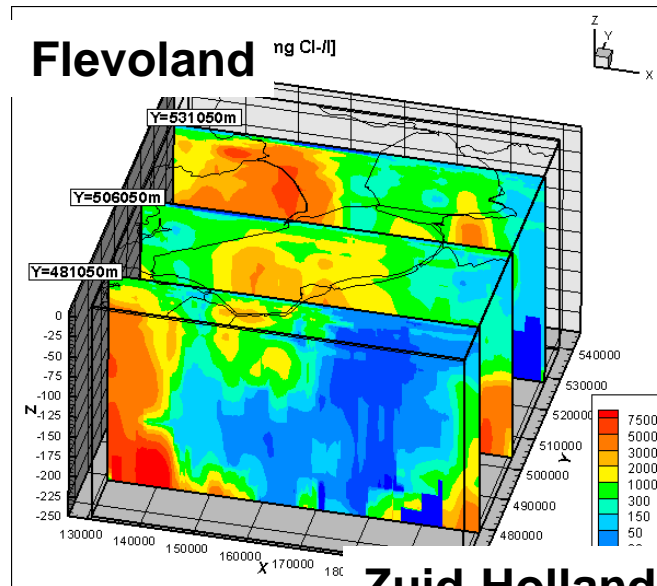
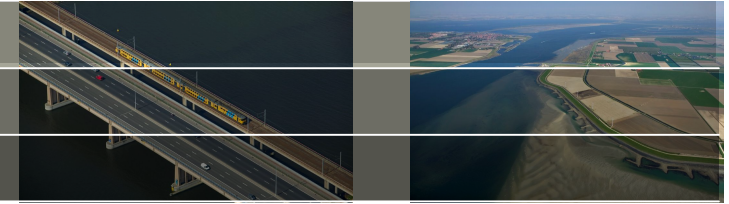


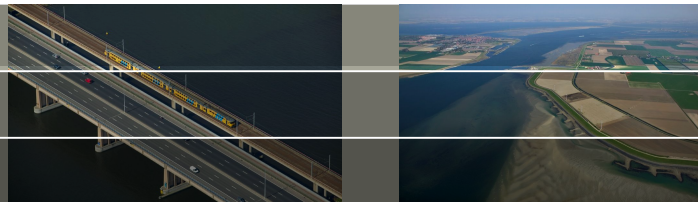
# Chloride concentraties in Zuid-Holland : ~6000 metingen

## Histogram



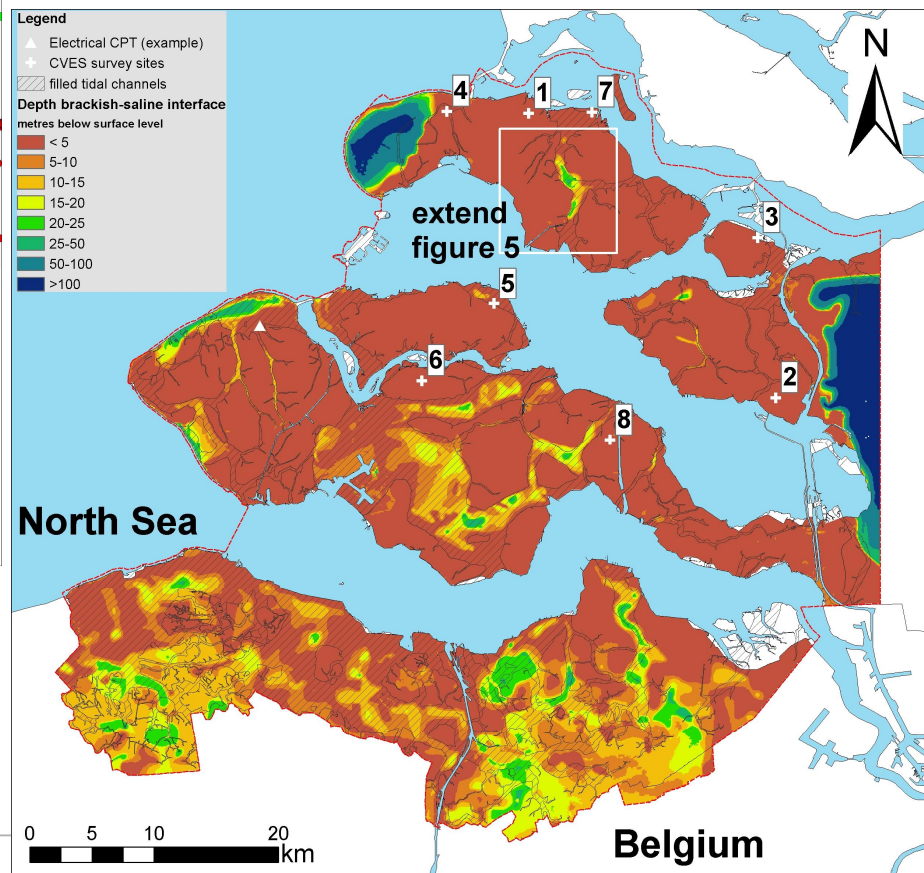
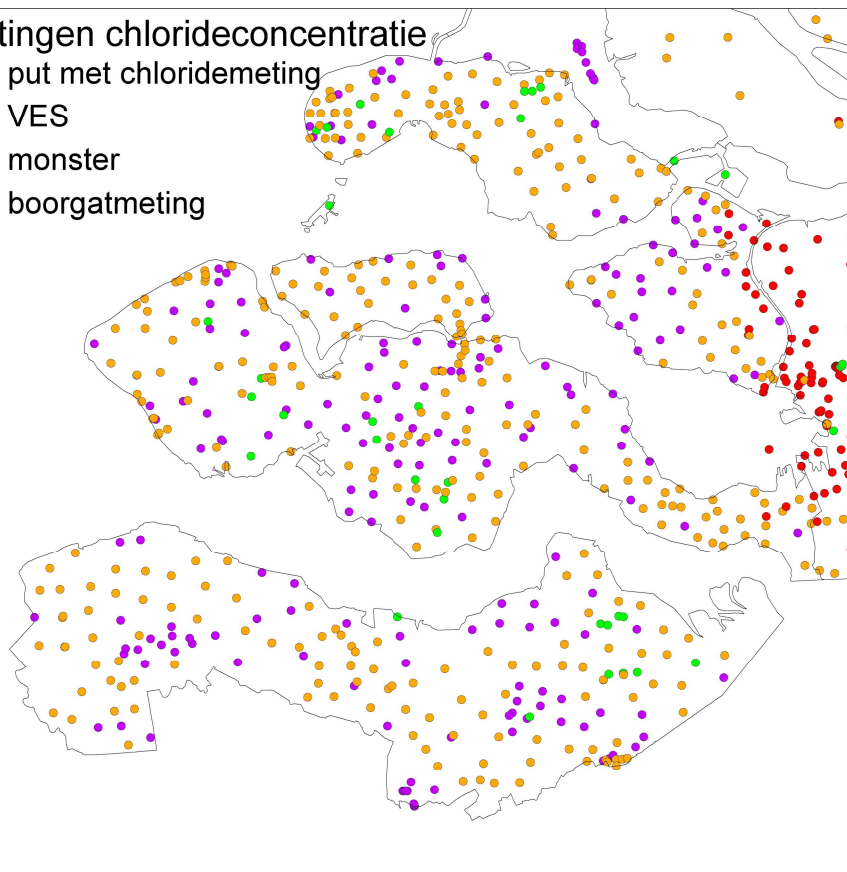
# Enkele regionale studies





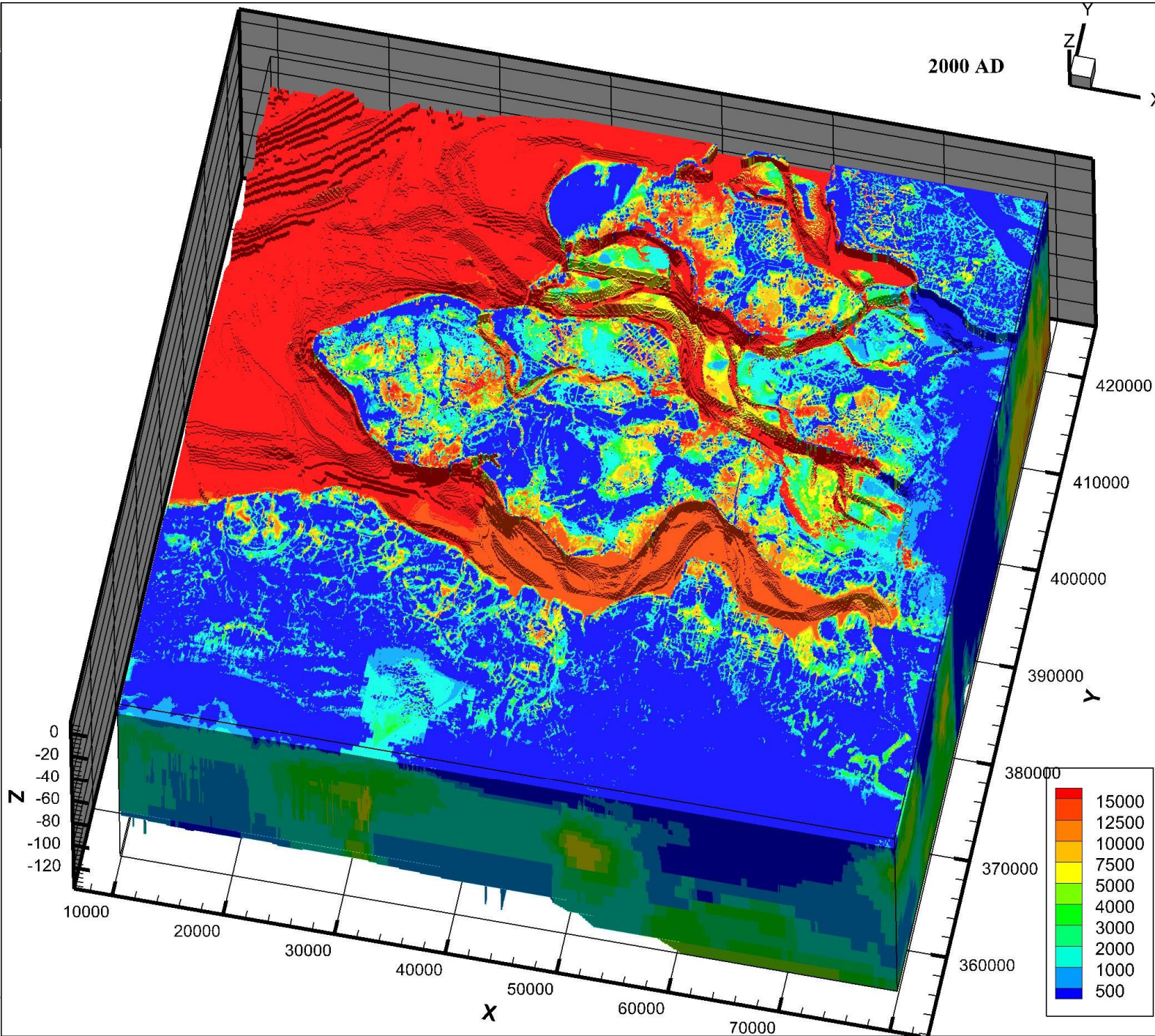
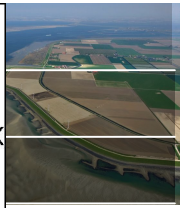
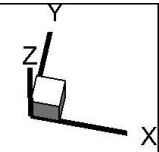
### metingen chlorideconcentratie

- put met chloridemeting
- VES
- monster
- boorgatmeting





2000 AD



res

# Modelling effect climate change on fresh-salt groundwater

## Modelling:

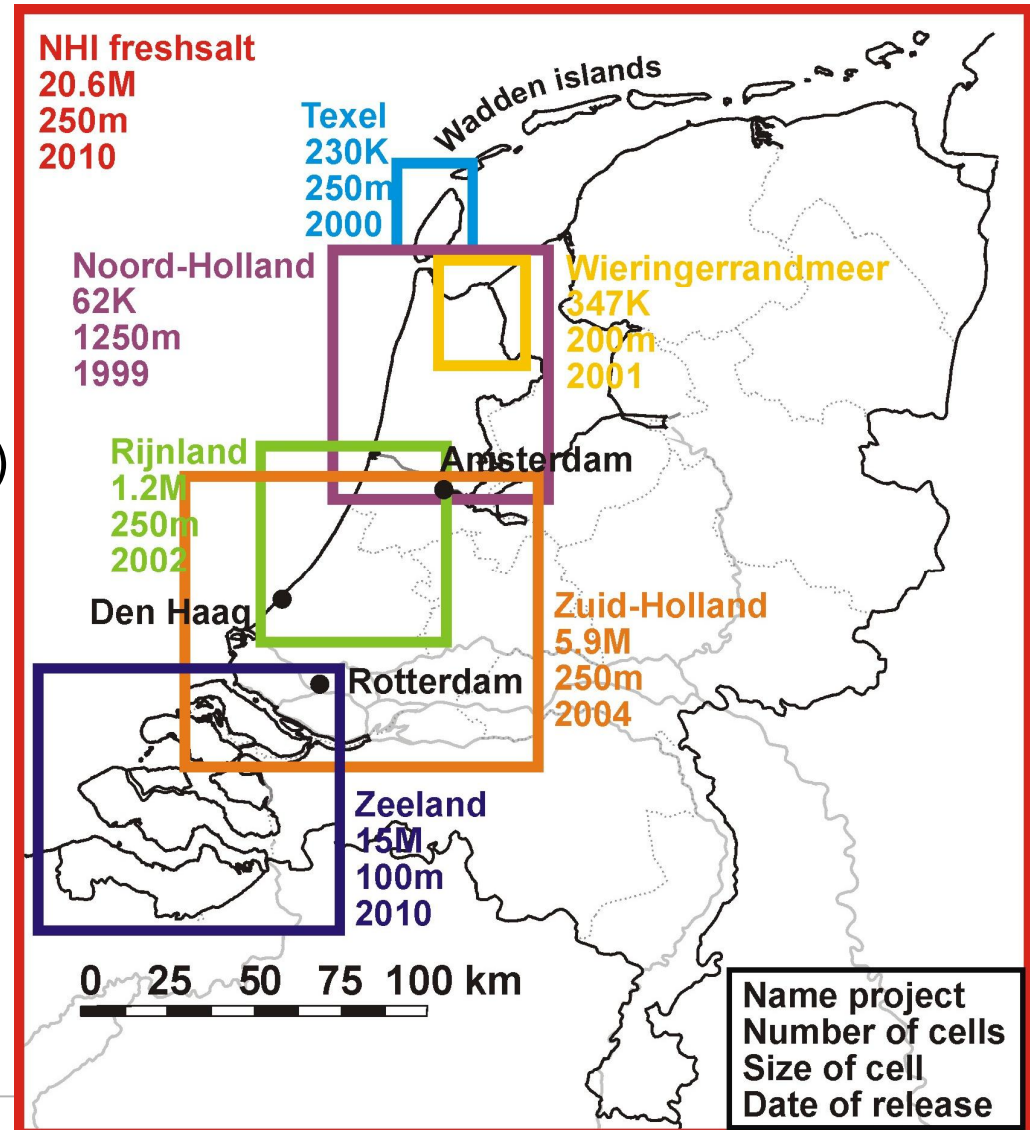
- variable-density
- 3D, non-steady
- groundwater flow
- coupled solute transport

## Code:

MOCDENS3D (MODFLOW family)  
similar to SEAWAT

## Assessing effects:

- autonomous salinisation
- sea level rise
- changing recharge pattern
- land subsidence
- changing extraction rates
- adaption measures



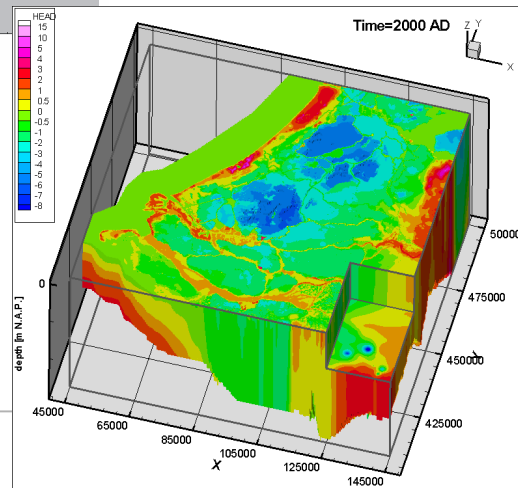
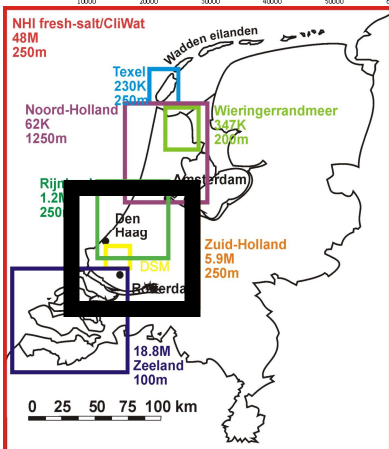
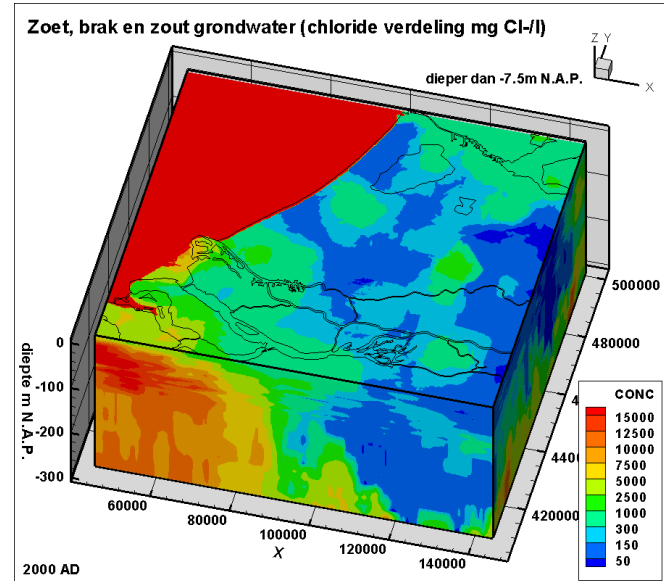
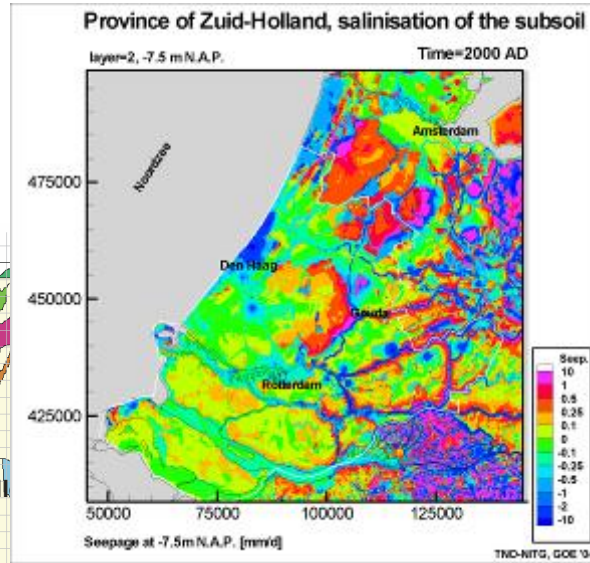
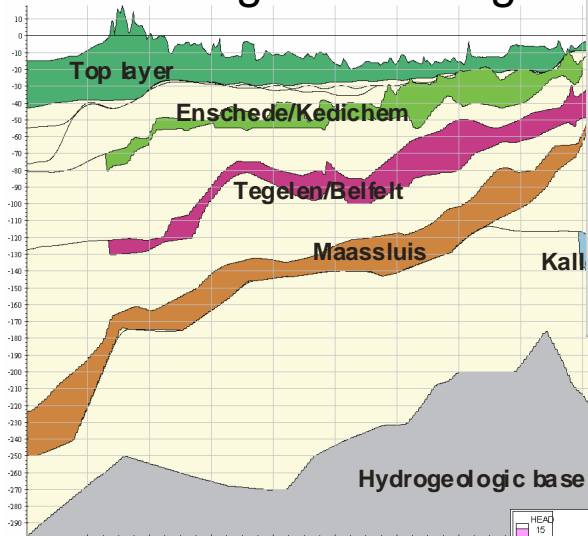




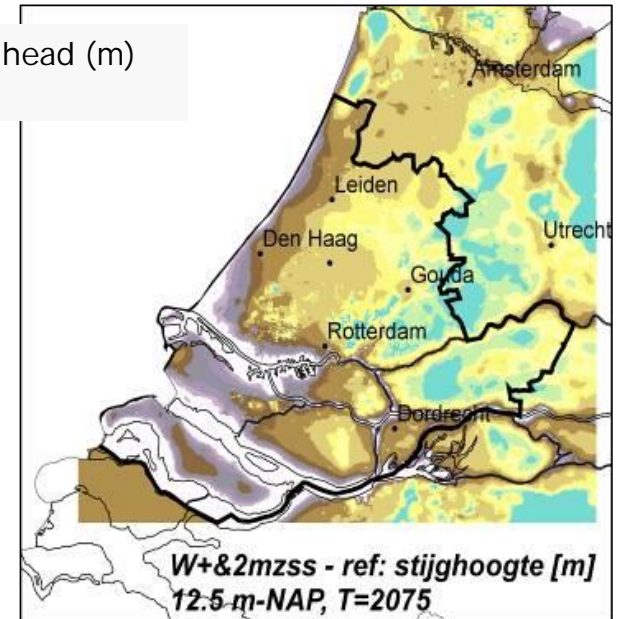
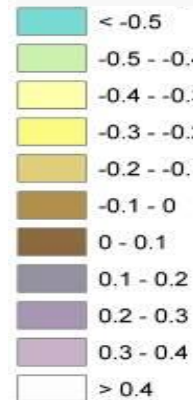
# Province Zuid-Holland (2004, update 2008, etc.)



- Land subsidence
- Sea level rise
- Change in recharge



difference in head (m)



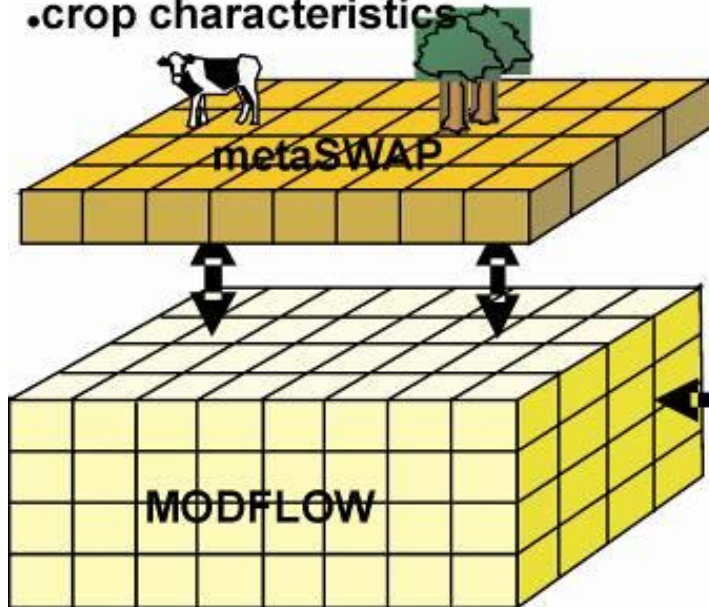


# Nationaal Hydrologisch Instrumentarium NHI

539486 UZ-units

data on:

- soil characteristics
- land use
- crop characteristics

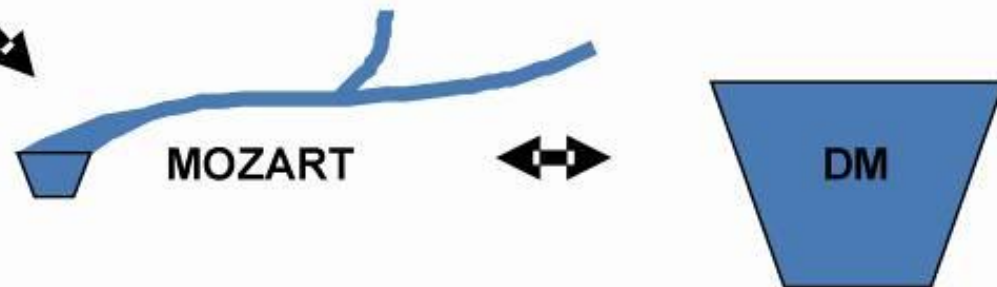


250 x 250 m,  
1300 x 1200 x 4 cells  
geohydrologic model REGIS

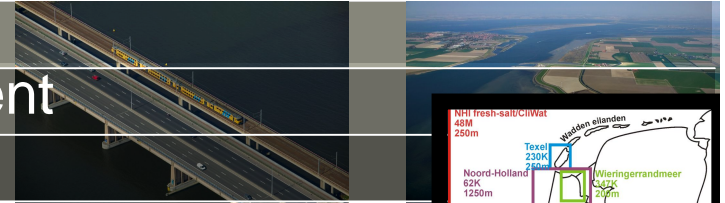
130 main surface water units,  
9000 local surface water units,  
60000 weir level units

data on:

- water level
- hydraulic capacities
- water management practices



# Netherlands Hydrological modelling Instrument for fresh-saline groundwater



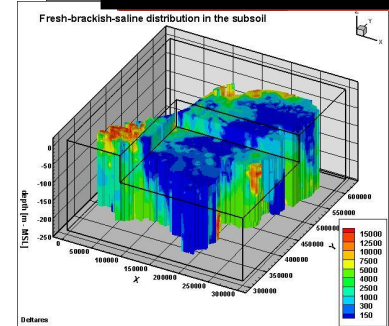
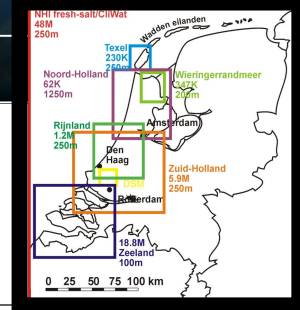
Using the national subsoil parametrisation

4 layers based on REGIS and NHI

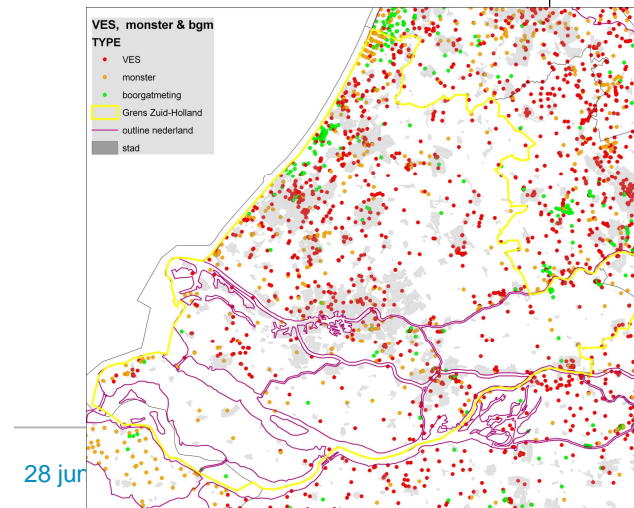
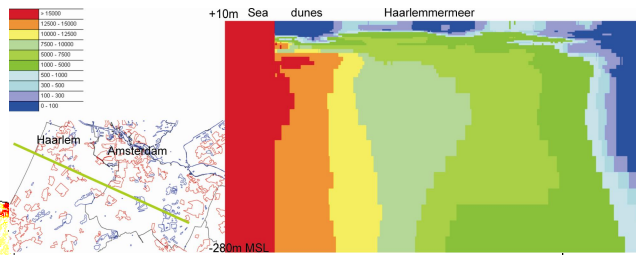
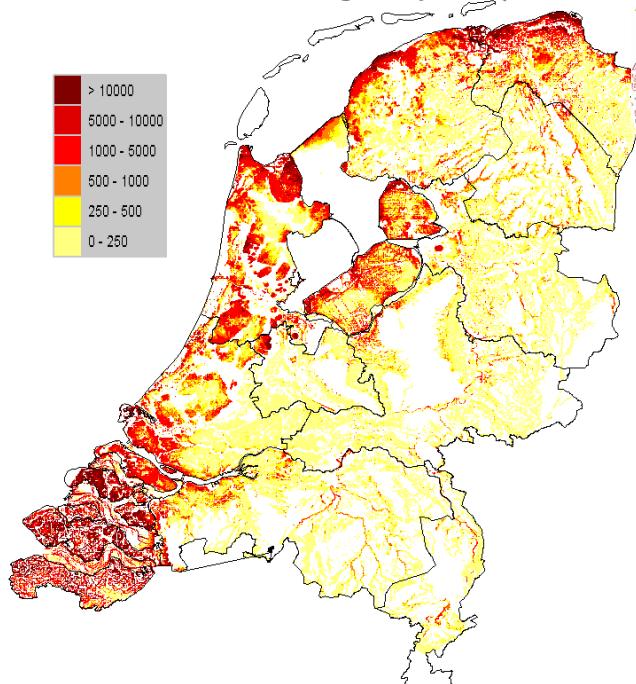
31 model layers with varying thicknesses: 2\*5m; 10\*2m; 8\*5m en 11\*20m  
cellsize 250x250m

Using the national 3D salt concentration in groundwater

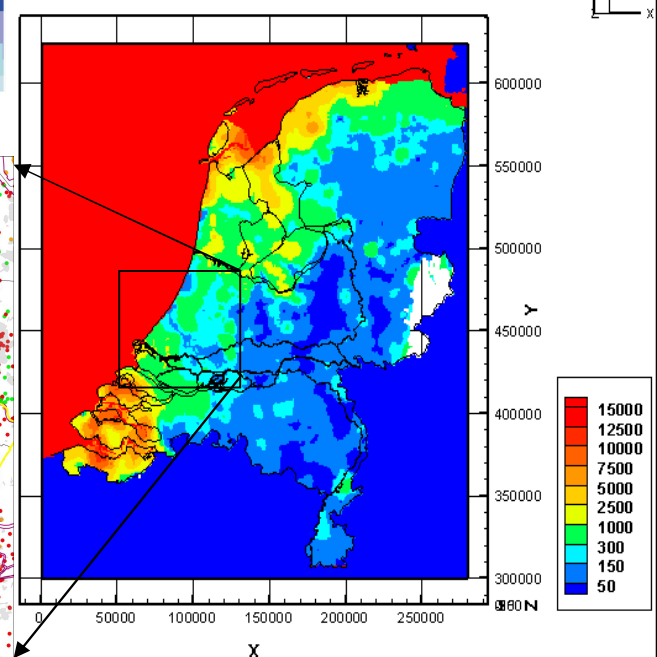
Fresh-saline REGIS: ~65000 measuring points (analyses, VES, Borehole)



Result: Salt load [kg/ha/yr]



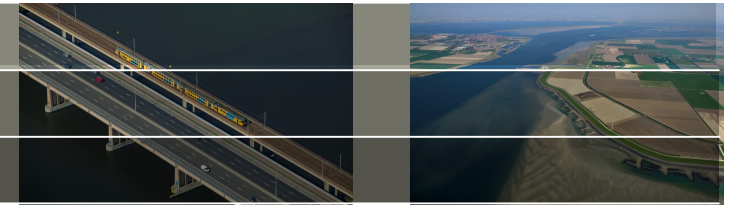
concentratie op -10 m N.A.P.



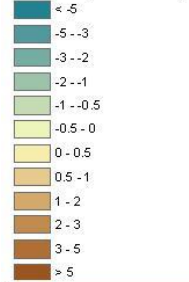
28 jun



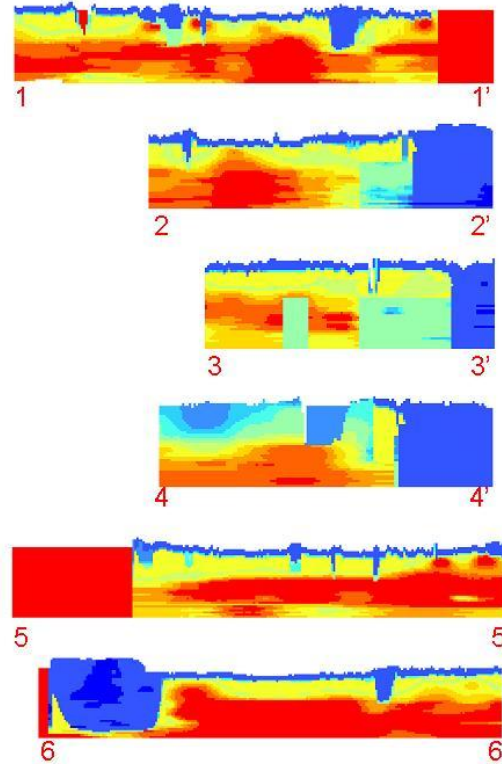
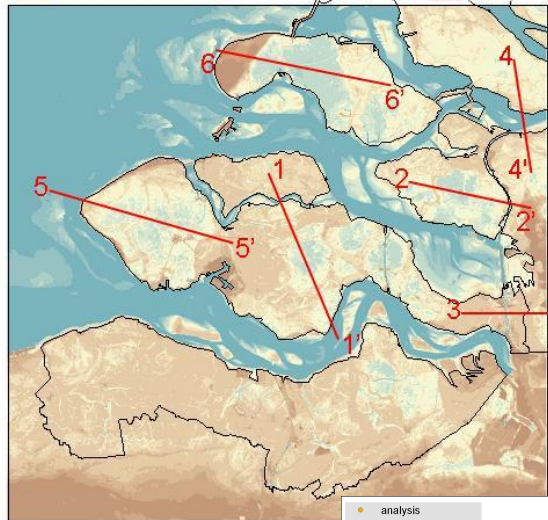
# 3D model Zeeland



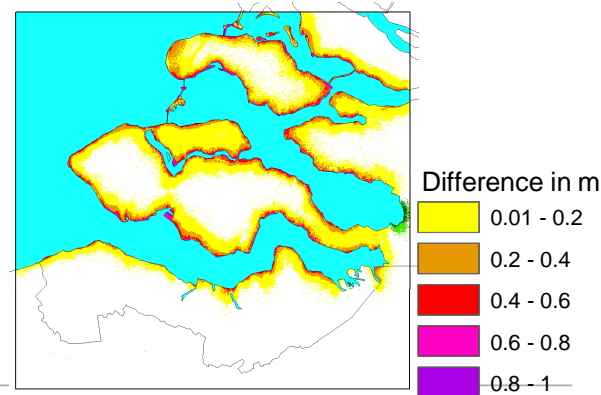
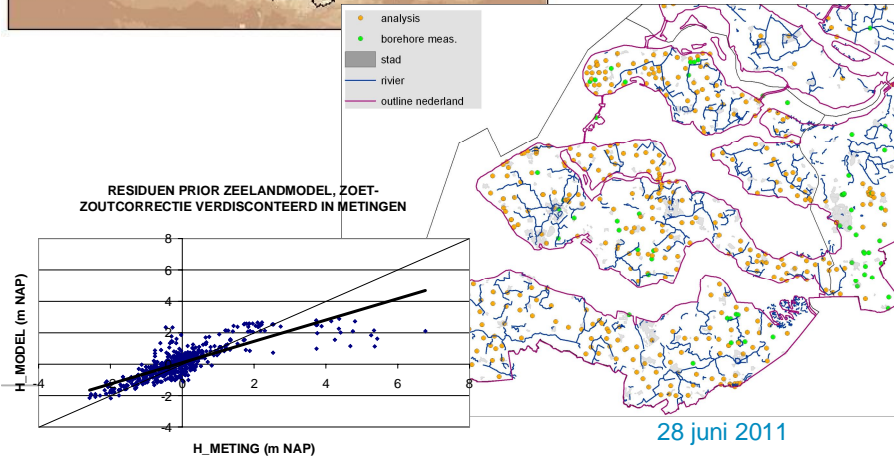
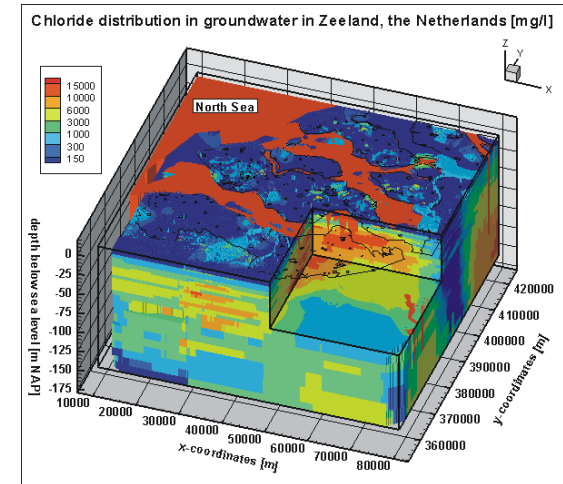
maaiveld (m tov NAP)



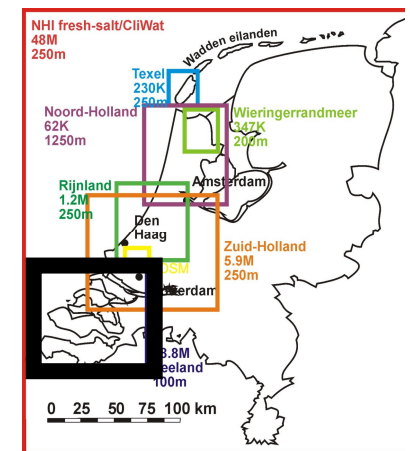
mg Cl/l



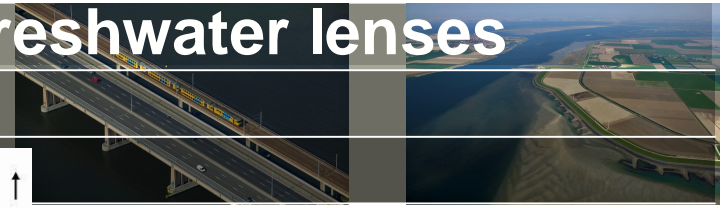
- Cels 100\*100 m<sup>2</sup>
- Top modellayers 0.5 m thickness
- Calibrated
- Improved 3D initial fresh saline field



Freshwater head sea level +1m

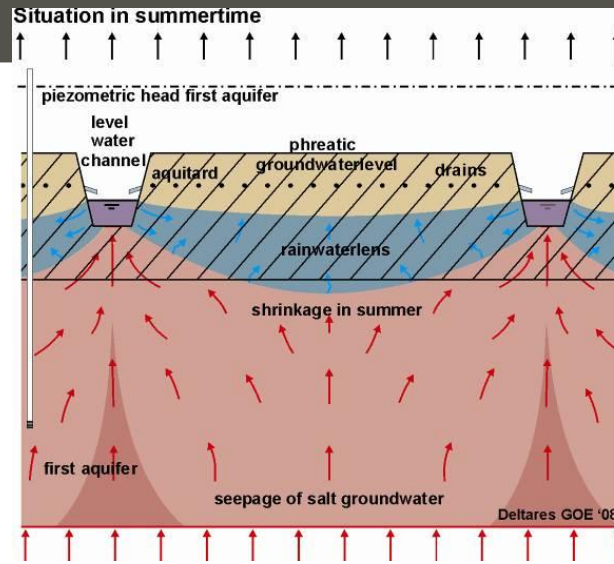


# Assessing climate change effect on freshwater lenses



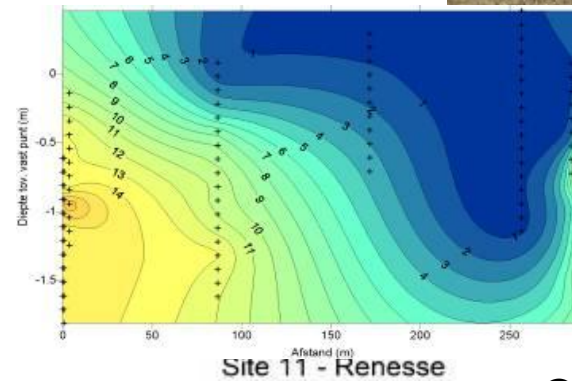
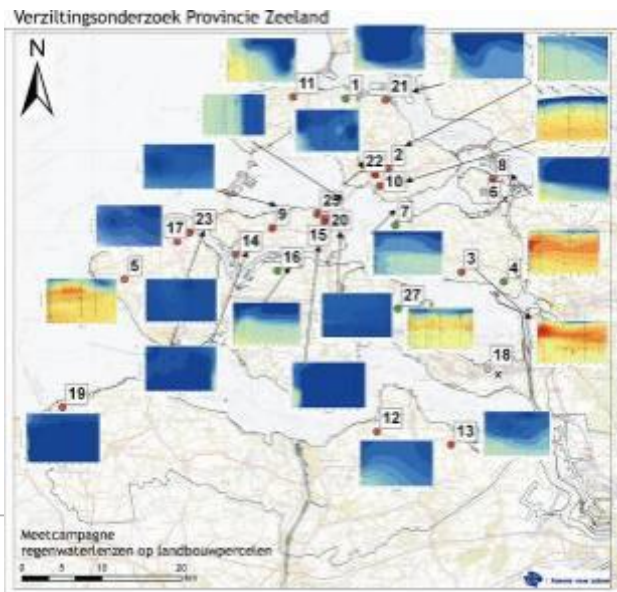
## Monitoring:

- Geology
- TEC-probe
- Sampling (head & conc)
- EM31
- CVES
- HEM

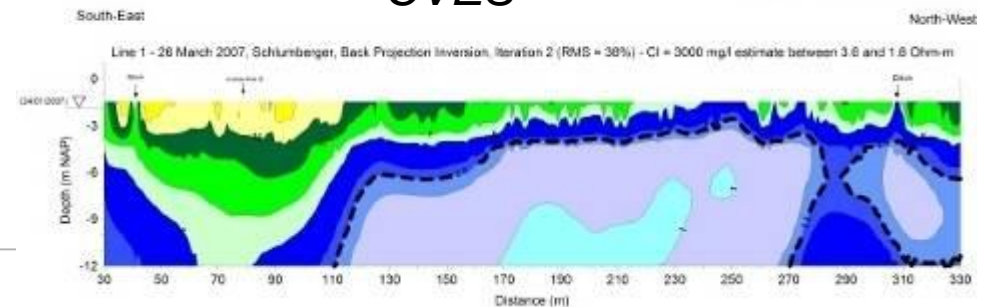


TEC probe

## Groundwater modeling

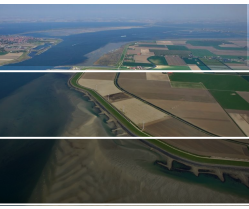
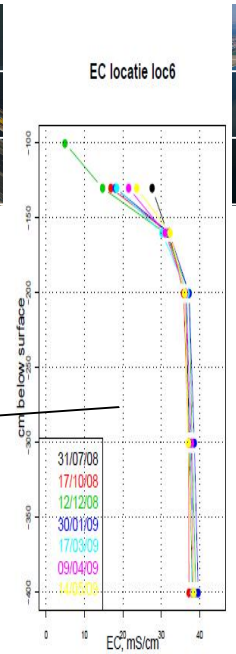
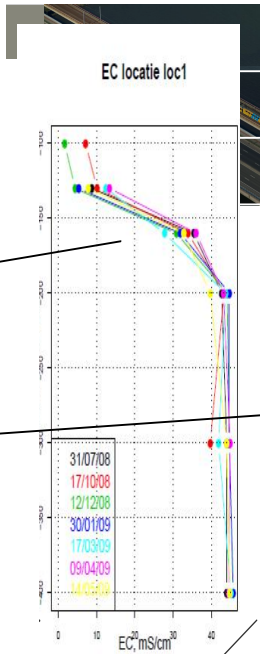
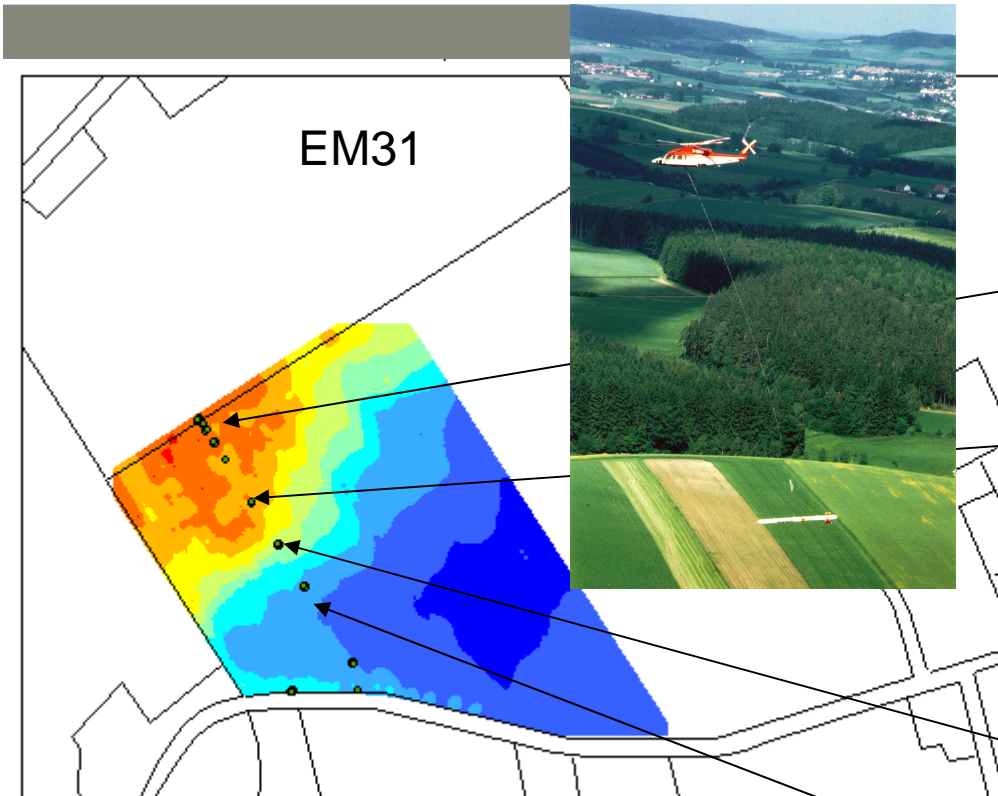


CVES

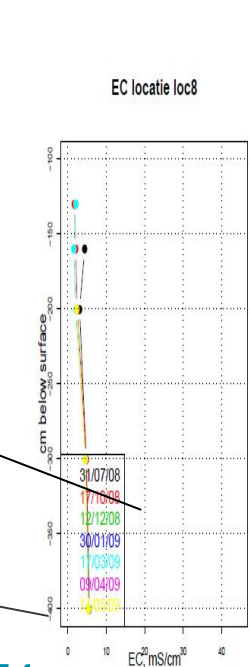
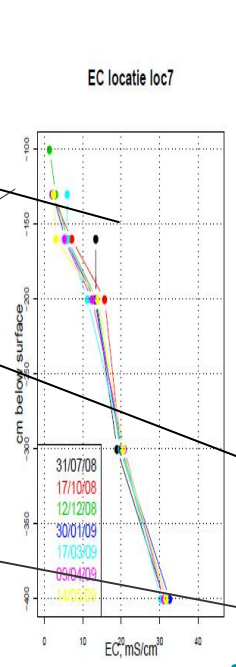
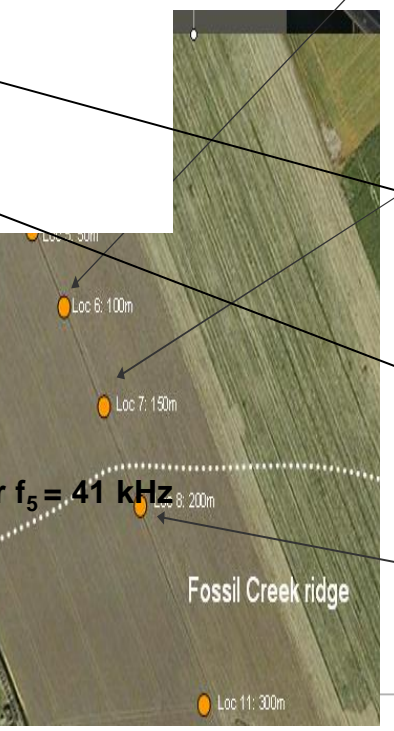
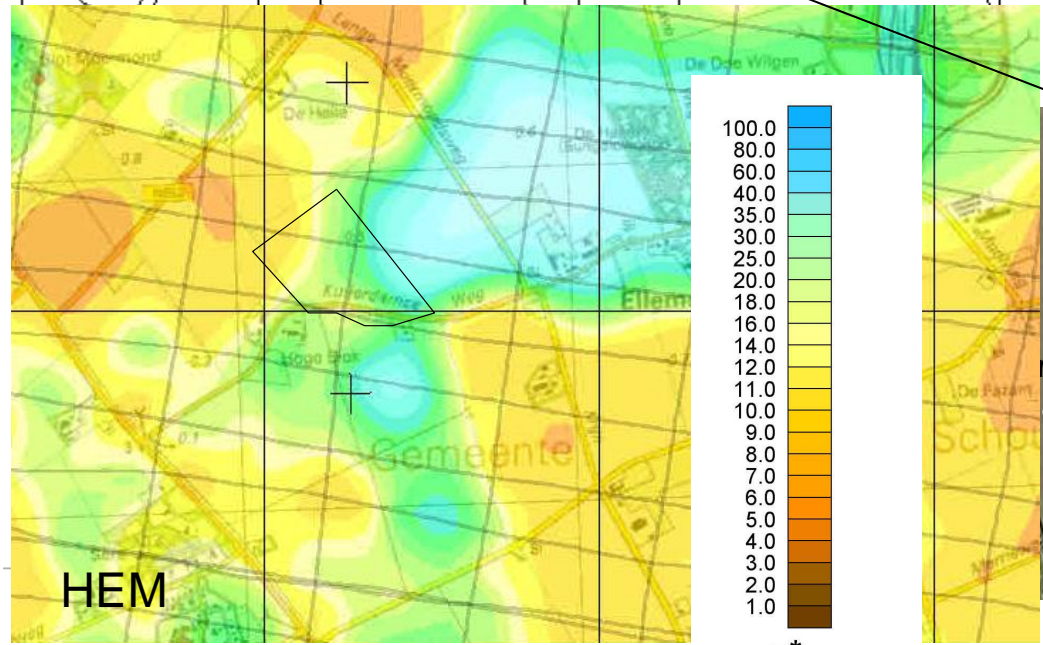


28 juni 2011





Sampling





---

28 juni 2011

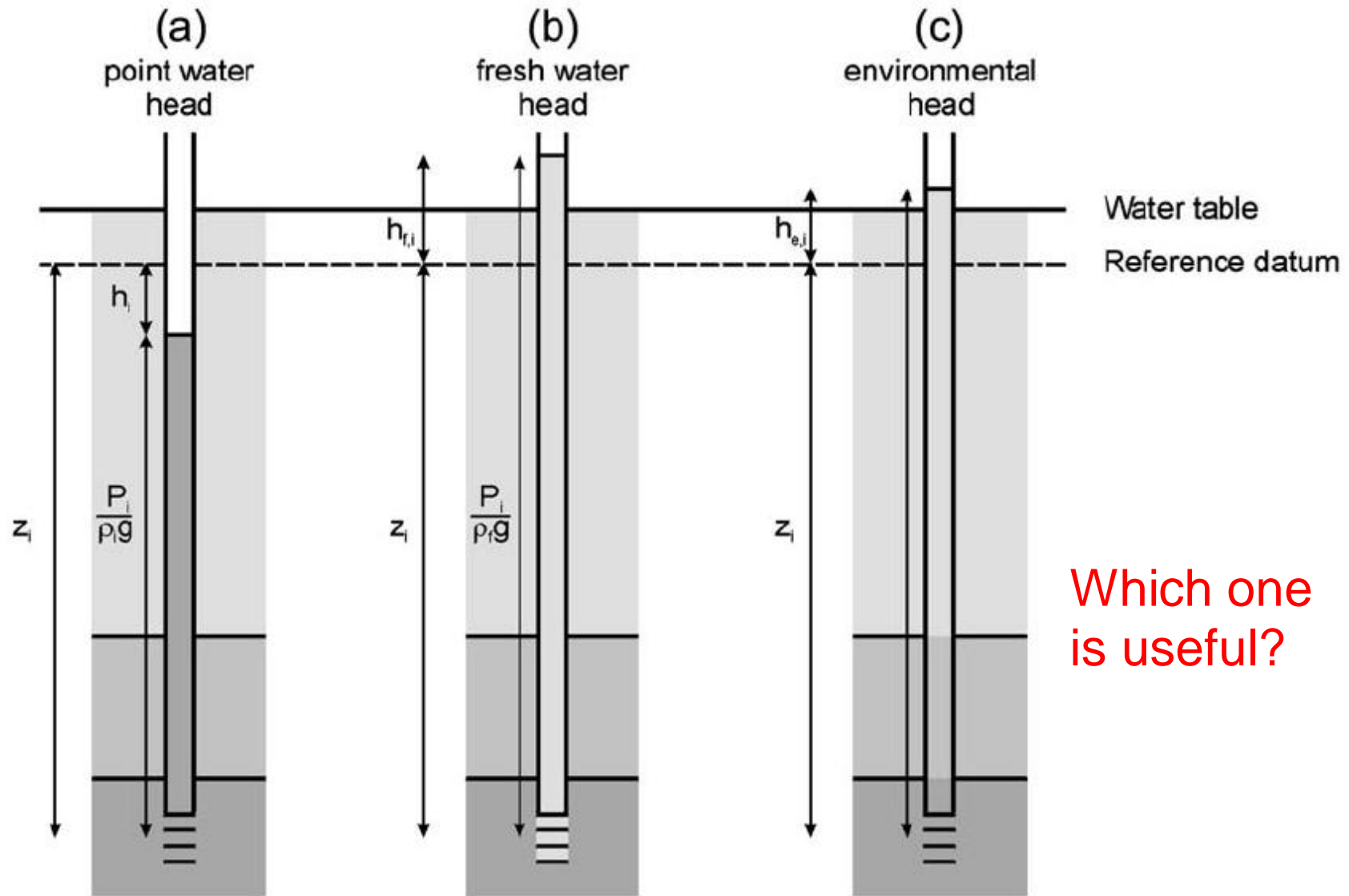
**Deltares**



# Zoetwaterstijghoogte

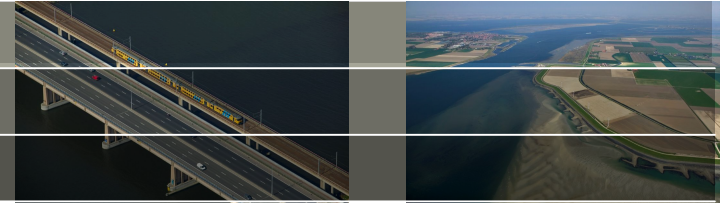


# There are three common ways to define head



Which one is useful?

# Zoetwaterstijghoogte $\phi_f$



$$\phi_f = \frac{p}{\rho_f g} + z$$

- Vergelijken grondwater met verschillende dichtheden
- Fictieve parameter
- In plaats van drukken
- Verwarrend (zoetwaterstijghoogte staat loodrecht op stroomlijnen)

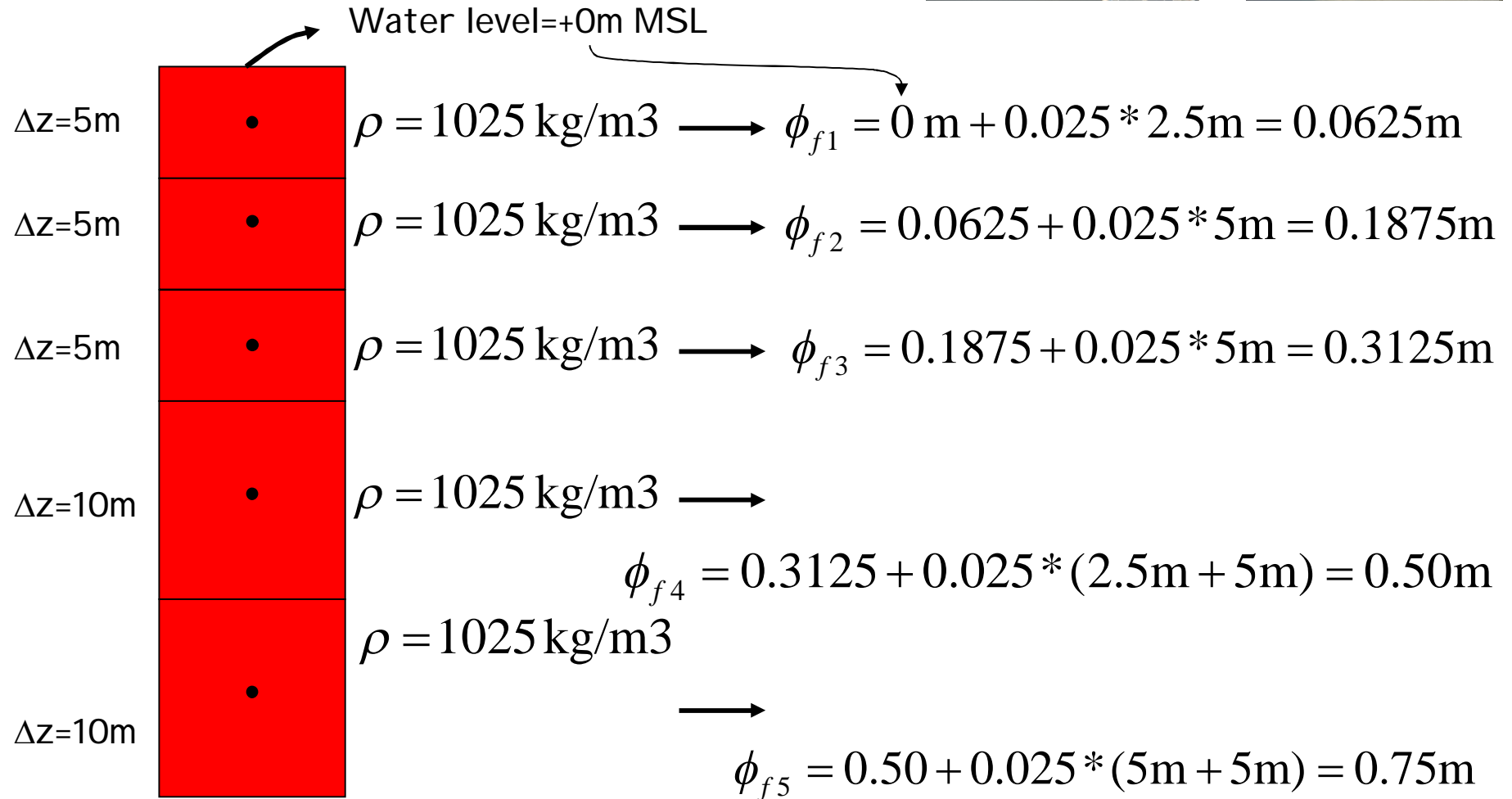
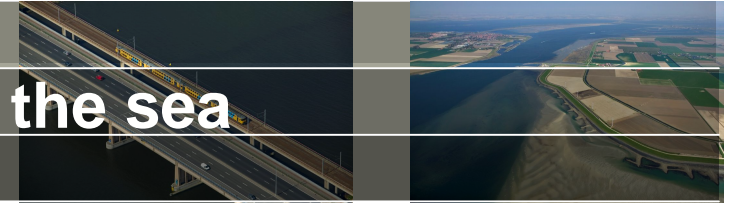


---

28 juni 2011

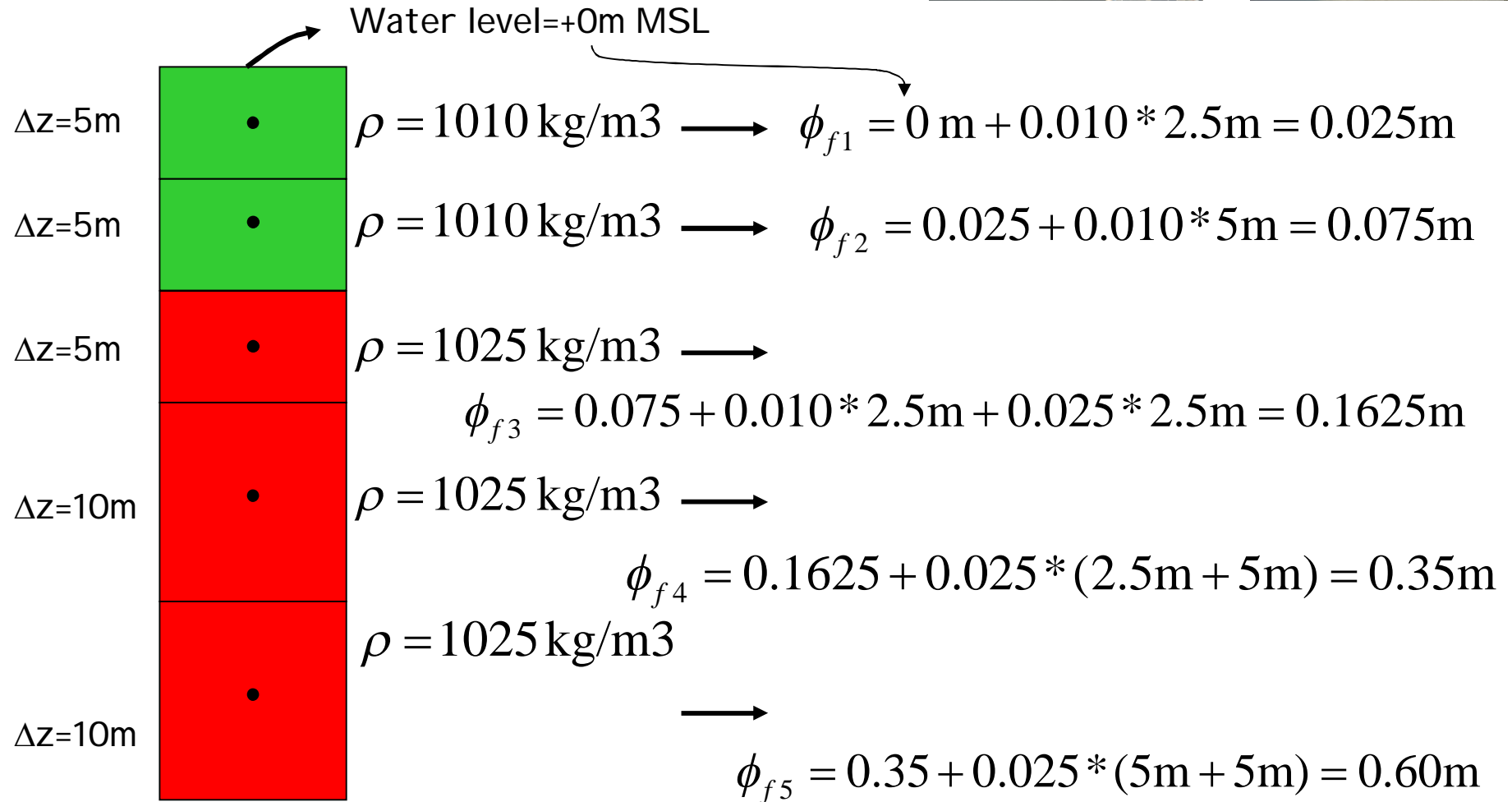
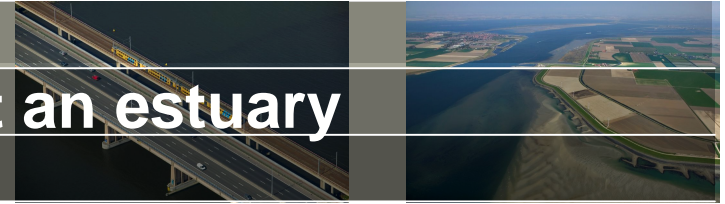
**Deltares**

# Hydrostatic boundary condition at the sea



$$\phi_{f2} = \phi_{f1} + \frac{\rho - \rho_f}{\rho_f} (\Delta z)$$

# Hydrostatic boundary condition at an estuary



$$\phi_{f2} = \phi_{f1} + \frac{\rho - \rho_f}{\rho_f} (\Delta z)$$



$\Delta z = 10\text{m}$

$\rho = 1000 \text{ kg/m}^3$

$p_1$  5m

$\Delta z = 10\text{m}$

$\rho = 1000 \text{ kg/m}^3$

$p_2$  10m

$\Delta z = 10\text{m}$

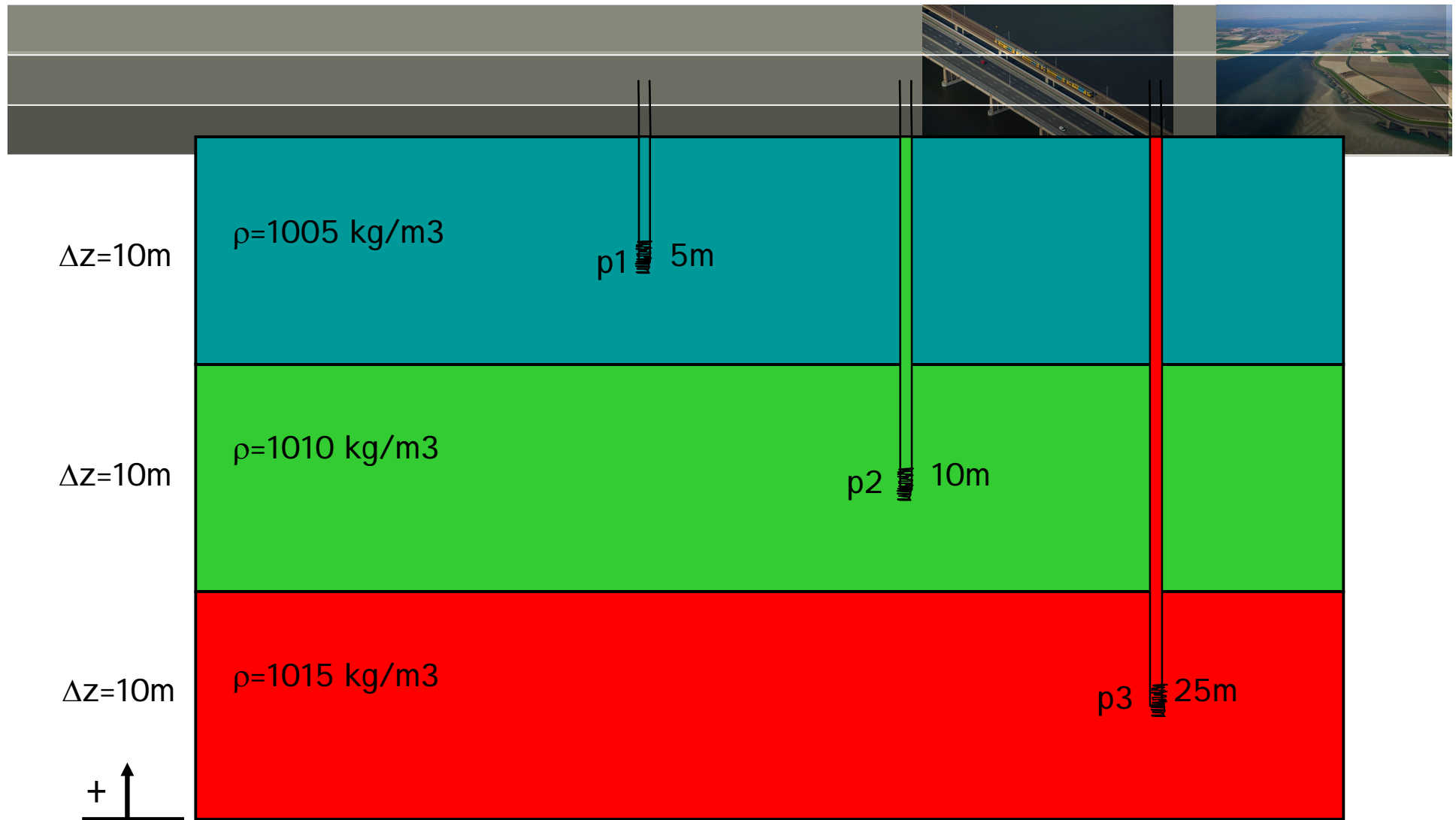
$\rho = 1000 \text{ kg/m}^3$

$p_3$  25m



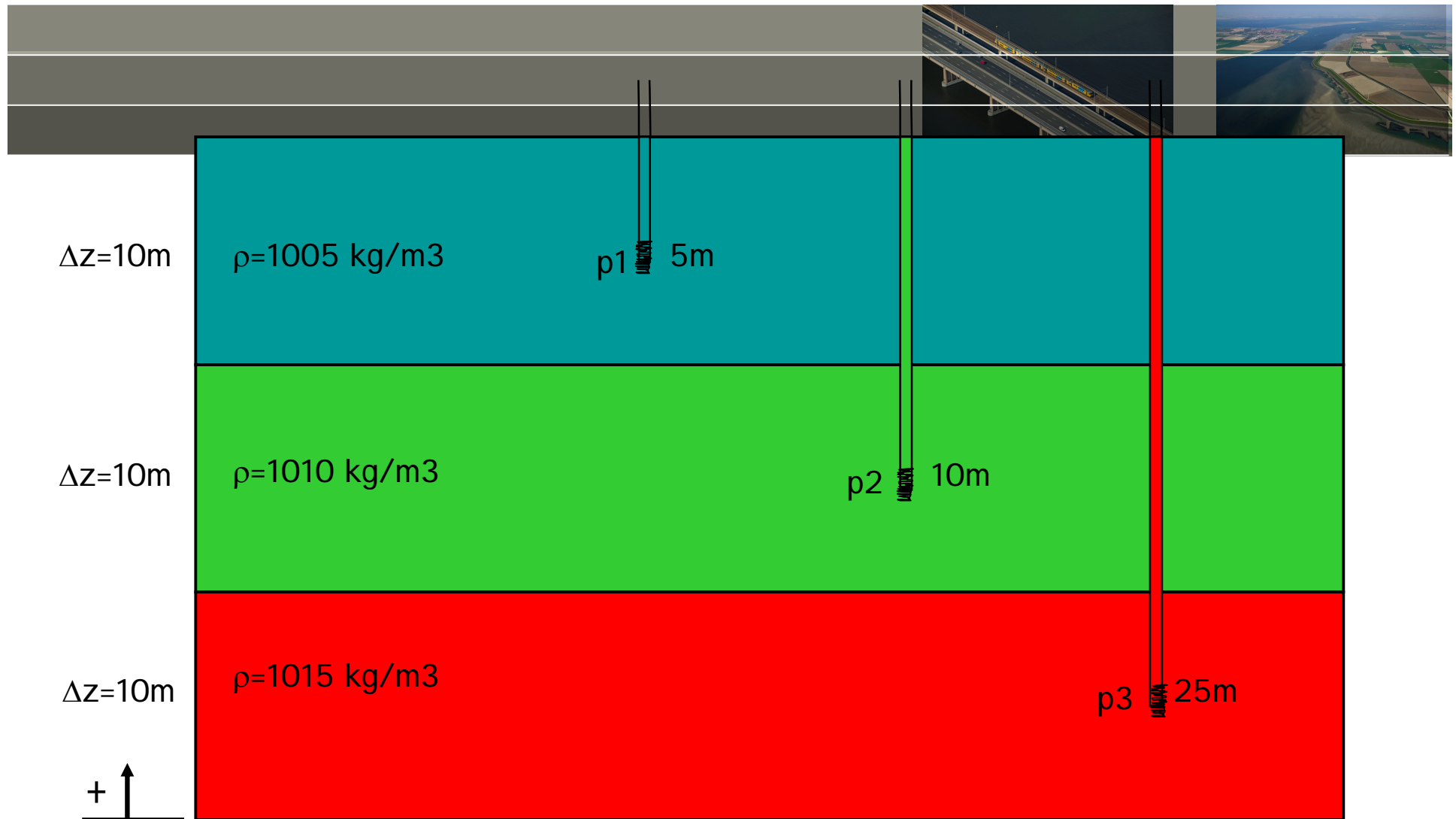
No flow





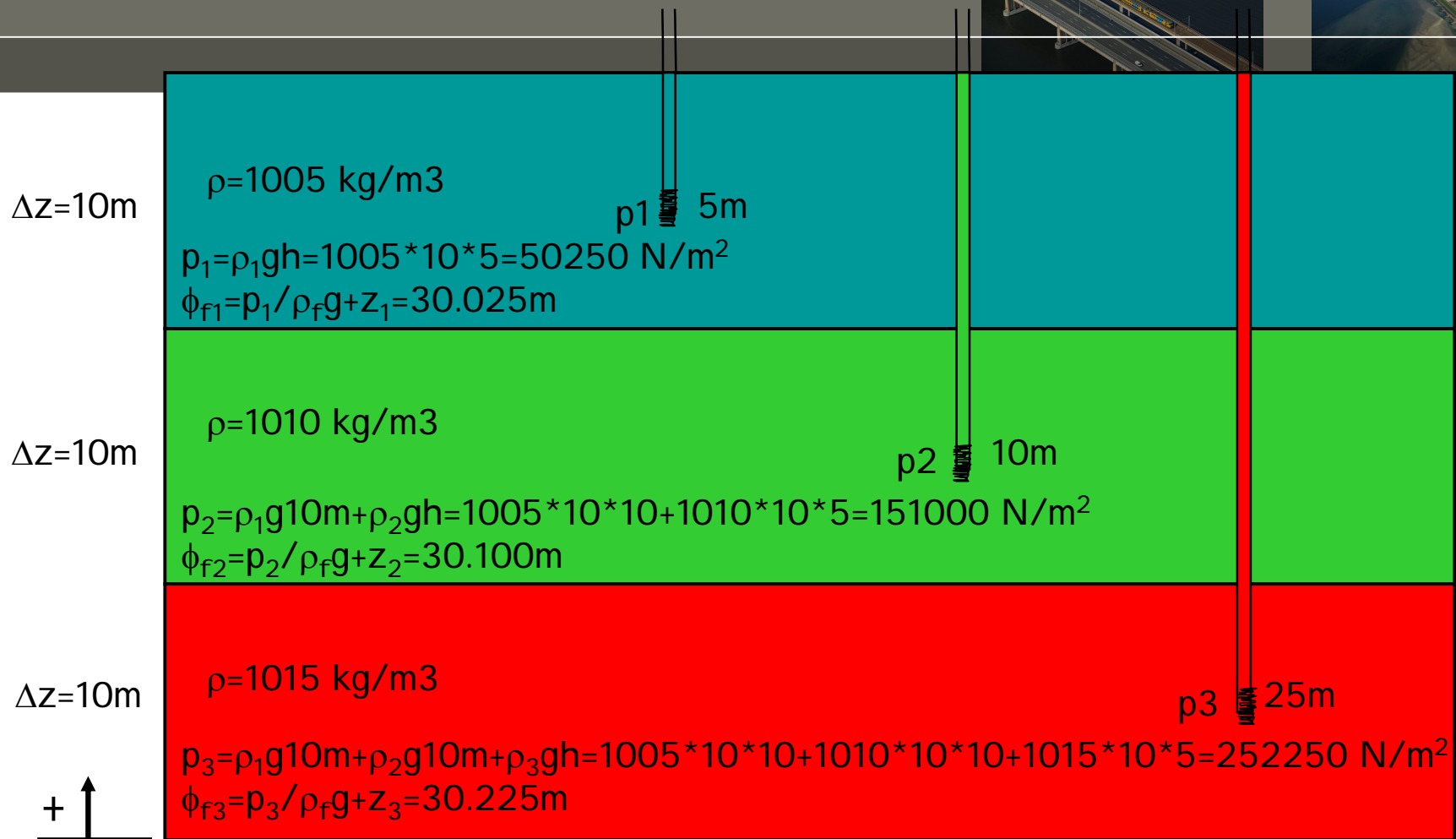
Flow or no flow? (if  $p \neq$  hydrostatic than flow)

Calculate to freshwater head!



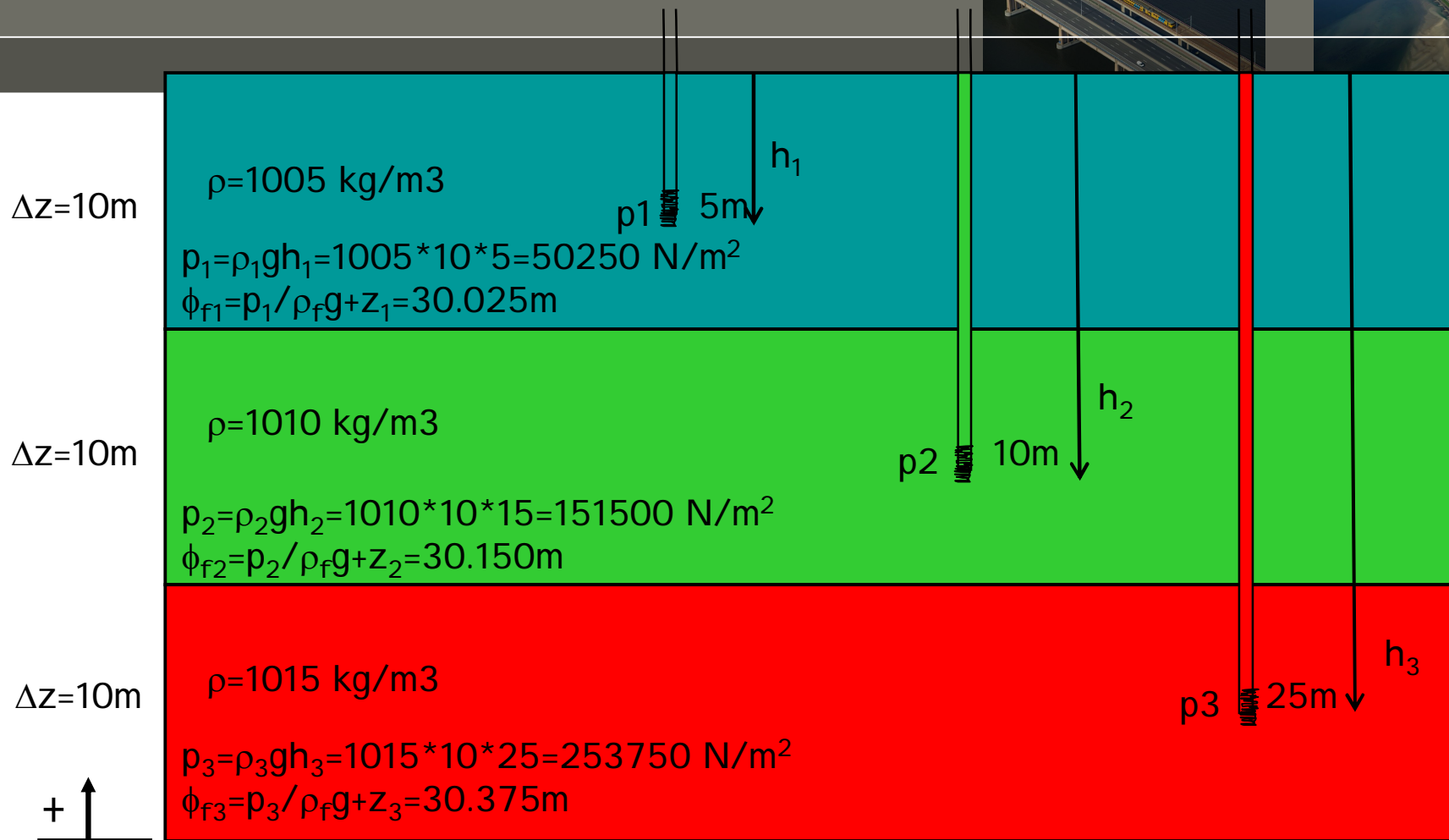
1. Determine hydrostatic pressure (and freshwater head  $f_f$ )
2. Determine pressure  $p$  in well! (and freshwater head  $f_f$ )

# 1. Determine hydrostatic pressure and fryhead



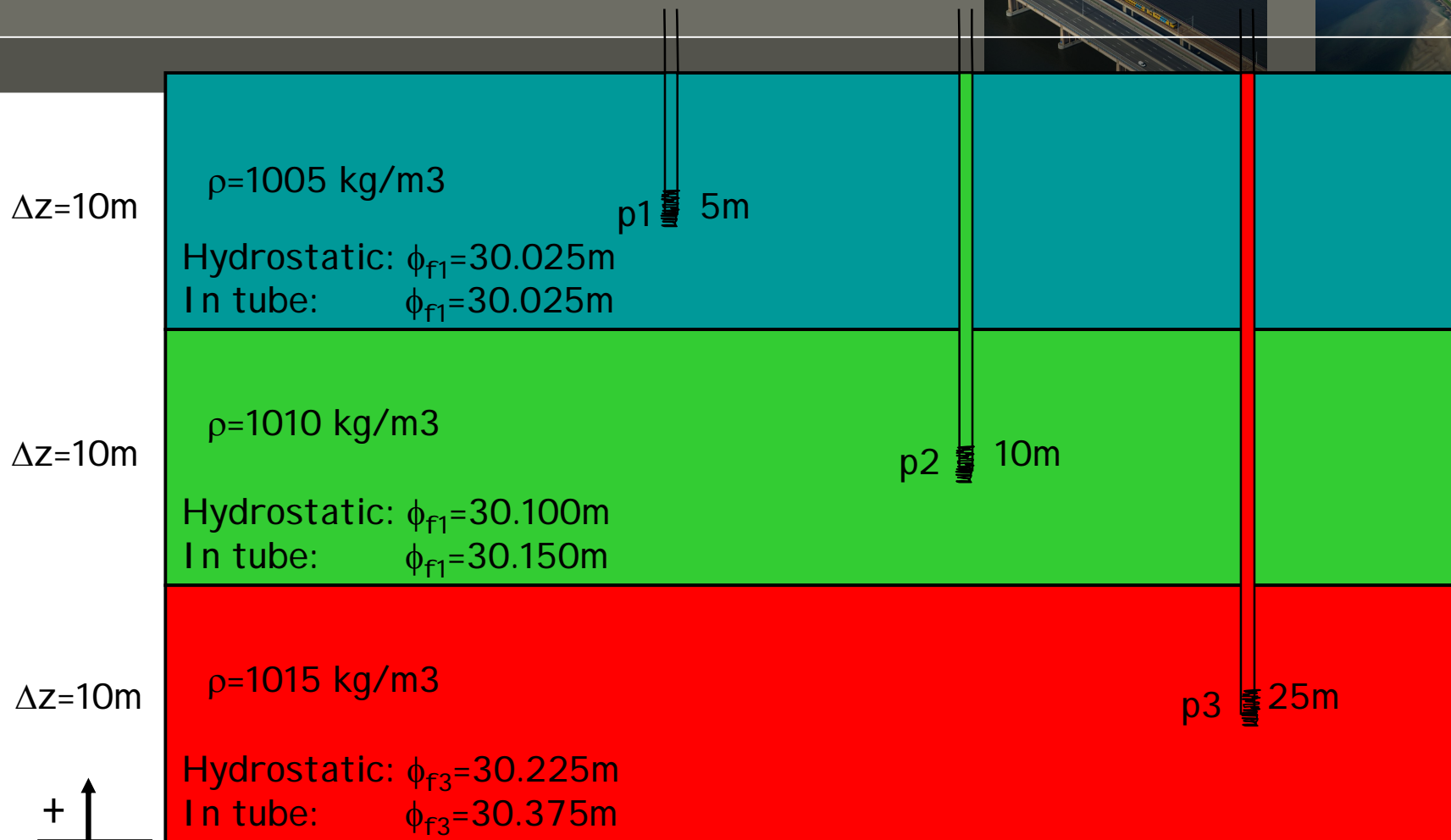
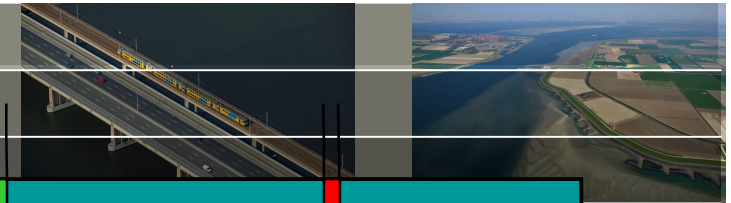
1. Determine hydrostatic pressure (and freshwater head  $f_f$ )
2. Determine pressure  $p$  in well! (and freshwater head  $f_f$ )

## 2. Determine pressure p in well and frwhead



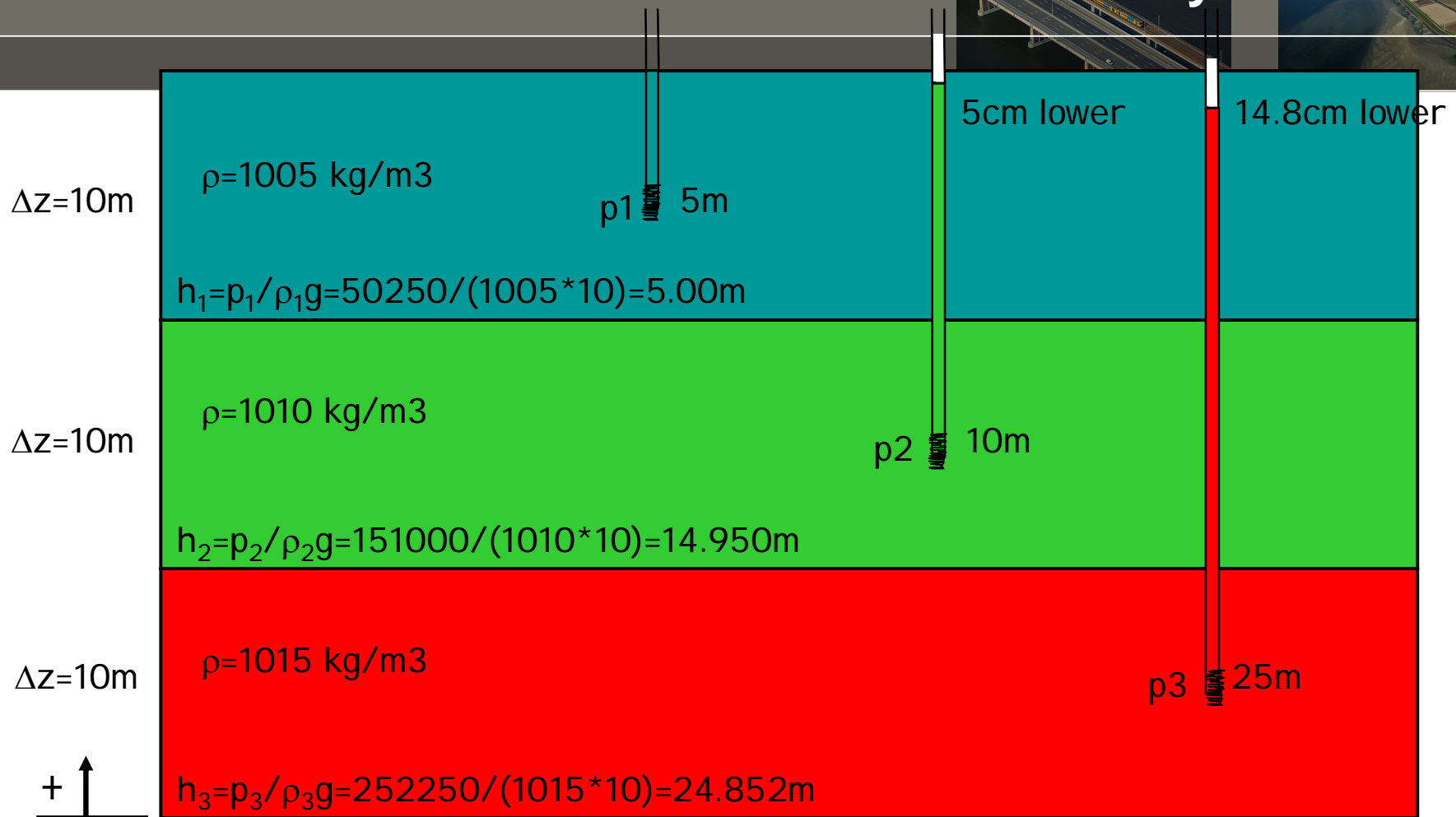
1. Determine hydrostatic pressure (and freshwater head  $f_f$ )
2. Determine pressure p in well! (and freshwater head  $f_f$ )

# Comparison



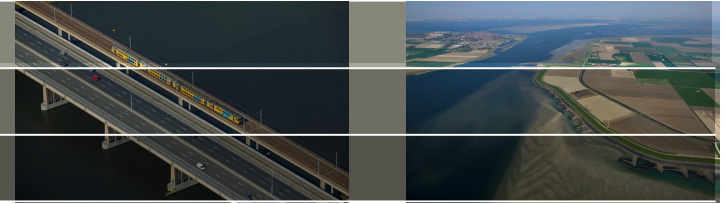
Conclusion: freshwater head not equal, so vertical upward flow!

# What would be the water level in the tube if hydrostatic?





# Hydrostatic pressure: $q_z=0$



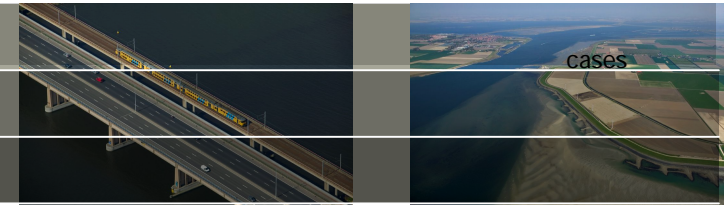
$$q_z = -\frac{\kappa_z \rho_f g}{\mu} \left( \frac{\partial \phi_f}{\partial z} + \frac{\rho - \rho_f}{\rho_f} \right)$$

$$0 = \left( \frac{\partial \phi_f}{\partial z} + \frac{\rho - \rho_f}{\rho_f} \right)$$

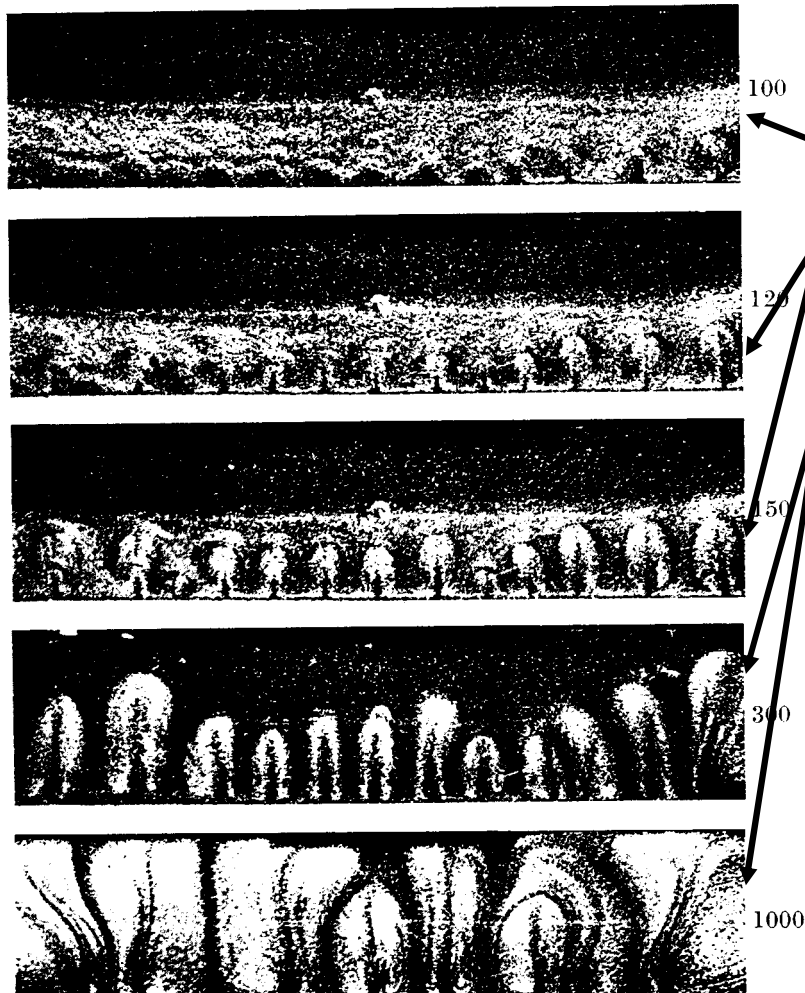
$$\partial \phi_f = -\frac{\rho - \rho_f}{\rho_f} \partial z$$

$$\phi_{f2} = \phi_{f1} - \frac{\rho - \rho_f}{\rho_f} (z_2 - z_1)$$

↓ +  $\phi_{f2} = \phi_{f1} + \frac{\rho - \rho_f}{\rho_f} (\Delta z)$



It is originally a heat transport problem

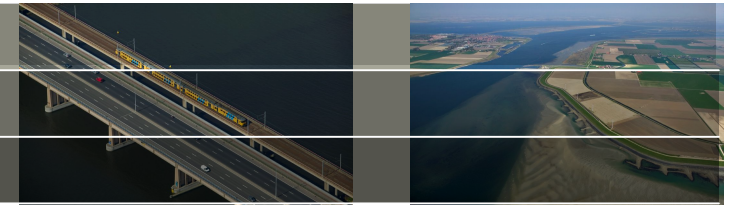


Phases:

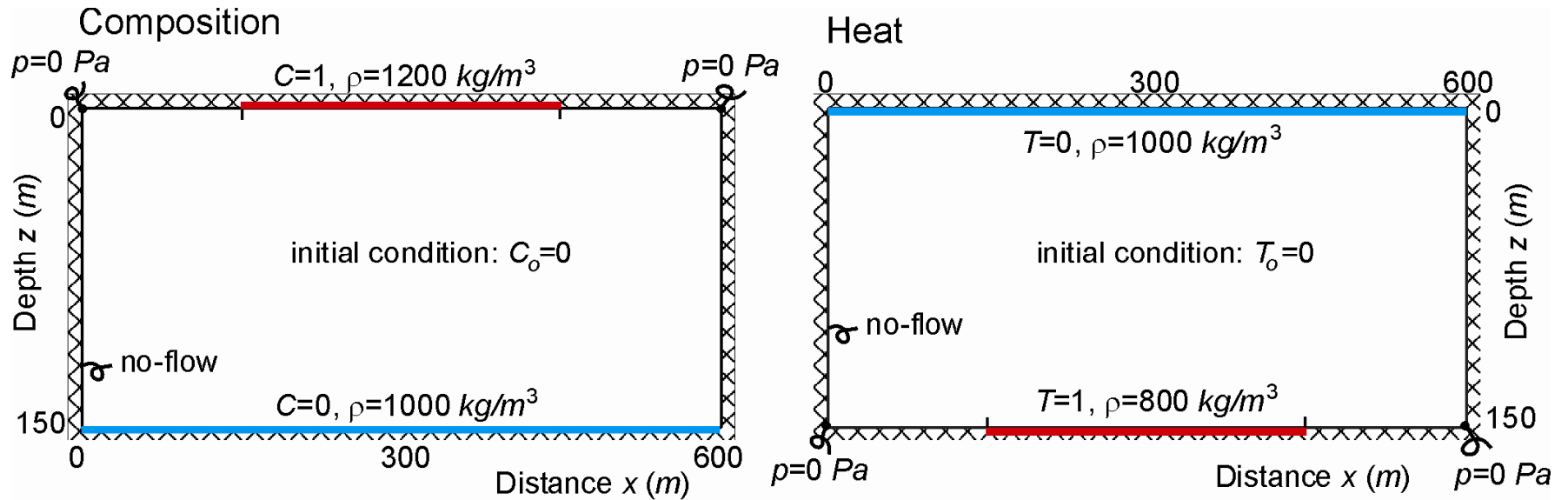
1. Stable growth diffusive boundary layer
2. Development flow cells embedded in boundary layer
3. Emergence of disturbances that grow into fingers

Convection of heat occurs when:

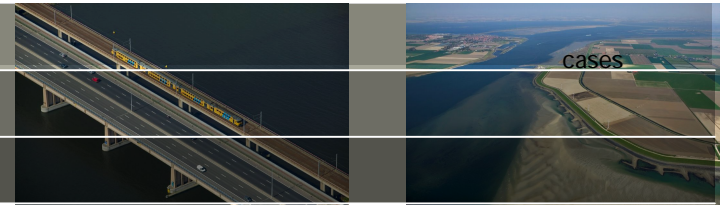
$$\text{Rayleigh number} > 4\pi^2$$



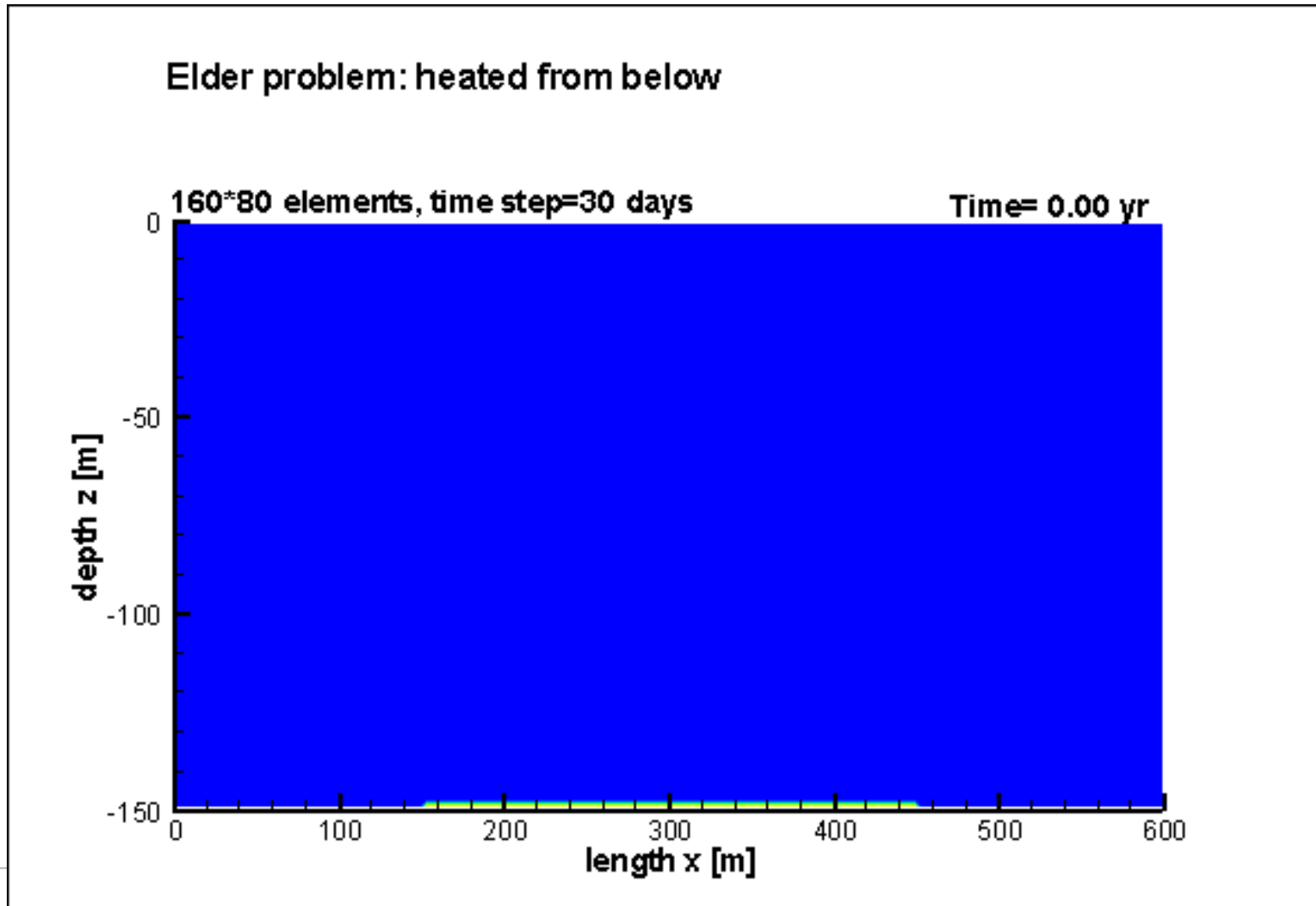
Analogy composition and heat



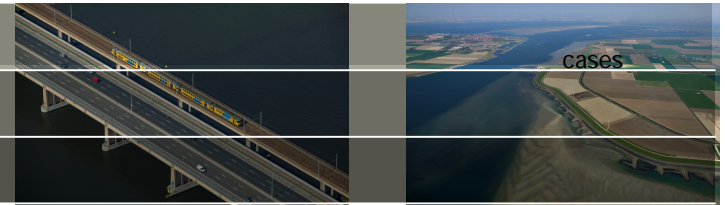
Elder problem (III)



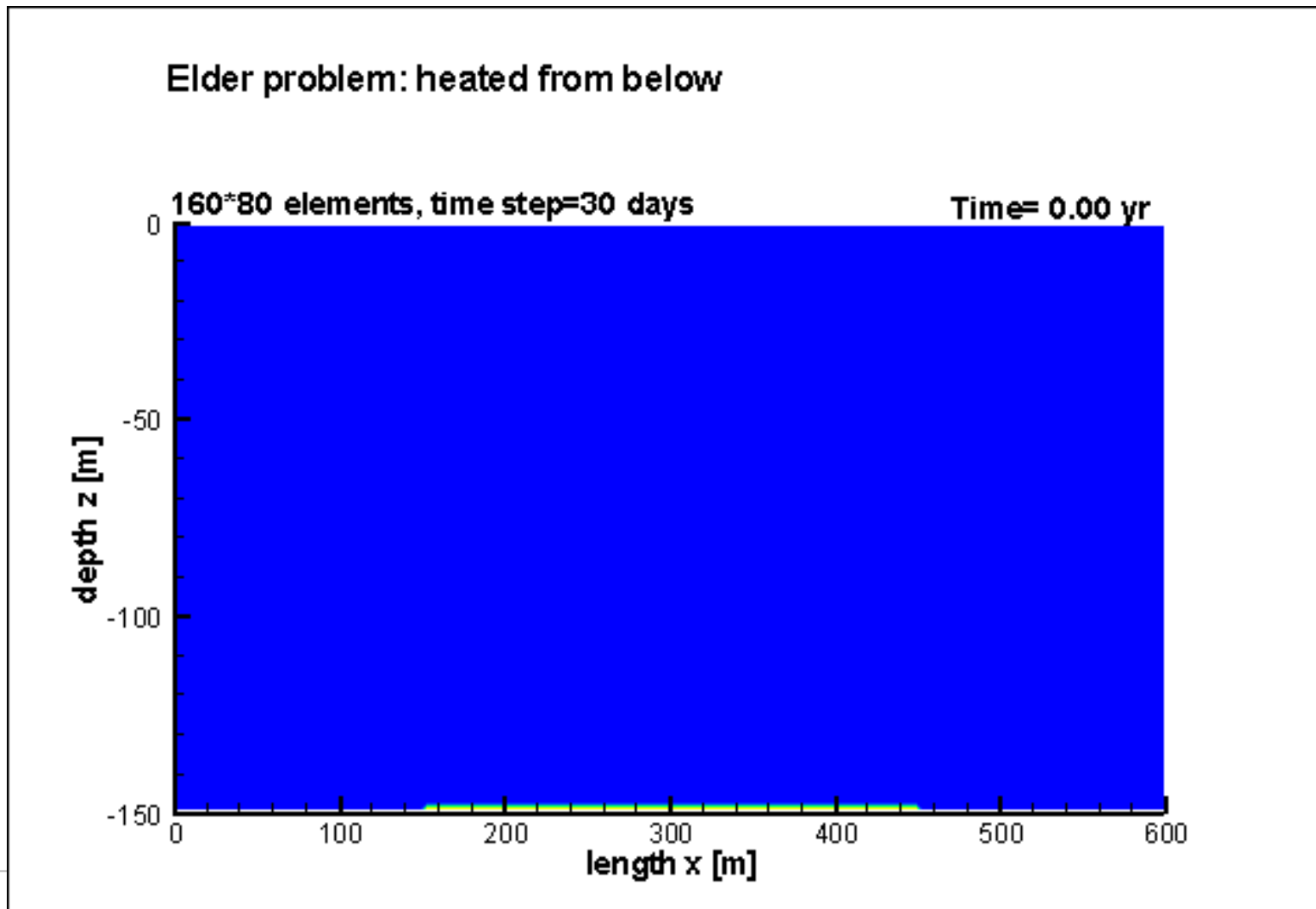
Development of convection cells (Rayleigh number=400)



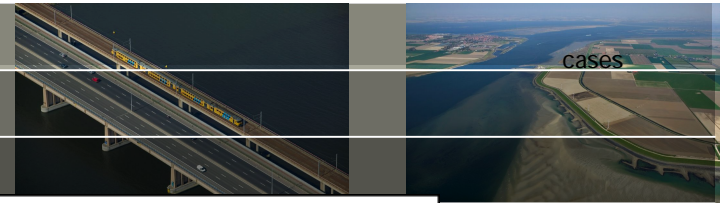
Elder problem (III)



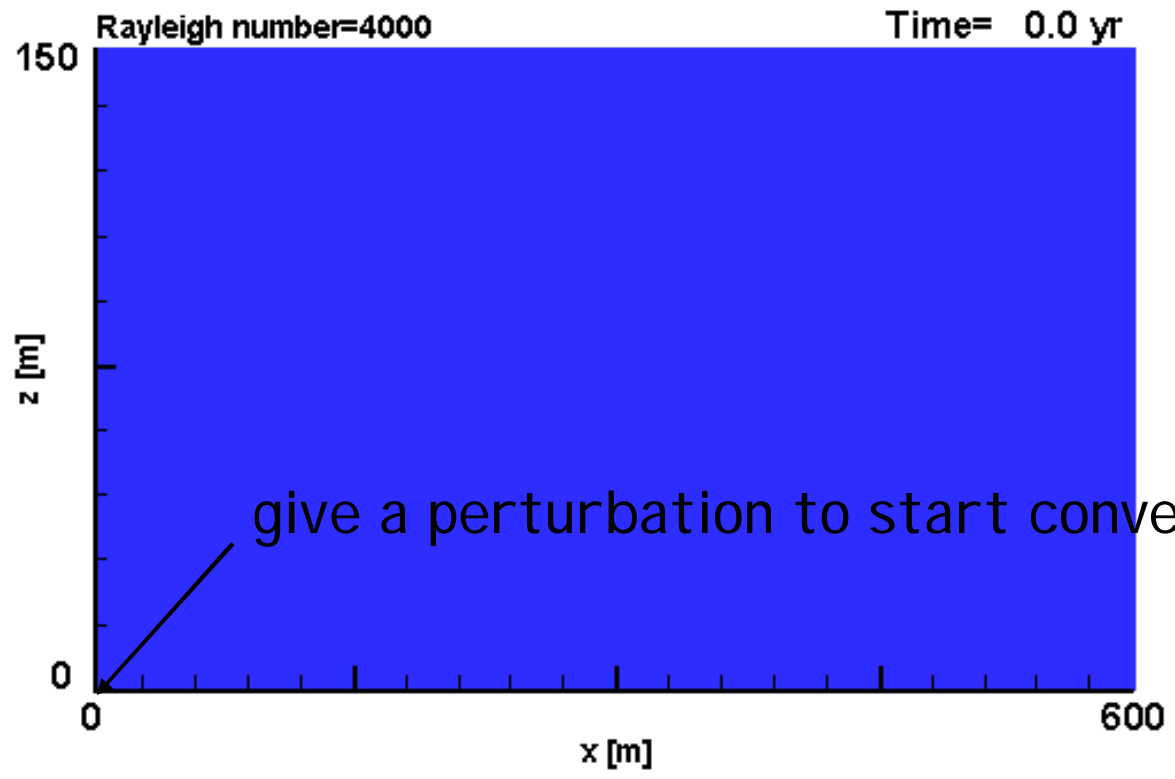
Development of convection cells (Rayleigh number=400)



Heat transport (Rayleigh number=4000)



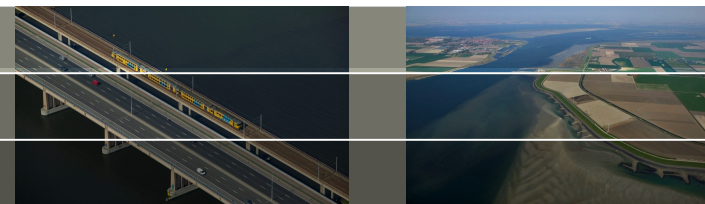
### Heat transport: conduction and convection



give a perturbation to start convection



# Conclusies



1. In kustgebieden kunnen dichtheidsverschillen een belangrijke rol spelen
2. Autonome proces heeft een grotere invloed op de verzilting van Nederland dan klimaatverandering
3. Het verschil tussen de gevolgen van de klimaatscenario's is klein voor het grondwater systeem
4. Regenwaterlenzen zijn kwetsbare systemen
5. Te weinig metingen kwaliteit in het kustgebied